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A STATISTICAL TROPICAL CYCLONE MOTION FORECASTING SYSTEM FOR THE GULF OF MEXICO

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CONTENTS

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ABSTRACT. The development of a statistical (climatology and persistence) model for the prediction of tropical cyclone motion over the southwestern Gulf of Mexico is described. The system complements an existing statistical model (CLIPER - CLImatology and PERsistence), which is currently in operation at the National Hurricane Center but which does not perform well over this region of the Gulf.

Because of the small size of the region and the high dissipation rate of cyclones making landfall in Mexico, problems relating to the changing statistical properties of the developmental data arise. Three different sets of developmental data and their advantages and disadvantages are discussed.

The final operational version of this system is based on a data set in which the tracks of tropical cyclones dissipating within 36-72 h are linearly extrapolated to overcome a bias against forecasts of westward motion present in CLIPER. The study concludes with examples of forecasts produced by the system.

1. INTRODUCTION

Despite the introduction of more sophisticated techniques, tropical cyclone motion prediction models based only on climatology and persistence are still widely used in the world's tropical cyclone basins. Particularly at short time periods of 36 h and less, persistence and climatology account for a large percentage of the total variance explained by any statistical tropical cyclone forecast model (Neumann and Lawrence 1975; i Neumann and Hope 1972). Climatology and persistence models are of two main types--analog and statistical. Analog models scan a data file containing all known tracks of tropical cyclones in a given basin, selecting as analogs those with motion, location, and date of occurrence similar to that of the cyclone to be forecast. These analogs are then combined to : produce the forecast track.

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Statistical models consist of pairs of regression equations for prediction of zonal and meridional displacements of the cyclone center for forecast intervals 12 through 72 h in 12-h increments. The first statistical climatology and persistence model, CLIPER (CLImatology-PERsistence), became operational in the Atlantic basin in 1972 (Neumann 1972). CLIPER predicts zonal and meridional displacements at 12-h intervals through 72 h, using as predictors the day number, latitude, longitude, current meridional and zonal motion, meridional and zonal motion 12 h before forecast time and maximum wind of the cyclone. Similar CLIPER-type models have been developed for the eastern North Pacific Ocean (Neumann and Leftwich 1977), the southwest Indian Ocean (Neumann and Randrianarison 1976), and the north Indian Ocean (Neumann and MandaI 1978).

Although CLIPER equals or outperforms synoptic and dynamic models at short ranges and at southern latitudes within the Atlantic basin, seven seasons of operational use have revealed some deficiencies as well. CLIPER forecasts for tropical cyclones in the Gulf of Mexico, especially those with initial motion to the west or northwest, tend to curve sharply to the right, giving the model a bias to the north. The work presented in this paper is a result of efforts to reduce or eliminate this bias and increase forecaster confidence in the model over the Gulf of Mexico.

The immediate cause of the bias is apparent from the map of mean tropical cyclone motion vectors in Figure 1. Over the large area of the Atlantic basin northwest of the dashed line, the mean motion vectors show a gradual veering from the northwest through northeast with increasing latitude. The "typical" track in this area is a broad parabola described as the cyclones recurve around the edge of the subtropical ridge into the westerlies. Southwest of the dashed line, in the southwest Gulf of Mexico and northwest Caribbean Sea, cyclone motion is best approximated by a straight line or even a slightly leftward-curving one, especially during July and August.

One solution to this bias against westward-moving forecast tracks in CLIPER is to redevelop the regression equations using a data set incorporating only those cases located in a region of straight or left-curving mean tracks. Unfortunately, the sample size available for the longer-range prediction equations decreases sharply as the cyclones move inland and dissipate. This dissipation is not random but, rather, is most likely for westward-moving cyclones carried into the mountains of Mexico. Cyclones with a more northerly track into the United States are usually carefully tracked and may persist for days over the relatively smooth terrain. This paper describes possible methods to compensate for this bias by manipulation of the developmental data. Three sets of developmental data and the performance characteristics of the resulting forecast equations presented are discussed.

