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# A STATISTICAL MODEL FOR THE PREDICTION OF WESTERN NORTH PACIFIC TROPICAL CYCLONE MOTION (WPCLPR)

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## **CONTENTS**

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# A STATISTICAL MODEL FOR THE PREDICTION OF WESTERN NORTH<br>PACIFIC TROPICAL CYCLONE MOTION (WPCLPR) PACIFIC TROPICAL CYCLONE MOTION (WPCLPR)

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# ABSTRACT

The derivation, implementation and operational utility of a<br>new statistical model for the prediction of western North Pacific tropical cyclone motion is described. The model uses regression equations to forecast tropical cyclone motion through 72h and incorporates predictors derived from climatology, persistence, and storm intensity. It is patterned after models that were developed for most of the other tropical cyclone basins. In addition to its for most of the other tropical cyclone basins. In addition to it usefulness for operational prediction, one model provides a con-<br>vaniant thrachold skill laval for avaluating the nerformance of venient threshold skill level for evaluating the perfonmance of other, more sophisticated models.

Developmental data consisted of western Pacific tropical<br>cyclone tracks and associated storm intensities for 1946 through 1980. The model was tested on independent data for 1981 and 1982 1980. The model was tested on independent data for 1981 and 1982 and on operational data for 1983 and 1984.

#### 1. INTRODUCTION

This report documents a recently developed statistical model (WPCLPR) for the prediction of western North Pacific (WESPAC) tropical cyclone motion. The the prediction of western north racific (westing tropical cyclone motion. The need prediction scheme is based on a series of regression equations. The preyear), persistence (average storm motion over the past 12 and 24h) and storm intensity (maximum sustained surface wind). Predictors derived from analyzed '. intensity (maximum sustained surface wind). Predictors derived from analyzed fields of environmental data (winds or geopotential heights) have explicit been omitted. Predictands are the meridional (north/south) and zonal (east/west) components of tropical cyclone motion in 12-h increments through 72h.

l Research accomplished while on temporary assignment to the National Hurricane<br>Center. Center.



Figure 1. Tracks of the 873 western North Pacific tropical storms and typhoons, 1946-1980. These storms were used as dependent data.

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This type of model, commonly referred to as a "CLIPER-class" model, has<br>been used for several years in other basins and is well-documented in the literature. References to the other basins include: Neumann (1972) for the Atlantic; Neumann and Randrianarison (1976) for the southwest Indian Ocean: Neumann and Leftwich (1977) for the eastern North Pacific; and Neumann and Mandal (1978) for the North Indian Basin. Because of this rather extensive documentation only those aspects of the model unique to WESDAC are described ... documentation, only those aspects of the model unique to MESPAC are described

#### 2. DEVELOPMENTAL DATA 2. DEVELOPMENTAL DATA

# 2.1 Historical Storm Tracks

Developmental data consist of the best tracks<sup>2</sup> of all recorded western<br>North Pacific tropical cyclones over the 35-y period 1946-1980. This data set (through 1975) originally had been obtained from the NOAA National Climatic Center, Asheville, North Carolina (tropical cyclone deck 993). Included were storm positions for every 12h and maximum winds for most storms. This original data set was extensively supplemented by storm positions and maximum winds at 6-hourly intervals as obtained from WESPAC storm summaries that are published annually by the Joint Typhoon Warning Center on Guam (for example, Annual Tropical Cyclone Report, 1984). Also, some missing storm intensities for the earlier years were obtained from records maintained by the People's Republic of China (Central Meteorological Bureau, 1972). The final data set. beginning in 1946, consists of storm positions and intensities at 6-hourly intervals. Through 1980, 873 storms are documented; these are depicted in Figure 1. The latter plot of storm tracks led to spatial bounds of the model being set at  $5^{\circ}$ -35°N latitude and west of  $150^{\circ}$ F longitude being set at 5°-35°N latitude and west of 150oE longitude.

In the temporal sense, cases were excluded if they occurred before 15 May<br>or after 15 December. As shown in Figure 2, this 8-month period comprises the or after 15 December. As shown in Figure 2, this 8-month period comprises the<br>bulk of the WESPAC season . Activating the program outside of these spatial bulk of the NESPAC season. Activating the program outside of these spatial to run the model (appendix) disallows running the program outside of these to run the model (appendix) disallows running the program outside of these temporal bounds of if a storm is influently beyond 35°N latitude. The  $\leq$  34 kt. Storms in existence for  $\leq$  36h are also inherently excluded from the developmental data set in that there is a requirement for past positions through at least  $-24h$  and a future storm nosition through at least  $+12h$ through at least -24h and a future storm position through at least +12h.

Storms that occurred in 1981 and 1982 were reserved for testing of the model in an independent data mode and the model, developed early in 1983, was subsequently tested in an operational mode for 1983 and 1984. Storms that subsequently tested in an operational mode for 1983 and 1984. Soon model in an operational model in a occurred during these latter 2-y periods are shown in Figures 3 and 4.

The 1946-1980 developmental data set is large enough (5,410 cases at 12h to 2,788 cases at 72h) that, even allowing for lost degrees of freedom through to 2,788 cases at 72h) that, even allowing for lost degrees of freedom, entough serial correlation, the classical significance testing exercise could probably

<sup>&</sup>lt;sup>2</sup>The best track is the accepted track of a storm after a post-analysis of all available data.



Figure 2. Daily frequency of typhoons (shaded area) and tropical storms and typhoons combined (nonshaded area) per 100 years based on the 39 year period 1946-1984. Data have been smoothed over 9-day period. Mean number of days per year with tropical storms or period. Mean number of days per year with tropical scorms of<br>typhoons is 149.5. Mean number of days per year with typhoons alone is 79.9.

4



Figure 4. Tracks of the 50 western North Pacific tropical storms and typhoons, 1983-1984. These storms were used in operational testing of program.

have been omitted and the 1981 and 1982 storms profitably could have been added to the developmental data. This option was considered, but not adopted.

#### 2.2 Definition of Predictors/Predictands

.From the basic developmental data set, 8 first-order predictors can be defined. These are: initial storm latitude, initial storm longitude, time of year (Julian day number), average meridional translational speed over past<br>12h, average zonal speed over past 12h, average meridional storm translational speed over past 24h, average zonal storm translational speed over past 24h and initial storm intensity. The assumption is made that each of the orthogonal components of projected motion  $(Y_+)$  is a function of these 8 predictors,

$$
Y_{t} = f(P_{1}, P_{2}, P_{3}, P_{4}, P_{5}, P_{6}, P_{7}, P_{8}). \qquad (1)
$$

When we developed CLIPER-class models for other basins, the above function was taken as a second- or third-order polynomial, with the order being determined by the size of the developmental data set and the geometri complexity of the basin. The very large data set available here and the parabolic nature of the tracks over WESPAC justify the use of a third-order polynomial. The number of possible predictors (excluding the "intercept" value) in the polynomial expansion of (1) is given by

 $T = (m+n)!/(m!n!) - 1,$  (2)

where n is the order of the polynomial and m is the number of basic predictors. From (2), it follows that the third-order polynomial, including the intercept value, will contain 165 terms. Accordingly, a master data file was structured, and contained, for each case, the 12 predictands (storm meridional and zonal motion displacements for 12, 24, 36, 48, 60 and 72h) and the 164 potential predictors. The additional predictors, 9 through 164, can be generated by considering all possible products and cross products of the 8 basic predictors. These are identified in the FORTRAN program listing beginning on page 22. The predictor indexing, however, is somewhat different in the program from that just described.

