

Observations: Cryosphere Supplementary Material

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4.SM.1 Supplementary Material for the Sea Ice Section

Most of the published studies on the large-scale variability and trends of the global sea ice cover that have been published in recent years were based primarily on results from analysis of passive microwave satellite data (Parkinson et al., 1999; e.g., Zwally et al., 2002; Stroeve et al., 2007; Comiso and Nishio, 2008; Stammerjohn et al., 2008; Cavalieri and Parkinson, 2012; Parkinson and Cavalieri, 2012)

The first satellite-borne imaging passive microwave sensor was the Nimbus-5/Electrically Scanning Microwave Radiometer (ESMR), which provided quantitative measurements of the extent and variability of the sea ice cover in both hemispheres from 1973 to 1976 (Zwally et al., 1983; Parkinson et al., 1987). However, with only one channel and a system that scanned a wide field of view (between -50 and $+50$ degrees off-nadir, causing changes in incidence angle and footprint size), gaps in the record from a few days to a few months and an unknown bias, these data have not been included in time series variability and trend analysis in Chapter 4.

The data that are most frequently used are those from the multichannel, conically scanning (i.e., constant incident angle) and dual polarized microwave radiometers that provide more accurate and more consistent ice concentration and hence ice extent and ice area (see Glossary for definitions) products. This series started with the Nimbus-7 Scanning Multichannel Microwave Radiometer (SMMR) which was launched in October 1978 and provided, for the first time, measurements that allowed unambiguous determination of sea ice concentration (Gloersen et al., 1993).

Several algorithms for deriving sea ice concentration using different techniques and utilizing different sets of channels have been developed (Svendsen et al., 1983; Cavalieri et al., 1984; Swift et al., 1985; Comiso, 1986; Steffen et al., 1992). The most commonly used techniques for sea ice studies are the Nimbus-7 National Aeronautics and Space Administration (NASA) Team algorithm (NT1, Cavalieri et al., 1984) and the Bootstrap algorithm (Comiso, 1986). SMMR was eventually succeeded by a series of Special Sensor Microwave/Imager (SSM/I) sensors and the two systems now provide a continuous set of data from November 1978 to the present. Subsequently, a more capable and improved sensor, Advanced Microwave Scanning Radiometer - Earth Observing System (AMSR-E), was launched on board the NASA/Aqua satellite and this has provided higher resolution and improved sea ice data from May 2002 to October 2011. The algorithms currently used for this sensor are the AMSR-E Bootstrap Algorithm (ABA) and the NASA Team Algorithm, Version 2 (NT2) as discussed in Markus and Cavalieri (2000). With some enhancements ABA was also adapted and used to reprocess SMMR and SSM/I data, and called SSM/I (or SMMR) Bootstrap Algorithm (SBA) as discussed in Comiso and Nishio (2008).

Using the SBA, AMSR-E data were used as the baseline and basis for improving the SMMR and SSM/I ice data sets used in Chapter 4. NT2 addressed some of the problems associated with NT1, such as erroneously low ice concentrations within the pack caused by unpredictable polarization ratios. Comparisons of NT2 and Bootstrap data have shown good agreement (Comiso and Parkinson, 2008; Parkinson and

Comiso, 2008) in analyzed trends and variability. NT2 data, however, could not be used for the entire historical data because it requires the use of 89 GHz data, which is not available in SMMR and in the early part of the SSM/I time series. The time series that has been used as an alternative to the NT2 time series has thus been the NT1 time series which provides values different from NT2. In the meantime, the Hadley Centre, in collaboration with National Oceanic and Atmospheric Administration (NOAA), constructed another sea ice record referred to as HadISST_Ice. The data set has been assembled together with sea surface temperature (SST) at a relatively coarser resolution (1° latitude by 1° longitude) and for a longer time series. The data set as described in Rayner et al., (2003) made use of atlases, *in situ* data and ice centre data (as in, Walsh and Chapman, 2001) for the pre-satellite era. Starting in 1979, satellite data including those from NT1 have been used. The use of satellite data in the Hadley data set, however, apparently has had some problems of consistency because, apparently spurious, sudden increases in ice concentration from one year to another have been identified (e.g., Screen, 2011), making the data set difficult to use for variability and trend analysis. Also, as Screen (2011) pointed out, there are large differences in ice extent and ice area distributions derived from NT1 data compared with those derived from NT2 data.