2. DEVELOPMENTAL DATA CONSIDERATIONS

Three developmental data sets were considered for use in developing the CLIPER-GULF motion prediction equations for the western Gulf of Mexico, To obtain the stratification desired, all cases to the northeast of the dashed

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Figure 1. Mean motion vectors (resultant speeds and directions) for the period 1 May through 30 November for tropical cyclones with winds greater
than 34 knots. The dashed line is the boundary of the CLIPER-GULF stra-The dashed line is the boundary of the CLIPER-GULF stratification region.

line in Figure 1 were immediately rejected. The different developmental data sets were then specified as subsets of this restricted set of cases from
1886-1979. In addition, 10 storms, selected at random, were removed for In addition, 10 storms, selected at random, were removed for use as independent data. The homogeneous (HOM) data set uses only those cases for which at least 72 h of future displacements exist. Because motion 12 h before forecast time is included as a predictor in CLIPER and was to be included in CLIPER-GULF as well, any tropical cyclone with less than 84 h (3 1/2 days) of best-track positions is automatically excluded, as are cases in which dissipation occurs within 72 h. The non-homogeneous (NHOM) data
set uses all available cases at each prediction time. 12 through 72 h. In set uses all available cases at each prediction time, 12 through 72 h. this set (made up of 6 subsets, corresponding to the 6 forecast times), the number of cases and the statistical attributes of the predictors and predictands change with increasing forecast time. The third data set, extrapolated non-homogeneous (ENH), consists of the NHOM set for the 12,24 and 36 h forecast equations, with linearly extrapolated displacements through 72 h
for qualones which persist longer than 36 but less than 72 h. These three for cyclones which persist longer than 36 but less than 72 h. developmental data sets are discussed in more detail below.

The characteristics of the homogeneous (HOM) data set are shown in Table A homogeneous data set is used in CLIPER and in the statistical clima-1. tology-persistence model for the southwest Indian Ocean (Neumann and Randrianarison 1976). The advantages of using a homogeneous data set are: (1) the sample means and standard deviations are constant for all predictor
and predictands, resulting in smooth and "reasonable" forecast tracks, and (2) the constant statistical properties of the developmental data render the regression and screening process more efficient since the covariance; matrix need only be computed once. The main disadvantage of the HOM data set is an inherent bias towards the longer duration tropical cyclones. This bias becomes critical over a small, landlocked area such as the Gulf of Mexico because of the high number of dissipating cases and short tracks. The sample size for the HOM data set is reduced by about 40 percent when the 72 h displacement restriction is imposed. Cyclones most likely to be included in the sample are those moving slowly or on a northerly course and failing to strike the mountains of Mexico where dissipation is rapid and almost certain. Figure 2 shows the total dissipation for 2.5 degrees latitude-longitude areas over the Gulf of Mexico region from best-track data, 1886-1979. ~

Table 1. Characteristics of the homogeneous (HOM) developmental data set for the CLIPER-GULF stratification region. All velocities and displacements are given in kt and nmi respectively, with north and east considered positive. , ." ..

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Non-Homogeneous Data Set **B.**

The characteristics of the non-homogeneous (NHOM) data set are given in able 2. As stated earlier, this set is actually 6 sets, each composed of ases with displacements through the corresponding forecast time, $12-72$ h. ion-homogeneous data sets are used for CLIPER-type systems for the eastern orth Pacific (Neumann and Leftwich 1977) and north Indian Ocean (Neumann ind Mandal 1978) tropical cyclone basins. Note that, with increasing orecast time, the mean motion becomes slower and more northerly and the ean initial location shifts southeastward. The primary advantage of a HOM data set is optimized "point verification" (minimum root mean square gror) and zero bias for the developmental data as all possible cases are sed.

able 2. Same as Table 1 except for the non-homogeneous (NHOM) developmental ata set for CLIPER-GULF.

