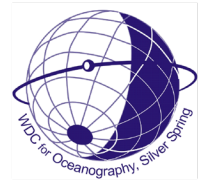


# NOAA Atlas NESDIS 97



## WORLD OCEAN DATABASE 2023

Alexey V. Mishonov, Tim P. Boyer, Olga K. Baranova, Courtney N. Bouchard, Scott L. Cross, Hernan E. Garcia, Ricardo A. Locarnini, Christopher R. Paver, James R. Reagan, Zhankun Wang, Dan Seidov, Alexandra I. Grodsky, and James G. Beauchamp

Technical Editor: Courtney N. Bouchard

*National Centers for Environmental Information  
Ocean Climate Laboratory*

Silver Spring, MD  
October 2024

### **U.S. DEPARTMENT OF COMMERCE**

Gina M. Raimondo, Secretary

### **National Oceanic and Atmospheric Administration**

Richard W. Spinrad, Under Secretary of Commerce for Oceans and Atmosphere & NOAA Administrator

### **National Environmental Satellite, Data, and Information Service**

Stephen Volz, Assistant Administrator

Additional copies of this publication, as well as information about NCEI data holdings and services, are available upon request directly from NCEI.

NOAA/NESDIS  
National Centers for Environmental Information  
SSMC3, 4th floor  
1315 East-West Highway  
Silver Spring, MD 20910-3282  
U.S.A.

Telephone: +1 (828) 271-4800  
E-mail: NCEI.Info@noaa.gov  
WEB: <http://www.ncei.noaa.gov/>

For updates on the data, documentation and additional information about WOD23 please refer to:

<https://www.ncei.noaa.gov/products/ocean-climate-laboratory>

This publication should be cited as:

Mishonov, A.V., T.P. Boyer, O.K. Baranova, C.N. Bouchard, S.L. Cross, H.E. Garcia, R.A. Locarnini, C.R. Paver, J.R. Reagan, Z. Wang, D. Seidov, A.I. Grodsky, J.G. Beauchamp (2024). World Ocean Database 2023. C. Bouchard, Technical Ed. *NOAA Atlas NESDIS 97*, DOI <https://doi.org/10.25923/z885-h264>

# CONTENTS

CONTENTS .....	3
LIST OF TABLES.....	6
LIST OF FIGURES.....	8
LIST OF ACRONYMS .....	12
PREFACE.....	14
ACKNOWLEDGMENTS .....	15
CHAPTER 1: INTRODUCTION.....	16
ABSTRACT .....	16
1.1. INTRODUCTION .....	16
1.1.1. Purpose.....	17
1.1.2. Contributors .....	18
1.1.3. Size and shape .....	18
1.1.4 Data Organization .....	19
1.1.5. WOD Datasets.....	21
1.1.6. Economic and scientific justification for maintaining archives of historical oceanographic data: the value of stewardship.....	28
1.1.7. Data assembly .....	28
1.1.8. Distribution media.....	28
1.1.9. Application software interfaces.....	29
1.1.10. Units.....	29
1.2. COMPARISON OF WOD23 WITH PREVIOUS RELEASES .....	29
1.3. DATA SOURCES .....	31
1.3.1. IODE .....	31
1.3.2. The World Data System.....	31
1.3.3. IOC Global Oceanographic Data Archaeology and Rescue Project .....	31
1.3.4. Near-real time data sources.....	31
1.3.5. International Research Projects Data .....	32
1.3.6. ICES Contribution.....	32
1.3.7. Declassified Naval Data Sets .....	32
1.3.8. Integrated Global Ocean Service - Volunteer Observing Ship programs.....	34
1.4. QUALITY CONTROL FLAGS .....	34
1.4.1. Levels of Quality Control.....	35
1.5. OUTLOOK FOR FUTURE ACQUISITIONS OF HISTORICAL OCEAN PROFILE AND PLANKTON DATA AND INTERNATIONAL COOPERATION IN THE “WORLD OCEAN DATABASE PROJECT” .....	36
1.6. LAYOUT OF THE REST OF THIS DOCUMENT .....	37
1.7. REFERENCES AND BIBLIOGRAPHY .....	37
CHAPTER 2: OCEAN STATION DATA (OSD), LOW-RESOLUTION CTD, LOW-RESOLUTION EXPENDABLE XCTD, AND PLANKTON .....	40
2.1. INTRODUCTION .....	40
2.2. COMMONLY USED LOW AND LARGE VOLUME WATER COLUMN SAMPLERS.....	41
2.3. VARIABLES AND METADATA INCLUDED IN THE OSD DATASET .....	44
2.4. OSD DATA COVERAGE.....	49
2.5. PARAMETERS AND METADATA NOT INCLUDED IN THE OSD DATASET .....	50
2.6. PROSPECTS FOR THE FUTURE.....	51
2.7. REFERENCES AND BIBLIOGRAPHY .....	69
CHAPTER 3: CONDUCTIVITY-TEMPERATURE-DEPTH DATA (CTD).....	72
3.1. INTRODUCTION .....	72
3.2. CTD ACCURACY .....	73
3.3. CTD CAST DISTRIBUTIONS .....	73
3.5. REFERENCES AND BIBLIOGRAPHY .....	77

CHAPTER 4: EXPENDABLE BATHYTHERMOGRAPH DATA (XBT) .....	78
4.1. INTRODUCTION .....	78
4.2. XBT ACCURACY .....	79
4.3. XBT TIME-DEPTH EQUATION ERROR.....	80
4.4. CORRECTIONS TO XBT TIME-DEPTH EQUATION ERRORS.....	81
4.5. SURFACE DATA ACQUIRED CONCURRENTLY WITH XBT CASTS .....	83
4.6. XBT PROFILE DISTRIBUTIONS .....	83
4.7. REFERENCES AND BIBLIOGRAPHY .....	89
CHAPTER 5: EXPENDABLE CONDUCTIVITY-TEMPERATURE-DEPTH DATA (XCTD).....	92
5.1. INTRODUCTION .....	92
5.2. XCTD PRECISION AND ACCURACY .....	92
5.3. XCTD FALL-RATE ERROR.....	93
5.4. XCTD CAST DISTRIBUTIONS .....	94
5.5. RELEVANT WEB SITES .....	97
5.6. REFENCES AND BIBLIOGRAPHY .....	98
CHAPTER 6: PROFILING FLOATS DATA (PFL).....	99
6.1. INTRODUCTION .....	99
6.2. PREDECESSORS OF PROFILING FLOATS .....	100
6.3. FIRST GENERATION PROFILING FLOATS .....	100
6.4. PRESENT FLOAT TECHNOLOGY .....	101
6.4.1. The Argo Project .....	102
6.5. SENSOR ACCURACY .....	103
6.6. DATA PROBLEMS.....	104
6.6.1. Sensor problems .....	104
6.6.2. Data-Stream Errors.....	106
6.7. ORIGINATORS FLAGS.....	106
6.8. PFL DATA DISTRIBUTIONS .....	107
6.9. RELEVANT WEB SITES .....	111
6.10. REFERENCES AND BIBLIOGRAPHY .....	111
CHAPTER 7: MECHANICAL BATHYTHERMOGRAPH DATA (MBT) .....	115
7.1. INTRODUCTION.....	115
7.2. MBT ACCURACY.....	115
7.3. SURFACE DATA ACQUIRED CONCURRENTLY WITH MBT CASTS.....	116
7.4. MBT PROFILE DISTRIBUTION .....	117
7.5. REFERENCES AND BIBLIOGRAPHY .....	121
CHAPTER 8: DIGITAL BATHYTHERMOGRAPH (DBT) PROFILES .....	123
8.1. INTRODUCTION.....	123
8.2. DBT ACCURACY.....	123
8.3. DBT PROFILE DISTRIBUTIONS .....	123
8.4. REFERENCES AND BIBLIOGRAPHY .....	126
CHAPTER 9: MOORED BUOY DATA (MRB).....	127
9.1. INTRODUCTION .....	127
9.2. MRB DATA PRECISION AND ACCURACY .....	131
9.3. MRB PROFILE DISTRIBUTIONS .....	131
9.4. RELEVANT WEB SITES .....	134
9.5. REFERENCES AND BIBLIOGRAPHY .....	135
CHAPTER 10: DRIFTING BUOY DATA (DRB) .....	137
10.1. INTRODUCTION .....	137
10.1.1. Arctic Ocean Buoy Program .....	137
10.1.2. Global Temperature-Salinity Profile Program (GTSPP).....	138



