



A Double Seismic Array Experiment on Mt. Etna

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Abstract

On September 1999 two seismic antennas (array) and a profile of 3-D stations equipped with short period seismometers were installed on Mt. Etna; Aims of the experiment were to investigate the structure and the polarisation parameters of the volcanic tremor wavefield radiated in eruptive conditions, and to measure the seismic velocities and attenuation in the shallow structure beneath the arrays. The first array was installed close to Pizzi Deneri Volcanological Observatory, East of the crater area; the second array was located close to Torre del Filosofo site, south of the crater area. The profile was set up close to the Pizzi Deneri array. It had a length of 600 meters and consisted of 16 short period 3-component stations. In addition, a circular array of 8 short period seismometers was set up around the main crater area (Fig. 1). Several tremor samples and explosion quakes were recorded in 10 days of operation .

Spectrograms show predominant energy into 1-4 Hz frequency band. Slowness spectra indicate a predominant source of the tremor located at surface and coincident with the crater area. Polarisation analysis show that the direction of predominant motion is transverse with respect to the direction pointing to the active craters.

Introduction

Seismic arrays have been widely used to study the wavefield associated to volcanic tremor and explosion quakes (*Goldstein & Chouet, 1994; Ferrazzini et al., 1991; Chouet et al., 1997; Del Pezzo et al., 1997*). They are essential tools to track the seismic source of the volcanic signals as Long Period events and volcanic tremor. As the insurgence of this kind of events is very often related to the occurrence of volcanic eruptions, the real-time multichannel analysis applied to array data assumes increasing importance in volcano monitoring.

The knowledge of the wavefield structure of the tremor is the first goal to achieve together with the knowledge of the elastic properties of the shallow layers through which seismic waves propagate. For this reason, a series of experiments was conducted over the last 6 years in the framework of an international co-operation and with the main financial support of Italian GNV and Spanish Antarctic Project. These experiments were aimed at a testing of array techniques and on investigating the properties of volcanic signals (tremor and explosion quakes) on several volcanoes like *Stromboli (O.V. Open File Report n. 1-98)*, *Vesuvius (O.V.*

Open File Report n. 1-99), Deception Island (*O.V. Open File Report n. 3-99 and n. 1 2000*), Teide (*Almendros et al., 2000*) and the present one at Mt. Etna.

The present report provides information about the array set up and operation, instrument calibration and data format.

Site selection

A preliminary survey was carried out on the end of May, 1999, in order to select the sites for the installation of the two seismic antennas and the profile. We searched two flat and geologically homogeneous zones, about 300 meters wide, positioned at about 90° apart with respect to the position of the source, which was assumed to be located in close proximity of the crater area (fig.1). Using this geometry, the estimate of the spatial location of the source is optimised, as shown in *O.V open file report n. 1-98 and in Saccorotti et al.,1998*. The area around Pizzi Deneri Volcanological Observatory (area1) and the area around Torre del Filosofo site (area 2) resulted the most suitable for the installation. Area 1 was selected also for the deployment of the seismic profile, due to the easier accessibility.

Instrument positioning and installation. Array configuration.

1. Time schedule of the instrument installation

September 8. Installation of array P (fig. 2) at Pizzi dei Neri site

September 9 – September 10 . Installation of array abc+L (fig. 3) at Pizzi Deneri site

September 10 – September 11. Installation of array def (fig. 4)at Torre del Filosofo

September 16 – September 17. Installation of array abc+U(fig. 5) at Pizzi Deneri

September 17 – “Circular” (fig. 6) array around craters.

- #### 2. Co-ordinates of all the array stations except the array “Circular” were measured using a total EDM station and referred to an absolute GPS benchmark located nearby. Array “Circular” co-ordinates were obtained directly by GPS positioning. Uncertainties are less than 5 cm for the co-ordinates measured with the total station, and about 100 meters for the circular array station co-ordinates. In Table 1 is reported the list of the co-ordinates of all the array configurations.

Data collection and storage.

Array abc and def are composed by MARK L-4 C sensors cable-connected to 8 channel data loggers. The instruments were designed by Ramon Ortiz and G. Alguacil, and

described in *Del Pezzo et al. (1997)* where the original references (in spanish) can be found. Details about the electronic scheme are reported in *Ortiz et al. (1997)*. Seismic signals are amplified, sampled with a 16 bits A/D converter and anti-alias filtered before the storage onto the hard disk of a portable P.C. Time synchronisation is performed by a GPS antenna connected to the data logger.

Array P,U and “Circular” are equipped with 3-channel, MARS LITE data loggers, and 3-D, 4.5 electronically extended to 1 Hz sensors.

The time operation of the instruments is reported in Fig. 7 and 8.

Transfer function of the instruments.

Calibration of Mars Lite seismometers has been checked using a Teledune Geotech 3-D S13 seismometer as reference. A long noise sample has been recorded both at Mars-Lite seismometer and S-13 using identical data loggers. The ratio between the Fourier Transforms of the noise samples at homologue components is the ratio of the transfer functions of the sensors. S-13 (critically damped) has been calibrated through the access to the moving mass (see S-13 operation and maintenance manual) . Results show that S-13 and MARS-Lite seismometers have the same amplitude response at frequencies higher than 1 Hz, as shown in fig. 9 where the amplitude spectral ratio between the two vertical components is plotted as a function of frequency. Phase responses of the two instruments are compared in fig. 10. Velocity transduction results of 400 V/m/s, coincident with that reported in the Mars-Lite data sheet.

Mark L4 seismometers were damped at 65% critical damping (arrays abc and def). Amplitude and phase response curves are reported in fig. 11.

Description of the Etna eruptive activity during the experiment.

Etna volcano was active during the experiment . This eruptive phase started on January 1999 with a moderate activity at Bocca Nuova. Seven eruptive episodes (January 5 – February 4) occurred at S-E crater and terminated with the opening of a fracture at the cone base, directed toward S-E. Strombolian activity together with lava fountains accompanied a lava flow started on February 4, which rapidly reached Valle del Bove. The lava flow continued for the whole period of the experiment. On September 4 new activity started at Voragine crater, characterised by important lava fountains (information and maps on line at <http://193.204.162.114/gnv/index-etna.html>)

Purposes and Methods.

The experiment was aimed at a) the investigation of the kinematic properties of the volcanic tremor wavefield from frequency-slowness analysis; b) the constrain of the location of its source, c) the measurement of the seismic attenuation and d) the study of the polarisation properties of the tremor wavefield.

Slowness spectra can be calculated using time or frequency domain techniques. In time domain we use the so called Zero-Lag Cross-correlation technique (*Del Pezzo et al., 1997*); in frequency domain we use MUSIC algorithm (*Goldstein and Archuleta, 1987*).

The double array configuration will allow the space tracking of the source using the probabilistic method described in *Saccorotti et al., 1998*.

The Profile oriented toward the craters (L-array of fig. 3) will permit the estimate of the spatial attenuation of tremor energy, and hence the seismic quality factor of the earth medium.

Circular array will allow the study of the particle motion pattern of the tremor waves around the craters. Polarisation parameters of the waves composing the tremor wavefield will be fully investigated using the 3-D stations (*La Rocca, 2000*).

Application of the correlation method of *Aki (1957, 1965)* to both Pizzideneri and Torre Del Filosofo arrays will allow to derive the dispersive properties of Rayleigh and Love waves, from which a shallow velocity structure beneath the arrays may eventually be obtained.

Preliminary results

An example of data recorded at L005 station is shown in fig. 12.

The spectrogram of a sample of tremor recorded at array A shows the highest energy concentrated between 1 and 4 Hz, with some additional bursts between 4 and 6 Hz (fig. 13).

Preliminary frequency-slowness analysis carried out on a tremor sample recorded at array p, shows a predominant back-azimuth in the direction of the crater area (fig.14).

Polarisation analysis shows a predominant particle motion direction around 110 degrees, approximately normal to the back-azimuth. The incidence is around 90 degrees, showing an almost horizontal wave motion. Taken together with the high values of the rectilinearity coefficient, these observations are suggestive of a wavefield dominated by SH-Love (fig.15).

The correlation coefficients averaged over the array stations as a function of the frequency have been calculated for array abc for receiver spacing of 50, 100, and 150 m (Fig.

16). The smooth thin lines superimposed to the data represent the theoretical correlation coefficients derived from eq. 4 of *Chouet et al., (1998)*, using the dispersion curve shown in the last panel of the same figure. The procedure is described in *Chouet et al., 1998*. From this dispersion curve the velocity model can be inferred with an inversion procedure.

Future developments

We have planned to extend the above approaches to the whole data set in order to characterise the wavefield structure of the volcanic tremor at Etna volcano. Moreover we will extend the same methods to the explosion quakes which in many cases were superimposed to the background tremor radiation. The aim of this second analysis is twofold: first we could constrain the space position of the source of the explosion quake; second, we could compare the spectral characteristics of the volcanic tremor with those of the explosion quakes, in order to contribute to the understanding of the explosion quake source processes. *Ferrucci et al. (1990)*, analyse the polarisation properties of the tremor at Etna, using a 3 stations array located onto the south-western flanks of the volcano. They tentatively link the tremor source to the main dyke-system of Etna and exclude the confinement of the tremor source to the only active vent. A check of this result may be obtained by the analysis of the present data, and particularly from the array “circular”, which would better enlighten the polarisation pattern of the tremor around the crater area.

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Figure Captions

Fig. 1 . Sketch map of the Etna area around craters. In the zones of Torre del Filosofo and Pizzi Deneri the array configurations are plotted together. In particular at Pizzi Deneri we overplot 3 different configurations, used separately in different time periods. Note that the array represented by black triangles, (see details in the next Fig. 6) , is in reality composed by 8 stations: the northernmost and southernmost stations are not represented by black triangles as their symbol would have covered part of the semicircular array configurations.

Fig. 2 . Sketch map of array P at Pizzi Deneri site. Mars Lite stations were used in this configuration.

Fig. 3 . Array abc and L (profile) at Pizzi Deneri site. L array was made by Mars Lite. abc is composed by 3 8-channels data loggers, connected to MARK L4 sensors (see text).

Fig. 4 . Array def, at Torre del Filosofo site. Configuration is similar to and instruments are the same of array abc.

Fig. 5 . Array abc and U. Mars lite stations were deployed in a tripartite profile configuration. This configuration was called MEDUSA.

Fig. 6 . Array "circular". This configuration was obtained using Mars-Lite stations.

Fig. 7 . Time table of the overall operation for array abc and def

Fig. 8 . Time table for the overall operation of a) array P, b) array L, c) array R and d) array U.

Fig. 9 . Spectral ratio between long duration noise samples recorded at the same site, using the same data logger connected to a S13 couple (1 vertical and 1 horizontal) and to a 3-D Mars-lite seismometer.

Fig. 10 . Phase response of Mars lite and vertical S13, reported for comparison.

Fig. 11 Phase and amplitude response for the overall instrument chain used for abc and def arrays.