The results presented in Section 4.2 make use of mainly SBA data for consistency with those presented in AR4. To assess the robustness of the conclusions of Section 4.2 to changes in the data sets used, a comparative analysis of results from SBA, NT1 and Hadley (i.e., HadISST1_Ice) data is presented.

Time series plots of monthly anomalies in ice extent as derived from SBA, NT1 and Hadley for the period November 1978 to December 2011 are presented in Figure 4.SM.1. Although data up to December 2012 are presented in Chapter 4, data available for HadISST1_Ice and NT1 for this comparison goes up to December 2011 only. The NT1 data set used is an update version of that presented in Cavalieri and Parkinson (2012) and Parkinson and Cavalieri (2012). The plots in Figure 4.SM.1a and 4.SM.1b for SBA and NT1, respectively show very similar patterns, but the Hadley plot (Figure 4.SM.1c) shows deviations from the other two, especially from 1984 to 1986 and from 2007 to 2012, where the amplitudes of the interannual variation are significantly higher. Using linear regression, the derived trends are estimated to be -3.73% per decade for SBA, -4.22% per decade for NT1 and -2.0% per decade for Hadley. The trends from SBA and NT1 differ slightly but provide similar trend information but that from Hadley is about half the other two, with a much more modest decline. The difference in the trend values for the Hadley data is likely due to the anomalous deviation of the data from the other two data sets as indicated in the preceding text. The corresponding estimates for the trends in sea ice area data are -4.35% , -4.71% and -2.8% per decade, respectively, all relatively higher than those for ice extent but again providing similar trend results for SBA and NT1 but a significantly lesser rate of decline for Hadley.

In the Antarctic, the trends are more modest and in the opposite direction as depicted in Figure 4.SM.2. Again, the patterns of the interannual variability are very similar for all three, with the Hadley data being the most different, exhibiting higher short-term fluctuations and a more positive trend. Trend results are $+1.44\%$, $+1.34\%$ and $+2.48\%$ per decade for SBA, NT1 and Hadley, respectively. Again, the results

from SBA and NT1 are slightly different but provide similar trend information while Hadley provides a considerably higher trend. The corresponding trends in ice area are $+2.07\%$, $+1.57\%$ and $+3.07\%$ per decade, for SBA, NT1 and Hadley, respectively. The discrepancy in the trends for ice area between the SBA and NT1 results is in part due to lower concentration averages for NT1 compared with SBA data as indicated in the preceding text. However, they provide similar conclusions about the changes in the ice cover. The Hadley trend is again substantially higher than those of the other two.

For a more detailed comparison, September monthly sea ice extents for SBA, NT1 and Hadley for all years from 1979 to 2011 are plotted together in Figure 4.SM.3a. For completeness, ABA and NT2 ice extents using AMSR-E data are also shown for the period 2002–2011. The values from all three primary data sources are very similar, with the SBA showing the highest values and Hadley normally lowest. Values from AMSR-E data using ABA and NT2 are relatively lower because of higher resolution as discussed in Comiso and Nishio (2008), but otherwise there is good consistency. There are greater discrepancies among the three data sources when sea ice areas are plotted (Figure 4.SM.3b), with NT1 and Hadley showing good agreement up to 1997 and significant disagreement after that. The higher ice area values for SBA are associated with higher ice concentration values derived from the data than the other two, as discussed previously. The values for ABA and NT2 from AMSR-E (which have been used as the baseline) are in good

