

A seismic array on Mt. Vesuvius

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> Osservatorio Vesuviano Open file report 1999 n° 1

Abstract

In November 1997 a seismic antenna (array) of short period seismometers was installed on the south-western flank of Mt. Vesuvius; aim of the experiment was to test the use of non-conventional devices for the seismic monitoring of this volcano. In 7 months local seismicity, regional earthquakes and samples of seismic noise were recorded by the array and organised in a data base.

Local earthquakes and seismic noise have been analysed with array techniques to investigate the spectral, kinematic and polarization properties of the wavefield. Preliminary results show that the backazimuth of local earthquakes is oriented in the direction of the crater area. For some events, the source location has been constrained using a simplified back propagation in a 2-D velocity structure.

 The noise wavefield is characterized by the predominance of a sustained low frequency component (< 1Hz) whose source is located S-SE of the array. This low frequency signal has been interpreted as associated to the sea-loading in the gulf of Naples.

Introduction

Seismic arrays have been widely used (Goldstein & Chouet, 1994; Ferrazzini et. al., 1991; Dietel et al.,1994; Del Pezzo et. al, 1997) to study the wavefield associated to volcanic activity and many analyses have shown how array techniques are useful for the discrimination of the different seismic phases. In volcanic areas, the seismograms are very complex due to the topography and the heterogeneity of the medium in which seismic waves propagate: for this reason it is difficult to discriminate seismic phases in the wavefield associated to earthquakes. As a consequence of this problem, the S-wave arrival time picking is affected by a large error which causes an inaccurate and approximate location of the earthquakes. Moreover the misidentification of the seismic phases may lead to an inexact interpretation both of the source excitation mechanisms and of the later arrivals associated to the propagation in heterogeneous media.

In this framework, the application of array techniques to the seismicity recorded by the digital antenna can improve the results of the analyses (such as location of the source and polarization analysis) which are routinely performed on the data recorded by the Osservatorio Vesuviano seismic network.

Array techniques are also applied to some samples of seismic noise recorded during the experiment to investigate the spectral characteristics and the directional properties of the

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noise wavefield. This kind of analysis aims at the discrimination of possible insurgent volcanic tremor which can be a probable precursor of volcanic eruptions.

Instruments and data recording

 The array was located in the National Park of Mt. Vesuvius in the Baracche Forestali area, at a distance of about 1.5 Km from the crater (fig. 1). It was formed by 18 vertical geophones Mark L15B (600 ohm coil resistance) and 2 three component geophones Mark L15B (380 ohm coil resistance). The natural frequency (4.5 Hz) of the seismometers was electronically extended to 1 Hz. Sensors were buried in approximately 30 cm deep holes. The 3-D sensors were deployed with the horizontal components oriented in the N-S and E-W directions. Sensor coordinates were measured using differential GPS positioning, with a precision of about 30 cm in absolute sensor location(fig.1 , tab.1).

In the last week of operation four sensors were moved to the centre of the array to obtain a new configuration (fig.2, tab 2) more suitable for the study of the seismic noise.

The altitude difference among the sensors was less than 80 meters and the maximum aperture was about 500 meters. Although the array was deployed on the volcano flank, the average slope was about 10 degrees; this particular deployment allows the planar geometry approximation.

Station	Lat.	Long.	Lat. (m)	Long. (m)	$Elev.$ (m)
1B	40 48 41.976 N	14 24 41.909 E	307.6	325.9	658
1 ^C	40 48 37.892 N	14 24 36.662 E	81.7	203	645
1 _D	40 48 39.692 N	14 24 38.969 E	237.2	257.1	652
2D	40 48 37.734 N	14 24 33.438 E	176.8	127.4	633
3D	40 48 39.462 N	14 24 35.656 E	230.1	179.4	645
4B	40 48 40.678 N	14 24 40.096 E	267.6	283.4	654
4C	40 48 35.302 N	14 24 31.739 E	101.8	87.7	624
4D	40 48 40.306 N	14 24 28.350 E	256.1	8.2	630
5B	40 48 44.945 N	14 24 46.544 E	399.1	434.5	686
5C	40 48 34.481 N	14 24 34.927 E	76.5	162.3	622
5D	40 48 43.328 N	14 24 35.764 E	349.3	181.9	651
6B	40 48 40.713 N	14 24 46.076 E	268.6	423.6	672
6C	40 48 37.129 N	14 24 38.144 E	158.1	237.7	637
6D	40 48 42.421 N	14 24 38.282 E	321.3	241	661
7B	40 48 37.448 N	14 24 41.261 E	168	310.8	650
7C	40 48 32.725 N	14 24 36.572 E	22.4	200.9	613
7D	40 48 46.447 N	14 24 38.667 E	445.4	250	662
8B	40 48 36.385 N	14 24 46.803 E	135.2	440.6	661
8C	40 48 35.098 N	14 24 42.041 E	95.5	329	634
8D	40 48 45.059 N	14 24 40.621 E	402.6	295.8	702