### 3. PREDICTOR SELECTION

Experience from the development of other CLIPER-class models led to a modified procedure to determine which of the 164 potential predictors were to be retained in the final prediction equations. Typically, predictors are systematically selected until the incremental variance reduction drops below some preset value, often taken as 1 or 1/2%. The problem with this classical approach in the development of CLIPER-class models is that some predictors, which may be working in combination (as is often the case in nth-order polynomials), may be overlooked in the screening process. Another, even more serious, problem is that predictor selection from one period to another is done independently. This gives rise to the generation of meandering tracks that impart a certain degree of skepticism to the forecast.

To alleviate these problems, 20 "best" predictors were selected for each<br>of the 12 regression equations (meridional and zonal components for each of six forecast periods). Trial-and-error screening runs suggested that this retention of 20 predictors was about optimal in assuring that all predictors acting in combination were selected. There were some differences here, depending upon projection or component, but, in the interest of simplicity, these differences were ignored. In this connection, the large sample size guarantees that if worthless predictors are included in the program, the partial correlations and, thus, the regression coefficients, will be peak partners and, the regressions and, the regression . Thus, will be nearly be nearly be nearly be nearly be near

Next, we searched for predictors that were used at least once for any of<br>the six meridional time periods, 12 through 72h. As a result, we obtained 32 of the 164 possible predictors. This sorting was also carried out for zonal motion and, coincidentally, 32 predictors (not necessarily the same ones) were  $\mathbb{R}$  identified. To avoid the meandering track problem referred to earlier, the .: identified. To avoid the meandering track problem referred to earlier, the program was structured about the 32 predictors. The 32 predictors is a 32 predictors.

 $T_{\rm eff}$  for the prediction equations is:  $T_{\rm eff}$  the prediction equations is:

$$
D = C_0 + \sum_{i=1}^{i=32} C_i P_i,
$$
 (3)

given period,  $C_0$  is the intercept value and  $C_3$  is the 32 regression coefficients corresponding to the 32 predictors  $P_i$  for that given forecas period and orthogonal component. period and orthogonal component.

the data cards following the FORTRAN program listing given in appendix (beginning on page  $28)$ . The predictand/predictor numbering conven  $\frac{1}{2}$  and  $\frac{1}{2}$  . The program is: The predictandlpredictor numbering convention  $\frac{1}{2}$ tion in the program is:

 $\frac{1}{2}$  miles (predictands) for each of the six projections. 12 through 72h. miles (predictands) for each of the six projections, 12 through 72h.

'- P3 is the initial storm latitude.

 $P_A$  is the initial storm longitude.

- $P<sub>5</sub>$  is the current Julian day number.
- $P_6$  is the average meridional speed (knots) over the past 12h.<sup>3</sup>

<sup>&</sup>lt;sup>3</sup>It was intended that  $P_6$  and  $P_7$  be in knots. However, through oversight, the equations were derived using  $1/2$  of this amount. Compensation for this oversight is included in the program definition of  $P_c$  and  $P_7$  and is transparen  $\frac{1}{2}$  is included in the program definition of  $P$  and  $P$ 

 $P<sub>7</sub>$  is the average zonal speed (knots) over the past 12h.

 $P<sub>A</sub>$  is the average meridional speed (knots) over the past 24h.

 $P<sub>q</sub>$  is the average zonal speed (knots) over the past 24h.

 $P_{10}$  is the storm intensity in knots.

P<sub>11</sub> through P<sub>166</sub> are additional predictors generated by the cubic products and cross products of P<sub>3</sub> through P<sub>10</sub>.

It can be noted in the data cards that specify the predictors and regression coefficients that there are 12 nine-card sets of 32 predictor numbers and. associated regression coefficients, each preceded by an intercept value.<br>These 12 sets are in the order 12h meridional, 12h zonal, 24h meridional...72h zonal. For example, the intercept value for 12h meridional motion is 82.43, while the first predictor is number 29 and the associated regression coefficient is 0.1673843. As noted on page 25, predictor number 29 is defined as the product of  $P_4$  and  $P_6$  or the product of initial storm longitude and average meridional speed over the past 12h. These predictor/ regression coefficient sets are listed in the order that they were selected in the screenin program. In the example under discussion, subsequent predictor numbers are 141,154, 113, 133, etc.

For each of the 12 prediction equations, the first and most important predictor turned out to be a function of average motion over the past 12h. This characteristic points out the importance of the persistence factor in the prediction scheme and, as discussed in section 6, great care must be exercised in detenming this motion.

### 4. PERFORMANCE ON DEPENDENT, INDEPENDENT, AND OPERATIONAL DATA

Tables 1 and 2 depict, respectively, the performance of the model on dependent and independent data. The dependent data forecast errors are somewhat greater for the short-term projections and somewhat less for the long-term projections than for the Atlantic counterpart of the model (Neumann, 1972). Comparison with still other basins shows that the WESPAC dependent data errors are higher for all periods. The explanation here is probably related to the degrees of forecast difficulty one encounters in going from one basin to another or to parts of the same basin. The concept is discussed by Pike (1985).

Comparison of Table 1 with Table 2 shows, for the most part, that the model performed better on the 2-y independent sample than on the 35-y developmental data set. Typically, the reverse is true. For example, in structuring a ClIPER-class model for the southwest Indian Ocean, Neumann and Randrianarison (1976) found about a 20% increase in forecast error when running the model on an independent sample. The explanation probably lies partially in that the data set used in developing WPCLPR was unusuall large. Also, the sample of storms used to test the model for 1981 and 1982 (Figure 3) showed more adherence than normal to persistence and climatology.