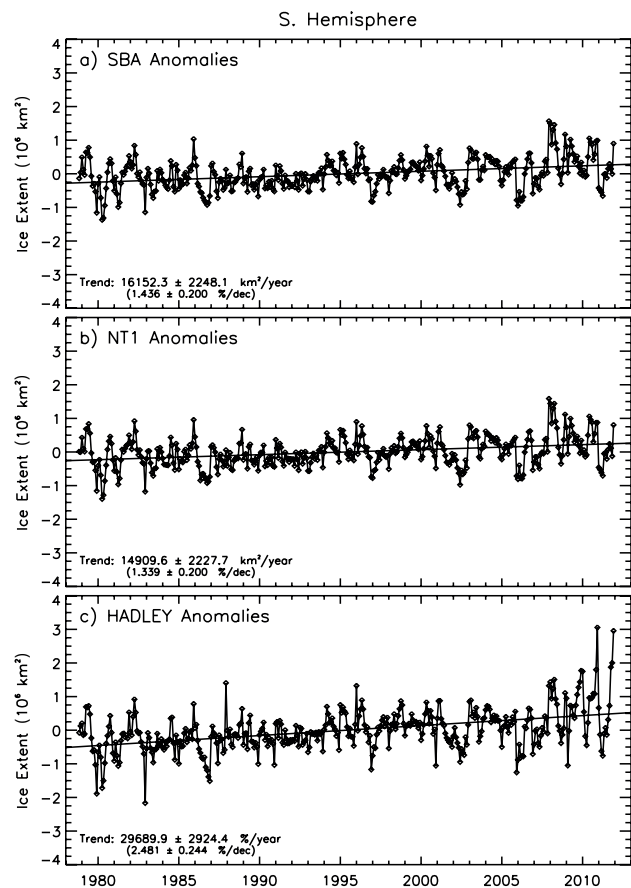


Figure 4.SM.2 | Monthly anomalies of ice extent from November 1978 to December 2011 using (a) SBA, (b) NT1 and (c) Hadley data in the Southern Hemisphere. Trends are shown with uncertainty calculated at 1 standard deviation (1σ).

agreement and also agree well with the SBA values. Large difference between NT1 and NT2 values are evident, as has been observed by Screen (2011). Nevertheless, the trends in extent for SBA and NT1 are -10.2% and -10.5% per decade, respectively, basically providing the same conclusion, while the trend in extent for the Hadley data is -8.0% per decade. The trends for ice area are also similar enough at -11.3% (SBA) and -12.5% per decade (NT1) while the trend for Hadley data is -10.2% per decade.

Another data set that is available and has been used for sea ice studies is that from the Arctic Radiation and Turbulence Interaction Study (ARTIST) provided by the University of Bremen. The data make use of only the 89-GHz channels (horizontal and vertical polarized data) to generate relatively high resolution data from AMSR-E (5 to 6 km). High resolution is needed in many mesoscale studies. However, the 89-GHz radiation is very sensitive to weather and changes in the snow cover conditions, and the data should be used with care because they may be partly contaminated by incorrect values, especially under adverse weather conditions. Also, the ARTIST time series data from Aqua/AMSR-E is, as yet, too short for meaningful sea ice variability and trend studies.

Of the three satellite data sets that are currently available for sea ice variability and trend analysis, SBA and NT1 provides basically the same

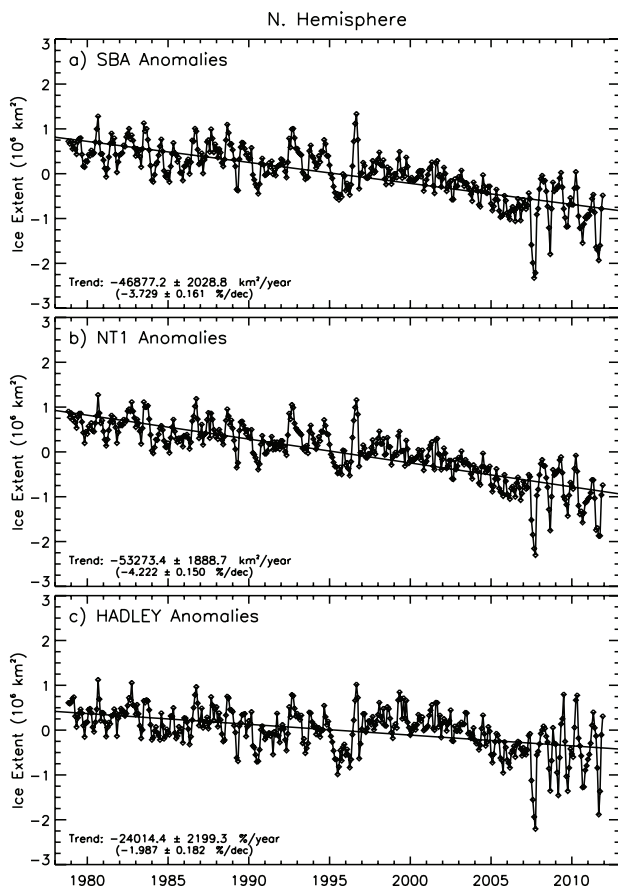


Figure 4.SM.1 | Monthly anomalies of ice extent from November 1978 to December 2011 using (a) SBA, (b) NT1 and (c) Hadley data in the Northern Hemisphere. Trends are shown with uncertainty calculated at 1 standard deviation (1σ).