Table Errore. L'argomento parametro è sconosciuto. - Sensor coordinates (configuration A).

Fig. 1 - Configuration A of the seismic array.

Fig. 2 - Configuration B of the seismic array.

Station	Lat.	Long.	Lat. (m)	Long. (m)	Elev. (m)
1C	40 48 37.892 N	14 24 36.662 E	81.7	203	645
1 _D	40 48 39.692 N	14 24 38.969 E	237.2	257.1	652
2D	40 48 37.734 N	14 24 33.438 E	176.8	127.4	633
3D	40 48 39.462 N	14 24 35.656 E	230.1	179.4	645
4B	40 48 40.678 N	14 24 40.096 E	267.6	283.4	654
4C	40 48 35.302 N	14 24 31.739 E	101.8	87.7	624
5B	40 48 41.976 N	14 24 41.909 E	307.6	325.9	658
5C	40 48 34.481 N	14 24 34.927 E	76.5	162.3	622
5D	40 48 43.328 N	14 24 35.764 E	349.3	181.9	651
6 _B	40 48 39.996 N	14 24 41.505 E	268.6	423.6	653
6C	40 48 37.129 N	14 24 38.144 E	158.1	237.7	637
6D	40 48 42.421 N	14 24 38.282 E	321.3	241	661
7B	40 48 37.448 N	14 24 41.261 E	168	310.8	650
7C	40 48 34.968 N	14 24 38.945 E	22.4	200.9	631
7D	40 48 40.748 N	14 24 36.602 E	445.4	250	655
8C	40 48 35.098 N	14 24 42.041 E	95.5	329	634
8D	40 48 45.059 N	14 24 40.621 E	402.6	295.8	702

Table 2 - Sensor coordinates (configuration B). The geophones 6B, 7C and 7D have been moved from the old position, while the 3-D station 1B was substituted by the geophone 5B.

The sensors, cased in an aluminium box together with the amplifier and the circuit for the dynamical extension, were cable-connected to the data loggers. As the acquisition system can record up to 8 channels, the array was divided into 3 subarrays controlled by 3 distinct data loggers. Each data logger was formed by a 16 bits A/D converter, an anti-aliasing filter and a GPS control circuit connected to a GPS antenna for absolute timing. These components

were cased in a plastic box and connected via parallel and serial port to a portable PC. The array acquisition system is a non-commercial one designed by R. Ortiz and G. Alguacil; more information about the instruments as well as details about the electronic schemes can be found in Olmedillas and Ortiz (1990).

The block-diagram of the device and the characteristic parameters of the geophones and pre-amplifier are reported in the following figure and table.

Fig. 3 - Block scheme of the device used for the experiment.

Parameter	L15B 380 Ohm	L15B 600 Ohm
d (damping factor)	3.90	4.06
G (trasduction constant) [Vs/m]	36.06	45.32
M (mass) $[Kg]$	0.023	0.023
R_c (geophone resistance) [Ohm]	380	600
R (total resistance of the circuit $geophone + pre-amplifier)$ [Ohm]	275.7	422.2
Av (voltage amplification)	-71.11	-79.02
ω_0 (geophone natural frequency) [Hz]	4.5	4.5

Table 3 - Characterisic parameters of the geophones and pre-amplifier.