					MEANS FOR FORECAST PERIOD OF LENGTH		STANDARD DEVIATIONS FOR FORECAST PERIOD OF							
QUANTITY	SYMBOL	12 _b	24 _b	36 _h	48 h	60 h	72 _h	$\frac{12}{43.4}$	$\frac{24}{43}$.8	$\frac{36}{44}$. 2	48 h	60 _h	72 _b	
Day Number	P1	248.3	248.5	248.6	248.7	248.9	249.1				44.8	45.5	46.4	
Initial Latitude	P ₂	21.5	21.3	21.2	20.9	20.6	20.3	4.7	4.6	4.6	4.5	4.4	4.3	
Initial Longitude	P3	89.2	88.7	88.3	37.9	87.5	87.2	5.3	5.1	5.0	4.9	4.9	4.9	
Current Meridional Speed	P4	4.0	4.2	4.3	4.4	4.5	4.6	4.1	4.0	3.9	3.8	3.6	3.5	
Current Zonal Speed	P5	-4.6	-4.4	-4.2	-3.9	-3.7	-3.5	5.1	5.0	5.0	4.9	4.7	4.6	
12 h Past Meridional Speed	26	3.8	3.9	4.0	4.1	4.2	4.2	3.8	3.7	3.6	3.6	3.5	3.4	
'12 h Past Zonal Speed	P7	-5.2	-5.0	-4.8	-4.6	-4.4	-4.2	4.9	4.9	4.9	4.8	4.7	4.6	
Maximum Wind	P36	57.7	58.7	59.0	58.8	58.2	57.3	24.1	24.4	24.6	24.5	24.2	23.7	
12 h Meridional Displacement	DY12	48.6	51.5	53.9	55.1	56.1	56.7	52.7	50.0	48.2	46.4	44.2	42.5	
24 h Meridional Displacement	DY24	------	105.3	110.7	113.9	116.0	117.3	-----	102.0	97.0	92.4	87.3	83.1	
36 h Meridional Displacement	DY36	-----	-----	169.7	175.5	179.3	181.3	-----	-----	148.0	139.0	130.2	122.9	
48 h Meridional Displacement	DY48	-----		-----	239.2	245.1	248.2	-----	-----	------	188.8	173.3	162.4	
60 h Meridional Displacement	DY60				-----	313.6	318.0	-----			مستنسب	219.2	202.8	
72 h Meridional Displacement	DY72	-----				-----	391.0					------	247.0	
12 h Zonal Displacement	DX12	-50.5	-47.9	-45.1	-42.1	-39.4	-36.4	63.5	62.4	61.6	60.1	57.9	55.9	
24 h Zonal Displacement	DX24	-----	-85.6	-79.8	-73.8	-68.1	-62.1	-----	128.2	125.7	122.3	117.5	113.1	
36 h Zonal Displacement	DX36	------	------	-103.5	-94.6	-86.0	-76.5	-----	-----	195.8	188.9	180.6	173.4	
48 h Zonal Displacement	DX48	-----			-104.4	-92.6	-80.0	------	------	-----	262.4	249.2	238.7	
60 h Zonal Displacement	DX60	-----			-----	-88.0	-71.9	-----			-----	324.3	308.9	
72 h Zonal Displacement	DX72					-----	-53.5					-----	386.5	
Mumber of Cases	N	3246	2940	2638	2337	2048	1797							

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Difficulty arises, however, when a NHOM set is used in a basin such as the Gulf of Mexico where many dissipations occur and the statistical attributes of the sample vary with changing forecast time. As the forecast period lengthens, the bias resulting from dissipations becomes more pronounced and, at 72 h, the NHOM and HOM sets are identical. The result is an acceleration in the forecast tracks in the direction of the mean track favored by the dissipation-induced bias, which, in the Gulf of Mexico, is a northward-moving track. In extreme cases, the forecast track may actually double back on itself. Figure 3 shows two forecasts for Hurricane Anita of 1977, using HOM and NHOM-based equations. Although the NHOM track actually verifies better, the spurious "kink" in the track would likely result in the track being disregarded by the operational forecaster. The necessity of providing the forecaster with a "believable" track is discussed by Neumann (1972) and Neumann and Mandal (1978) in connection with retention of insignificant predictors. It is felt that the same'reasoning also applies here.