10.1.3. JAMSTEC Buoys.....	138
10.1.4. Ice-Tethered Profiling Buoys (ITP) .....	140
10.1.5. PAICEX Russian Drifting Ice Camp data.....	141
10.2. DRB ACCURACY .....	141
10.3. DRB PROFILE DISTRIBUTIONS .....	143
10.4. RELEVANT WEB SITES.....	145
10.5. REFERENCES AND BIBLIOGRAPHY .....	146
CHAPTER 11: UNDULATING OCEAN RECORDER DATA (UOR).....	148
11.1. INTRODUCTION .....	148
11.2. UOR DATA PRECISION AND ACCURACY.....	151
11.3. UOR PROFILE DISTRIBUTIONS.....	152
11.4. RELEVANT WEB SITES.....	154
11.5. REFERENCES AND BIBLIOGRAPHY .....	155
CHAPTER 12: AUTONOMOUS PINNIPED BATHYTHERMOGRAPH DATA (APB) .....	157
12.1. INTRODUCTION .....	157
12.2. DATA SOURCES.....	158
12.3. INSTRUMENTATION AND ACCURACY .....	158
12.4. GEOGRAPHICAL, TIME, AND DEPTH DISTRIBUTION OF DATA.....	160
12.5. REFERENCES AND BIBLIOGRAPHY .....	162
CHAPTER 13: MICRO BATHYTHERMOGRAPH DATA (MICRO BT) .....	164
13.1. INTRODUCTION .....	164
13.2. MICRO BT ACCURACY.....	164
13.3. MICRO BT PROFILE DISTRIBUTIONS.....	164
CHAPTER 14: SURFACE-ONLY DATA (SUR).....	167
14.1. INTRODUCTION .....	167
14.2. SUR ACCURACY .....	167
14.3. DATA COVERAGE.....	168
14.4. REFERENCES AND BIBLIOGRAPHY .....	172
CHAPTER 15: GLIDER DATA (GLD) .....	174
15.1. INTRODUCTION .....	174
15.2. GLD ACCURACY .....	175
15.3. GLIDER DESIGN AND OPERATION .....	176
15.4. GLD PROFILES DISTRIBUTIONS.....	177
15.5. RELEVANT WEB SITES.....	183
15.6. REFERENCES AND BIBLIOGRAPHY .....	183
CHAPTER 16: PLANKTON DATA .....	186
16.1. INTRODUCTION .....	186
16.2. BASIC QUALITY CONTROL.....	190
16.3. DATA SOURCES.....	191
16.4. PLANKTON DATA DISTRIBUTIONS.....	195
16.5. PLANKTON CONTENT .....	198
16.5.1. Abundance .....	198
16.5.2. Total Biomass .....	202
16.6. REFERENCES AND BIBLIOGRAPHY .....	205

## LIST OF TABLES

TABLE 1.1. INSTRUMENT TYPES IN THE WOD23 .....	22
TABLE 1.2. METEOROLOGICAL AND SEA-STATE PARAMETERS STORED IN THE WOD23 .....	27
TABLE 1.3. COMPARISON OF THE NUMBER OF OCEANOGRAPHIC CASTS IN WOD PRODUCTS SERIES .....	33
TABLE 2.1 DEPTH DEPENDENT <i>IN SITU</i> MEASURED VARIABLES.....	46
TABLE 2.2. THE NUMBER OF OCEAN STATION DATA (OSD) CASTS AS A FUNCTION OF YEAR ....	53
TABLE 2.3. NUMBER OF OSD CASTS PROVIDED BY EACH COUNTRY (ISO, INTERNATIONAL ORGANIZATION FOR STANDARDIZATION) .....	55
TABLE 3.1. LIST OF ALL VARIABLES AND PROFILE COUNTS IN THE WOD23 CTD DATASET. ....	73
TABLE 3.2. THE NUMBER OF HIGH-RESOLUTION CTD CASTS IN WOD23 AS A FUNCTION OF YEAR. ....	74
TABLE 3.3. NATIONAL CONTRIBUTIONS OF HIGH-RESOLUTION CTD CASTS IN WOD23.....	76
TABLE 4.1. CHARACTERISTICS OF EXPENDABLE PROBES PRODUCED BY LOCKHEED MARTIN SIPPICAN. ....	79
TABLE 4.2. THE NUMBER OF ALL XBT PROFILES AS A FUNCTION OF YEAR IN WOD23. ....	84
TABLE 4.3. NATIONAL CONTRIBUTION OF XBT PROFILES IN WOD23. ....	87
TABLE 5.1. THE NUMBER OF XCTD CASTS IN WOD23 AS A FUNCTION OF YEAR.....	94
TABLE 5.2. NATIONAL CONTRIBUTIONS OF XCTD CASTS IN WOD23.....	95
TABLE 6.1. VARIABLES AND PROFILE COUNTS IN THE PFL WOD23 DATASET. ....	99
TABLE 6.2. NATIONAL CONTRIBUTION OF PFL CASTS IN WOD23. ....	109
TABLE 6.3. THE NUMBER OF PROFILING FLOAT DATA (PFL) CASTS AS A FUNCTION OF YEAR IN WOD23. YEAR 2023 IS A PARTIAL YEAR FROM JANUARY TO APRIL. ....	110
TABLE 7.1. NUMBER OF ALL MBT PROFILES AS A FUNCTION OF YEAR IN WOD23. ....	118
TABLE 7.2. COMPARISON OF OBSERVATIONS TAKEN WITH MECHANICAL BATHYTHERMOGRAPHS AND REVERSING THERMOMETERS.....	116
TABLE 7.3. NATIONAL CONTRIBUTIONS OF MECHANICAL BATHYTHERMOGRAPH PROFILES IN WOD23.....	120
TABLE 8.1. THE NUMBER OF DIGITAL BATHYTHERMOGRAPH (DBT) PROFILES AS A FUNCTION OF YEAR IN WOD23.....	124
TABLE 8.2. NATIONAL CONTRIBUTIONS OF DIGITAL BATHYTHERMOGRAPH (DBT) PROFILES IN WOD23.....	124
TABLE 9.1. NATIONAL CONTRIBUTIONS OF MRB PROFILES IN WOD23 .....	130
TABLE 9.2. THE NUMBER OF MRB PROFILES IN WOD23 AS A FUNCTION OF YEAR.....	132
TABLE 10.1. THE NUMBER OF DRB PROFILES IN AS A FUNCTION OF YEAR IN WOD23. ....	143
TABLE 10.2. NATIONAL CONTRIBUTIONS OF DRB CASTS IN WOD23.....	144
TABLE 11.1. PROFILE COUNT FOR MAJOR VARIABLES IN THE WOD23 UOR DATASET. ....	149
TABLE 11.2. THE NUMBER OF ALL UOR CASTS AS A FUNCTION OF YEAR IN WOD23.....	152
TABLE 11.3. NATIONAL CONTRIBUTIONS OF UOR CASTS IN WOD23. ....	153
TABLE 12.1. PROJECTS CONTRIBUTING TO THE WOD23 APB DATASET AND NUMBER OR PROFILES SUBMITTED. ....	159
TABLE 12.2. THE NUMBER OF ALL APB CASTS AS A FUNCTION OF YEAR IN WOD23. ....	160
TABLE 13.1. THE NUMBER OF ALL MICRO BT PROFILES AS A FUNCTION OF YEAR IN WOD23. .	165
TABLE 13.2. NATIONAL CONTRIBUTIONS OF MICRO BATHYTHERMOGRAPH PROFILES IN WOD23. ....	165
TABLE 14.1. LIST OF PARAMETERS AND NUMBER OF OBSERVATIONS IN THE SUR DATASET OF WOD23.....	168
TABLE 14.2. THE NUMBER OF ALL SUR OBSERVATIONS AS A FUNCTION OF YEAR IN WOD23. .	171
TABLE 14.3. NATIONAL CONTRIBUTIONS OF OBSERVATIONS, AND NUMBER OF CRUISES BY COUNTRY OF ORIGIN IN THE SUR DATASET IN WOD23.....	172
TABLE 15.1. DIFFERENT GLIDER PLATFORM CAPABILITIES <sup>1</sup> AND NUMBER OF PLATFORMS & CASTS IN WOD23. ....	176
TABLE 15.2. THE NUMBER OF ALL GLIDER (GLD) CASTS AS A FUNCTION OF YEAR IN WOD23. .	179
TABLE 15.3. NATIONAL CONTRIBUTION OF GLIDER (GLD) PROFILES IN WOD23.....	182