8

Forecast period (hours)	Component	Sample size	Multiple corr. coef.	Standard error	Forecast error
12	Meridional Zonal	5410	0.92 0.83	40.6 (78.9) 37.3 (72.5)	44.0 (85.5)
24	Meridional Zonal	4894	0.90 0.78	88.8 (172.5) 80.5(156.4)	97.5(189.4)
36	Meridional Zonal	4342	0.87 0.72	(280.5) 144.4 127.2 (247.1)	157.7(306.3)
48	Meridional Zona <sub>1</sub>	3784	0.83 0.65	(399.2) 205.5 172.1 (334.3)	219.7(426.8)
60	Meridional Zonal	3276	0.80 0.60	267.7 (520.0) 210.7 (409.3)	278.1(540.2)
72	Meridional Zonal	2788	0.76 0.56	(637.5) 328.2 244.9 (475.7)	334.9(650.6)

Table 1. Performance of the model on best-track independent data. Period of record is  $1946-1980$ . Errors are in n.mi. (km).

Table 2. Performance of the model on best-track independent data. Period of record is 1981-1982. EFTURS are in n.mil. (Km).



Regardless of a model's performance on dependent or independent data, it must be tested on operational data where marked degradation over dependent or even independent data is not unusual. In the latter modes, initial input data is derived from the best track of the storm, whereas in an operational mode, a best-track scale of motion can only be estimated from warning time positions. As is noted in section 5, the model is particularly sensitive to uncertainties in the specification of the average motion over the past 12h.

During the last part of the 1983 season and throughout the 1984 season, the model was run operationally at JTWC. Verification statistics are presented in the Annual Typhoon Report, 1984 (JTWC, 1984). On page 164 of thi report, it can be noted that the model's performance met expectations. That  $i$  in comparison with other models, best performance was observed at the  $\cdot$ shorter range projections. At the more extended projections, models sensitive to environmental forcing were superior.

#### 5. PERFORMANCE CHARACTERISTICS

In this section, examples of model performance under controlled initialization are presented. As stated, input data to the model consist of 8 predictors -- initial storm latitude, initial storm longitude, time of year, average meridional translational speed over past 12h, average zonal speed over past 12h, average meridional translational speed over past 24h, average zonal translational speed over past 24h, and maximum storm intensity. Speeds are computed within the program from current, 12h- and 24h-old warning time positions.

How sensitive is the model to inaccuracies in operational specification of these predictors? This question is best answered by holding certain predictors constant and varying others.

#### 5.1 Time of Year

For a storm at a given location, which has a given intensity and for which past motion characteristics have been determined, the expected track, in the climatological sense, is a function of the time of year. This, of course, is merely a reflection of a normal climatological shift in the environmental<br>steering forces. The model's ability to sense these average forces is steering forces. The model's ability to sense these average forces is<br>demonstrated in Figure 5. Here, all input data were held constant, except for the Julian day number. The resultant shift in track is clearly noted. In accordance with climatological prediction, recurvature within 72h can be expected early and late in the season, but not during mid-season when the maximum westerly component occurs near mid-August.

#### 5.2 Initial Latitude

In the climatological sense, storms initially in the deep tropics are more likely to remain embedded in the easterlies (move with a westward component of motion) through 72h than are storms initially at a more polewar location. Controlled WPCLPR forecasts, as illustrated in Figure 6, agree with this expectation. However, the model sensitivity to errors in initial



Sensitivity of WPCLPR to time of year. Shown are 72-h forecast Figure  $5.$ tracks on fifteenth day of each month, May through December, with other predictors being held constant. Storm intensity was set at 100 kt.



Sensitivity of WPCLPR model to initial latitude. Shown are 72-h Figure 6. forecast tracks with different initial latitudes and with other predictors being held constant. Date and storm intensity are set at 15 September and 100 kt, respectively.

error, the downstream effect of even a  $1^{\circ}$  or  $2^{\circ}$  error in latitude is r erious. The downstream effect of even a 10 or 20 error in latitude is not 20 error in latitude is not 20 error

#### 5.3 Initial Longitude .5.3 Initial Longitude of the United States of the United States of the United States of the United States of

Figure 7 shows the effect of varying the initial longitude and holding constant the other seven input parameters. Here, the sensitivity is even less than for initial latitude, although there is some tendency for storms that are initially closer to the western edge of the basin to have a smaller northerly component in 72h. component in 72h.

Two predictors (average meridional and zonal speed over the past 12h) are involved here. The model computes these orthogonal components from the present and the 12h-old positions of the storm. As noted in Figure 8, there is much model sensitivity here, with errors in the 12h-old position having rather marked downstream effect. In this example, if the 12h-old position is to the north, the 72h forecast position will be to the south. Similarly, if the 12h-old position is to the south, the 72-h forecast position will be to the north. Further tests (not shown here), show even greater sensitivity to differences in present position. Accordingly (section 6), great care must be taken in specification of present and 12h-old warning time positions. Collectively, these two positions should reflect the forecaster's best estimate of average storm motion over the past 12h.

#### 5.5 Average Motion Over the Past 24h 5.5 Average Motion Over the Past 24h

mate of average storm motion over the past 12h.

In contrast to model sensitivity to average motion over the past 12h, model sensitivity to average motion over the past 24h (as obtained from the present and the 24h-old positions) is considerably less. This is depicted in Figure 9. It can be noted that the downstream effects are rather small.  $\mathcal{F}_{\mathcal{F}}$ 

#### 5.6 Storm Intensity

n component than small storms. Although the WPCLPR does not directl address storm size, it does consider storm intensity and there is a we positive statistical relationship between storm size (as measured by the outer closed surface isobar) and storm intensity (Merrill, 1982). Also, weak storms tend to be steered more by the lower troposphere and intense storms more by a deep laver throughout the troposphere (Simpson, 1971). The net result of these factors, and probably others, is that the more intense storms tend to have a larger poleward component than do the weaker storms. Also, there is slight increase in the westerly component with increasing storm intensity. The effect is illustrated in Figure 10. The effect is illustrated in Figure 10. The effect is in Figure 10. The effect is in Figure 10. The effect is i



Shown are 72-h Figure 7. Sensitivity of WPCLPR model to initial longitude. Shown are 72-h forecast tracks with different initial longitudes and with other predictors being held constant. Date and storm intensity are set at 15 September and 100 kt, respectively.



Figure 8. Sensitivity of WPCLPR model to 12h-old position. Shown are 72-h forecast tracks with three 12h-old positions and with other predictors being held constant. Date and storm intensity are set at 15 September and 100 kt, respectively.



Figure 9. Sensitivity of WPCLPR model to 24h-old position. Shown are 72-h forecast tracks with different 24h-old positions with our predictors being held constant. Date and storm intensity are set at 15 September and 100 kt, respectively.



Figure 10. Sensitivity of WPCLPR model to initial storm intensity. Shown<br>are 72-h forecast tracks with three initial intensities and with are 72-h forecast tracks with three initial intensities and with other predictors being held constant. Date has been set at 15 September.