Figure 3. CLIPER-GULF forecasts for Hurricane Anita of 1977 using equations based on HOM (triangles) and NHOM (squares) developmental data sets. The circles are best-track positions at 12 h intervals.

c Extrapolated Non-Homogeneous Data Set

The extrapolated' non-homogeneous (ENH) data set represents an effort to retain the optimized verification properties of the NHOM data, particularly at short range, while at the same time reducing the dissipation-induced bias and providing a realistic forecast track. For the first 36 h, this set consists of the NHOM data. Beyond 36 h, the observed displacements are used if available. If not, the remaining displacements at 48, 60 and 72 h are computed by linear extrapolation of the cyclone motion during the 12 h preceding the final position available. An example of this linear track extrapolation is shown in Figure 4, and the statistical properties of the ENH data set are presented in Table 3.

The justification for the use of these linearly extrapolated tracks is based on two main premises. First, tropical cyclone best tracks are sometimes terminated due to lack of either data or operational need even though an identifiable circulation exists. Such situations arise over mainland China (Jarrell and Wagoner 1973) and along the Mexican Gulf coast, although

Figure 4. Example of linear track extrapolation procedure used in preparing The circles are available best-track positions and the squares ENH data set. and data set. The circles are available best-tiack positions and the straps
are extrapolated positions computed using the observed displacement from
36 to 48 h.
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6 to 48 h.
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Gables One Tower ~. 1320 South Dixie Highway. Room 520 .c ., Coral Gabies. Florida 33145 not nearly as often in the latter area in recent years. One possible, but very time-consuming, solution is a review of the records in an attempt to extend the best tracks for some of the cases. Another justification for track extrapolation arises because statistical models are inherently unable to predict dissipation. A discrete region with a high incidence of dissipations is statistically interpreted as an area into which tropical cyclones; ះនៅ seldom move. What is needed for forecasting cyclone motion is an estimate \mathbb{R} of the steering influences carrying the cyclone into such an area--the $\frac{1}{2}$ question of whether or not the cyclone survived the journey is irrelevant as far as motion prediction is concerned. Track extrapolation is an effort $\frac{1}{2}$ to represent the steering influence predominating in an individual case, $\frac{1}{2}$ even though the cyclone itself may have dissipated. Linear extrapolation \rightarrow is adopted as a quick "first guess" and is also justifiable in that the mean motion field for the western Gulf of Mexico is approximately linear.

The advantages of such a data set are optimized point verification at the $\frac{1}{2}$ s riti s al 12, 24 and 36 h f s re s ast times, and a smooth, representative forecast track beyond 36 h. The main disadvantage is that the system will have higher average forecast errors at 48-72 h than HOM or NHOM equations because many of its "better" forecasts (westward into a high-dissipation area) will not be verified.

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3. SYSTEM, DEVELOPMENT .

Using the Atlantic basin tropical cyclone best-track data file (Jarvinen and Caso 1978), seven basic predictors (see Table 4) were read or computed for each tropical cyclone position within the CLIPER-GULF stratification region. In addition to these predictors, which have become almost standard for CLIPER-type models (Neumann and Manda1 1978; Neumann and Randrianarison 1976; and Neumann and Leftwich 1977), wind speed of the cyclone was added as a supplementary predictor. Means and standard deviations of these predictors for the ENH data set are listed in Table 3. In order to partially account for nonlinear effects (Neumann and Mandal 1978), predictors 9

Table 3. Same as Table 1 except for the extrapolated non-homogeneous (ENH) developmental data set for CLIPER-GULF.

through 35 are defined as the products and cross-products of the basic predictors. Predictands are zonal and meridional displacements in nautical miles, either computed from the best-track positions or by linear extrapolation.