Fig. 12 . Example of 4-hours continuous data recording. Filtering between 0.1 and 10 Hz is applied in order to reduce the high frequency noise. Vertical lines are the minute marks.

Fig. 13 . Spectrogram of a tremor sample. The spectral power is normalised to the maximum. The grey scale evidences along the sample a maximum power between 1.5 and 4.5 Hz.

Fig. 14 Slowness spectrum obtained using MUSIC method applied to a tremor sample recorded at array abc. Power is reproduced by the grey scale. Maximum power is compatible with the direction of craters.

Fig. 15 . Polarisation parameters as a function of lapse time along the same tremor sample of figure 13. First panel from the top shows the particle motion dominant direction. The second represents the dominant incidence angle. The third reports Rectilinearity and the largest eigenvalue of the co-variance matrix. All quantities are plotted as a function of the lapse time along the sample.

Fig 16 . Correlation coefficient averaged over azimuth for different station spacing. This data permitted to estimate the dispersion curve, shown in the panel below, from which a velocity model can be obtained using an inversion scheme.

Table 1

Station	lat. deg.	lat. min.	lat. sec.	lon. deg.	lon. min.	lon. sec	Cartesian Y axis (km)	Cartesian X axis(km)
u05	37	45	53,8193	15	0	54,6124	4180,58	1336,53
u04	37	45	54,8492	15	0	53,0056	4180,61	1336,49
u03	37	45	55,8518	15	0	51,4092	4180,64	1336,45
u02	37	45	56,868	15	0	49,8093	4180,67	1336,41
u01	37	45	58,049	15	0	47,9244	4180,71	1336,36
u010	37	45	59,6008	15	0	54,5815	4180,76	1336,53
u09	37	45	59,4498	15	0	52,9472	4180,75	1336,49
u08	37	45	59,0378	15	0	51,2238	4180,74	1336,45
u07	37	45	59,1614	15	0	49,6994	4180,74	1336,41
u06	37	45	59,0241	15	0	48,2369	4180,74	1336,37
za	37	45	58,873	15	0	46,6129	4180,74	1336,33
4a	37	45	59,1751	15	0	44,6045	4180,74	1336,28
zb	37	46	0,39734	15	0	45,9332	4180,78	1336,32
4c	37	45	57,28	15	0	46,2387	4180,69	1336,32
5c	37	45	55,7007	15	0	45,8714	4180,64	1336,31
6c	37	45	54,0939	15	0	45,4937	4180,59	1336,3
zc	37	45	57,7606	15	0	45,1263	4180,7	1336,3
7c	37	45	56,6345	15	0	43,6466	4180,67	1336,26
8c	37	45	55,5222	15	0	42,1635	4180,63	1336,22
5a	37	45	58,461	15	0	42,5686	4180,72	1336,23
6a	37	45	58,2413	15	0	40,5396	4180,72	1336,18
7a	37	46	0,4248	15	0	43,0115	4180,78	1336,24
8a	37	46	1,20758	15	0	41,2296	4180,81	1336,2
7b	37	46	1,92169	15	0	45,2499	4180,83	1336,3
8b	37	46	3,47351	15	0	44,5599	4180,88	1336,28
4b	37	46	0,47974	15	0	46,9906	4180,78	1336,34
5b	37	46	2,07275	15	0	47,3579	4180,83	1336,35
6b	37	46	3,6795	15	0	47,7322	4180,88	1336,36
L001	37	46	3,95416	15	0	51,6975	4180,89	1336,46
L002	37	46	2,93793	15	0	50,6847	4180,86	1336,43
L003	37	46	1,92169	15	0	49,6616	4180,83	1336,41
L004	37	46	0,90546	15	0	48,6488	4180,8	1336,38
L005	37	45	59,903	15	0	47,6292	4180,77	1336,36
L006	37	45	58,873	15	0	46,6129	4180,74	1336,33
L007	37	45	57,8568	15	0	45,6001	4180,7	1336,31
L008	37	45	56,8543	15	0	44,5805	4180,67	1336,28
L009	37	45	55,8243	15	0	43,5642	4180,64	1336,26
L010	37	45	54,8218	15	0	42,5446	4180,61	1336,23
L011	37	45	53,8055	15	0	41,5318	4180,58	1336,21
L012	37	45	52,7756	15	0	40,5087	4180,55	1336,18
L013	37	45	51,7731	15	0	39,4959	4180,52	1336,16
L014	37	45	50,7568	15	0	38,4762	4180,49	1336,13
L015	37	45	49,7406	15	0	37,46	4180,45	1336,11
L016	37	45	48,7244	15	0	36,4403	4180,42	1336,08
zd	37	43	57,6526	15	0	29,7455	4177,01	1335,92
6f	37	43	59,8224	15	0	24,4103	4177,07	1335,78
5f	37	43	59,0945	15	0	26,1887	4177,05	1335,83
4f	37	43	58,3804	15	0	27,9568	4177,03	1335,87

2f	37	43	59,3417	15	0	27,6203	4177,06	1335,86
3f	37	44	0,4541	15	0	26,2127	4177,09	1335,83
8f	37	44	1,89606	15	0	27,0298	4177,14	1335,85
7f	37	44	0,4953	15	0	27,919	4177,1	1335,87
zf	37	43	59,0808	15	0	28,8288	4177,05	1335,89
2d	37	44	1,60767	15	0	28,9593	4177,13	1335,9
6d	37	44	2,41791	15	0	30,6931	4177,15	1335,94
5d	37	44	0,82489	15	0	30,3806	4177,11	1335,93
4d	37	43	59,1083	15	0	30,669	4177,05	1335,94
7d	37	43	59,9597	15	0	32,6363	4177,08	1335,99
8d	37	44	1,08582	15	0	34,0679	4177,11	1336,02
3d	37	43	59,4928	15	0	34,2773	4177,06	1336,03
8ee	37	43	57,9135	15	0	35,8704	4177,02	1336,07
7ee	37	43	57,8311	15	0	33,8173	4177,01	1336,02
ze	37	43	57,7487	15	0	31,7608	4177,01	1335,97
4ee	37	43	56,9247	15	0	31,572	4176,99	1335,96
5ee	37	43	56,2106	15	0	33,3435	4176,96	1336
6ee	37	43	55,4553	15	0	35,1872	4176,94	1336,05
3ee	37	43	57,0346	15	0	34,7443	4176,99	1336,04
p101	37	45	58,873	15	0	46,6129	4180,74	1336,33
p102	37	45	57,8293	15	0	45,5692	4180,7	1336,31
p103	37	45	56,8131	15	0	44,5118	4180,67	1336,28
p104	37	45	55,7968	15	0	43,4681	4180,64	1336,25
p107	37	45	57,3898	15	0	42,0399	4180,69	1336,22
p106	37	45	57,8842	15	0	43,5642	4180,7	1336,26
p105	37	45	58,3786	15	0	45,0817	4180,72	1336,29
p108	37	45	58,7769	15	0	44,9856	4180,73	1336,29
p109	37	45	58,6807	15	0	43,3376	4180,73	1336,25
p110	37	45	58,6121	15	0	41,6966	4180,73	1336,21
p113	37	46	0,15015	15	0	41,9128	4180,77	1336,22
p112	37	45	59,7107	15	0	43,4784	4180,76	1336,26
p111	37	45	59,2987	15	0	45,0577	4180,75	1336,29
p114	37	45	59,7794	15	0	45,4216	4180,76	1336,3
p115	37	46	0,69946	15	0	44,2165	4180,79	1336,27
p116	37	46	1,60584	15	0	43,0252	4180,82	1336,24
R00A	37	45,96	0	15	0,84	0	4180,7	1336,43
R001	37	45,87	0	15	0,035	0	4180,53	1335,23
R002	37	45,8	0	14	59,365	0	4180,4	1334,24
R003	37	45,564	0	14	59,14	0	4179,97	1333,9
R004	37	45,12	0	14	59,13	0	4179,15	1333,89
R005	37	44,64	0	14	59,18	0	4178,26	1333,96
R006	37	44,38	0	14	59,75	0	4177,78	1334,81
R007	37	44,015	0	15	0,43	0	4177,11	1335,82
CRATER POSITION	37	45	0	14	59	30	4178,93	1334,44