TABLE 16.1. MEASUREMENT TYPE AND/OR GROUPS AND THEIR CORRESPONDING CBV UNIT.	190
TABLE 16.2. WOD23 BROAD GROUP-BASED RANGES FOR PLANKTON ABUNDANCE. ....	190
TABLE 16.3. WOD23 BROAD GROUP-BASED RANGES FOR BIOMASS. ....	191
TABLE 16.4. NATIONAL CONTRIBUTIONS OF PLANKTON CASTS IN WOD23. ....	192
TABLE 16.5. PROJECT CONTRIBUTIONS OF PLANKTON CASTS SORTED BY PERCENT CONTRIBUTION FROM EACH PROJECT. ....	194
TABLE 16.6. NUMBER OF PLANKTON CASTS IN WOD23 AS A FUNCTION OF YEAR. ....	197
TABLE 16.7. WOD23 ABUNDANCE MEASUREMENTS CONTENT. ....	198
TABLE 16.8. WOD23 BIOMASS MEASUREMENTS CONTENT. ....	202

# LIST OF FIGURES

FIGURE 1.1. THE LOCATION AND WOD STATION NUMBERS OF THE EARLIEST DATA IN WOD23.17	
FIGURE 1.2. WOD23 OBSERVATIONS DENSITY (PROFILES PER 1-DEGREE SQUARE). .....	19
FIGURE 1.3. DATA HOLDING GROWTH OVER THE YEARS OF THE WOD DEVELOPMENT. ....	30
FIGURE 2.1 GEOGRAPHIC POSITIONS OF <i>IN SITU</i> SEA SURFACE TEMPERATURE MEASUREMENTS COLLECTED IN THE 1772-1800 TIME PERIOD AVAILABLE IN THE WOD.....	42
FIGURE 2.2. PICTURE OF ORIGINAL <i>IN SITU</i> SEA SURFACE TEMPERATURE AND METEOROLOGICAL MEASUREMENTS COLLECTED BETWEEN FEBRUARY 28 AND JUNE 3, 1800 DIGITIZED AS PART OF THE GODAR PROJECT. ....	43
FIGURE 2.3. TIME SERIES OF THE NUMBER OF AVAILABLE OSD CASTS IN WOD23 (1900-2022). THERE ARE 3,256,037 CASTS (1900-2022) AND 105,973 CASTS (1772 TO 1900, NOT SHOWN). ....	44
FIGURE 2.4. NUMBER OF TEMPERATURE OBSERVATIONS IN WOD23 AS A FUNCTION OF 102 STANDARD DEPTH LEVELS. ....	45
FIGURE 2.5. NUMBER AND DISTRIBUTION OF OSD CASTS IN WOD23 BY ONE-DEGREE SQUARES	50
FIGURE 2.6 NUMBER OF OSD CASTS IN NCEI (PREVIOUSLY NODC)/WDS DATABASES AS A FUNCTION OF TIME. ....	51
FIGURE 2.7. TIME SERIES OF THE NUMBER OF TEMPERATURE PROFILES IN WOD23 OSD DATASET .....	57
FIGURE 2.8. TIME SERIES OF THE NUMBER OF SALINITY PROFILES IN WOD23 OSD DATASET....	57
FIGURE 2.9. TIME SERIES OF THE NUMBER OF DISSOLVED OXYGEN PROFILES IN WOD23 OSD DATASET.....	58
FIGURE 2.10. TIME SERIES OF THE NUMBER OF PHOSPHATE PROFILES IN WOD23 OSD DATASET	58
FIGURE 2.11. TIME SERIES OF THE NUMBER OF SILICATE PROFILES IN WOD23 OSD DATASET ..	59
FIGURE 2.12. TIME SERIES OF THE NUMBER OF NITRATE PROFILES IN WOD23 OSD DATASET ...	59
FIGURE 2.13. TIME SERIES OF THE NUMBER OF PH PROFILES IN WOD23 OSD DATASET .....	60
FIGURE 2.14. TIME SERIES OF THE NUMBER OF CHLOROPHYLL PROFILES IN WOD23 OSD DATASET .....	60
FIGURE 2.15. TIME SERIES OF THE NUMBER OF ALKALINITY PROFILES IN WOD23 OSD DATASET	61
FIGURE 2.16. TIME SERIES OF THE NUMBER OF PARTIAL PRESSURE OF CARBON DIOXIDE PROFILES IN WOD23 OSD DATASET .....	61
FIGURE 2.17. TIME SERIES OF THE NUMBER OF DISSOLVED INORGANIC CARBON PROFILES IN WOD23 OSD DATASET.....	62
FIGURE 2.18. TIME SERIES OF THE NUMBER OF TRITIUM PROFILES IN WOD23 OSD DATASET....	62
FIGURE 2.19. TIME SERIES OF THE NUMBER OF HELIUM PROFILES IN WOD23 OSD DATASET....	63
FIGURE 2.20. TIME SERIES OF THE NUMBER OF $\Delta$ -HELIUM-3 PROFILES IN WOD23 OSD DATASET	63
FIGURE 2.21. TIME SERIES OF THE NUMBER OF $\Delta$ -CARBON-14 PROFILES IN WOD23 OSD DATASET .....	64
FIGURE 2.22. TIME SERIES OF THE NUMBER OF $\Delta$ -CARBON-13 PROFILES IN WOD23 OSD DATASET .....	64
FIGURE 2.23. TIME SERIES OF THE NUMBER OF ARGON PROFILES IN WOD23 OSD DATASET .....	65
FIGURE 2.24. TIME SERIES OF THE NUMBER OF NEON PROFILES IN WOD23 OSD DATASET .....	65
FIGURE 2.25. TIME SERIES OF THE NUMBER OF CHLOROFLUOROCARBON-11 PROFILES IN WOD23 OSD DATASET .....	66
FIGURE 2.26. TIME SERIES OF THE NUMBER OF CHLOROFLUOROCARBON-12 PROFILES IN WOD23 OSD DATASET .....	66
FIGURE 2.27. TIME SERIES OF THE NUMBER OF CHLOROFLUOROCARBON-113 PROFILES IN WOD23 OSD DATASET .....	67
FIGURE 2.28. TIME SERIES OF THE NUMBER OF $\Delta$ -OXYGEN-18 PROFILES IN WOD23 OSD DATASET .....	67
FIGURE 2.29. TIME SERIES OF THE NUMBER OF PLANKTON TOWS IN WOD23 OSD DATASET.....	68
FIGURE 3.1. TEMPORAL DISTRIBUTION OF HIGH-RESOLUTION CTD CASTS IN WOD23. ....	74
FIGURE 3.2. GEOGRAPHIC DISTRIBUTION OF HIGH-RESOLUTION CTD CASTS IN WOD23. ....	75
FIGURE 3.3. DISTRIBUTION OF HIGH-RESOLUTION CTD CASTS AT STANDARD DEPTH LEVELS IN WOD23.....	75