The feedback connection between the geophone and the pre-amplifier provides an input signal which is the sum of the ground motion plus an electric signal proportional to the geophone output. This configuration improves the stability of the circuit because the geophone damping factor is incremented by the quantity:

$$
\Delta d = \frac{G^2}{2MR\omega_0}
$$

and the total damping of the system geophones + pre-amplifier is $D = d + \Delta d$. The transfer function of the circuit geophone + pre-amplifier is:

$$
F_1(\omega) = GA_v \frac{(i\omega)^2}{(i\omega)^2 + 2D\omega_0 i\omega + (\omega_0)^2}
$$

where A_{ν} is the voltage amplification.

The pre-amplifier is connected to the amplification circuit which has the following transfer function:

$$
F_2(\omega) = \left(1 + \frac{R_4}{R_5}\right) \frac{i\omega}{\frac{i\omega}{\omega} + 1}
$$

with:

$$
\omega_1 = \frac{1}{R_5 C} = 3.61 \text{ rad/s} \qquad \omega_2 = \frac{R_4 + R_5}{R_4 R_5 C} = 227 \text{ rad/s}
$$

\n
$$
R_4 = 9.51 \text{ K}\Omega \qquad R_5 = 590 \text{ K}\Omega \qquad C = 470 \text{ nF}
$$

\nIf $\omega \to \infty$ the function $F_2(\omega) \to 1$, while if $\omega \to 0$ then
\n
$$
F_2(\omega) = 1 + R_4 / R_5 \approx 63.
$$
 For this reason, the amplification circuit extends the observable

frequency range to values lower than 4.5 Hz which is the natural frequency of the geophone.

 The anti-aliasing filter is formed by a Butterworth filter and an amplifier and its cut-off frequency is $f_3 = 49.1$ Hz. The transfer function of this circuit (filter + amplifier) is:

$$
F_3(\omega) = \frac{a}{\prod_{j=1}^4 \left[\frac{(i\omega)^2}{\omega_3^2} + 2d_j\frac{i\omega}{\omega_3} + 1\right]}
$$

with $d_1 = 0.98$, $d_2 = 0.82$, $d_3 = 0.59$, $d_4 = 0.20$, $a = 6.687$.

The analogical signal is converted into a digital signal by the A/D converter which gives its contribution to the whole transfer function by a factor:

$$
A_d = \frac{2^{16} - 1}{8192}
$$
 counts/Volt

Therefore, the total transfer function of the device geophone + acquisition system is:

$$
F(\omega) = F_1(\omega) F_2(\omega) F_3(\omega) A_d
$$

The modulus of $F(\omega)$ is shown in fig. 4: the particular shape of this function makes the instrument response constant in the 1-50 Hz frequency range.

Fig. 4 - The total instrument transfer function.

Seismic signals were recorded using the trigger algorithm (LTA/STA), sampled at 200 samples/s and stored on the PC hard disk. Batteries were changed twice a week and a check of the instruments was made at the same time. Data recorded by the array were transferred once a week from the 3 acquisition PCs on a portable PC and then, after a selection in the seismological laboratory of Università di Salerno, were stored on PC-Zip diskettes. Data are available in PICFASE format (described in Del Pezzo et. al, 1998) and can be read by PICFASE program (Guirao, 1991).

 The array operated from November 1997 to June 1998, recording about 450 local earthquakes (among them there were two seismic swarms occurred on January 11th 1998 and March 2nd 1998), some regional earthquakes and some artificial explosions. An example of a local event is shown in fig.6. The number of local earthquakes (fig. 7) recorded by the array in the period November 1997 - June 1998 is higher than that recorded by the Osservatorio Vesuviano seismic network, demonstrating that the digital antenna improves the signal to noise ratio and decreases the detection threshold for local microearthquakes. Samples of seismic noise (fig. 5) were recorded using 215, 120-s-long programmed windows. Of these, 143 were recorded with the geophones deployed in configuration A (fig. 1) and 72 in configuration B (fig.2). A complete list of local and regional earthquakes and seismic noise samples recorded by the array is reported in appendix A.

Fig. 5 - 120 s time window of seismic noise recorded by the array.

Fig. 6 - Local earthquake 0110355 recorded by the array.

Fig 7 - Distribution of the local earthquakes recorded by the array in the period November 1997 - June 1998.