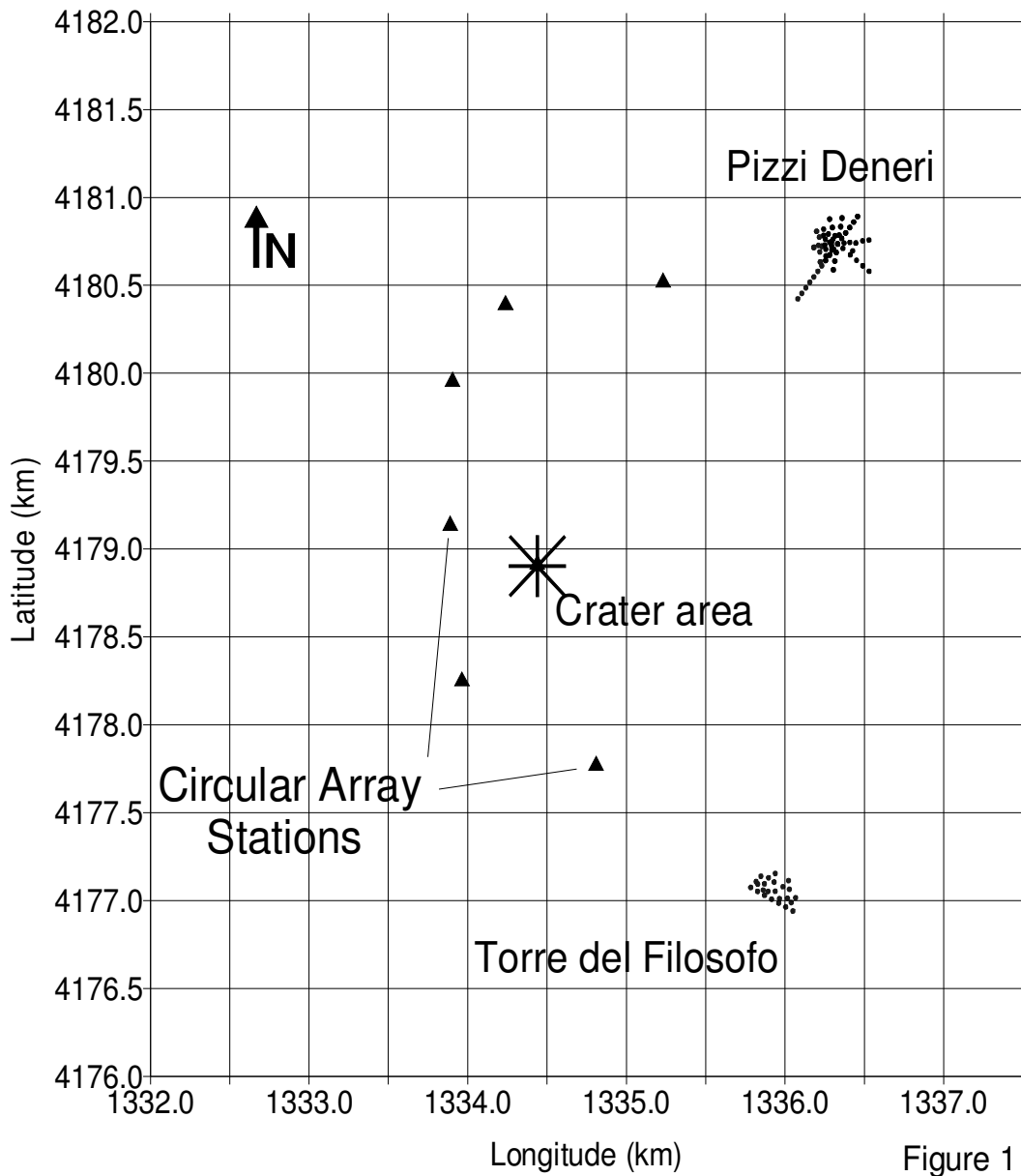


Figure 1

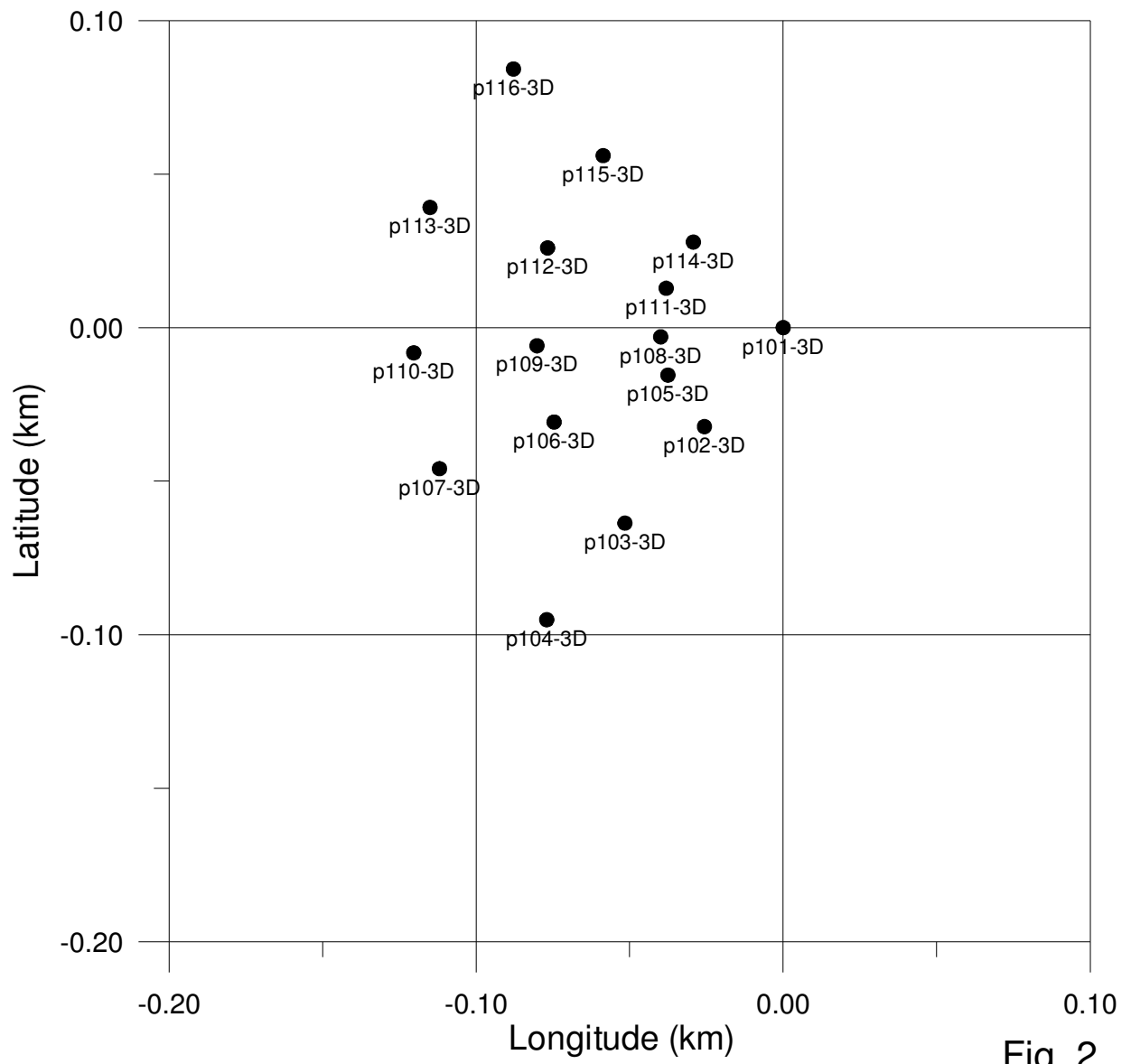


Fig. 2

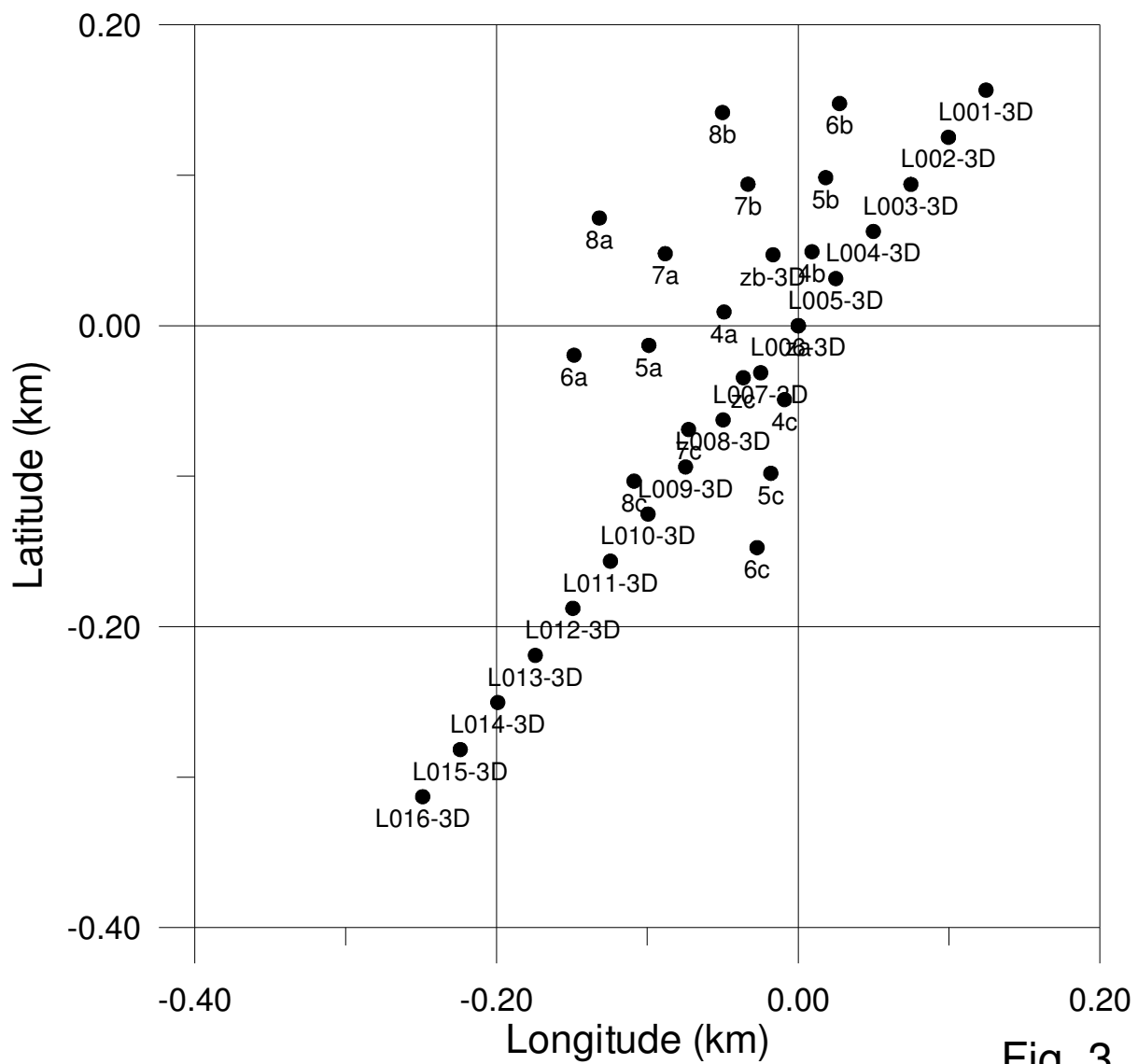


Fig. 3