#### 6. OPTIMIZING MODEL PERFORMANCE

#### 6.1 Initial Motion Vectors

In the preceding section, it was noted that the model is very sensitive to the average motion vector over the past 12h as defined by the current and the 12h-old storm positions. The forecaster must make every effort to assure that these positions reflect a best-track scale of motion. The methodology to accomplish this varies from one forecast center to another. A pitfall is the unqualified use of storm positions that reflect small-scale, perhaps trochoidal, oscillations of the storm center, which are not really representative of the larger scale, more conservative motion of the entire storm envelope.

In this connection, the current position of a storm need not automatically be the 12h-old position of a storm 12h hence; similarly, the current 12h-old position of a storm need not automatically become the 24h-old position 12h hence. The three sets of positions (now, 12h and 24h ago) might requir continuous adjustment so as to best reflect current motion trends.

#### 6.2 Model Limitations

As stated, operational use of the model is limited to storms that are initially at 5°-35°N and westward from 150°E longitude through the Asian mainland. In the temporal sense, the model should not be activated on storms occurring before 15 Mayor after 15 December. Finally, the system must 'be of at least tropical storm intensity. Violation of these spatial, temporal, and wind restrictions will result in performance degradation. For example, Figure 11 illustrates a predicted 72-h track on a storm that is initially near the northern boundary of the dependent data set (35°N). The model is acutely biased toward storms that moved slowly; faster moving storms having been dropped from the master storm data file.

Activating the model on storms that were initially east of 150°E apparently does not have serious effects on the model performance. Figure 7 shows one such forecast on a storm, initially at 15°N, 160°E. The 72-h track does not appear to be out of line with the other tracks that are within the spatial bounds of the model.

#### .' 7. FURTHER COMMENTS

The model described here is designed to make optimum use of climatology and persistence in WESPAC tropical cyclone prediction and provides a good first estimate of future storm motion. However, the third-order polynomia representation of the storm tracks does not allow for small local deviation from the large-scale climatology. Thus, track deviations as storms cross mountainous areas, such as Taiwan or the Philippine Islands (Brand and Blelloch, 1972, 1973) are not well-handled by the model. These areas would have to be modeled separately and blended into the larger scale patterns.



Figure 11. Example of WPCLPR model performance on a storm initially located near northern boundary of developmental (dependent) data set Date and storm intensity were set at 15 September and 90 kt respectively. Storm symbols give positions every 12h.

Through knowledge of current and future steering forces, it should be possible to refine model performance. Indeed, the model can be used as input to more sophisticated models that are sensitive to the existing environmental conditions. However, experience has shown that the model is subject to degradation if these synoptic steering forces are not known with sufficient precision (Neumann, 1980).

In addition to its use as a "first-guess" to the projected track, or as input to more sophisticated models, the WPCLPR model has other potential uses. Some of these are:

1) Establishment of a benchmark skill level with which to assess the real skill of more sophisticated models (Neumann and Pelissier, 1981).

2) Establishment of a "Forecast Difficulty Level," which can be used to assess long-term trends in tropical cyclone prediction (Neumann, 1981). When the model is run in this mode, best-track, rather than operational input, data are used.

3) Simulation studies that use Monte-Carlo techniques (Neumann, 1975; Jarrell et al.,  $1984$ ).

4) Normalization of WESPAC tropical cyclone forecasts to those of other basins (Pike, 1985).

16

#### 8. COMPUTER PROGRAM LISTING

Listing of a recommended FORTRAN IV computer program to run the program is given in the appendix. The program was written for an IBM 32-bit (4-byte) word-size machine and some of the statements may not .be compatible with non-IBM compilers. Also, the job control language has been omitted; this must be user-supplied.

When the program is run, two sets of data cards are read in at execution time; the regression coefficient set and the stonm data card set. The former consists of 110 cards, the first and last of which are dummies and read as such by the program. The 108 cards that contained the coefficients could probably be stored elsewhere or entered through a block data subroutine.

.Following the regression coefficient cards are the storm data cards; there is no limit to the number of storms that can be run in a single job step. The program senses the last storm data card that goes through as endof-file-marker; however, a "sentinel" card with 9999999 punched in columns 1 though 8 for the integer variable YMDH could alternately be incorporated with minor program modification. The specific formats (FORMAT statement 20 of the MAIN program) of the data card are:

Columns 1 through 8 -- Date/time in integer format (i.e., 85081706 represents August 17,1985,0600 GMT). Columns 9 and 10 -- leave blank. Columns 11 through 15 -- initial storm latitude. Columns 16 through 20 -- initial storm longitude. Columns 21 through 25 -- storm latitude 12h earlier. Columns 26 through 30 -- storm longitude 12h earlier. Columns 31 through 35 -- storm latitude 24h earlier. Columns 36 through 40 -- storm longitude 24h earlier. Note: Above latitudes and longitudes are in F5.1 format. Column 41 -- leave blank. Columns 42 through 44 -- wind in whole knots (integer format). Note: If wind is < 100 kt, the two-digit entry must be right-adjusted. Columns 45 through 56 -- storm name (Alphanumeric format). Columns 57 through 80 -- leave blank.

Two sample storm data cards are included on the final page of the program. Program output generated by each data card is:

72H WPCLPR FORECAST ON STORM TEST STORM1<br>BEGINNING OF FORECAST PERIOD YR/MO/DA/HR (GMT) IS 85051500 BEGINNING OF FORECAST PERIOD YR/MO/DA/HR (GMT) IS 85051500 MAXIMUM WIND IS 100 KNOTS  $T = T$   $T = 0$   $T = 0$ 



and the second is:

72H WPCLPR FORECAST ON STORM TEST STORM2<br>BEGINNING OF FORECAST PERIOD YR/MO/DA/HR (GMT) IS 85091500 MAXIMUM WIND IS 100 KNOTS MAXIMUM WIND IS 100 KNOTS<br>THIS IS RUN NUMBER 2  $T_{\text{max}}$  is non-non-permitting  $\frac{1}{2}$ 



mese foresast tracks (for 15 May and 15 September) were among those<br>trated in Figure 5. It is recommended that the program he tested of illustrated in Figure 5. It is recommended that the program be tested on these two cases.