Data analysis and preliminary results

Local earthquake analysis

The wavefield associated to local earthquakes has been analysed with the zero-lag cross-correlation (ZLCC) technique (Frankel et al., 1991; Del Pezzo et. al, 1997) to estimate horizontal slowness (the inverse of apparent velocity) and backazimuth. An example is shown in fig. 8 where we report the slowness spectrum of the P-wave onset for 8 local events belonging to the seismic swarms mentioned above. In those spectra, the array-averaged zerolag cross-correlation coefficient is contoured as a function of the two components of the horizontal slowness. The vector which locate the main peak with respect to the spectrum origin gives the apparent velocity and backazimuth of the signal according to the relations:

$$
v_{app} = \frac{1}{\sqrt{p_x^2 + p_y^2}} \qquad \qquad \varphi = \frac{\pi}{2} - \tan^{-1} \frac{p_y}{p_x}
$$

where p_x and p_y are the coordinates of the peak. All the backazimuths associated to the displayed spectra point to the crater area.

 To discriminate the seismic S-phase we performed a polarization analysis using the covariance matrix technique (Jurkevics, 1988). The result of the analysis on the event 0612240 is shown in fig. 9. The S-wave has been identified with the phase coming 0.7 seconds after the P-wave onset because it has an apparent velocity which is about 1.75 times lower than the P-wave velocity and a γ angle of about 90°.

Fig. 8 - Results of the ZLCC analysis for local earthquakes belonging to the seismic swarm of January 11th (first 4 panels) and March $2nd$ (last 4 panels). The slowness components are reported on the x - y axes. On the contour lines the cross-correlation function assumes the same values.

Fig. 9 - ZLCC and polarization analysis on the event 0612240. In the first panel the seismogram is represented. In the next 3 panels we plot backazimuth φ, apparent velocity Va and correlation C as a function of time. In the last 5 panels we show the polarization attributes: RL is the rectilinearity, ψ is the angle between the projection of the polarization vector **p** on the surface and the North, $β$ is the angle that **p** forms with the vertical direction, $γ$ is the angle between **p** and the wave vector **k** and α is the angle that the **k** forms with the vertical direction (incidence angle).

Fig. 10 - The same of fig. X for the event 0110244. The polarization analysis does not evidence the S-phase of this earthquake.

We remind that γ is the angle between the wave vector direction and the direction of the eigenvector corresponding to the highest eigenvalue of the covariance matrix. For a more detailed explanation see Del Pezzo et al. (1997) and the caption of fig. 9.

 In some events the S-phase is evidenced by the polarization analysis, however there are other local earthquakes for which this phase is not clear (see for example fig. 10); this could be due to reflected and/or converted seismic phases in the wavefield which hide the Sarrival.

To constrain the depth of the earthquake sources we used a simple procedure based on a simplified 2-D ray-tracing technique and on the hypothesis that earthquake epicenters are confined in the crater area. The ray tracing technique is a modification of the Thurber (1983) procedure. It assumes that the ray path connecting source and receiver can be approximated by a second order polynomial which follows the Fermat principle. The velocity model is given on a 2-dimensional velocity grid. The search for the minimum time path is carried out with a trial and error procedure. The advantage of this approach is that the ray is expressed in an analytical form which results useful for further analyses. The computer program is reported in Appendix B.

Fig. 11 - Source depth as a function of the ray parameter.

To estimate the depth of local earthquakes we proceeded in the following way: first we fixed the epicenter at the crater centre, then we estimated the polynomials associated to the

rays coming at the array from sources with different depths. Furthermore we evaluated the incidence angle simply calculating the space derivative of the polynomial at the array point and hence we calculated the ray parameter (slowness at the array) for each ray. The plot of source depth as a function of the ray parameter was then fitted with a $4th$ order polynomial, which constitutes a nomogram (fig. 11) for the determination of the source depth.

We calculated the apparent velocity applying the ZLCC technique to a time window centred around the arrival time of the P-wave onset. From the knowledge of the apparent velocity and hence of the ray parameter, we can read on the nomogram the corresponding source depth along the crater axis. Moreover if we can estimate the S-P time, we can fix the ray path length and hence locate the source position along the ray (fig 12).

Fig. 12 - Rays obtained with the simplified 2-D ray-tracing technique. The empty circles represent the hypocenters (determined from the nomogram of fig. 11) of 40 earthquakes analysed with ZLCC technique. For 5 earthquakes it was possible to estimate the S-T time and to fix the source position along the ray (full circles).