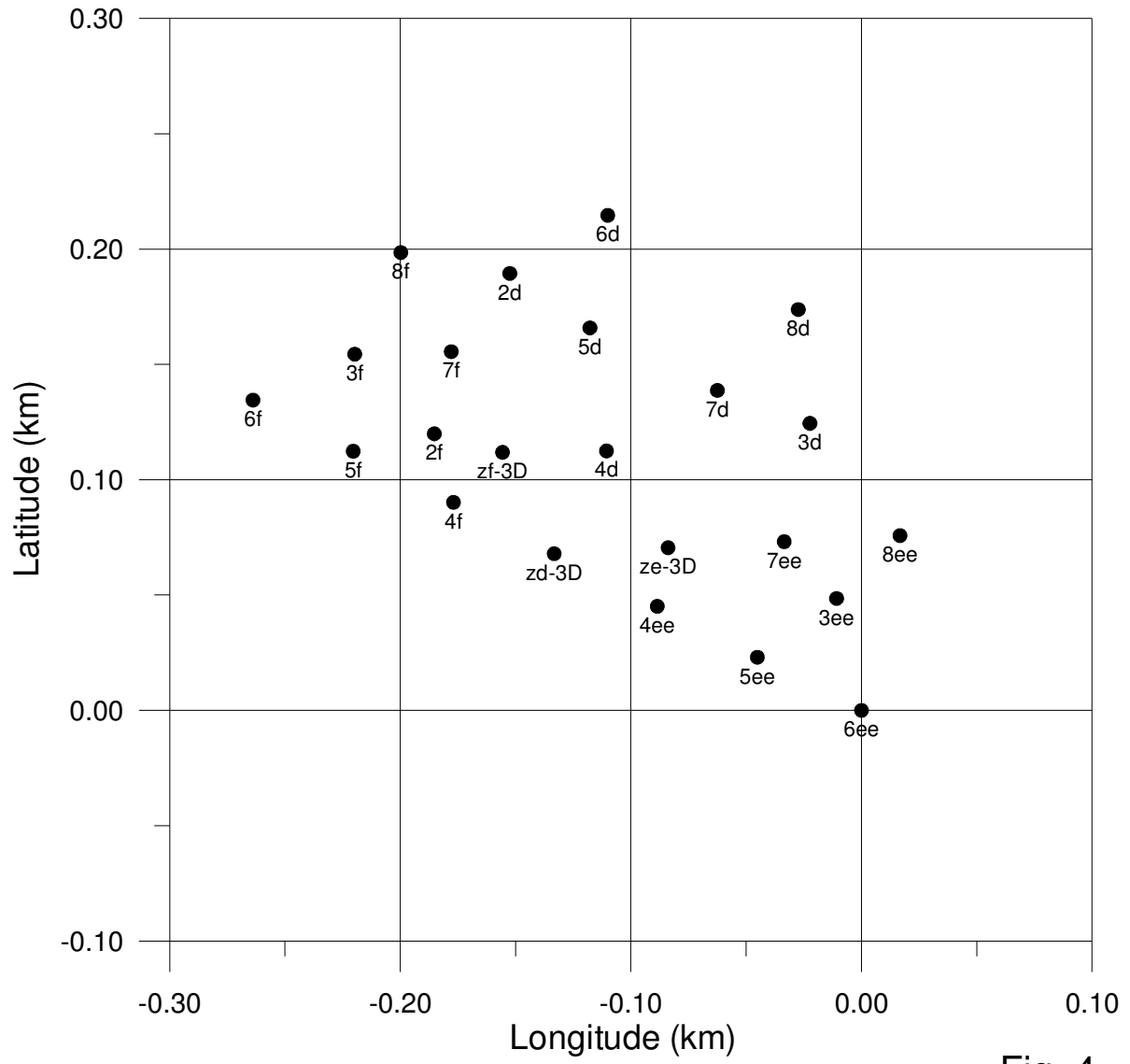


Fig. 4

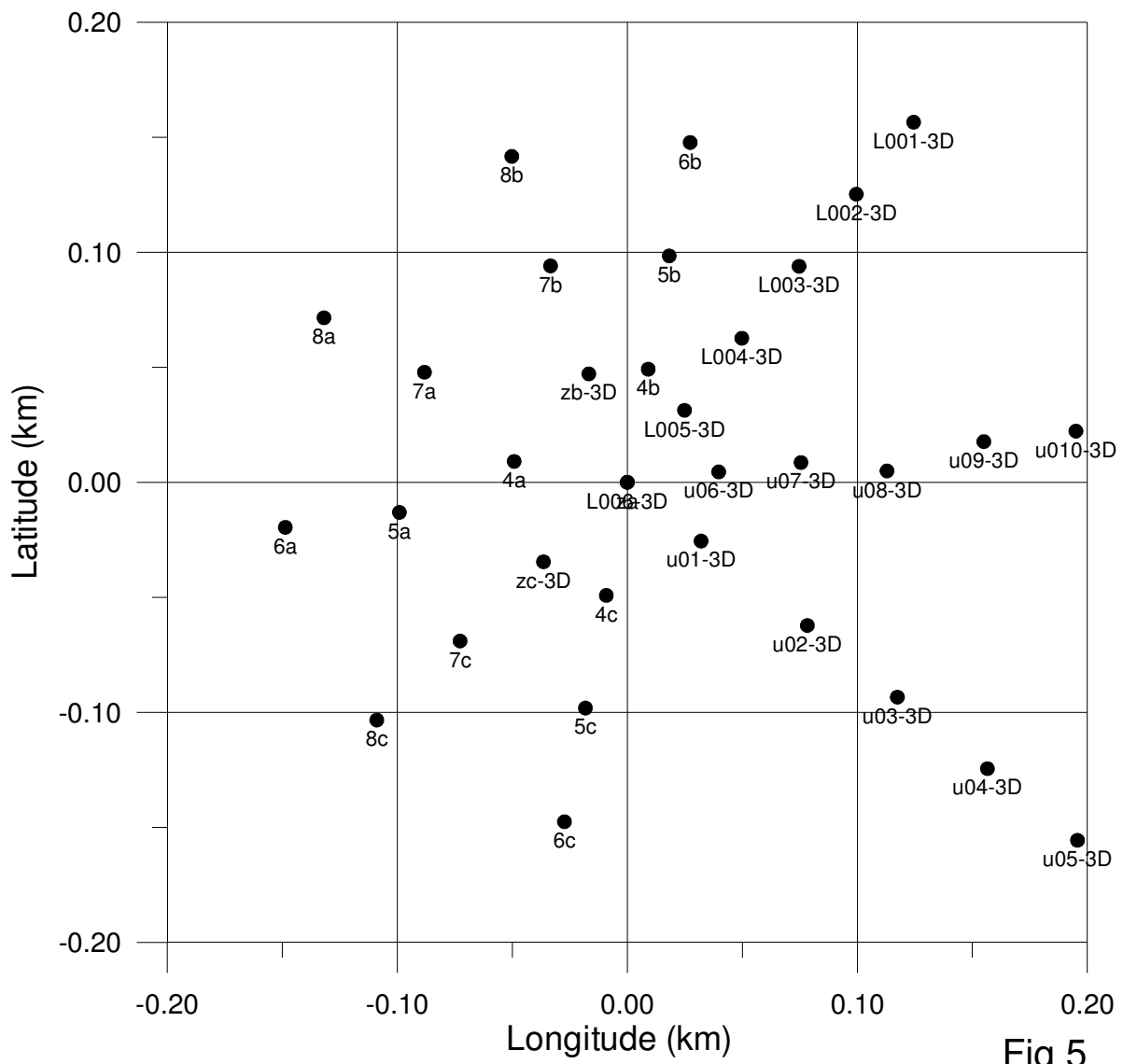


Fig 5

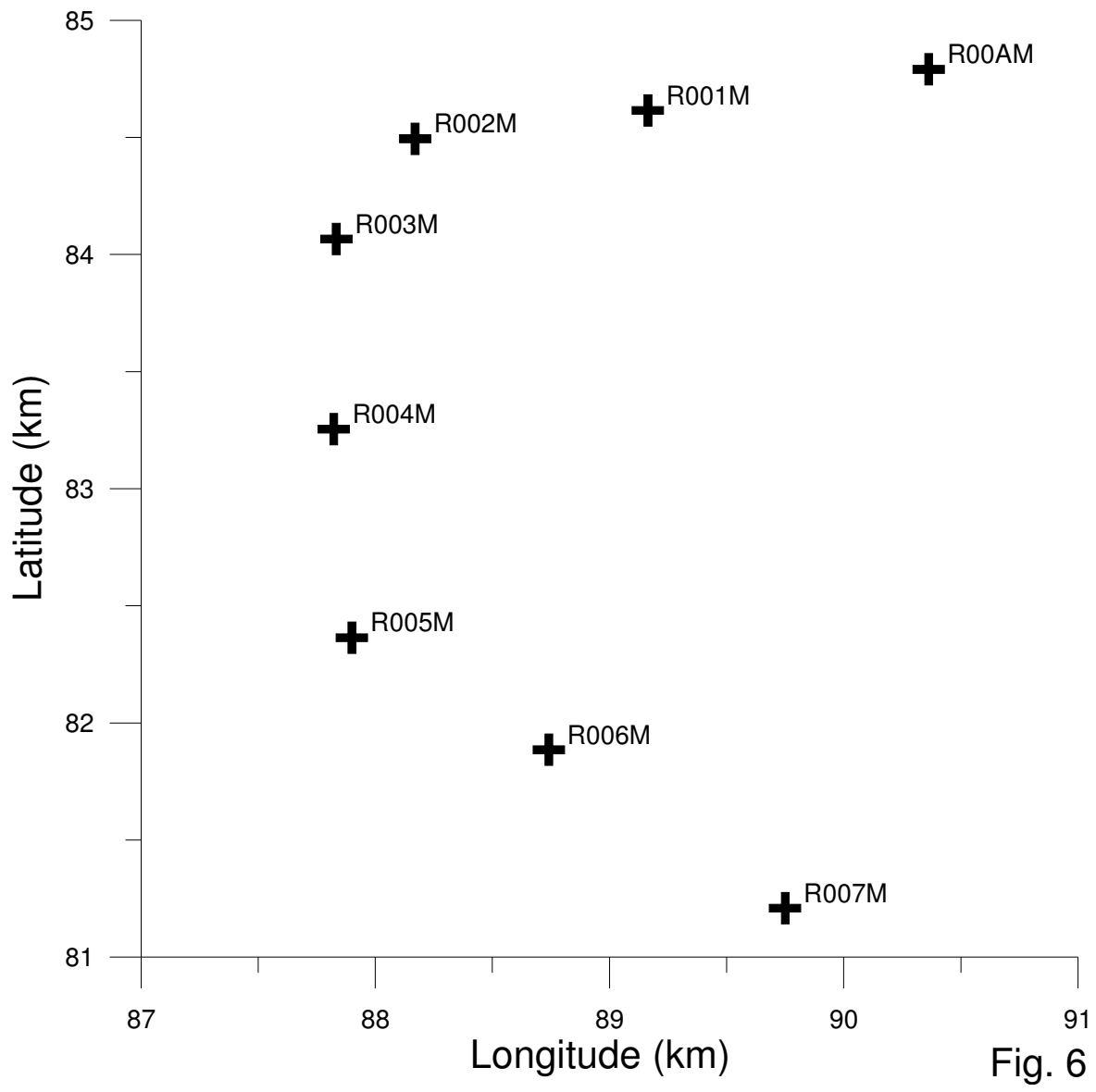


Fig. 6