#### 9. ACKNOWLEDGMENTS

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#### REFERENCES

- Brand, S. and J. Blelloch, 1972: Changes in the characteristics of typhoon crossing the Philippines. Naval Environmental Prediction Research
	- Brand, S. and J. Blelloch, 1973: Changes in the characteristics of typhoons crossing the Island of Taiwan. Naval Environmental Prediction Research<br>Facility Research Report TP 8-73, 21 pp.
	- Central Meteorological Bureau, 1972: Northwest Pacific typhoon track maps, 1949-1969, Peking, PRC, 378 pp. (in Chinese).
	- Jarrell, J. D., S. Brand and P. F. Krumpe, 1984: Tropical cyclone threat estimates: Where are we? Postprints, Fifteenth Conference on Hurrican and Tropical Meteorology, January 9-13, 1984, Miami, Florida. Americ Meteorological Society, Boston, J29-J34.
	- Joint Typhoon Warning Center, 1985: Annual Typhoon Report, 1984. Joint Typhoon Warning center (JTWC), Guam, Mariana Islands, 222 pp.
	- Merrill, R. T., 1982: A comparison of large and small tropical cyclones. University of Colorado, Department of Atmospheric Science Paper No. 352, Ft. Collins, 75 pp.
- Neumann, C. J., 1972: An alternate to the HURRAN tropical cyclone forecast system. NOAA Technical Memorandum NWS SR-62, Miami, Florida, 25 pp.
	- Neumann, C. J., 1975: A statistical study of tropical cyclone positioning errors with economic applications. NOAA Technical Memorandum NWS SR-62, " .Miami, Florida, 21 pp.
		- Neumann, C. J., 1980: The prediction of tropical cyclone motion An nustic Colected Papers. Thirteenth Technical Conference on Hurricane and Tropical Meteorology, Miami Beach, Florida, December 1-5, 1980. American Meteorological-Society, Boston, 68-78.
		- Neumann, C. J. and P. W. Leftwich, 1977: Statistical guidance for the prediction of eastern North Pacific tropical cyclone motion - Part 1. NOAA Technical Memorandum NWS WR-124, Miami, Florida.
- Neumann, C. J. and G. S. Mandal, 1978: Statistical prediction of tropical cyclone motion over the Bay of Bengal and Arabian Sea. Indian Journal of Meteorology, Hydrology, and Geophysics, 29, 487-500.
- Neumann, C. J. and J. M. Pellissier, 1981: An analysis of Atlantic tropical cyclone forecast errors, 1970-1979. Monthly Weather Review, 109, 1248- 1266.
- Neumann, C. J. and E. A. Randrianarison, 1976: Statistical prediction of tropical cyclone motion over the southwest Indian Ocean. Monthly Weather Review, 104, 76-85.
- Pike, A. C., 1985: The variation of track forecast difficulty among tropical cyclone basins. Extended Abstracts, Sixteenth Conference on Hurricanes and Tropical Meteorology, May 14-17, 1985, Houston, Texas. American what in up real increasingly that is the state of the contract of the contract of the state of the material society.
- Simpson, R. L., 1971: The decision process in hurricane forecasting. NOAA<br>Technical Memorandum NWS SR-63, Miami, Florida, 35 pp.