FIGURE 4.1. GEOGRAPHIC DISTRIBUTION OF XBT PROFILES IN WOD23: NUMBER OF PROFILES BY ONE-DEGREE SQUARES.....	84
FIGURE 4.2. TEMPORAL DISTRIBUTION OF EXPENDABLE BATHYTHERMOGRAPH (XBT) PROFILES IN WOD23.....	85
FIGURE 4.3. DISTRIBUTION OF EXPENDABLE BATHYTHERMOGRAPH (XBT) DATA AT STANDARD DEPTH LEVELS IN WOD23.....	85
FIGURE 4.4. XBT DATA CONTRIBUTION BY COUNTRIES IN WOD23. TOTALS FOR PANAMA AND LIBERIA INCLUDE DATA OBTAINED FROM MERCHANT SHIPS IN THE SHIP OF OPPORTUNITY PROGRAM (SOOP).....	86
FIGURE 5.1. TEMPORAL DISTRIBUTION OF EXPENDABLE CONDUCTIVITY, TEMPERATURE AND DEPTH (XCTD) CASTS IN WOD23 CTD AND OSD DATASETS.....	95
FIGURE 5.2. GEOGRAPHIC DISTRIBUTION OF EXPENDABLE CONDUCTIVITY, TEMPERATURE AND DEPTH (XCTD) CASTS IN WOD23.....	96
FIGURE 5.3. DISTRIBUTION OF EXPENDABLE CONDUCTIVITY, TEMPERATURE AND DEPTH (XCTD) DATA AT STANDARD DEPTH LEVELS IN WOD23.....	97
FIGURE 6.1. CASTS FROM DIFFERENT TYPES OF PROFILING FLOATS (PFL) IN WOD23.....	101
FIGURE 6.2. GEOGRAPHIC DISTRIBUTION OF PROFILING FLOATS (PFL) CASTS FOR THE PERIOD 1994-2022 IN WOD23.....	108
FIGURE 6.3. PROFILING FLOATS (PFL) DATA CONTRIBUTION BY COUNTRIES IN WOD23.....	108
FIGURE 6.4. TEMPORAL DISTRIBUTIONS OF PROFILING FLOAT DATA (PFL) CASTS IN WOD23.....	110
FIGURE 6.5. DISTRIBUTION OF PROFILING FLOAT DATA (PFL) DATA AT STANDARD DEPTH LEVELS BETWEEN 0 AND 2000M IN WOD23.....	111
FIGURE 7.1. TEMPORAL DISTRIBUTION OF MECHANICAL BATHYTHERMOGRAPH (MBT) PROFILES IN WOD23.....	117
FIGURE 7.2. GEOGRAPHIC DISTRIBUTION OF MECHANICAL BATHYTHERMOGRAPH (MBT) PROFILES IN WOD23.....	118
FIGURE 7.3. DISTRIBUTION OF MECHANICAL BATHYTHERMOGRAPH (MBT) DATA AT STANDARD DEPTH LEVELS IN WOD23.....	119
FIGURE 8.1. TEMPORAL DISTRIBUTION OF DIGITAL BATHYTHERMOGRAPH (DBT) PROFILES IN WOD23.....	124
FIGURE 8.2. GEOGRAPHIC DISTRIBUTION OF DIGITAL BATHYTHERMOGRAPH PROFILES IN WOD23.....	125
FIGURE 8.3. DISTRIBUTION OF DIGITAL BATHYTHERMOGRAPH DATA AT STANDARD DEPTH LEVELS IN WOD23.....	126
FIGURE 9.1. NUMBERS OF MRB DAILY PROFILES BY OBSERVING PROGRAM.....	128
FIGURE 9.2. TEMPORAL DISTRIBUTION OF THE MOORED BUOYS (MRB) PROFILES IN WOD23.....	133
FIGURE 9.3. GEOGRAPHIC DISTRIBUTION OF THE MOORED BUOYS (MRB) PROFILES COLLECTED BY MAJOR RESEARCH PROGRAMS IN WOD23 BY ONE-DEGREE SQUARES.....	133
FIGURE 9.4. DISTRIBUTION OF THE MOORED BUOYS (MRB) DATA AT STANDARD DEPTH LEVELS IN WOD23.....	134
FIGURE 10.1. DISTRIBUTION OF THE DRIFTING BUOY DATA IN WOD23 AMONG MAJOR RESEARCH PROGRAMS.....	143
FIGURE 10.2A. GEOGRAPHIC DISTRIBUTION OF THE DRIFTING BUOY (DRB) DATA (GLOBAL OCEAN) BY ONE-DEGREE SQUARES IN WOD23.....	144
FIGURE 10.3. TIME SERIES OF THE DRIFTING BUOY (DRB) CASTS AS A FUNCTION OF YEAR IN WOD23.....	145
FIGURE 10.4. DISTRIBUTION OF THE DRIFTING BUOY (DRB) DATA AT STANDARD DEPTH LEVELS IN WOD23.....	145
FIGURE 11.1. DISTRIBUTION OF THE UOR DATA IN WOD23 AMONG THE KNOWN CONTRIBUTING INSTITUTIONS.....	150
FIGURE 11.2. DISTRIBUTION OF THE UOR DATA IN WOD23 AMONG THE KNOWN CONTRIBUTING PROJECTS.....	150
FIGURE 11.3. TEMPORAL DISTRIBUTION OF UNDULATING OCEAN RECORDERS (UOR) CASTS IN WOD23.....	152
FIGURE 11.4. DISTRIBUTION OF THE UNDULATING OCEAN RECORDERS (UOR) DATA IN WOD23 AMONG THE CONTRIBUTING COUNTRIES.....	153

FIGURE 11.5. GEOGRAPHIC DISTRIBUTION OF UNDULATING OCEAN RECORDERS (UOR) CASTS IN WOD23 BY ONE-DEGREE SQUARES.....	154
FIGURE 11.6. DISTRIBUTION OF UNDULATING OCEAN RECORDERS (UOR) DATA AT STANDARD DEPTH LEVELS IN WOD23.....	155
FIGURE 12.1. TEMPORAL DISTRIBUTION OF APB CASTS IN WOD23.....	160
FIGURE 12.2. GEOGRAPHICAL DISTRIBUTION OF THE AUTONOMOUS PINNIPED BATHYTHERMOGRAPH (APB) DATA IN WOD23 BY ONE-DEGREE SQUARES.....	161
FIGURE 12.2. DISTRIBUTION OF THE AUTONOMOUS PINNIPED BATHYTHERMOGRAPH (APB) DATA AT STANDARD DEPTH LEVELS IN WOD23.....	161
FIGURE 13.1. TEMPORAL DISTRIBUTION OF MICRO BATHYTHERMOGRAPH DATA IN WOD23..	165
FIGURE 13.2. GEOGRAPHIC DISTRIBUTION OF MICRO BATHYTHERMOGRAPH DATA IN WOD23.	166
FIGURE 13.3. DISTRIBUTION OF MICRO BATHYTHERMOGRAPH DATA AT STANDARD DEPTH LEVELS IN WOD23.....	166
FIGURE 14.1. TEMPORAL DISTRIBUTION OF SURFACE (SUR) OBSERVATIONS IN WOD23. ....	169
FIGURE 14.2. GEOGRAPHIC DISTRIBUTION OF SURFACE (SUR) OBSERVATIONS BY ONE-DEGREE SQUARES IN WOD23. ....	<b>ERROR! BOOKMARK NOT DEFINED.</b>
FIGURE 14.3. GEOGRAPHIC DISTRIBUTION OF SURFACE (SUR) OBSERVATIONS BY ONE-DEGREE SQUARES COLLECTED BY EUROPEAN POLAR RESEARCHERS IN ARCTIC BEFORE 1900 YEAR IN WOD23.....	170
FIGURE 15.1 SEAGLIDER LICENSED TO KONGSBERG. PHOTO SOURCE .....	176
FIGURE 15.2. GEOGRAPHICAL DISTRIBUTION OF GLIDER (GLD) DATA IN WOD23: NUMBER OF CASTS PER 1°SQUARE.....	178
FIGURE 15.3. TEMPORAL DISTRIBUTION AND MAJOR SOURCES OF GLIDER (GLD) DATA IN WOD23.....	178
FIGURE 15.4. CONTRIBUTION OF GLIDER (GLD) DATA BY DIFFERENT SOURCES IN WOD23. ....	180
FIGURE 15.5. GEOGRAPHICAL DISTRIBUTION OF GLIDER (GLD) DATA SUBMITTED BY DIFFERENT PROGRAMS IN WOD23.....	181
FIGURE 15.6. DISTRIBUTION OF GLIDER (GLD) DATA AT STANDARD DEPTH LEVELS IN WOD23: A) FROM SURFACE TO 1500M DEPTH, B) DEEP GLIDERS (UW) – FROM 1000M TO 7200M DEPTH. ....	181
FIGURE 15.7. GLIDER (GLD) DATA CONTRIBUTION BY COUNTRIES IN WOD23. ....	182
FIGURE 16.1. AN EXAMPLE OF A PLANKTON CAST IN WOD23 (USING PROVIDED OUTPUT SOFTWARE). ....	188
FIGURE 16.2. AN EXAMPLE OF A PLANKTON CAST IN ‘CSV’ OUTPUT FILE AVAILABLE ON-LINE THROUGH THE WODSELECT.....	189
FIGURE 16.3. GEOGRAPHIC DISTRIBUTION OF PLANKTON (245,059 CASTS) IN WOD23. ....	196
FIGURE 16.4. TEMPORAL DISTRIBUTIONS OF PLANKTON CASTS IN WOD23 AS A FUNCTION OF YEAR. ....	196
FIGURE 16.5 CONTRIBUTIONS OF PLANKTON CASTS BY MEASUREMENT TYPE. ....	198
FIGURE 16.6. GEOGRAPHIC DISTRIBUTION OF ZOOPLANKTON NUMERICAL ABUNDANCE (46,224 CASTS) IN WOD23.....	200
FIGURE 16.7. GEOGRAPHIC DISTRIBUTION OF PHYTOPLANKTON NUMERICAL ABUNDANCE (37,961 CASTS) IN WOD23.....	200
FIGURE 16.8. GEOGRAPHIC DISTRIBUTION OF ICHTHYOPLANKTON NUMERICAL ABUNDANCE (54,286 CASTS) IN WOD23. ....	201
FIGURE 16.9. GEOGRAPHIC DISTRIBUTION OF BACTERIOPLANKTON NUMERICAL ABUNDANCE (1,986 CASTS) IN WOD23. ....	201
FIGURE 16.10. GEOGRAPHIC DISTRIBUTION OF TOTAL DISPLACEMENT VOLUME (125,022 CASTS) IN WOD23.....	203
FIGURE 16.11. GEOGRAPHIC DISTRIBUTION OF TOTAL SETTLED VOLUME (9,926 CASTS) IN WOD23.....	203
FIGURE 16.12. GEOGRAPHIC DISTRIBUTION OF TOTAL WET MASS (34,075 CASTS) IN WOD23. ..	204
FIGURE 16.13. GEOGRAPHIC DISTRIBUTION OF TOTAL DRY MASS (2,554 CASTS) IN WOD23.....	204
FIGURE 16.14. GEOGRAPHIC DISTRIBUTION OF TOTAL ASH-FREE DRY MASS (446 CASTS) IN WOD23.....	205