Noise analysis

 To investigate the spectral properties of the seismic noise we calculated the average spectra, which were derived for samples recorded in different days and times. An example of spectra is shown in fig. 13.

The spectra show a very clear peak at 0.5-0.7 Hz, whose amplitude does not change during day and night, two peaks at 3 and 10 Hz and other minor peaks at frequencies higher than 1-1.5 Hz. The amplitude of these last peaks depends on the day-time of the records and it is predominant in the morning hours, suggesting that these components of noise are due to the antropic activity. It is noteworthy that even though the response curve amplification at 0.5 Hz is a factor 3 lower than that at 3 Hz, the 0.5 Hz peak is clearly predominant over all the others.

 MUSIC technique (Schmidt, 1981, 1986) was applied using a focusing frequency of 0.58 Hz to calculate azimuth and slowness of the sustained low frequency component. Results obtained for noise samples recorded during different day-time (fig. 14) show that the slowness is around 0.7 s/Km and the backazimuth points toward the South, in the opposite direction of the crater area. These results suggest that the low-frequency component may be due to the sealoading in the gulf of Naples.

12

6

0.1 1.0 10.0 Hz

 $\overline{\mathbb{T}}$

मामा

08:59

Fig. 13 - Average spectra of seismic noise recorded at different day-time on June 14th.

Fig. 14 - Slowness spectra obtained with MUSIC for noise samples recorded during the time interval 00:59 - 08:59 on April 21 $\mathrm{^{st}}$.

Conclusions

The preliminary results can be summarised in the following statements:

- 1) The local earthquake source is in the direction of the crater area.
- 2) The use of array techniques allows the identification of the S-phase for more than the 50% of the analysed events.
- 3) The source location obtained from the array analysis is consistent with the location obtained using conventional techniques (Hypo71, computer program) to data recorded by the monitoring network (fig. 15).

Fig. 15 - Location obtained with Hypo of some earthquakes of the seismic swarms recorded by the seismic network.

- 4) The source depth for the events with non-readable S-P times is constrained by the array estimate of slowness to be in the range $1.5 - 3$ Km, if an epicenter in correspondence of the crater area is assumed (fig. 12).
- 5) A low frequency spectral peak is present in the noise wavefield. It corresponds to a coherent signal which is interpreted as related to the sea-loading in the gulf of Naples.
- 6) The analysis of the seismic noise shows that there is no evidence of coherent signal in the frequency band $(1 - 5 Hz)$ characteristic of the volcanic tremor.

 The results achieved during this experiment encourage to carry on further analyses on the collected data, both on local earthquakes and noise. The above results confirm that the array is an useful complementary tool for the seismic monitoring of the volcanic area. For this reason we are planning a new research project aiming both at the installation of a new multichannel digital seismic antenna to be used for the seismic signal acquisition and at the development of real-time analysis techniques.

Acknowledgements

Part of this report is based on the Physics degree thesis of Danilo Galluzzo and Alessandra Cirillo. The instruments were assembled in the laboratory of Observatorio de Cartuja - Universidad de Granada.

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Appendix A

Table A1 - List of the local earthquakes recorded at the array. The letters in the first column indicate which subarrays recorded the earthquake (for example, BCD means that all the 3 subarrays triggered). The X in the 5th column (N) means that the event was recorded by the seismic network too.

Event	Date	GMT	Location	
3221308 C	18/11/97	13:08	Greece	
			$M = 6.1$	
3221315 BC	18/11/97	13:15	Greece	
3221501 C	18/11/97	15:01	Greece	
3221524 BC	18/11/97	15:24	Greece	
3221805 B	18/11/97	18:05	Greece	
3281904 BCD	24/11/97	19:04	Sannio-Matese	
			$M = 3.8$	
3290137 BCD	25/11/97	01:37	Sannio-Matese	
			$M = 3.5$	
3442109 BC	10/12/97	21:09	Sannio-Matese	
0100435 CD	10/01/98	04:35	Regional earthquake	
0101923 BCD	10/01/98	19:23	Regional earthquake	
0262317 BCD	26/01/98	23:17	Regional earthquake	
			$M = 3.7$	
0660327 CD	07/03/98	03:27	Regional earthquake	
0851626 BCD	26/03/98	16:26	Umbria-Marche	
0972137 BCD	07/04/98	21:37	Lavello-Venosa	
			$M = 3.9$	
1122339 BCD	22/04/98	23:39	Cervinara	
			Valle Caudina	
			$M = 3.0$	
1190332 BC	29/04/98	03:32	Greece	
			$M = 5.4$	
1381719 BCD	18/05/98	17:19	Tyrrhenian Sea	