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APPENDIX: FORTRAN Computer Program and Associated Data Needed for<br>WPCLPR model

```
C THIS IS MAIN PROGRAM
      INTEGER YMOH.WIND
      REAL LAO.LOO.LAM12.LOM12.LAM24.LOM24
      DIMENSION CI(121.M(32.12).COF(32.12).OISP(2.6).FP(2.6)
      DIMENSION IDIR(3.8).SPD(8)
      NRUNS=0
C
C READ IN REGRESSION COEFFICIENTS AND CORRESPONDING PREDICTOR NUMBERS
Ċ
      CALL READROICOF.M.CII
   10 READ(5.20.END=50) YMDH.LAO.LOO.LAM12.LOM12.LAM24.LOM24.WIND
                  .NAME1.NAME2.NAME3
   20 FORMAT(18.2X.6F5.1.1X.13.3A4)
      NRUNS=NRUNS+1
C
C PREPARE FORECAST
\mathsf{C}CALL WPCLIP (YMDH.LAO.LOO.LAM12.LOM12.LAM24.LOM24.WIND.CI.M.COF.
     $DISP.FPIC
  OBTAIN PAST AND FORECAST MOTIONS
\mathsf{C}C
      CALL DIRSPO(LAO.LOO.LAM12.LOM12.LAM24.LOM24.FP.IDIR.SPD)
C
  WRITE OUTPUT TO PRINTER
c
Ċ
       WRITE(6.23)
   23 FORMAT (///////5X.61(1H*))
       WRITE(6.25)NAME1.NAME2.NAME3.YMDH.WIND.NRUNS
   25 FORMAT(5X. 72H WPCLPR FORECAST ON STORM :. 3A4/SX. BEGINNING OF FOR
      $ECAST PERIOD TR/MO/DA/HR (GMT) IS 'I8/5X. MAXIMUM WIND IS 'I3.
      $' KNOTS'/SX.'THIS IS RUN NUMBER '. I4/J
       WRITE(6.26)
    26 FORMAT(IHO.29X.'DISPLACEMENT (NMI) MOTION (DIR/SPD)'/
                                                           OVER LAST 12H'I
                              LONG
                                       N + / S -E+7W-$5X. PROJECTION LATO
       WRITE(6.27)LAM24.LOM24
                                                                       --/--
    27 FORMAT (8X, '-24H', 4X, F4, 1, 1HN, F6, 1, 'E
      $--']WRITE(6.28)LAM12.LOM12.IDIR(1.1).IDIR(2.1).IDIR(3.1).SPD(1)
                                                  - - ---- 7X.311.28 FORMAT(8X.'-12H'.4X.F4.1.1HN.F6.1.'E
      $1H/.F4.1.' KTS')
       WRITE(6.29)LAO.LOO.IDIR(1.2).IDIR(2.2).IDIR(3.2).SPD(2)
                                                            0'.8X.311.29 FORMAT(8X. 0H'.4X.F4.1.1HN.F6.1. E
                                                   \mathsf{n}$1H/.FH.1. KTS')
       DO 35 J=1.6
       KHRS:12*J
       WRITE(6.30)KHRS.FP(1.J).FP(2.J).DISP(1.J).DISP(2.J).
      $IDIR(1.J+2).IOIR(2.J+2).IOIR(3.J+2).SPD(J+2)
    30 FORMAT(8X.1H+.12.1HH.4X.F4.1.1HN.F6.1.1HE.2F9.1.6X.3I1.1H/.
      $F4.1.' KTS'I
    35 CONTINUE
```

```
WRITE(6,38)
   38 FORMAT(5X.61(1H*)) -
      G0 I0 I050 CONTINUE
      STOP
      END
SUBROUTINE READRCICOF.M.CI)
      DIMENSION COF(32.12).M(32.12).CI(12).RDUMY(4).IDUMY(4)
C READ 108 CARDS CONTAINING REGRESSION COEFFICIENTS. THERE ARE 12 SETS
C OF 9 CARDS EACH. FIRST SET IS FOR 12H MERIDIONAL MOTION. SECOND SET IS
C FOR 12H ZONAL MOTION. THIRD SET IS FOR 24H MERIDIONAL MOTION. ETC.
C
 ARRAY COF HOLDS REGRESSION COEFFICIENTS. 32 COEFFICIENTS PER SET
 ARRAY M HOLDS CORRESPONDING PREDICTOR NUMBER
c
 ARRAY CI IS INTERCEPT VALUES. ONE PER SET. THIS IS PUNCHED ON FIAST
C
C CARD. CARDS 2 THRU 6 OF EACH SET HOLD PREDICTOR NUMBER AND REGRESSION
C COEFFICIENTS.
\mathsf{C}CARDS ARE SELF INDEXING.... THEY CAN BE OUT OF ORDER
Ċ.
      READ (5.6) DUMMY
                        ł
    6 FORMAT(A4)
                        \bar{\Sigma}DO 30 I=1.108
      READ(5,10) INDEX. (IDUMY(J). RDUMY(J). J=1.4)
   10 FORMAT(13.1X.4(14.E15.7))
      J = (IMDEX + 8)/9IF(MOD(INDEX-1,9), EQ.0)GO TO 25
      INIT = (INDEX - 9*(J-1))*4-7LAST=INIT+3
     N = 0DO 20 L=INIT.LAST
     N = N + 1M(L.J)=IDUMY(N)
   20 COF (L.J) =ROUMY (N)
     GO TO 30
   25 CI(J)=ROUMY(1)
   30 CONTINUE
     READ (5.6) DUMMY
     RETURN
     END
C ж ж ж ж ж
      SUBROUTINE DIRSPD(LAO.LOO.LAM12.LOM12.LAM24.LOM24.FP.IDIR.SPD)
C COMPUTE APPROXIMATE HEADING AND SPEED AVERAGED OVER 12H INTERVALS
     REAL LAO.LOO.LAMI2.LOMI2.LAM24.LOM24
     DIMENSION FP(2.6).Q(2.9).IDIR(3.8).SPD(c).LDIR(8)
     DQ_5 I = 1.200 \t4 \tJ = 1.64.0(1.3+3)=FP(1.3)5 CONTINUE
     Q(1,1) = LAM24
     Q(2.1) = QM24Q(1.2)=LAM12
     Q(2.2) = LOM12
```
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 $Q(1.3)$  = LAO  $0(2.3)$  = LOO T:.0087266 DO 20 J=1.8  $DT = Q(1, J+1) - Q(1, J)$ DX=(Q(2.J+1)-Q(2.J))\*COS((Q(1.J+1)+Q(1.J))\*T) SPO(J)=SQRT(DY=DY+DX=DX)=5. IF(SPD(J).EQ.0.0)GO TO 10 OIR=ATAN2(DX.DY) #57.29578  $\cdot$ IF(DIR.LT.O.O)DIR=DIR+360.  $LDIR(J) = DIR + .5$ IF(LOIR(J).EQ.O)LDIR(J)=360 GO TO 20 10 LDIR(J)=0 20 CONTINUE DO 30 J=1.8  $IDIR(1, J) = LDIR(J)/100$ IDIR(2.J)=(LOIR(J)-IDIR(1.J)#100)/10 30 IDIR(3. J) = LOIR(J) - IDIR(1. J) = 100 - IDIR(2. J) = 10 **RETURN END** Сжжжж SUBROUTINE WPCLIP(YMDH.LAO.LOO.LAMI2.LOMI2.LAM24.LOM24.WIND. \$ CI.M.COF.OISP.FPI INTEGER YMOH.WIND REAL LAO.LOO.LAM12.LOM12.LAM24.LOM24 DIMENSION CI(12).M(32.12).COF(32.12).DISP(2.6).FP(2.6) A WEST PACIFIC CLIMATOLOGY-PERSISTENCE METHOD FOR FORECASTING STORM DISPLACEMENTS THROUGH 72H AT 12 HRLY VALIO FROM 15 MAY THRU 15 DECEMBER ONLY AND FOR INCREMENTS. STORMS INITIALLY AT OR SOUTH OF 35N LATITUDE AND WEST OF 150E. **ARGUMENTS:** ON INPUT YMOH--DATE(YEAR.MONTH.DAY.HOUR). 18. (6/1/83.00Z-83060100) LAO--INITIAL LATITUDE. DEGREES LOO--INITIAL LONGITUDE. DEGREES LAM12--LATITUDE AT -12 HOURS LOMI2--LONGITUDE AT -12 HOURS LAM24--LATITUDE AT -24 HOURS LOM24--LONGITUDE AT -24 HOURS WIND--INITIAL MAXIMUM WIND. KNOTS CI--REGRESSION INTERCEPTS M--REGRESSION VARIABLE NUMBERS COF--REGRESSION COEFFICIENTS ON OUTPUT DISP--DISPLACMENTS 12Z.. 12M.. 24Z.. 24M.. 36Z.. 36M.. 48Z.. 48M..60Z..60M..72Z..72M.. **NM**  $Z$ .--  $E$  TO W NEG.

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C

 $\mathsf{C}$ M.-- S TO N POS.  $\mathsf{C}$ FP--FORECAST POSITIONS DEGREES DIMENSION PIIGGI C. POTENTIAL PREDICTORS ARE NUMBERED 3 THRU 166. ONLY 32 OF THESE ARE C USED FOR EACH OF THE 12 REGRESSION EQUATIONS.  $P(3)$  = LAO IFILAO.GT.35.1GO TO 2 GO TO 4 2 WRITE(6.3JLAO 3 FORMAT(//SX. CURRENT LATITUDE OF 'F4.1.' IS NORTH OF 35.0. PROGRAM \$ BEING TERMINATED'//) STOP  $4$   $P(4)$ = $100$ IT=TMOH/1000000 IM=YMDH/10000-IY\*100 ID=YMOH/100-IM=100-IY=10000 IH=YMDH-IY\*1000000-IM\*10000-ID\*100 P(5)=3055\*(IM+2)/100-(IM+10)/13\*2-91+10 IF(P(5).LT.134..OR.P(5).GT.350.1GO TO 5 **GO TO 7** 5 WRITE(6.6) YMOH 6 FORMAT(//SX. PROGRAM RESULTS NOT VALID BEFORE 15 MAY OR AFTER 15 D \$ECEMBER. CURRENT ORTETIME IS : IIO/IX. PROGRAM BEING TERMINATED / ) **STOP** 7 P(6)=(LAO-LAM12) \*2.5 P (7) = (LOO-LOM12) \*2.5\*COS((LAO+LAM12) \*0.0087267) C UNIT NM/30 MINS  $P(8) = (LAO-LAM24)*2.5$ P(9)=(LOO-LOM24)\*2.5\*COS((LAO+LAM24)\*0.0087267) С UNIT KNOT  $P(10)$  =WIND  $P(11)=P(3)*P(3)$  $P(12)=P(3)*P(3)*P(3)$  $P(13)=P(3)=P(4)$  $P(14)=P(3)*P(3)*P(4)$  $P(15)=P(4)=P(4)$  $P(16):P(3)*P(4)*P(4)$  $P(17)=P(4)=P(4)=P(4)$  $P(18)=P(3)*P(5)$  $P(19)=P(3)*P(3)*P(5)$  $P(20)=P(4)=P(5)$  $P(21)=P(3)*P(4)*P(5)$  $P(22)=P(4)=P(4)=P(5)$  $P(23)=P(5)=P(5)$  $P(24)=P(3)*P(5)*P(5)$  $P(25) = P(4) * P(5) * P(5)$  $P(26) = P(5) * P(5) * P(5)$  $P(27)=P(3)=P(6)$  $P(28)=P(3)*P(3)*P(6)$  $P(29)=P(4)=P(6)$  $P(30)=P(3)=P(4)=P(6)$  $P(31)=P(4)=P(4)=P(6)$ 