## LIST OF ACRONYMS

<b>Acronym</b>	<b>Expanded Term</b>
APB	Autonomous Pinniped Bathythermograph
AWI	Alfred Wegener Institute for Polar and Marine Research
BAMS	Bulletin of the American Meteorological Society
Argo	Core Argo program floats
BCG-Argo	BioGeoChemical Argo float, an extension of the Argo core program
CFA	Continuous Flow Analyzer
CMIP	Coupled Model Inter-comparison Project
CRM	Certified Reference Material
CSV	Comma-Separated Value
CTD	Conductivity Temperature Depth dataset in the WOD
DBT	Drifting Bathythermograph
DIVA	Data-Interpolating Variational Analysis
DNA	Deoxyribonucleic Acid
DOC	Department of Commerce
DOE	Department of Energy
DRB	Drifting Buoy dataset in the WOD
EEl	Earth Energy Imbalance
ENSO	El Niño-Southern Oscillation,
ERL	Earth Research Laboratory
EOV	Essential Ocean Variable
ETOPO2	Earth Topography 2 arc minute
EVR	Extended Vertical Resolution
FAIR	Findable, Accessible, Interoperable, and Reusable
GIS	Geographic Information System
GLD	Glider dataset in the WOD
GMT	Greenwich Mean Time, or Generic Mapping Tools
GLODAP	Global Ocean Data Analysis Project
GODAR	Global Ocean Data Archaeology and Rescue
GOBAl-O <sub>2</sub>	Gridded Ocean Biogeochemistry from Artificial Intelligence for O <sub>2</sub>
GPP	Gross Primary Production
GTSP	Global Temperature-Salinity Profile Program
IAPSO	International Association for the Physical Sciences of the Oceans
IOC	Intergovernmental Oceanographic Commission of UNESCO
IODE	International Oceanographic Data Exchange of IOC
IRI	International Research Institute for Climate and Society
IUPAC	International Union of Pure and Applied Chemistry
JAMSTEC	Japan Agency for Marine-Earth Science and Technology
JPOTS	Joint Panel on Oceanographic Tables and Standards
LDEO	Lamont-Doherty Earth Observatory
MAST	Marine Science and Technology
MBT	Mechanical Bathythermograph in the WOD
MEDAR	Mediterranean Data Archeology and Rescue
MRB	Moored Buoy dataset in the WOD
NAO	North Atlantic Oscillation
NASA	National Aeronautics and Space Administration
NATO	North Atlantic Treaty Organization
NPP	Net Primary Production
NCEI	National Centers for Environmental Information



<b>Acronym</b>	<b>Expanded Term</b>
NESDIS	National Environmental Satellite, Data, and Information Service
NOAA	National Oceanic and Atmospheric Administration
NODC	National Ocean Data Center
OA	Objective Analysis
OSD	Ocean Station Data dataset in the WOD
O2S	Dissolved oxygen saturation (%)
O2I	Dissolved oxygen inventory
OCL	Ocean Climate Laboratory Team
ODV	Ocean Data View (product developed by Dr. Reiner Schlitzer, AWI)
OMZ	Oxygen Minimum Zones
PFL	Profiling Float dataset in in the WOD
PIRATA	Prediction and Research Moored Array in the Tropical Atlantic
PDO	Pacific Decadal Oscillation
PSS	Practical Salinity Scale
QC	Quality Control
QCF	Quality Control Flags
RAMA	Research Moored Array for African-Asian-Australian Monsoon Analysis and Prediction
RNA	Ribonucleic Acid
SST	Sea Surface Temperature
SME	Subject Matter Expert
SOCCOM	Southern Ocean Carbon and Climate Observations and Modeling project
SUR	Surface
TAO/TRITON	Tropical Atmosphere Ocean moored buoy array
TSK	Tsurumi-Seiki Company
UN	United Nations
UNESCO	United Nations Educational, Scientific and Cultural Organization
UOR	Undulating Oceanographic Recorder in the WOD
USA	United States of America
USN	United States Navy
WDC	World Data Center
WDS	World Data System
WDS Oceanography	World Data Service for Oceanography of WDS hosted at NCEI
WOA	World Ocean Atlas series (WOA1994, 1998, 2001, 2005, 2009, 2013, 2018, and 2023)
WOA23	World Ocean Atlas 2023 (WOA23F and WOA23N))
WOA23F	World Ocean Atlas 2023 (1965-2022)
WOA23N	World Ocean Atlas 2023 Climate Normal (1971-2000)
WOCE	World Ocean Circulation Experiment
WOD	World Ocean Database series (WOD1994, 1998, 2001, 2005, 2009, 2013, 2018, and 2023)
WOD18	World Ocean Database 2018 used for WOA18
WOD23	World Ocean Database 2023 used for WOA23
XBT	Expendable Bathythermograph in the WOD
XCTD	Expendable Conductivity Temperature Depth in the WOD

## PREFACE

The oceanographic databases described by this atlas series expands on the *World Ocean Database 2023* (WOD23) product and its predecessors. WOD23 has expanded by including substantial amounts of both recent and historical data not previously available. Previous versions of the oceanographic databases developed by the National Centers for Environment Information (former Oceanographic Data Center, NODC)/World Data System (WDS), and products derived from these databases, have proven to be of great utility to the international oceanographic and climate research, and operational environmental forecasting communities. In particular, the objectively analyzed fields of temperature and salinity derived from these databases have been used in a variety of ways. These include use as boundary and/or initial conditions in numerical ocean circulation models, validation and verification of numerical simulations, analysis and reanalysis of the ocean, as a form of "sea truth" for satellite measurements such as altimetric observations of sea surface height among others. Increasingly, nutrient fields are being used to initialize and/or verify biogeochemical models of the world ocean. In addition, NCEI/WDS products are critical for support of international assessment programs such as the Intergovernmental Program on Climate Change (IPCC) of the United Nations.

It is well known that the amount of carbon dioxide (CO<sub>2</sub>) in the earth's atmosphere increased compared to pre-Industrial Revolution time. It is necessary that the scientific community has access to the most complete historical oceanographic databases possible in order to study climate change and its variability, ecosystem response to climate change, and for other scientific and environmental problems. Moreover, in situ data gathered at great expense should be available for future use, in line with the open science principles.

In the acknowledgment section of this publication we have expressed our view that creation of global ocean databases is only possible through the cooperation of scientists, data managers, and scientific administrators throughout the international community. In addition, we thank the scientists and technicians of the National Centers for Environmental Information (NCEI/NOAA) as well as of the NOAA Cooperative Institutions involved for their dedication to the project leading to publication of this atlas series. Their commitment has made this database possible.

The production of oceanographic databases is a major undertaking. Such work is due to open data sharing and the input of many individuals and organizations. We have tried to structure the data sets in such a way as to encourage feedback from experts who have knowledge that can improve the data and metadata contents of the database. It is only with such feedback that high-quality global ocean databases can be prepared. Just as with scientific theories and numerical models of the ocean and atmosphere, the development of global ocean databases is not carried out in one giant step, but proceeds in an incremental fashion.

*Ocean Climate Laboratory*  
NCEI  
Silver Spring, MD  
July 2024

## ACKNOWLEDGMENTS

This work was made possible by a grant from the NOAA Climate and Global Change Program which enabled the establishment of a research group, the Ocean Climate Laboratory (OCL), at the National Oceanographic Data Center, or NODC (now National Center of Environmental Information, or NCEI). The purpose of the OCL is to develop research quality oceanographic databases, as well as to compute objective analyses of, and diagnostic studies based on these databases. This work is funded in partnership with the NOAA's Global Ocean Monitoring and Observing Program (GOMO).

We acknowledge the scientists, technicians, and programmers who have submitted data to national and regional data centers as well as the managers and staff at the various data centers. Our databases allow to preserve the provenance information by storing a metadata which includes information about Principal Investigators to recognize their efforts.

The OCL expresses thanks to those who provided comments and helped develop an improved *World Ocean Database 2023* (WOD23) product. Any errors in WOD23 are the responsibility of the Ocean Climate Laboratory.