I Following past experience with such models (Neumann and Randrianarison 1976), it was elected to fit least-squares polynomials to all the predictors, rather than using only selected predictors obtained through a step wise screening method. The 37 normal equations for the 36 predictor co efficients and the intercept for each forecast equation were solved using the stepwise screening program available at NHC, with the minimum acceptable reduction of variance set at zero. The values of the coefficients for the l2 prediction equations based on the ENH developmental data set are given in the Appendix.

Tables 5 and 6 show the reduction of variance contributed by specific predictors and the total reduction of variance of each regression equation. Only those predictors contributing at least 0.5 percent or more are listed, Jlthough all predictors are included in the forecast equations. Most of the variance is, quite reasonably, explained by a term involving initial motion (P4 or P5) although, unlike all other CLIPER-class models inspected, a non-Linear term provides the highest reduction of variance at the 48, 60 and 72 h eridional displacements. Also note that the linear term failed to contri-;)ute-significantly at these times because of a high correlation between P4 and 'P4xP3.

PERFORMANCE CHARACTERISTICS

Preliminary testing of the models based on each of the three data sets was begun by verifying against the NHOM developmental data set. Performance characteristics for the full basin CLIPER (Neumann 1972), HOM CLIPER-NHOM CLIPER-GULF and ENH CLIPER-GULF are shown in Table 7. The dissipation bias associated with system development also occurs in the verification data and, thus, must be taken into account when evaluating

Table 4. The basic and supplementary predictors for CLIPER-GULF, the symbols representing them in this paper and their units.

Table 5. Percent reduction of variance contributed by specific predictors to the CLIPER-GULF meridional displacement equations for the ENH data The predictors listed contributed 0.5 percent or more. Totals in parentheses are the percentages of the total variance reduced at each forecast time by the final forecast equations which include all 36 predictors.

As discussed earlier, NHOM. CLIPER-GULF has the best verification (minimum average forecast error), with HOM, ENH and CLIPER having progresgively greater average errors. ENH and NHOM are, of course, identical at 12, 24 and 36 h. NHOM also possesses the lowest standard deviation of forecast error. Biases are somewhat differently distributed than was expected, with CLIPER actually showing a southward bias at 60 and 72 h. NHOM shows expected near-zero biases, as the verification sample is identical to the NHOM developmental data. The southwest bias of ENH forecasts results the dissipation of some of the westward-moving tropical cyclones in the verification sample.

The four equation sets were also verified against a 10-storm, randomly selected, independent data set with results shown in Table 8. Mean forecast errors are higher than for the developmental data beyond 24 h, but the four models can still be ranked in the same order (NHOM, HOM, ENH and CLIPER) the basis of mean forecast error. The full basin CLIPER here reveals a pronounced bias to the northeast, and all but ENH have a bias to the north. observed in developmental data testing, ENH again shows a bias to the West at long forecast periods.

Because of the reliance of these models on persistence for most of their : reduction, use of operational data rather than best-track or postanalysis data inevitably degrades system performance. Operational input for CLIPER has been archived since the beginning of the 1972 season, and was used to verify the four equation sets. The forecast errors are computed

Table 7. Comparative performance of full basin CLIPER and HOM, NHOM and ENH CLIPER-GULF on dependent data. All distances are given in nmi, with north and east considered positive.

Table 8. Same as Table 7 except for a 10-storm independent data sample.

Table 9. Same as Table 7 except for operational data, 1972-1979.