P(321:PISJ-PI6J  $P(33)=P(3)*P(5)*P(6)$  $P(34)=P(4)=P(5)=P(6)$  $P(35):P(5)*P(5)*P(6)$ PI361:PI6J-PI61  $P(37) = P(3) * P(6) * P(6)$ P(38)=P(4)\*P(6)\*P(6)  $P(39)=P(5)=P(6)=P(6)$  $P(40) = P(6) = P(6) = P(6)$ <br> $P(41) = P(3) = P(7)$ P(42) =P(3) \*P(3)\*P(7)  $P(43)=P(4)*P(7)$ <br> $P(44)=P(3)*P(4)*P(7)$  $P(45)=P(4)*P(4)*P(7)$ P(46J:PISJ-PI71  $P(47)=P(3)*P(5)*P(7)$  $P(48)=P(4)*P(5)*P(7)$  $P(49)=P(5)*P(5)*P(7)$  $P(50)=P(6)*P(7)$  $P(51)=P(3)*P(6)*P(7)$  $P(52)=P(4)*P(6)*P(7)$  $P(53)=P(5)*P(6)*P(7)$  $P(54)=P(6)*P(6)*P(7)$  $P(55)=P(7)*P(7)$ P(56)=P(3)\*P(7)\*P(7) P(57)=P(4)\*P(7)\*P(7) P(58)=P(5)=P(7)=P(7)  $P(59)=P(6)*P(7)*P(7)$  $P(60.1=P(7)*P(7)*P(7)$  $P(61)=P(3)*P(8)$  $P(62)=P(3)*P(3)*P(8)$  $P(63) = P(4) * P(8)$  $P(64) = P(3) * P(4) * P(8)$  $P(65)=P(4)*P(4)*P(8)$  $P(66) = P(5) * P(8)$  $P(67) = P(3) * P(5) * P(8)$  $P(68)=P(4)=P(5)=P(8)$  $P(69) = P(5) = P(5) = P(8)$  $P(70)=P(6)=P(8)$ P(711=P(3)\*P(6)\*P(8)  $P(72)=P(4)*P(6)*P(8)$ P(73)=P(5)\*P(6)\*P(8)  $P(74)$ :  $P(6)$   $P(6)$   $P(8)$  $P(75)=P(7)*P(8)$  $P(76)=P(3)=P(7)=P(8)$  $P(77)=P(4)=P(7)=P(8)$  $P(78) = P(5) * P(7) * P(8)$  $P(79)=P(6)*P(7)*P(8)$  $P(80)=P(7)*P(7)*P(8)$ .  $P(81)=P(8)=P(8)$ P(821=P(3)\*P(8)\*P(8)  $P(83):P(4)*P(8)*P(8)$ 

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 $P(84) = P(5) * P(8) * P(8)$  $P(85)=P(6)*P(8)*P(8)$  $P(86):P(7):P(8):P(8)$  $P(87) = P(8) * P(8) * P(8)$  $P(881=P(3)*P(9)$  $P(89) = P(3) * P(3) * P(9)$  $P(90) = P(4) * P(9)$  $P(91)=P(3)*P(4)*P(9)$  $P(92)=P(4)*P(4)*P(9)$  $P(93)=P(5)*P(9)$  $P(94) = P(3) * P(5) * P(9)$  $P(95) = P(4) * P(5) * P(9)$ P(96)=P(5)\*P(5)\*P(9)  $P(971; P(6)*P(9)$  $P(98) = P(3) * P(6) * P(9)$  $P(99) = P(4) * P(6) * P(9)$  $P(100)=P(5)*P(6)*P(9)$  $P(101) = P(6) * P(6) * P(9)$  $P(102)=P(7)=P(9)$  $P(103)=P(3)*P(7)*P(9)$  $P(104)=P(4)*P(7)*P(9)$ P(105)=P(5)\*P(7)\*P(9)  $P(106) = P(6) * P(7) * P(9)$  $P(107) = P(7) * P(7) * P(9)$  $P(108)=P(8)*P(9)$  $P(109) = P(3) * P(8) * P(9)$  $P(110)=P(4)*P(8)*P(9)$ P(1111=P(5)\*P(8)\*P(9)  $P(112)=P(6)*P(8)*P(9)$  $P(113)=P(7)*P(8)*P(9)$ P(114)=P(8) \*P(8) \*P(9)  $P(115)=P(9)=P(9)$ P(116)=P(3)\*P(9)\*P(9)  $P(117)=P(4)*P(9)*P(9)$  $P(118)=P(5)*P(9)*P(9)$  $P(119)=P(6)*P(9)*P(9)$  $P(120) = P(7) * P(9) * P(9)$  $P(121)=P(8)*P(9)*P(9)$  $P(122)=P(9)*P(9)*P(9)$  $P(123)=P(3)*P(10)$  $P(124) = P(3)*P(3) * P(10)$  $P(125)=P(4)*P(10)$  $P(126) = P(3) * P(4) * P(10)$  $P(127) = P(4) * P(4) * P(10)$  $P(128)=P(5)*P(10)$ P(129)=P(3)\*P(5)\*P(10)  $P(130) = P(4) = P(5) = P(10)$  $P(131)=P(5)=P(5)=P(10)$  $P(132)=P(6)*P(10)$  $P(133)=P(3)*P(6)*P(10)$  $P(134)=P(4)*P(6)*P(10)$  $P(135) = P(5) = P(6) * P(10)$ 