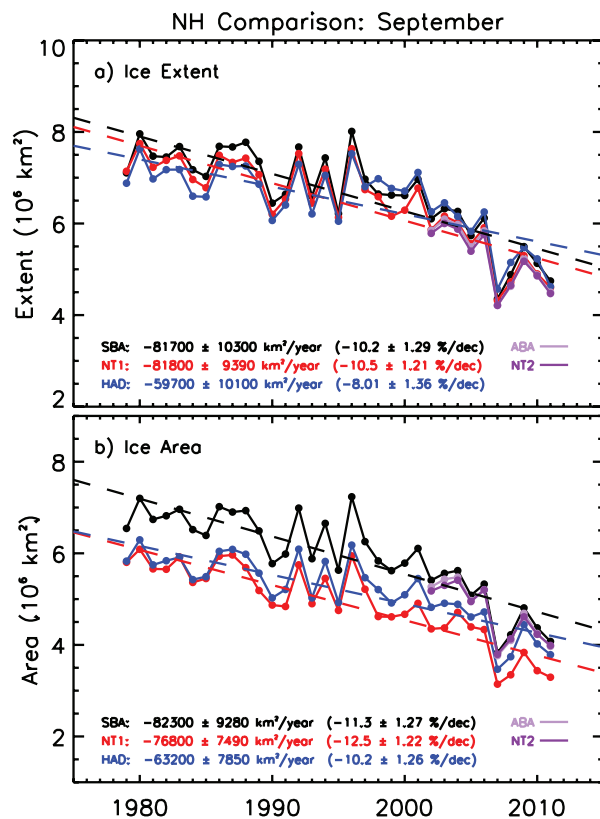


Figure 4.SM.3 | Monthly estimates of (a) sea ice extent and (b) sea ice area for the month of September from 1979 to 2011 for the Northern Hemisphere. Trends are shown with uncertainty calculated at 1 standard deviation (1σ).

patterns of interannual variability and approximately similar trends. The distributions are similar enough to provide basically the same information and conclusions about the trend of the changing sea ice cover. The patterns of variability provided by the Hadley data set are generally similar to those of SBA and NT1 but there are years when the data are suspect, as has been identified by Screen (2011). This is the primary reason for the discrepancies in the trends of the Hadley data compared with those of the SBA and NT1 data.

4.SM.2 Details of Studies of Glacier Area Change

Table 4.SM.1 provides an overview of studies reporting glacier area changes over entire mountain ranges or larger regions. Where available, area change rates are also given for sub-periods that have been extracted from the respective publications, partly using a linear extrapolation of given change rates to determine values for a common year in all sub-regions. In some studies, the glacier count is not given.

Table 4.SM.1 | Overview of studies presenting glacier area changes. **Bold** values in the column 'Change rate' indicate values shown in Figure 4.10. Values in *italics* refer to the entire period with 'Area covered' giving the value for the last year. Change rates are given for some regions to three decimals to avoid overlap of lines in Figure 4.10. For three regions (5, 9, and 12) studies on area changes were not found.