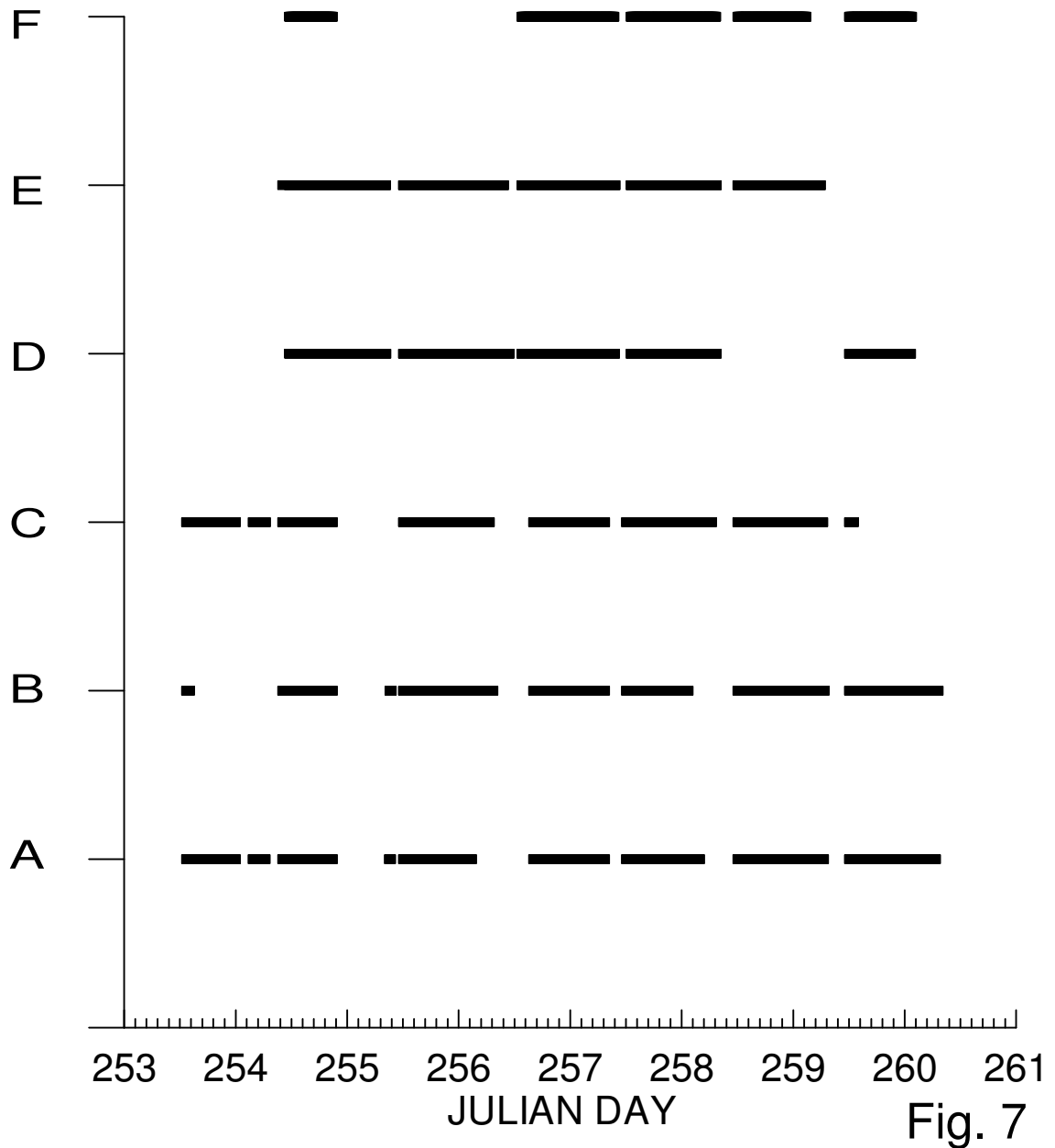


Fig. 7

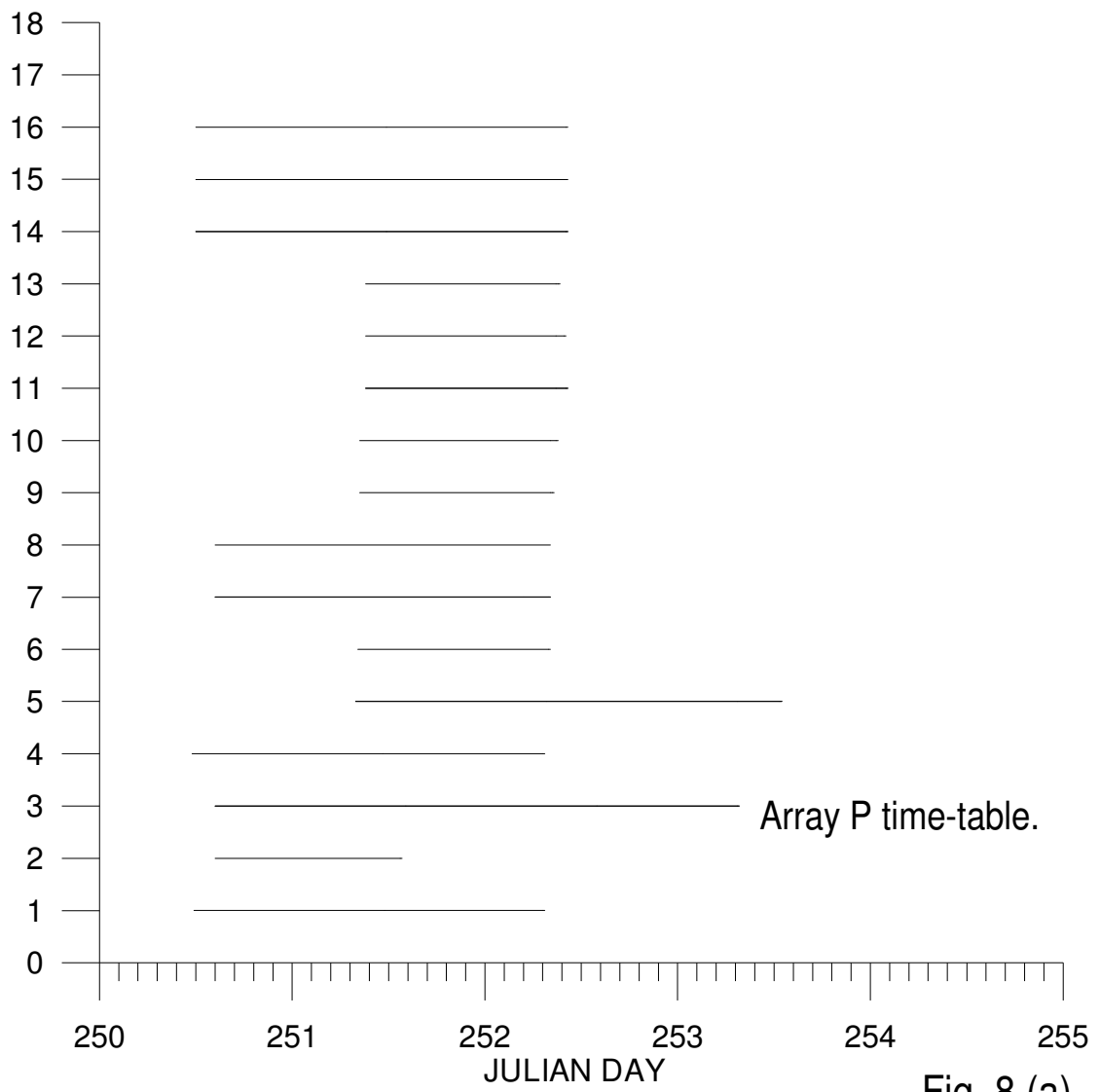


Fig. 8 (a)

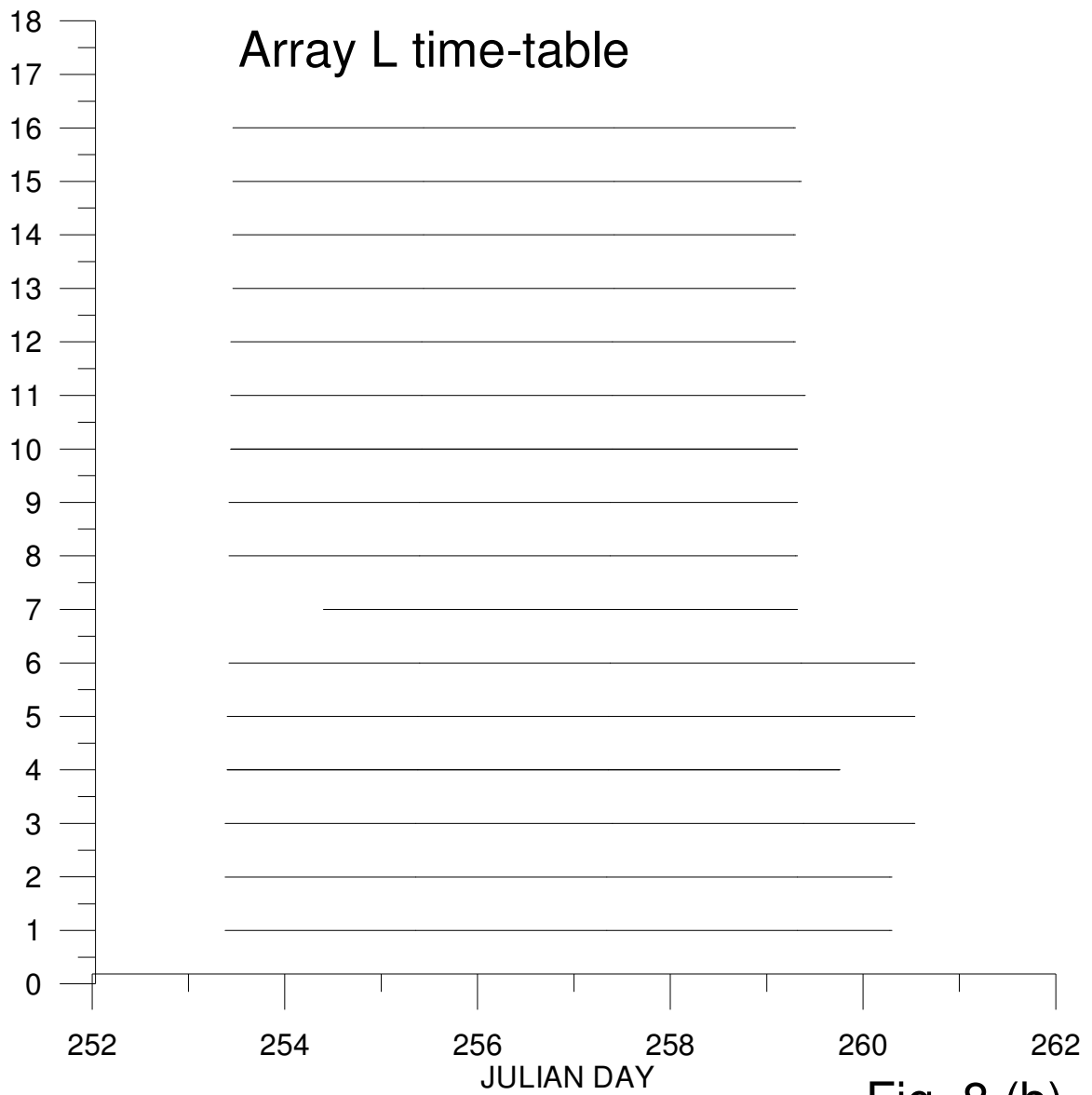


Fig. 8 (b)