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P(1361=P(6)\*P(6)\*P(10)  $P(137)=P(7)*P(10)$  $P(138)=P(3)*P(7)*P(10)$ P(139)=P(4)\*P(7)\*P(10)  $P(140) = P(5) = P(7) = P(10)$  $P(14112P(6)*P(7)*P(10)$  $P(142)=P(7)*P(7)*P(10)$  $P(143)=P(8)*P(10)$  $P(144)=P(3)*P(8)*P(10)$  $P(145)$  =P(4) \*P(8) \*P(10) P(146)=P(5)\*P(8)\*P(10)  $P(147)=P(6)*P(8)*P(10)$  $P(148)=P(7)=P(8)=P(10)$  $P(149)=P(8)*P(8)*P(10)$  $P(150) = P(9) * P(10)$  $P(151)=P(3)*P(9)*P(10)$  $P(152)=P(4)*P(9)*P(10)$  $P(153) = P(5) * P(9) * P(10)$  $P(154)=P(6)*P(9)*P(10)$  $P(15512P(7)*P(9)*P(10)$  $P(156) \pm P(8) \neq P(9) \neq P(10)$  $P(157)=P(9)*P(9)*P(10)$  $P(158)=P(10)*P(10)$  $P(159) = P(3) * P(10) * P(10)$ P(160)=P(4)\*P(10)\*P(10) P(161)=P(5)\*P(10)\*P(10) P(162)=P(6)\*P(10)\*P(10) P(163)=P(7)\*P(10)\*P(10)  $P(164)=P(8)*P(10)*P(10)$  $P(165)=P(9)*P(10)*P(10)$ P(166)=P(10)\*P(10)\*P(10)  $(P(I), I=3, 166)$ WRITE(6.9) 9 FORMAT(25HOLIST OF PREDICTORS---- .8E12.4/10(E12.4)) DO 30 K=1.6 00 20 J=1.2  $KJ=(K-1)*2+J$ DISP(J.K)=CI(KJ)  $0010$   $1 = 1.32$ L=M(I.KJ) 10 DISP(J.K)=DISP(J.K)+P(L)\*COF(I.KJ) 20 CONTINUE FP(1.K)=DISP(1.K)/60.0+P(3) FP(2.K)=DISP(2.K)/60.0/COS((FP(1.K)+P(3))#0.0087266)+P(4) 30 CONTINUE **RETURN END** PERMANENT OATA CARDS (THIS CARD IS CONSIDERED PART OF SET) 0.8243047E 02  $\mathbf{I}$ 0.2086875E-01 154 -0.3422998E-02 113 -0.3635096E-02 0.1673843E 00 141 29 2

C

65-0.9927880E-04 148-0.5865134E-02 47 0.1775683E-03  $0.2699498E - 02$ 3 133

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.NWS NHC 6 A Tropical Cyclone Data Tape for the North Atlantic Basin, 1886-1977: Contents, limitations, and Uses. Brian R. Jarvinen and Eduardo l. Caso - June 1978 (PB285504/AS) NWS NHC 7 The Deadliest, Costliest, and Most Intense United States Hurricanes of the Century (and Other Frequently Requested Hurricane Facts). Paul J. Hebert and Glenn Taylor -August 1978 (PB 286753/AS) NWS NHC 8 Annual Data and Verification Tabulation of Atlantic Tropical Cyclones 1977. Miles B. Lawrence, Paul J. Hebert and Staff, NHC - March 1979 (PB295702) NWS NHC 9 Annual Data and Verification Tabulation of Atlantic Tropical Cyclones 1978. Paul J. Hebert and Staff, NHC - April 1979 (PB296323) NWS NHC 10 Statistical Forecasts of Tropical Cyclone Intensity for the North Atlantic Basin. Brian R. Jarvinen and Charles J. Neumann - April 1979 (PB297185) NWS NHC 11 A Guide to Atlantic and Eastern Pacific Models for the Prediction of Tropical Cyclone Motion. Charles J. Neumann - April 1979 (PB297141/AS) NWS NHC 12 Modification of NMC Analyses and Prognoses for Use in Statistical Tropical Cyclone Prediction Models. Preston W. Leftwich, Jr. - May 1979 (PB297190)<br>NWS NHC 13 Annual Data and Verification Iabulation Atlantic Text in 1979 Annual Data and Verification Tabulation Atlantic Tropical Cyclones 1979. Paul J. Hebert and Staff, NHC - June 1980 NWS NHC 14 A Statistical Tropical Cyclone Motion Forecasting System for the Gulf of Mexico, Robert T. Merrill - August 1980 NWS NHC 15 Annual Data and Verification Tabulation Atlantic Tropical Cyclones 1980. Glenn Taylor and Staff, NHC -June 1981 NWS NHC 16 A Compilation of Eastern and Central North Pacific Tropical Cyclone Data. Gail M. Brown and Preston W. leftwich, Jr, -August 1982 (PB83115444) NWS NHC 17 Annual Data and Verification Tabulation Atlantic Tropical Cyclones 1981. Staff, NHC - November 1982 NWS NHC 18 The Dead1iest, Costliest, and Most Intense United States Hurricanes of this Century (and Other Frequently Requested Hurricane Facts). Paul J. Hebert and Glenn Taylor, NHC -January 1983 (PB83-163527) NWS NHC 19 Annual Data and Verification Tabulation Atlantic Tropical Cyclones 1982. Gilbert B. Clark and Staff, NHC - February 1983 (PB83184077) NWS NHC 20 The Miss/Hit Ratio - An Estimate of Reliability for Tropical Cyclone Track Predictions, Preston W. Leftwich, Jr. - April 1983 NWS NHC 21 Annual Data and Verification Tabulation Atlantic Tropical Cyclones 1983 **NWS NHC 22 A Tropical Cyclone Data Tape for the North Atlantic Basin, 1886-1983:** Contents, Limitations, and Uses. Brian R. Jarvinen, Charles J. Neumann, and Mary A. S. Davis - March 1984 NWS NHC 23 Frequency and Motion of Western North Pacific Tropical Cyclones. Zongyuan Xue<br>I and Charles J. Neumann - May 1984 (PP85106466) and Charles J. Neumann - May 1984 (PB85106466)<br>NWS NHC 24 Hurricane Experience Levels of Coastal County Populations - Texas to Maine -June 1984 (PB85111383) NWS NHC 25 A Tropica1 Cyclone Data Tape for the Eastern and Central North Pacific Basins, 1949-1983: Contents, limitations, and Uses -September 1984 (PB85110054) NWS NHC 26 Annual Data and Verification Tabulation Atlantic Tropical Cyclones 1984. Gilbert B. Clark and Robert A. Case, NHC - February 1985.<br>NWS NHC 27 A Storm Surge Atlas for Corpus Christi Jexas, Brian B. A Storm Surge Atlas for Corpus Christi, Texas. Brian R. Jarvinen, A. Barry Damiano, and Gloria J.D. Lockett - August 1985 .NWS NHC 28 A Statistical Model for the Prediction of Western North Pacific Tropical Cyclone Motion (WPCLPR). Yiming Xu and Charles J. Neumann - November 1985

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