A special acknowledgement to Syd Levitus who initiated the World Ocean Database project and guided the OCL from its inception in early 1990 until 2013.

We are thankful to Dr. Viktor Gouretski, Senior Researcher at the Chinese Academy of Sciences (Beijing, China) for the review of the XBT chapter and providing his constructive comments and suggestions which significantly improved it.

We are grateful to Dr. Simona Simoncelli, Researcher at Istituto Nazionale di Geofisica e Vulcanologia (INGV- Bologna, Italy) and Dr. Rachel Killick, Ocean Observation Scientist at the Met Office (Exeter, Devon, GB) for reviewing this Atlas and providing their constructive comments and suggestions which significantly improved this document.

The views, opinions, and findings contained in this report are those of the authors, and should not be construed as an official NOAA or U.S. Government position, policy, or decision.

# CHAPTER 1: INTRODUCTION

*Alexey V. Mishonov, Tim P. Boyer, Olga K. Baranova, Courtney N. Bouchard, Scott L. Cross, Hernan E. Garcia, Ricardo A. Locarnini, Christopher R. Paver, James R. Reagan, Zhankun Wang, Dan Seidov, Alexandra I. Grodsky, and James Beauchamp*

*Ocean Climate Laboratory  
National Centers for Environmental Information  
Silver Spring, MD*

## ABSTRACT

The World Ocean Database 2023 (WOD23) is a collection of scientifically quality-controlled ocean profile and plankton data gathered from 1772 to 2022 all over the world ocean, which is updated with the newly collected data every three months in order to keep it as contemporary as possible. Data includes measurements of temperature, salinity, beam attenuation, oxygen, phosphate, nitrate, silicate, chlorophyll, alkalinity, pH, pCO<sub>2</sub>, TCO<sub>2</sub>, Tritium, Δ<sup>13</sup>Carbon, Δ<sup>14</sup>Carbon, Δ<sup>18</sup>Oxygen, Freon, Helium, Δ<sup>3</sup>Helium, Neon, and plankton. WOD23 is a global ocean database that can be used internationally without any restriction for studying a wide variety of environmental problems.

## 1.1. INTRODUCTION

### The beginning

The earliest data in the World Ocean Database (see insert at the right and Figure 1.1) were collected during Second Voyage round the World (WOD cruise number GB-11003) of Captain James Cook (WOD PI code 2061) onboard of the *HMS Resolution* (WOD Platform code 8152) and comprise of three stations each containing observations at the sea surface and 100 fathoms (183m) depth. Although there is no mention of the instruments used for these measurements in a cruise report of the days (Forster, 1777, pp 95-106), most likely these records were obtained by Hale apparatus with thermometer (WOD sampler code 775) as it has been done before during Capt. Cook's First Voyage round the World (according to J. Prestwich, 1875). These data were collected in the Southern Ocean and stored in the WOD's Ocean Station Data (OSD) dataset.

```
----- OSD -----
Longitude Latitude Year Month Day Time Cruise# CC Prof_#
22.000 -55.133 1772 12 15 ---- 11003 GB 1
  Num Depth Temp
  1 0.00 -1.100
  2 183.00 1.100

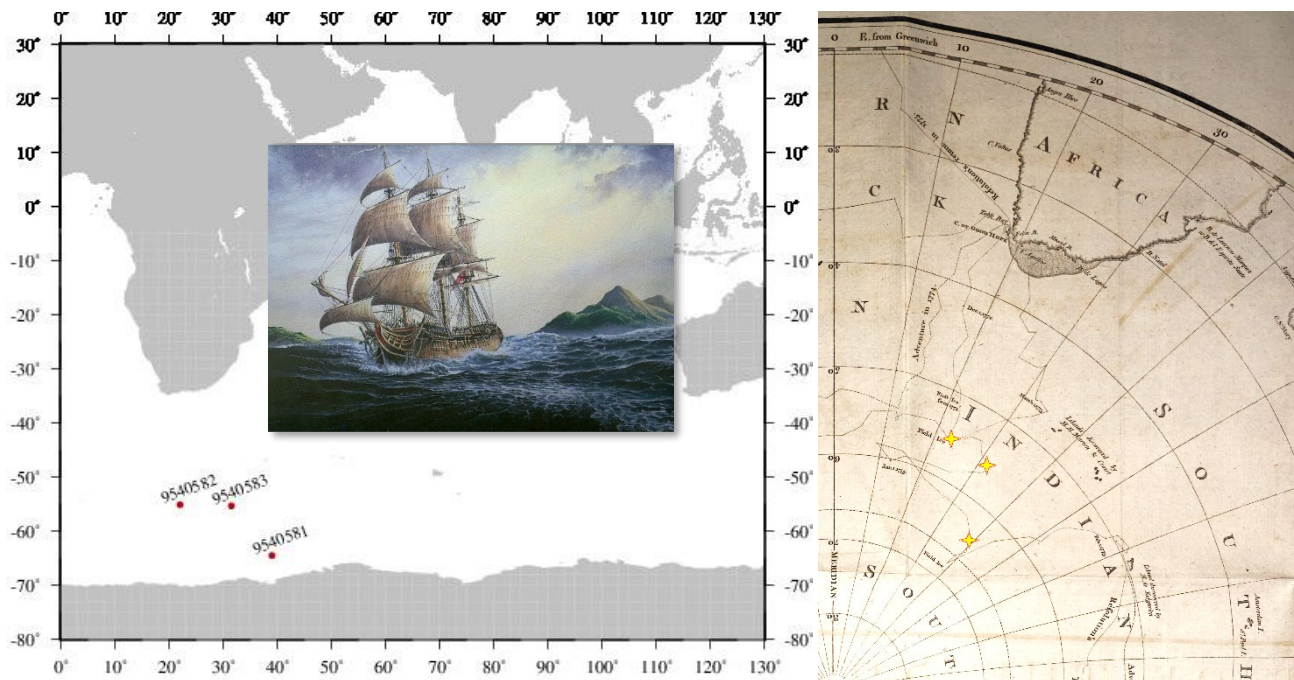
Access# 571
Platform 8152
Station_Number 9540582
T-S_Probe 7.000
NODCorig 3.000
water_sampler 775.000

31.500 -55.333 1772 12 23 ---- 11003 GB 2
  Num Depth Temp
  1 0.00 0.000
  2 183.00 1.300

Access# 571
Platform 8152
Station_Number 9540583
T-S_Probe 7.000
NODCorig 3.000
water_sampler 775.000

39.000 -64.500 1773 1 13 ---- 11003 GB 3
  Num Depth Temp
  1 0.00 0.800
  2 183.00 0.000

Access# 571
Platform 8152
Station_Number 9540581
T-S_Probe 7.000
NODCorig 3.000
water_sampler 775.000
```



**Figure 1.1. The location and WOD station numbers of the earliest data in WOD23.**

Locations of the seawater temperature measurements performed during *HMS Resolution* 1772-75 voyage shown on the WOD-generated map with WOD unique station numbers displayed (left) and superimposed over the expedition map cut-off from Forster, 1777 (right) with stations marks along the ship track. *HMS Resolution* artwork is reproduced from <https://i.pinimg.com/originals/a1/90/1f/a1901f851accdc13701840d05188f12d.jpg>

### 1.1.1. Purpose

The World Ocean Database (WOD) was first conceived as a way to provide reproducibility for the World Ocean Atlas (WOA) series. The WOA series is a continuation of the Climatological Atlas of the World Ocean (Levitus, 1982), which is a set of global gridded climatological mean fields of oceanographic variables on one- and quarter-degree longitude-latitude spatial grids at standard depth levels in the ocean. The WOA fields are to be used, among other things, as initial and boundary conditions for ocean and climate models. In order to produce comprehensive mean fields of oceanographic variables, aggregation of subsurface oceanographic measurements from many different sources, collected for many different reasons, using different instrumentation, different methods, different levels of calibration, quality monitoring, recording, formatting, metadata, units, and media delivery, were aggregated, converted to a uniform form, and quality controlled. The observation's original values and depths are preserved in the WOD along with quality flags assigned both by the data originator and by the WOD quality control procedures. Values interpolated from observed depths to standard levels are also included in the WOD, as these standard level interpolated values are the direct input into the WOA fields. Quality flags on standard level interpolated values are also included. These latter flags include subjective flags from the inspection and iterative calculation of the WOA gridded fields. Designated WOD quality flags on original observations denote whether the observation was used to interpolate to standard levels. Designated WOD quality flags on standard levels denote whether an interpolated value was used to calculate WOA gridded fields. Thus, WOD can be used to reproduce WOA. WOD has many further uses beyond reproducing climatological mean fields. For this reason, WOD is (since 2007)

updated on a quarterly basis in a preliminary state. The fully quality controlled WOD is released in conjunction with each release of the WOA.

