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HAND CALCULATOR PROGRAM TO COMPUTE PARCEL  
THERMAL DYNAMICS

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# HAND CALCULATOR PROGRAM TO COMPUTE PARCEL THERMAL DYNAMICS

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## I. INTRODUCTION

This program computes the temperature of an air parcel raised dry adiabatically to the lifted condensation level, LCL, and then pseudo-adiabatically thereafter. The procedure is accomplished without aid of an energy diagram (Skew-T, Pseudo-adiabatic chart, etc.) using iterative computations on a Hewlett-Packard 67 hand computer. The iterative computations involve using finite difference to solve the energy balance equation for the ascending parcel. The program can be used in lieu of an energy diagram for computation of atmospheric stability indices (e.g., determination of a thunderstorm gust potential). A sample problem and a computer program are in Appendix I.

## II. VARIABLE DEFINITION

$t$	=	temperature - °C
$t_d$	=	dew-point temperature - °C
$D$	=	dew-point depression - °C
$T$	=	temperature - °K
$p$	=	pressure - mb
$L_v$	=	latent heat of vaporization - j/g
$e_s$	=	saturation vapor pressure of water at a given temperature - mb
$R_a$	=	gas constant for dry air - .287 j/g°K
$R_v$	=	gas constant for moist air - .461 j/g°K
$c_p$	=	specific heat at constant pressure for dry air - 1.003 j/g°K
$m_v$	=	molecular weight of water - 18 g/g-mole
$m_d$	=	apparent molecular weight of dry air - 28.9 g/g-mole
$\gamma_m$	=	pseudo-adiabatic lapse rate - °K/mb
$dT/dp)_m$	=	pseudo-adiabatic lapse rate (finite difference) - °K/mb
$\Delta P$	=	change in pressure - mb
$w$	=	saturation mixing ratio
$c$	=	specific heat of liquid water

### III. EQUATIONS

1. Use simplified Haurwitz' equation [1] to obtain LCL temperature.

$$T_{LCL} = t_d - (0.212 + .001571 t_d - .000436 t) D$$

2. Use Poisson's equation to determine LCL pressure.

$$P_{LCL} = P (T_{LCL}/T_0)^{c_p/R_a} \quad c_p/R_a = 7/2$$

$$T_0 = \text{initial temperature} - ^\circ\text{K}$$

$$P = \text{initial pressure} - \text{mb}$$

3. Formula #184, p. 376 [2] is

$$c_p \frac{dT}{T} - R_a \frac{d(p-e_s)}{p-e_s} + d\left(\frac{wL}{T}\right) + cw \frac{dT}{T} = 0 \quad (\#184)$$

Differentiating #184 and applying some simplifying assumptions yields the equation for a lapse rate in a saturated atmosphere, equation #190, p. 377 [2].

$$\gamma_m = \Delta T / \Delta P)_m = \left( \frac{R_a T}{c_p p} \right) \frac{P + \frac{0.622 L v e_s}{R_a T}}{P + \frac{0.622 L v e_s}{c_p R_v T^2}} = ^\circ\text{K}/\text{mb} \quad (\#190)$$

4. Teton's formula #63, p. 343 [2] is used to derive the saturation vapor pressure,  $e_s$ .

$$e_s = 6.11 \times 10^{aT/b+T}$$

where  $a = 7.567$  and  $b = 239.7$  are based on the assumption that liquid water is present in the atmosphere below  $0^\circ\text{C}$ . This same assumption is used in the derivation of standard energy diagrams.

5. Since  $L_v$  changes with temperature, a linear interpolation is incorporated into the program to account for  $\Delta L_v / \Delta T$ . With  $L_v = 2358.12 \text{ j/g}$  at  $60^\circ\text{C}$ .,  $L_v = 2634.88 \text{ j/g}$  at  $-50^\circ\text{C}$ , the range of temperature =  $110^\circ\text{K}$ ,

and the difference in  $L_V$  between the two temperatures being 276.76 j/g, the proportion for the linear interpolation is established.

$$(333^\circ\text{K} - T_O) / 110^\circ\text{K} = X / (276.76 \text{ j/g}) \quad L_O = X + L_V$$

where  $T_O$  = temperature of the current iteration - °K

$L_V$  = 2358.12 j/g

$L_O$  = computed  $L_V$  for  $T_O$

$X$  = proportion resultant relating  $L_O$  and  $L_V$ .

6. Computational iterations are accomplished by finite differencing as follows:

$$a) \quad T_m^H = T_m^L + \Delta P (\chi)$$

where  $\chi$  = right hand side of equation #190, p. 377 [2].

$$b) \quad P_m^H = P_m^L + \Delta P$$

where  $T_m^H$  = new computed temperature (higher atmospheric level)

$T_m^L$  = old computed temperature (lower atmospheric level)

$P_m^H$  = pressure level of new temperature (higher atmospheric level)

$P_m^L$  = pressure level of old temperature (lower atmospheric level)

$$\Delta P = P_m^H - P_m^L$$

#### IV. LOGIC

1. As the program begins execution, the pressure and temperature of the LCL are computed and stored.
2. Computational iterations proceed adiabatically, checking the computed temperature and pressure against the LCL temperature and pressure.
3. At the LCL, the computer continues ascending parcel parameter computations but now accounts for diabatic effects.
4. The program stores temperature and pressure at the LCL as well as any other two levels as specified by the programmer.
5. The program terminates at the highest level at which temperature data are desired.

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