Table. A2 - List of the regional earthquakes recorded both at the array and at the seismic network

Event	Date	GMT	Type
3580157 BCD	24/12/97	01:57	Probable artificial explosion
3591736 BCD	25/12/97	17:36	Probable artificial explosion
3591756 BC	25/12/97	17:56	Probable artificial explosion
0461801 BC	15/02/98	18:01	Artificial explosion
0461829 BC	15/02/98	18:29	Artificial explosion
0491124 BCD	18/02/98	11:24	Artificial explosion
0502313 D	19/02/98	23:13	Artificial explosion
0590342 BCD	28/02/98	03:42	Artificial explosion
1121442 BCD	22/04/98	14:42	Under-sea explosion

Table A4 - List of the 215 seismic noise samples which were recorded programming time windows of 120 s. For each programmed recording, we report day and time of the first noise sample, day and time of the last sample, the number of time windows (the frequency of the programmed windows was 1/hour) and the configuration of the the number of time windows (the frequency of the programmed windows was 1/hour) and the configuration of the array..

Appendix B

We report the list of 2-D ray tracing program.

DECLARE FUNCTION ray! (a!, b!, c!, x!) DECLARE SUB interpol (ncolonne, nrighe, xval!, zval!, vinterp!) DIM trav(2000), a222(2000), a111(2000) ' programma di ray tracing su griglia INPUT "nome file modello di velocita", nome\$ INPUT "nome file uscita", nime\$ INPUT "coordinate stazione"; xr, zr INPUT "coordinate evento,z negativa"; xs, zs INPUT "numero passi deltax", nstep OPEN nome\$ FOR INPUT AS #1 OPEN nime\$ FOR OUTPUT AS #3 PRINT #3, xr, zr PRINT #3, xs, zs 'legge la griglia di velocit… INPUT #1, ncolonne, nrighe DIM x(ncolonne, nrighe), z(ncolonne, nrighe), v(ncolonne, nrighe) FOR $i = 1$ TO ncolonne FOR $i = 1$ TO nrighe **INPUT** #1, $x(i, j)$, $z(i, j)$, $v(i, j)$ 'PRINT $x(i, j)$, $z(i, j)$, $v(i, j)$, i, j NEXT j NEXT i '-- ' si opera una trasformazione di coordinate per porre a zero 'la coordinata xs ' PRINT "modello corretto" FOR $i = 1$ TO ncolonne FOR $i = 1$ TO nrighe $x(i, j) = x(i, j) - xs$ ' PRINT $x(i, j)$, $z(i, j)$, $v(i, j)$, i, j NEXT j NEXT i $xr = xr - xs$ $xs = 0$ '-- ' calcola il coeff. ang della retta stazione sorgente $a1 = (zr - zs) / xr$ $alfa1 = ATM(a1)$ $delta x = xr / nstep$ PRINT "a1,alfa1,deltax", a1, alfa1, deltax 'alfa1 in radianti $pig = 3.1415$ IF a $1 \geq 0$ THEN alfaincrem $= 2 *$ pig / 100 END IF IF a1 < 0 THEN

```
alfaincrem = 2 * \text{pi}/ 360END IF 
index = 0'----------------------------------- 
'------------------------------------- 
121 alfa1 = alfa1 - alfaincrem 
index = index + 1PRINT "nuova iterazione,index# ", index 
IF alfa1 \lt (-pig / 2) + ABS(3 * alfaincrem) THEN GOTO 2332
a11 = TAN(alfa1)' la seguente formula Š stata modificata rispetto alla versione precedente 
' RAYT2D2: al posto di xr ora compaiono le differenze (xr - xs) 