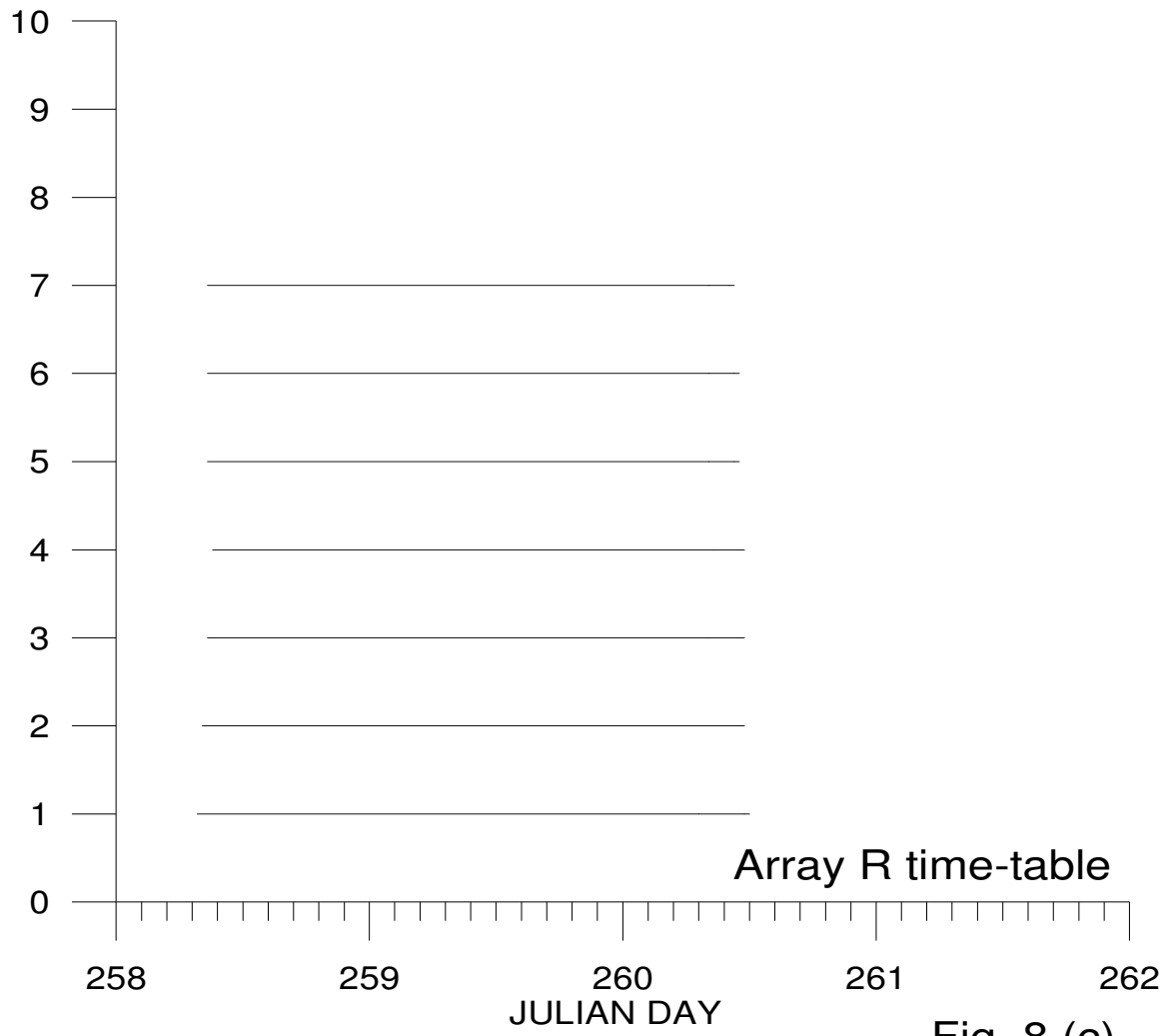


Fig. 8 (c)

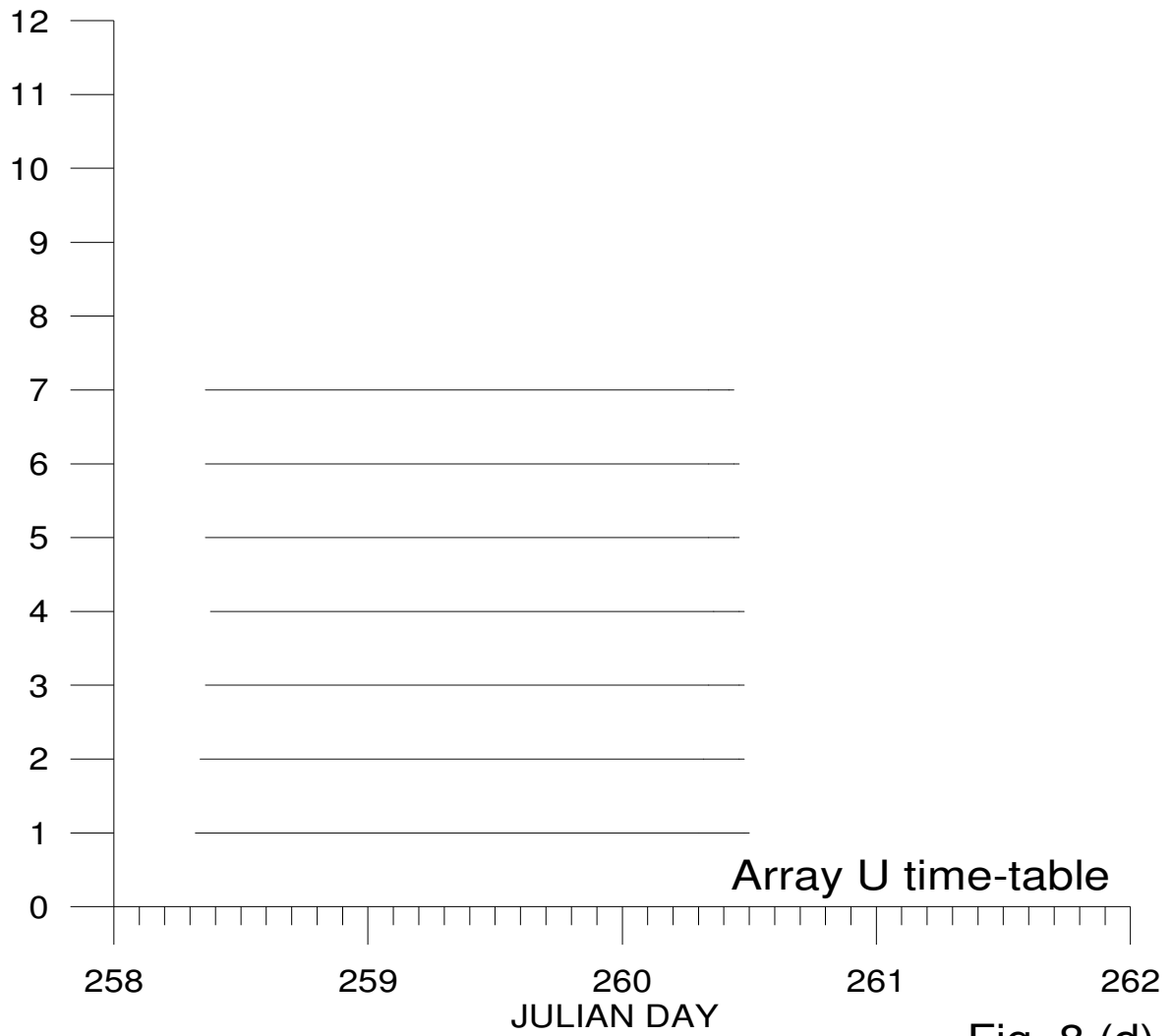


Fig. 8 (d)

Spectral Ratio between S13 and Mars-Lite
(Vertical component)