TIME	CASES	Full	MEAN ERROR BOH	KBOH	EXH			STANDARD DEVIATION Full HOM NHOM	ЮH			MORTHWARD BIAS Full HOM NHOM	m			EASTWARD BIAS Full BOM NHOM	œ	实体	
12	175		48 46	-47	- 47	30	- 29	30	- 30		-1 -2		$-3 - -3$	-2	-2	-6	-6	27.93 ل المي	
24	163	101	96	99	- 99	60	59	61	61	$\mathbf{1}$		-2 -7 -7		\mathbf{z}		$0 -13 -13$		$\pm 10^{\circ}$	
36	144		156 147	153	- 153	90	- 89		93 93	8.	\mathbf{z}		$-7 - -7$	5.		-2 -25 -25		. 기업 원 - 1	
48	126		226 211		215 222		124 123		126 130	20	- 11	4.	-5	14.	-4		$-29 - 41$	بالمناد \sim \sim	
72	96	370	- 331		331 365		188 205		205 228	66	43	43	- 12			$28 - 21 - 21 - 103$			

by preparing forecasts from the operational data, and then vectorially removing the initial positioning error. The result is a forecast displacement error rather than a forecast position error. Forecast displacement error is the standard measure of the accuracy of a tropical cyclone motion forecast used at NHC (Neumann and Pelissier 1980). These displacement errors are listed in Table 9.

As these comparative tests showed no serious reduction of accuracy by the track extrapolation technique, even at long range, and showed a reduction of the northward bias by the ENH equations, they were adopted as the forecast) equations for the operational CLIPER-GULF. The average errors of the ENH equations on developmental data and two sets of operational data are depicted graphically in Figure 5. The second operational data curve, that for

cast errors for the operational CLIPER-GULF based on the ENH data

tropical cyclones of hurricane strength (winds greater than 64 kt), illustrates the improvement of accuracy associated with the better position and motion estimates available for these more organized tropical cyclones. A comparison of the biases of CLIPER and CLIPER-GULF on operational data is shown in Figure 6. The westward bias of CLIPER-GULF is due, in part, to the fact that forecasts of westward-moving cyclones, on which CLIPER-GULF is designed to perform well. are often not verified because of cyclone dissipations.

5. 5. OPERATIONAL CONFIGURATION

ENH CLIPER-GULF has been incorporated into the NHC operational statistical guidance package as a subroutine accessed by the CLIPER model. CLIPER forecasts which serve as input to other models and are given directly to the forecaster now consist of a weighted average of CLIPER and CLIPER-GULF. The weighting function is described in Figure 7. CLIPER-GULF will be activated whenever a cyclone is positioned to the west of the 0.00 line; both individual and composite forecasts will then be included in a supplementary computer message.

 \mathcal{F} igure 6. Comparison of forecast biases for the full basin CLIPER (squares) :,~: and the operational CLIPER-GULF (circles) for operational data, 1972-1979. The outer circle is 100 nmi.

Figure 7. Weighting function for composite CLIPER forecast. The numbered contours give the weight assigned to the CLIPER-GULF system for a tropical cyclone located at that point.

Figures 8-11 illustrate some of the characteristics of the CLIPER-GULF system. Figure 8 depicts the forecast tracks resulting as day number is allowed to vary and other predictors are held constant. The maximum westward motion is to be expected in July and August, and the track shifts eastward in late season, reflecting the increased dominance of the midlatitude westerlies at lower latitudes. Figure 9 shows the effect of an increase in wind from 45 to 120 kt-~the stronger cyclone is forecast to move slightly faster. Also shown is a forecast for a stationary cyclone with 65 kt

Figure 8. Effect of variation of date on CLIPER-GULF forecast tracks. Tracks were run for the 15th day of the indicated month.

Effect of a variation in wind speed on CLIPER-GULF forecast *Figure 9.* tracks. Forecasts were run for the August cyclone in Figure 8 with winds of 45 and 120 kt. Also shown is the forecast track for a stationary cyclone with 65 kt winds on September 1.

 10^{10} winds on September 1, Figure 10 shows two forecast examples using 10^{10} nal data from the 1975 season. Hurricane Caroline moving weight tional data from the 1975 season. Hurricane Caroline, moving west-
northwestward at nearly constant speed, was handled well, while the accelerating motion of Hurricane Eloise resulted in a 72 h forecast error of over 1000 nmi. Systems such as CLIPER-GULF which rely only on persistence and climatology fail in situations of anomalous or rapidly changing synoptic conditions. The reduction of northward bias and the improved performance of CLIPER-GULF versus CLIPER for westward-moving cyclones are shown by the comparison forecasts for Hurricane Anita of 1977 (Figure 11).