RGI Region	Country / Region	Sub-region / Mountain Range	Start Year	End Year	Number of Years	Glacier Count	Area Covered (km ²)	Relative Change (%)	Change Rate (% a ⁻¹)	Reference
1	USA, AK	Chugach Mountains	1952	2007	55	347	1285.7	-23	-0.42	Le Bris et al. (2011)
2	USA	North Cascades	1958	1998	40	321	117.3	-7.0	-0.18	Granshaw and Fountain (2006)
2	Canada	Rocky Mountains	1985	2001	16	523	1056.7	-7.6	-0.48	Tennant et al. (2012)
			2001	2006	5	523	976.5	-9.9	-1.98	
			1985	2006	21	523	880.0	-16.7	-0.80	
2	USA	Wind River Range	1966	2006	40	n/a	45.9	-37.7	-0.94	Thompson et al. (2011)
2	Canada	Yukon	1959	2007	48	n/a	11622	-21.9	-0.456	Barrand and Sharp (2010)
2	Canada	Rocky Mountains ^a	1985	2005	20	14329	30063	-11.1	-0.555	Bolch et al. (2010)
2	Canada	Clemenceau Icefield	1985	2001	16	123	313	-13.4	-0.84	Jiskoot et al. (2009)
		Chaba Group	1985	2001	16	53	97	-28.9	-1.81	
2	Canada	Rocky Mountains	1952	2001	49	59	40	-15	-0.31	Debeer and Sharp (2007)
		Columbia Mountains	1952	2001	49	403	397	-5.0	-0.10	
		Coast Mountains	1964	2002	38	1053	2397	-5.0	-0.13	
3	Canada	Queen Elizabeth Island	1960	2000	40	1274	107071	-2.7	-0.07	Sharp et al. (2013)
3	Canada	North Ellesmere	1960	2000	40	473	27556	-3.4	-0.09	Sharp et al. (2013)
		Agassiz	1960	2000	40	296	21645	-1.3	-0.03	
		Axel/Meighen / Melville	1960	2000	40	165	12231	-1.7	-0.04	
		Prince of Wales	1960	2000	40	39	19558	-0.9	-0.02	
		South Ellesmere	1960	2000	40	187	10696	-5.9	-0.15	
		Devon Island	1960	2000	40	114	15344	-4.0	-0.10	
4	Canada	Bylot Island	1959	2001	42	n/a	5036	-5.0	-0.119	Dowdeswell et al. (2007)
4	Canada	Barnes Ice Cap	1958	2000	42	n/a	5995	-2.0	-0.048	Sharp et al. (2013)
		Penny Ice Cap	1959	2000	41	n/a	6604	-1.9	-0.046	
		Terra Nivea	1958	2000	42	n/a	197	-14.0	-0.33	
		Grinnel Ice Cap	1958	2000	42	n/a	135	-10.9	-0.26	
4	Canada	Baffin Island	1975	2000	25	264	2187	-12.5	-0.16	Paul and Svoboda (2009)
5	Greenland	n/a								
6	Iceland	Four ice caps	1998	2011	13	4	1004.5	-7.6	-0.58	Johannesson et al. (2013)
7	Svalbard	Glaciers >1 km ²	1990	2008	18	n/a	5204.7	-4.6	-0.26	König et al. (2013)
8	Norway	Jostedalbreen	1966	2006	40	297	725.1	-9	-0.225	Paul et al. (2011)
8	Norway	Jotunheimen	1965	2003	38	164	229.5	-12.4	-0.33	Andreassen et al. (2008)
8	Norway	Svartisen	1968	1999	31	300	518.0	-1.1	-0.04	Paul and Andreassen (2009)
9	Russian Arctic	n/a								
10	Russian Federation	Ural	1956	2000	44	30	9.17	-22.3	-0.51	Shahgedanova et al. (2012)
10	Russian Federation	Kodar Mountains	1995	2001	6	34	11.72	-18.7	-3.11	Stokes et al. (2013)
			2001	2010	9	34	9.53	-26.4	-2.94	
			1995	2010	15	34	7.01	-40.2	-2.68	
10	Russian Federation	Altai Chuya Ridges	1952	2004	52	126	284	-19.7	-0.38	Shahgedanova et al. (2010)
10	Russian Federation	Altai	1952	2008	56	1030	805	-10.2	-0.182	Narozhnyi and Zemtsov (2011)
11	Austria	Alps	1969	1998	29	925	567	-17.1	-0.59	Lambrech and Kuhn (2007)
11	Austria	Ötztaler Alps	1997	2006	9	81	116	-8.2	-0.9	Abermann et al. (2009)
11	Switzerland	Alps	1973	1999	26	938	1171.2	-16.1	-0.62	Paul et al. (2004)
			1985	1998	13	471	372.2	-18.0	-1.38	
11	Spain	Pyrenees	1982	1994	12	10	6.08	-20.9	-1.97	Gonzales Trueba et al. (2008)
			1994	2001	7	10	4.81	-39.7	-5.6	
			1982	2001	19	10	6.08	-52.3	-2.75	