### **1.1.2. Contributors**

The WOD is simply the final step in gathering oceanographic profile data together for public dissemination. The inventors, oceanographers, and engineers who conceived, designed, and tested the oceanographic instrumentation and measurement techniques are responsible for the plethora and variety of oceanographic data. The primary investigators, marine scientists and technicians, ship's crew, and volunteers who made and continue to make many of the oceanographic measurements, often under harsh conditions, are responsible for the quality and quantity of the oceanographic data. The institutions which maintain the platforms and the projects which plan, fund, and execute the field campaigns and operational ocean monitoring are responsible for the spatial and temporal coverage of the oceanographic profile data. Finally, the data managers are responsible for the preservation and reusability of the data. This is a vast network, maintained and updated over time which should receive the credit for the aggregated WOD. We have attempted to ensure that this credit is visible in the data itself. Every cast contains (when supplied) information on the instrumentation, platform, project, institution, and data management entity. The accession number, a number assigned to each data set received and archived at NCEI, is also found with each cast. More information about who submitted the data to NCEI and the original data themselves can be located using this accession number. The archive at NCEI and those who populate and maintain it also deserve credit for the continual availability of historical oceanographic data. Finally, international organizations such as the Intergovernmental Oceanographic Commission's (IOC) International Oceanographic Data and Information Exchange (IODE) and the World Data System (WDS) for Oceanography should be credited for creating and facilitating a global culture of data exchange and preservation.

The use of Data Object Identifiers (DOIs) and more specifically cascading DOIs shows great promise in a more succinctly and more easily accessible way to document credit for the different aspects of each oceanographic cast – from marine technicians, to primary investigators to institutions, projects, quality assurance, data managers, etc. In the future we intend to include a DOI (when possible) with each cast for a particular data set (right now represented by the accession number). This DOI will point to a set of related DOIs which document the different contributors and their role in the making of the given measurements.

### **1.1.3. Size and shape**

At the time of this release (i.e. 2024), WOD incorporates 27,299 different data sets received and archived at NCEI. The data were collected by different instrument types (see Table 1.1) and represent the results of 237,525 oceanographic cruises on 11,510 different platforms from 1,529 institutes around the world and 768 separate projects. The number of platforms is lower than might be expected, as some platform identifiers are generic (e.g. profiling float, moored buoy, drifting buoy, instrumented mammals, *etc.*) since the platform list is mainly used to identify ships and other named platforms (e.g. gliders). There are 3.13 billion individual profile measurements pairs (depth/pressure vs. measured variable) of temperature, 2.2 billion of salinity, 692 million of oxygen, and 4.5 million plankton measurements. These measurements make up the 18.6 million



oceanographic casts in the WOD23, which spread unevenly over the world ocean (Figure 1.2). There are an additional 22.7 million meteorological/sea state observations (see Table 1.2).

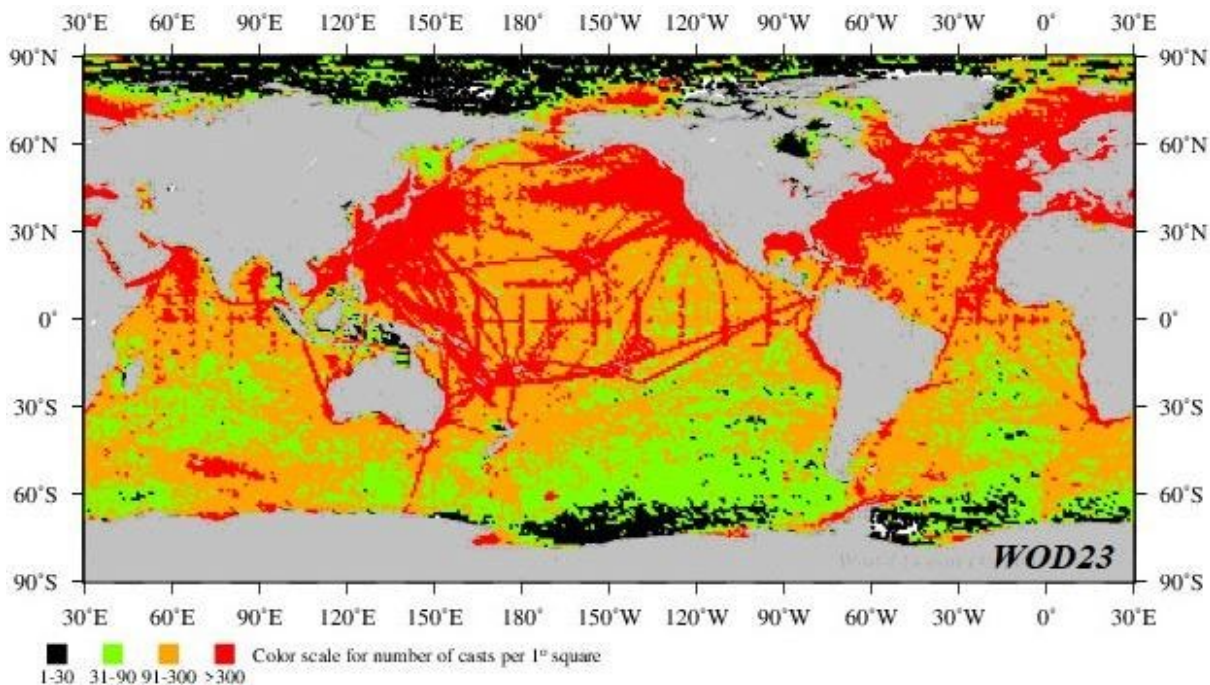


Figure 1.2. WOD23 observations density (profiles per 1-degree square).

### 1.1.4 Data Organization

Data in the WOD are organized using the following operational definitions:

**Profile:** A set of measurements for a single variable (temperature, salinity, *etc.*) at discrete depths taken as an instrument is being lowered or raised vertically in the water column. For surface-only data, the profile consists of measurements taken along a horizontal path. For moored buoys and drifting buoys, the instrument does not move vertically in the water column, so a profile is a discrete set of concurrent measurements from the instruments placed at different depths on a wire attached to the buoy.

**Cast:** A set of one or more profiles taken concurrently or nearly concurrently. Meteorological and other ocean data, *e.g.* Secchi disk depth data, are also included in a cast if measurements were taken concurrently with the profile(s). Observations and measurements of plankton from net-tows are included if taken concurrently or in close time proximity to profiles. If there are no profiles in close proximity, a net-tow by itself will constitute a cast. Each cast in the WOD is assigned a unique cast number. If the cast is subsequently replaced by higher quality data, the unique cast number is inherited. If any alteration is made to a cast, this information is noted in comments to the quarterly database update, referenced by the unique cast number. For surface-only data in dataset SUR, a cast is defined as a collection of concurrent surface measurements at discrete latitudes and longitudes over an entire cruise (see definition of cruise below). Latitude, longitude and Julian year-day values are included with each set of measured oceanographic variables.

**Station:** A particular geographic location at which one or more casts are taken.

**Cruise:** A cruise is defined as a specific deployment of a single platform for the purposes of a coherent oceanographic investigation. For an oceanographic research vessel, this deployment is usually well-defined with a unique set of scientific investigators collecting data for a specific project or set of projects. In some cases, different legs of a deployment with the same equipment and investigators are assigned different cruise numbers, as per the investigator's designation. In the case when merchant ships-of-opportunity (SOO) are used for data collection, a cruise is usually defined as the time at sea between major port calls. Profiling floats, instrumented pinnipeds, moored buoys, and drifting buoys are assigned the same cruise number for the life of the platform (life of the sensor package for pinnipeds). For gliders, each specific deployment is designated a cruise. For surface-only data in dataset SUR, a cast and cruise are the same, except for 27 cruises which were split into 2 casts each due to the large number of sets of measurement (> 24,000). A set of stations is grouped together if they fit the "cruise" definition.

In the WOD, a cruise identifier consists of two parts, the country code and the unique cruise number. The unique cruise number is only unique with respect to the country code. The country code is usually assigned based on the flag of the data collecting ship. If the platform from which data were collected was not a ship, (e.g. a profiling float, glider, drifting or moored buoy), the country of the primary investigator or institute which operates or releases the platform is used. The International Standards Organization (ISO) country codes are used (see <https://www.iso.org/publication/PUB500001.html>). For data for which no information on country is present, a country code of 99 is used. For data for which there is no way to identify a specific cruise, a cruise number of zero (0) is used.