Figure 10. Two forecasts using CLIPER-GULF based on operational data from the 1975 season. Hurricane Caroline (west Gulf) was well forecast, while Hurricane Eloise (central Gulf) was handled poorly.

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6. SUMMARY

Inis paper has described the development of a scatistical (climatology and , nersistence) tropical cyclone motion prediction method for use in the western Gulf of Mexico and northwest Caribbean Sea. The original objective was elim-
ination of the northward bias in the present CLIPER system over the region, but, during the course of this study, the importance of cyclone dissipation, the developmental data became apparent and methodology was developed within the developmental data became apparent and methodology was developed
to "tune" the model to obtain the best performance by using different developmental data sets. The resulting equation set, based on linearly extrapolated non-homogeneous tracks, is the first statistical cyclone motion $ext{repolved non-momogeneous tracts, is the result of the sum of fields of all matrices.}$ prediction scheme in which the developmental data were artificially modified. in an effort to correct for biases introduced by cyclone dissipation. The
CLIPER-GULF system will be operationally tested at NHC and it is hoped that CLIPER-GULF system will be operationally tested at NHC and it is hoped that similar approaches will be applied to other basins and additional documentation of and experience with this technique obtained. ;,: I';

Figure 11. Comparison of full basin CLIPER and operational CLIPER-GULF forecast tracks based on operational data for Hurricane Anita, 1977.

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REFERENCES

- Jarrell, J. D., and R. A. Wagoner, 1973: The 1972 typhoon analog program (TYFOON 72). U.S. Naval Environmental Prediction Research Facility Technical Paper 1-73.35 pp.
- Jarvinen, B. R., and E. L. Caso, 1978: A tropical cyclone data tape for the North Atlantic basin, 1886-1977: contents, limitations, and uses. NOAA Technical Memorandum NWS NHC-6, 19 pp.
- Neumann. C. J., 1972: An alternate to the HURRAN (Hurricane Analog) tropical cyclone forecast system. NOAA Technical Memorandum NWS SR-62, 25 pp.
- Neumann, C. J., and J. R. Hope, 1973: A diagnostic study on the statistical predictability of tropical cyclone motion. J. Appl. Meteor., 12, 62-73.
- Neumann. C. J., and M. B. Lawrence, 1975: An operational experiment in the statistical-dynamical prediction of tropical cyclone motion. Mon. Wea. Rev., 103, 665-673.
- Neumann. C. J., and P. W. Leftwich. 1977: Statistical guidance for the prediction of eastern North Pacific tropical cyclone motion - Part I. NOAA Technical Memorandum NWS WR-124, 32 pp.
- Neumann. C. J., and G. S. MandaI, 1978: Statistical prediction of tropical storm motion over the Bay of Bengal and Arabian Sea. Indian J. Met. Hydrol. Geophys., 29, 487-500.
- Neumann, C. J., and J. M. Pelissier, 1980: Models for the prediction of tropical cyclone motion over the North Atlantic: an operational evaluation. Submitted to Mon. Wea. Rev., 29 pp.
- Neumann, C. J., and E. A. Randrianarison, 1976: Statistical Prediction of tropical cyclone motion over the southwest Indian Ocean. Mon. Wea. Rev., 104, 76-85.

APPENDIX I, TABLES OF COEFFICIENTS FOR PREDICTION EQUATIONS

The coefficients for the meridional and zonal displacement prediction equations, respectively, are listed on the following two pages. The predictors and their units are specified in Table 4 of the text (p. 9). The predictands are meridional and zonal displacements in nautical miles, with north and east considered positive.

The prediction equations are of the form:

$$
D = I + \sum_{i=1}^{36} C_i P_i,
$$

where C_i is the coefficient associated with predictor P_i and I is the intercept for a given displacement D.

COEFFICIENTS FOR MERIDIONAL DISPLACEMENTS

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