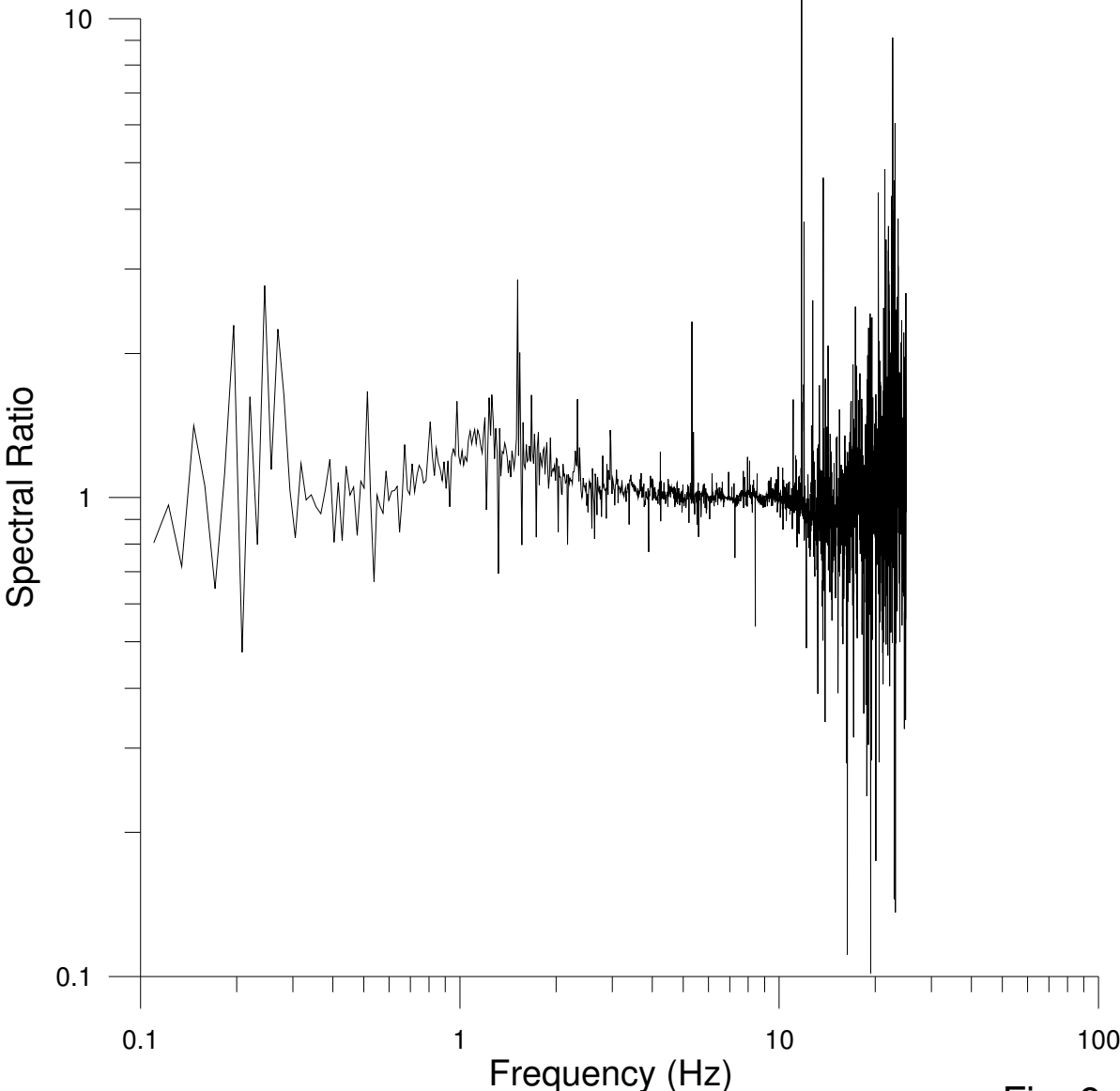


Fig. 9

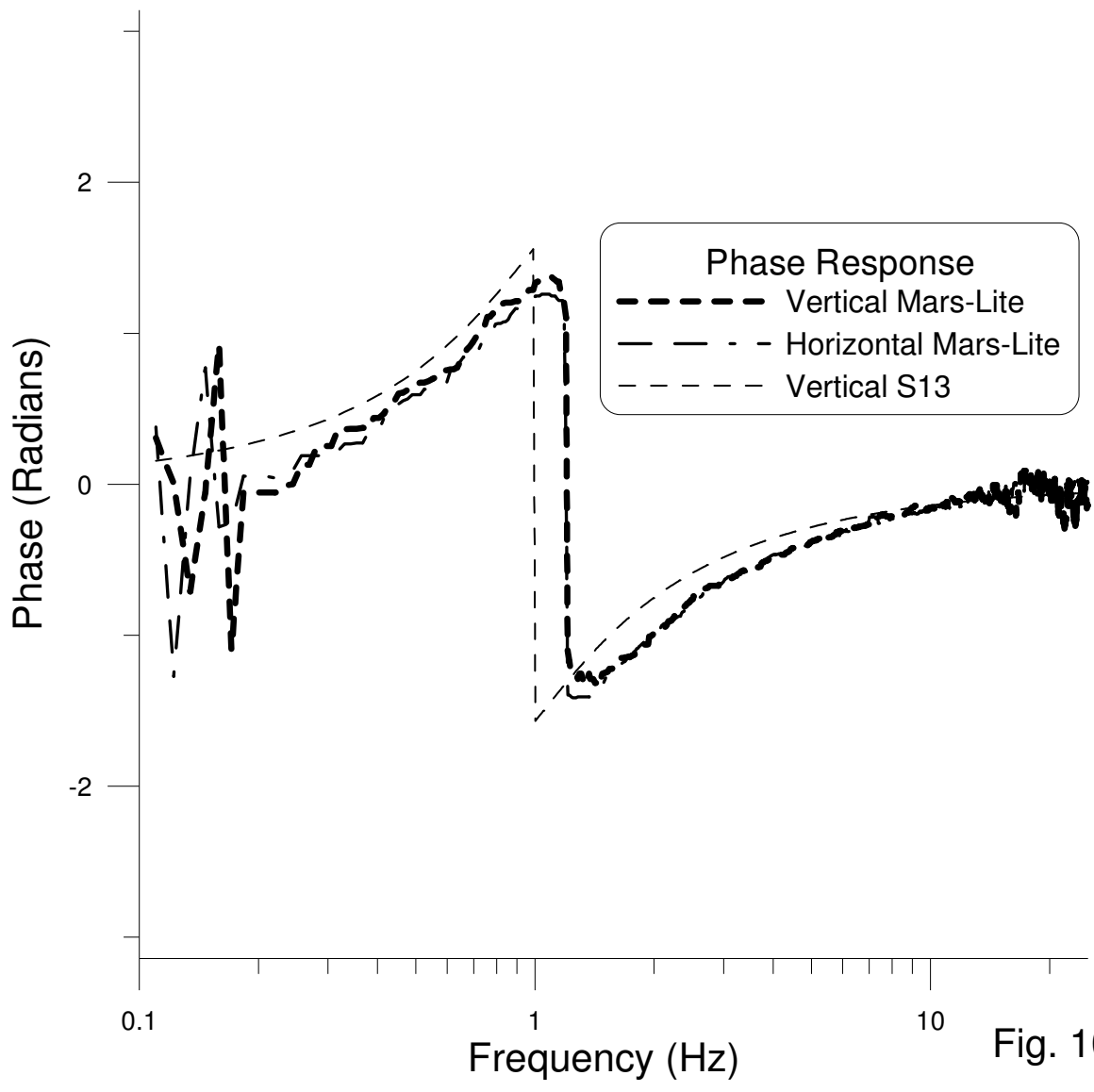


Fig. 10

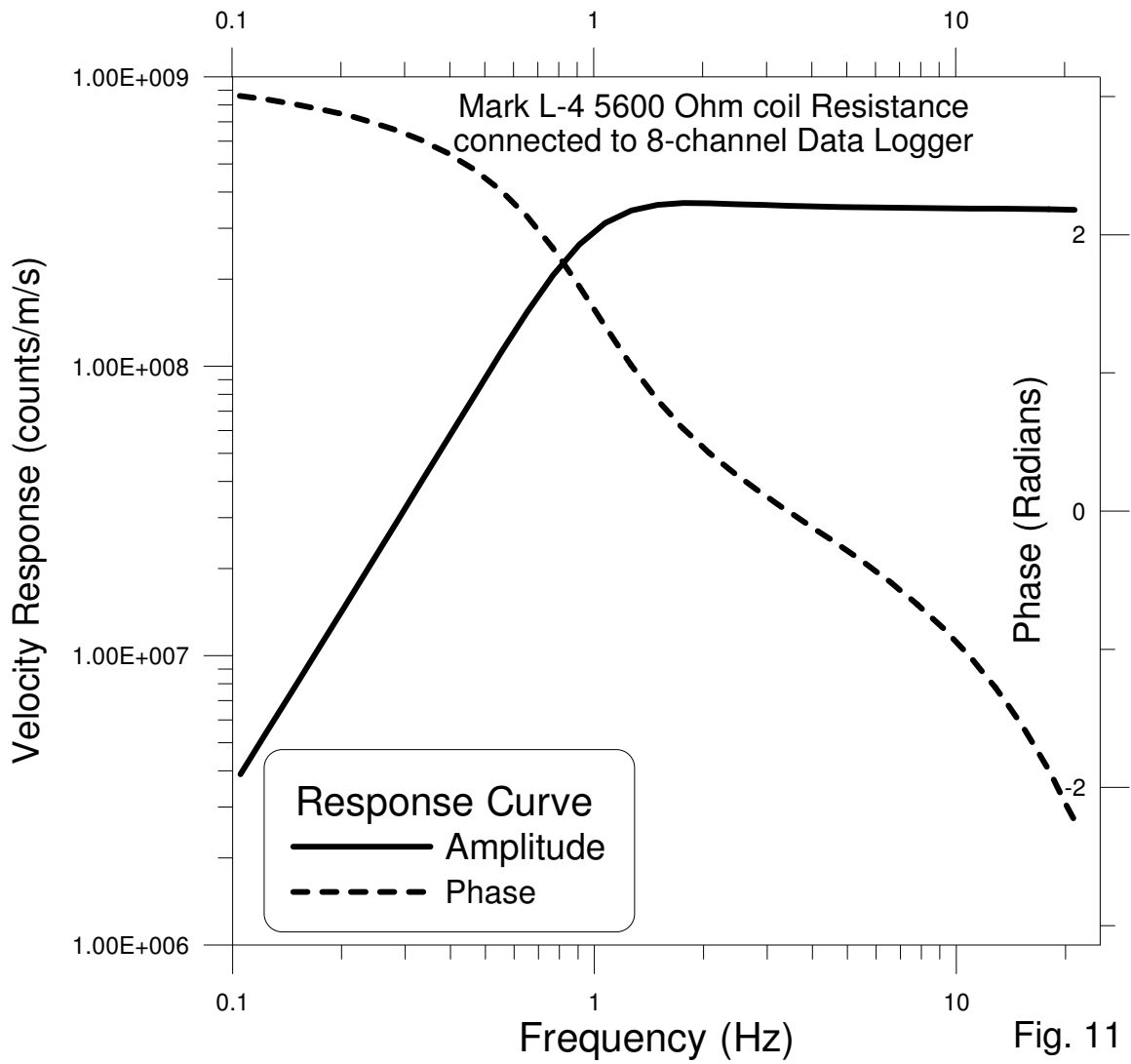


Fig. 11

09/13/1999 12:25 Band Pass Filter 0.1 – 10 Hz

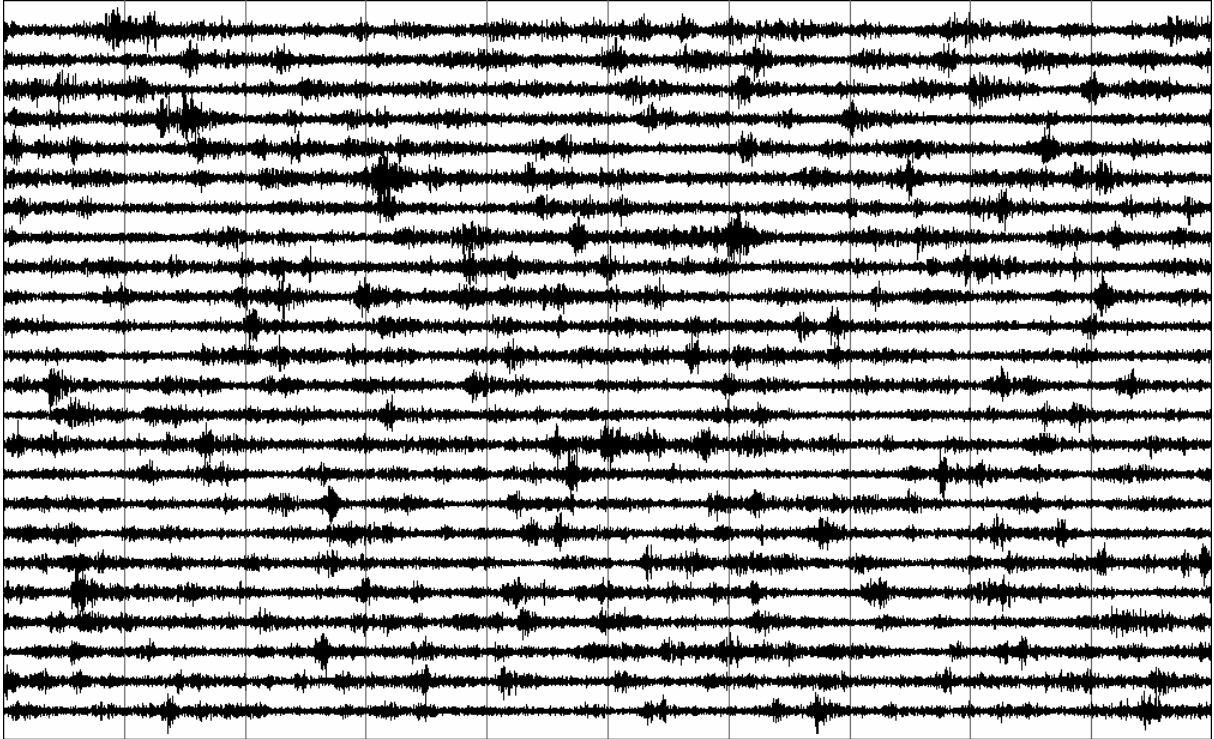


Fig. 12

ETNA 1999 - ARRAY PIZZI DENERI