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Table 4.SM.1 (continued)

RGI Region	Country / Region	Sub-region / Mountain Range	Start Year	End Year	Number of Years	Glacier Count	Area Covered (km ²)	Relative Change (%)	Change Rate (% a ⁻¹)	Reference
11	Italy	Aosta Valley	1975	1999	24	174	163.9	-16.7	-0.7	Diolaiuti et al. (2012)
			1999	2005	6	174	136.6	-12.4	-2.07	
			1975	2005	30	174	119.6	-27.0	-0.9	
11	Italy	South Tyrol	1983	1997	14	205	136.6	-19.7	-1.41	Knoll and Kerschner (2009)
			1997	2006	9	302	109.7	-11.9	-1.32	
			1983	2006	23	205	136.6	-31.6	-1.37	
11	Italy	Lombardy	1992	1999	7	249	117.4	-10.8	-1.54	Citterio et al. (2007)
12	Caucasus	n/a								
13	Mongolia	Altai Mountains	1989	2009	20	n/a	213	-4.2	-0.21	Krumwiede et al. (2013)
13	China	Muztag Ata and Konggur Mountains (East Pamir)	1962	1990	28	n/a	838.2	-3.4	-0.12	Shangguan et al. (2006)
			1990	1999	9	n/a	809.9	-4.6	-0.52	
			1962	1999	37		772.2	-7.9	-0.214	
13	China (Tarim Basin)	Quarqan	1977	2001	24	399	752.6	-3.4	-0.14	Shangguan et al. (2009)
		Keliya	1970	1999	29	731	1306.5	-3.1	-0.11	
		Hotan	1968	2000	32	2487	5131.8	-0.7	-0.02	
		Yarkant	1974	2001	27	1421	3170.6	-6.1	-0.23	
		Pamir	1964	2001	37	880	2085.4	-7.9	-0.21	
		Tianshan	1964	2000	36	1249	4039.2	-1.3	-0.04	
		Kaidu	1964	2000	36	498	405.8	-7.1	-0.20	
13	China (Tibet)	Gongga Mountains	1966	1989	23	74	257.7	-5.8	-0.25	Pan et al. (2012b)
			1989	2009	20	75	242.8	-5.9	-0.29	
			1966	2009	43	76	228.5	-11.3	-0.26	
13	Kyrgistan	Pskem	1968	2000	32	525	219.8	-19.47	-0.61	Narama et al. (2010)
			2000	2007	7	525	177.0	-6.69	-0.96	
		Ili-Kungoy	1971	1999	28	735	672.2	-12.18	-0.44	
			1999	2007	8	735	590.3	-4.12	-0.52	
		At-Bashi	1968	2000	32	192	113.6	-12.06	-0.38	
			2000	2007	7	192	99.9	-4.20	-0.60	
			1968	2000	32	306	190.1	-9.21	-0.29	
2000	2007	7	306	172.6	-0.52	-0.07				
13	China	Qilian Mountains ^a	1956	2003	47	910	397.4	-21.7	-0.462	Wang et al. (2011)
13	China	Karlik Shan	1971	1992	21	n/a	126	-2.63	-0.13	Wang et al. (2009)
			1992	2001	9	n/a	122.7	-2.67	-0.27	
			1971	2001	30	n/a	119.4	5.2	-0.17	
13	China	Lenglongling ^a	1972	2007	35	179	86.2	-28.3	-0.81	Pan et al. (2012a)
13	Kyrgistan	Akshirak	1977	2003	26	178	406.8	-8.6	-0.33	Aizen et al. (2007)
			1981	2003	22	48	40.62	-10.6	-0.48	
14	Himalaya	Ten basins mean ^a	1962	2004	42	1868	6332	-15.8	-0.38	Kulkarni et al. (2011)
14	India	Kang Yatze	1969	1991	23	121	96.4	-13.0	-0.56	Schmidt and Nusser (2012)
			1991	2010	18	121	83.9	-1.5	-0.09	
			1969	2010	41	121	96.4	-14.4	-0.35	
14	India	Gharwal Himalaya ^a	1968	2006	38	82	600	-4.6	-0.121	Bhambri et al. (2011)
15	Nepal	Khumbu Himal	1976	2006	30	n/a	3211.9	-15.6	-0.52	Nie et al. (2010)
15	Nepal	Khumbu Himal	1962	2005	43	n/a	92.3	-5.3	-0.123	Bolch et al. (2008)
15	Nepal	Sagarmatha National Park	1962	2001	39	29	403.9	-4.9	-0.126	Salerno et al. (2008)