a22 = (zr - zs - a11 * (xr - xs)) / (xr - xs) \wedge 2PRINT "a11,a22", a11, a22 
'il polinomio sara' z=zs+a11*x+a22*x^2 la funzione e' ray(a,b,c,x)
'che lo definisce con zs=a,a11=b,a22=c 
sum = 0FOR i = 1 TO nstep
  zray = ray(zs, a11, a22, (i - 1) * deltax)
  deltaz = ray(zs, a11, a22, i * deltax) - ray(zs, a11, a22, (i - 1) * deltax)
  CALL interpol(ncolonne, nrighe, (i - 1) * deltax, zray, vel)
   PRINT "velocita'", vel, "iterazione", i 
  p = 1 / veldeltas = SQR(deltax \land 2 + deltaz \land 2)
  sum = sum + p * deltas
NEXT i 
traveltime = sum 
PRINT "traveltime", traveltime 
trav/index) = traveltimea111(index) = a11a222(index) = a22PRINT #3, zs, a11, a22, traveltime 
 GOTO 121 
2332 mintrav = trav(1)indice = 0FOR k = 1 TO index - 1
       IF trav(k) \leq mintrav THEN
       mintrav = trav(k)indice = k
        END IF 
        NEXT k 
        PRINT "zs,a11,a22,trav time", zs, a111(indice), a222(indice), trav(indice) 
 END 
SUB interpol (ncolonne, nrighe, xval, zval, vinterp) 
SHARED x(), z(), v()'PRINT "xval,zval", xval, zval 
xmax = x(ncolonne, nrighe)
xmin = x(1, 1)zmax = z(1, 1)
```

```
zmin = z(ncolonne, mright)'PRINT "xmax,xmin,zmax,zmin", xmax, xmin, zmax, zmin 
FOR i = 1 TO ncolonne
FOR j = 1 TO nrighe
 IF x(i, j) < xval THEN
              IF x(i, j) \geq xmin THEN
              xmin = x(i, j)icol = i END IF 
   END IF 
   NEXT j 
   NEXT i 
  FOR i = 1 TO nrighe
  IF z(icol, j) > zval THEN
              IF z(icol, j) \leq zmax THEN
               zmax = z(icol, j)\text{irig} = \text{i} END IF 
   END IF 
  NEXT j 
IF icol = 0 THEN icol = 1
IF jrig = 0 THEN jrig = 1IF icol \geq ncolonne THEN icol = ncolonne - 1
IF jrig >= nrighe THEN jrig = nrighe - 1
'PRINT "icol, cioe' lindice di colonna prima del valore", icol 
'PRINT "jrig, cioe' l'indice di riga subito prima del valore", jrig 
' ora la interpolazione avviene tra v(icol,jrig),v(icol,jrig+1),v(icol+1,jrig), 
'e v(icol+1,jrig+1) con una media pesata per le distanze inverse 
dist1 = SQR((x(icol, ring) - xval) \land 2 + (z(icol, ring) - zval) \land 2)dist2 = \text{SOR}((x(icol, irig + 1) - xval) \wedge 2 + (z(icol, irig + 1) - zval) \wedge 2)dist3 = SQR((x(icol + 1, irig) - xval) \wedge 2 + (z(icol + 1, irig) - zval) \wedge 2)dist4 = SQR((x(icol + 1, jrig + 1) - xval) ^ 2 + (z(icol + 1, jrig + 1) - zval) ^ 2)IF dist1 < 10 ^ -6 THEN dist1 = .0001IF dist2 < 10 ^ -6 THEN dist2 = .0001IF dist3 < 10 ^ -6 THEN dist3 = .0001IF dist4 < 10 ^ -6 THEN dist4 = .0001denom = (1 / dist1) + (1 / dist2) + (1 / dist3) + (1 / dist4)'PRINT "v(icol, jrig)", v(icol, jrig) 
'PRINT "v(icol + 1, jrig)", v(icol + 1, jrig)
vinterp = (v(icol, jrig) / dist1) + (v(icol, jrig + 1) / dist2) + (v(icol + 1, jrig) / dist3) + (v(icol + 1, jrig))1, |rig + 1) / dist4)
vinterp = vinterp / denom
END SUB 
FUNCTION ray (a, b, c, x) 
ray = a + b * x + c * x \wedge 2END FUNCTION
```