10-09-99 13:49:45.00 GMT VERTICAL COMPONENT

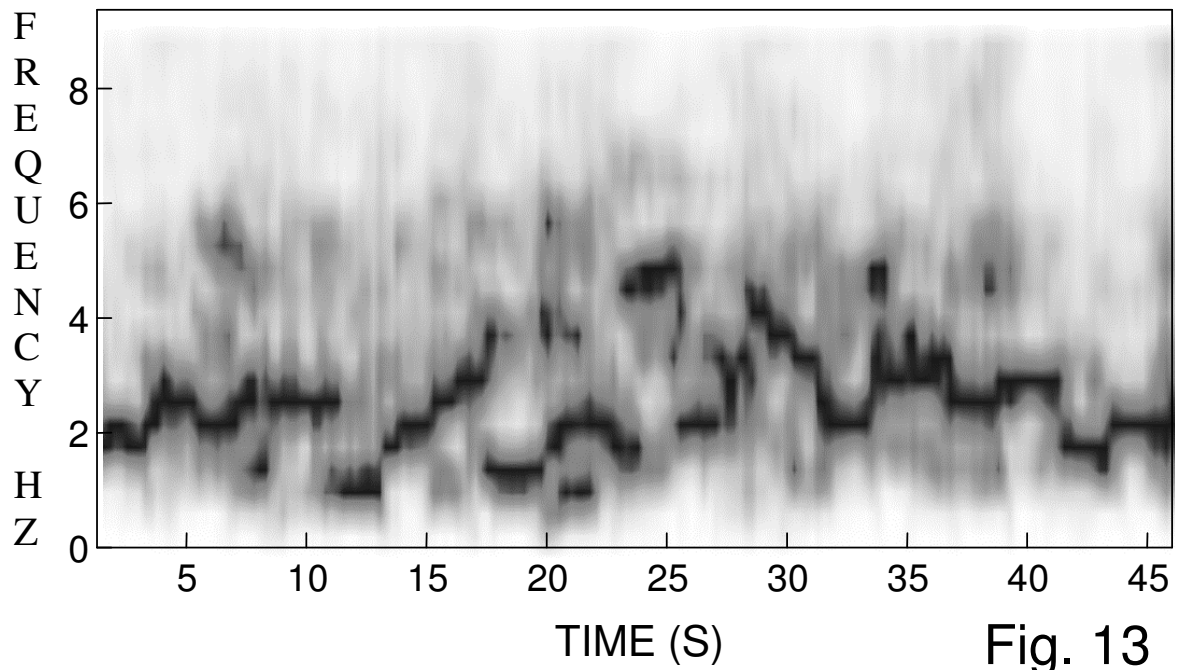
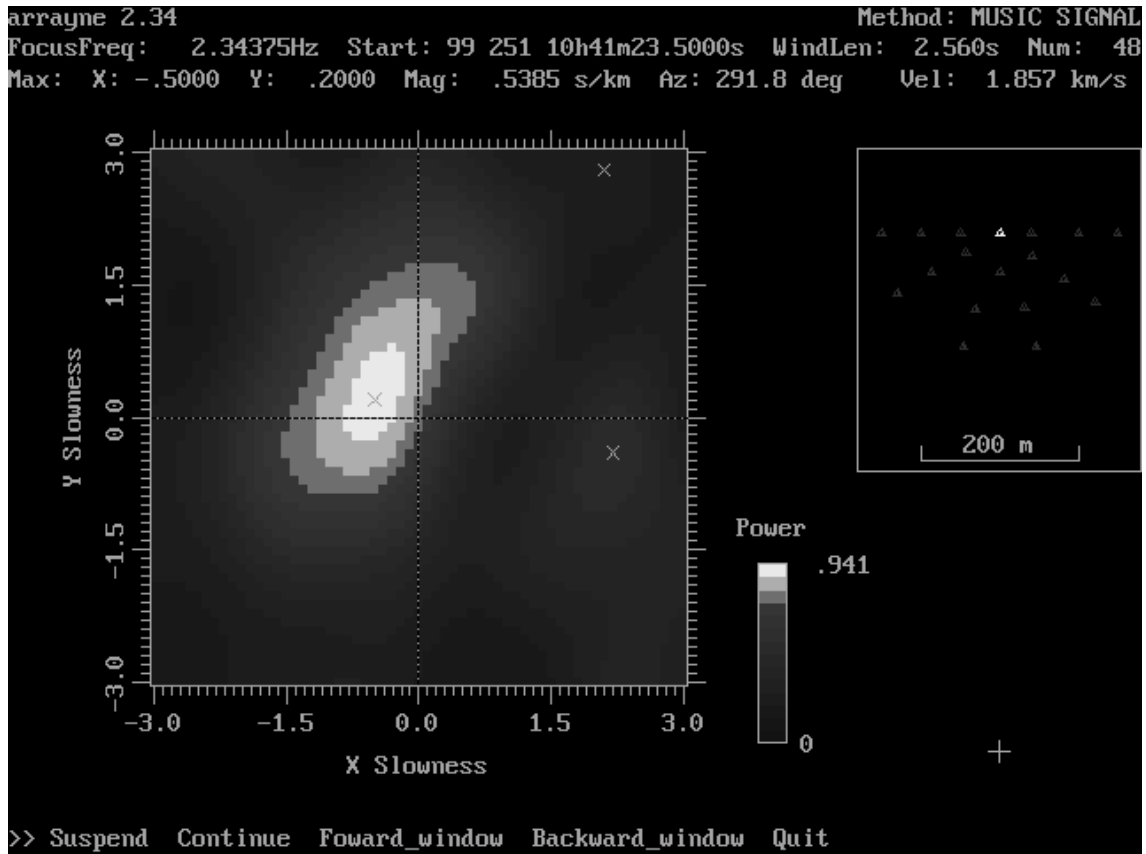


Fig. 13

Fig. 14



ETNA 1999 - ARRAY PIZZI DENERI

10-09-99 13:49:45.00 GMT ARRAY A

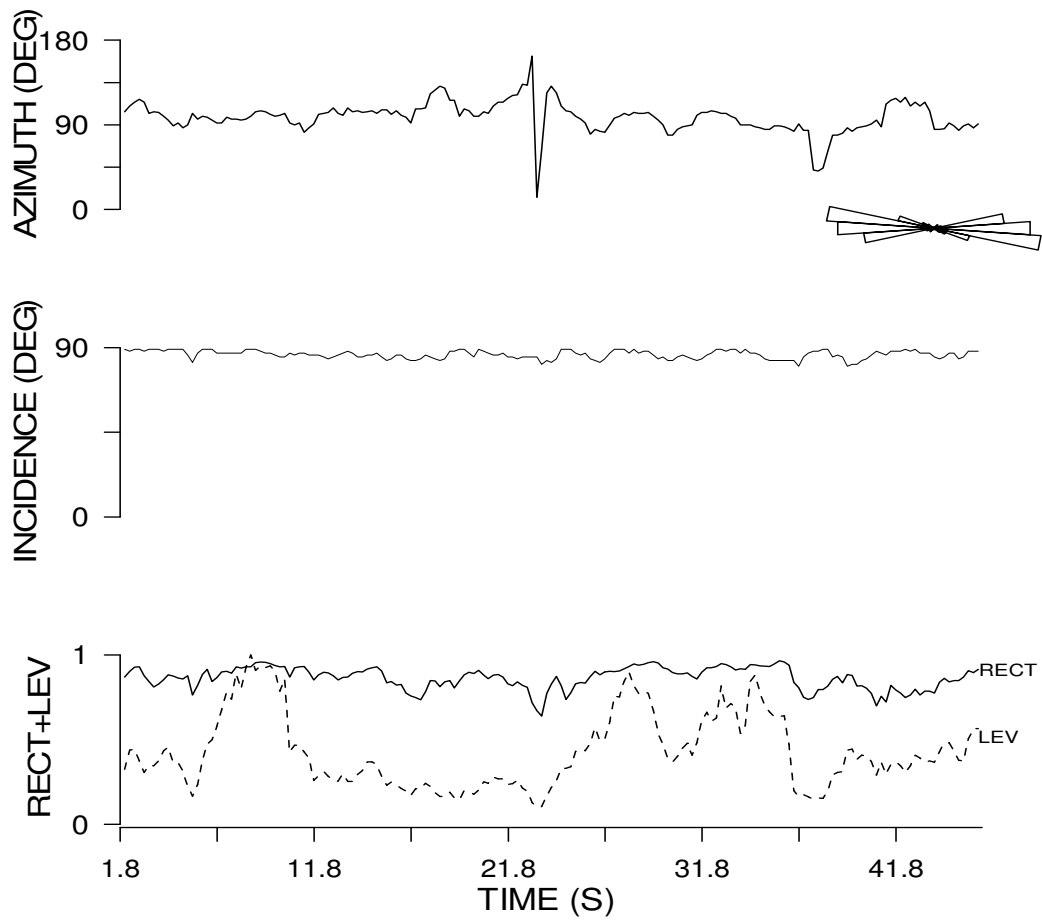


Fig. 15

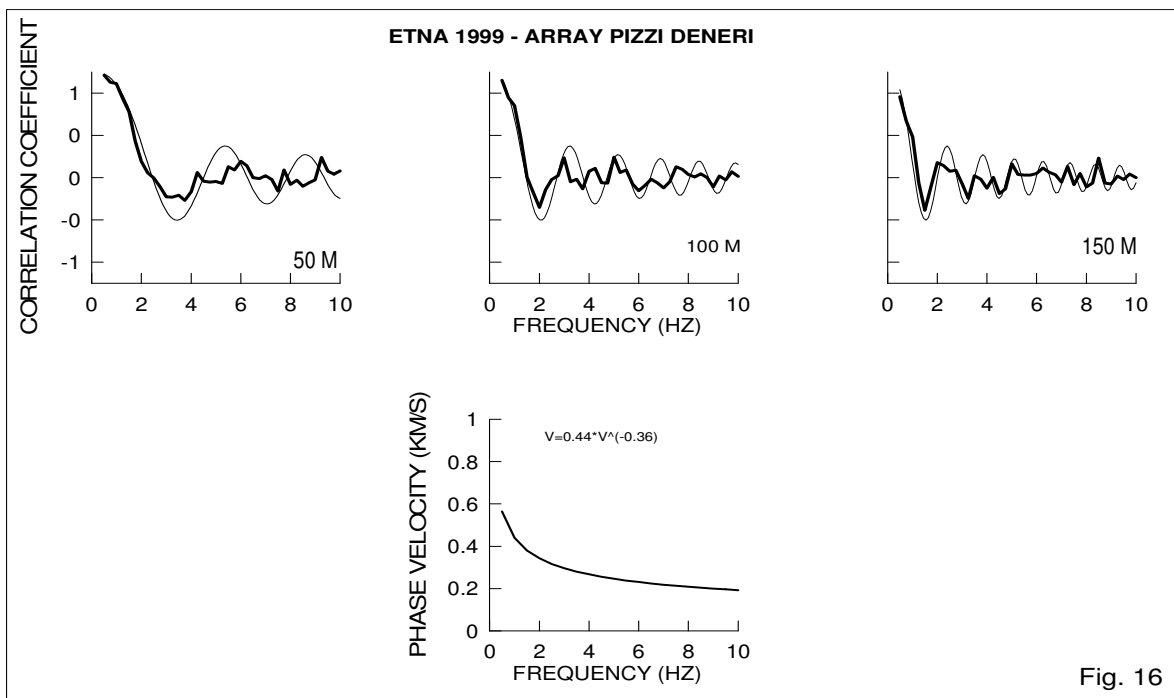


Fig. 16