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Table 4.SM.1 (continued)

RGI Region	Country / Region	Sub-region / Mountain Range	Start Year	End Year	Number of Years	Glacier Count	Area Covered (km ²)	Relative Change (%)	Change Rate (% a ⁻¹)	Reference
16	Peru	Cordillera Coropuna	1955	2003	48	711	123	-54	-1.125	Silverio and Jaquet (2012)Peru
16	Peru	Cordillera Blanca	1970	1990	20	n/a	190	-12.8	-0.64	Baraer et al. (2012)
			1990	2009	19		165	-17.4	-0.92	
			1970	2009	39		136.3	-28.0	-0.72	
16	Peru	Cordillera Vilcanota	1985	1996	11	n/a	444	-22.5	-2.05	Salzmann et al. (2013)
			1996	2006	10		344	-13.7	-1.37	
			1985	2006	21		297	-33.2	-1.58	
16	Peru	Quehcaya Ice Cap	1985	2000	15	n/a	55.7	-17.6	-1.17	Salzmann et al. (2013)
			2000	2009	9		45.9	-3.1	-0.34	
			1985	2009	24		42.8	-23.1	-0.96	
16	Indonesia	Puncack Jaya	1942	1972	30	10	9.9	-30.3	-1.01	Klein and Kincaid (2006)
			1972	2002	30	5	6.9	-66.2	-2.23	
			1942	2002	60		2.15	-78.3	-1.30	
16	Columbia	Six mountain ranges ^a	1959	1987	28	n/a	106.8	-21.9	-0.78	Ceballos et al. (2006)
			1987	2002	15		83.5	-33.5	-2.23	
			1959	2002	43		45.6	-48.1	-1.18	
16	Peru	Cordillera Blanca	1970	2003	33	445	665.1	-22.4	-0.68	Racoviteanu et al. (2008)
16	Tansania	Kilimandscharo ^a	1962	2011	49	n/a	7.32	-76.0	-1.55	Cullen et al. (2013)
17	Chile	Gran Campo Nevado	1942	2002	60	81	252.6	-14.4	-0.24	Schneider et al. (2007)
17	Chile / Argentina	San Lorenzo Mountains	1985	2000	15	213	239.0	-9.9	-0.66	Falaschi et al. (2013)
			2000	2008	8	213	215.4	-9.7	-1.21	
			1985	2008	23	213	206.9	-13.4	-0.58	
17	Chile / Argentina	Patagonia	1986	2001	15	183	23743	-2.2	-0.14	Davies and Glasser (2012)
			2001	2011	10	165	23229	-2.2	-0.22	
			1986	2011	25	183	22717	-4.3	-0.17	
17	Chile	Northern Patagonia Icefield	1979	2001	22	>70 ^b	4093	-3.4	-0.15	Rivera et al. (2007)
17	Chile	Aconcagua Basin	1955	2003	48		151	-19.9	-0.41	Bown et al. (2008)
18	New Zealand	Southern Alps	1978	2002	24	n/a	513	-16.6	-0.69	Gjermundsen et al. (2011)
19	Antarctica	Kerguelen Island ^a	1963	2001	38	n/a	703	-21	-0.55	Berthier et al. (2009)
19	Antarctica	King George Island	1956	1995	39	n/a	1250	-7.0	-0.179	Rückamp et al. (2011)
			2000	2008	8	n/a		-1.6	-0.20	

Notes:

(a) More detailed analyses (e.g., sub-regions, other periods) are available in the respective papers.

(b) Glaciers <0.5 km² were not counted separately.

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