All data grouped as a cruise are listed under one unique country code/unique cruise number combination. It is possible and desirable to get all bottle (OSD), high-resolution Conductivity-Temperature-Depth (CTD), bathythermograph (BT), expendable bathythermograph (XBT & XCTD), and towed-CTD (UOR) data for a cruise using one unique cruise identifier when those data were collected simultaneously during one cast (i.e. OSD and CTD), or between the cast on passage (XBT, BT, UOR) from the same ship during the same expedition. However, there are still cases for which BT data have a different cruise identifier. It is an ongoing project to match these BT data with the correct cruise where bottle and high-resolution CTD data were also gathered.

**Accession Number:** A group of stations received and archived at the NCEI. Each collection submitted to NCEI is given a unique "accession number". Using this number, a user can get an exact copy of the original data sent to NCEI as well as information about the data itself (i.e. metadata) through the NCEI Geoportal (<https://www.ncei.noaa.gov/metadata/geoportal/>). Cruises are not always subsets of accession numbers, as data from the same cruise may have multiple accession numbers. Each cast has an associated accession number (with a few exceptions). If data from a cast is replaced later by higher quality data, the new accession number will reflect the novel source of the data while the unique cast number will remain unchanged. If a profile for a variable not previously stored with a cast becomes available, the profile will be added to the existing cast, and a variable-specific accession number will be added to the station to record the source of the new profile.

**WOD Dataset:** All casts from similar instruments with similar resolution. For instance, all data acquired by bathythermographs (BTs) which are dropped over the side of a ship on a winch and recovered reside in the MBT dataset, all CTD data collected at high vertical depth resolution



(relatively small depth increments) are stored in the CTD dataset. For convenience, each dataset is stored in a separate file in WOD.

When requested, NCEI can add a DOI to new data being archived if one does not already exist. The [NCEI data archive](#) is searchable and spans a broad spectrum of scientific disciplines, archive methods, naming conventions, and file formats. Detailed information about submitting data to NCEI can be found on [Send2NCEI](#) web site. The data in the WOD data are Findable, Accessible, Interoperable, and Reusable (FAIR, Wilkinson *et al.* 2016).

### 1.1.5. WOD Datasets

The WOD datasets group together data acquired in a similar manner. So, bottle data and low vertical resolution CTD casts are grouped together since bottle casts often include temperature and salinity measurements from CTDs only at the depths at which bottles were tripped. High resolution CTD data are stored in a separate dataset because of their high volume. The low-resolution version of the CTD data is often available in bottle data casts. Cases where high and low-resolution CTD data are available in different datasets are identified as high-resolution pairs. The secondary header code 13 is used to link such profiles. For more details see Table 4a in Garcia *et al.* (2024).

The WOD datasets are briefly described below and in more detail in following chapters. WOD's instrument types and datasets are listed in Table 1.1. The three-letter notation for each dataset is the abbreviation used for the naming of the output data files. Note that not every particular instrument used for data acquisition has a dedicated separate dataset to hold the data, and that the three-letter dataset notation does not always reflect all diversity of instrumentation used for gathering the data found in the dataset. More detailed data descriptions and relevant oceanographic information can be found in chapters 2-16 of this document, and in the bibliographies and references provided for each chapter. For a description of the instrument codes as well as for other codes embedded in the data format, see Garcia *et al.* (2024).

The WOD includes oceanographic variables measured at “**observed**” depth levels as well as interpolated to a set of 137 “**standard**” depth levels. All climatic fields in the World Ocean Atlas (WOA) are produced based on “**standard**” depth levels data. Note that the 40 standard depth levels used in previous versions of WOD (before 2013) are all among the 137 standard depth levels used in WOD23, to provide continuity.

### **OSD Dataset – Ocean Station Data, low-resolution CTD, low-resolution XCTD, plankton tows**

#### *i.) Ocean Station Data*

Ocean Station Data has historically referred to measurements made from a stationary (anchored or drifting) research ship using reversing thermometers and water samples collected from bottles tripped at depths of interest in the water column. The water samples are analyzed to measure variables, including water salinity, oxygen, nutrients (phosphate, silicate, nitrate plus nitrite), chlorophyll, pCO<sub>2</sub>, TCO<sub>2</sub>, and tracers (Tritium,  $\Delta^{13}\text{Carbon}$ ,  $\Delta^{14}\text{Carbon}$ , Freons, Helium,  $\Delta^3\text{Helium}$ ,  $\Delta^{18}\text{Oxygen}$ , and Neon) concentrations. The two most commonly used bottle types are the [Nansen](#) and [Niskin](#) (see Chapter 2).

ii.) *Low-resolution CTD data*

Conductivity-Temperature-Depth (CTD) instruments are a combination of a pressure sensor (measured pressure is converted to depth), a resistance temperature measurement device (usually a platinum thermometer), and a conductivity sensor used to calculate salinity. CTDs are usually mounted on a metal frame ([a rosette](#)) and lowered through the water column suspended from a cable. The frame is often used to hang bottles for collecting water samples. Low-resolution here refers to a limited number of temperature and/or salinity measurements made along the vertical profile. Usually, but not always, these measurements are recorded at the depths at which bottles are tripped to collect water samples. This dataset also includes data from the older Salinity-Temperature-Depth (STD) instruments - the precursor to the CTD. About 5.66% of all data in the OSD dataset (184,156 of 3,256,037 profiles) are listed as containing low-resolution temperature and/or salinity data measured by CTD/STD (see Chapter 3).

iii.) *Low-resolution Expendable CTD (XCTD data)* – see description below under **CTD Dataset** and in Chapter 5.

iv.) *Plankton tow* – net tows or bottle casts from which plankton counts and/or biomass observations were taken (see Chapter 14).

**Table 1.1. Instrument types in the WOD23**

<b>DATASET</b>	<b>Instrument Type</b>
OSD	Bottle, low-resolution Conductivity-Temperature-Depth (CTD), low-resolution XCTD data, and plankton data
CTD	High-resolution Conductivity-Temperature-Depth (CTD) data and high-resolution XCTD data
MBT	Mechanical Bathythermograph (MBT) data, Digital BT (DBT), micro-BT ( $\mu$ BT)
XBT	Expendable (XBT) data
SUR	Surface only data (bucket, thermosalinograph)
APB	Autonomous Pinniped Bathythermograph - Time-Temperature-Depth recorders and CTDs attached to elephant seals
MRB	Moored buoy data mainly from the Equatorial buoy arrays -TAO
PFL	Profiling float data, mainly from the Argo program
DRB	Drifting buoy data from surface drifting buoys with thermistor chains and from ice-tethered profilers
UOR	Undulating Oceanographic Recorder data from a Conductivity/Temperature/Depth probe mounted on a towed undulating vehicle
GLD	Glider data

**CTD Dataset – High-resolution CTD (CTDs and XCTDs recorded at high depth/pressure frequency)**

i.) *High-resolution Conductivity-Temperature-Depth (CTD) data*

High-resolution CTD data consist of temperature and salinity profiles recorded at high frequency with respect to depth or pressure. These records are usually binned (averaged) in 1 to 5m depth interval mean values by the data submitter, although some means are calculated using

smaller depth intervals. Often the high-resolution CTD cast (usually recorded during descend of the rosette) has a low-resolution counterpart in the OSD dataset with accompanying records of different variables derived from the processing of the sea water samples taken by the bottle(s). Bottles are usually closed at the different depths during the ascent of the rosette and concurrent CTD readings recorded at each closing. In these cases, both the high-resolution CTD and the OSD data have a marker identifying these data as coming from the same station ('hi-res pair' - second header code # 13 in the WOD native format). High-resolution measurements of dissolved oxygen, chlorophyll (from a fluorometer), and beam attenuation coefficient (BAC) from a transmissometer are also included in this dataset when available. Note that in many cases the dissolved oxygen and chlorophyll data are uncalibrated and not of high quality. Information on whether these variables are calibrated is not usually supplied by the data submitter (see Chapter 3).

ii.) *High-resolution Expendable Conductivity-Temperature-Depth (XCTD) data*

Expendable Conductivity-Temperature-Depth ([XCTD](#)) probes are similar to XBT instruments (described below) - they are a torpedo-shaped device attached to a spool of copper wire. Along with the thermistor found in the XBT, a conductivity sensor is used to estimate salinity. XCTD instruments are produced by [Lockheed Martin/Sippican, Inc.](#) (Lockheed Martin, U.S.A.) and by [Tsurumi Seiki Co., Ltd.](#) (TSK, Japan). The standard XCTD has a manufacturer-specific drop-rate equation error (Johnson, 1995; Mizuno and Watanabe, 1998, Kizu *et al.*, 2008). Depth corrections for both manufacturers are incorporated in the standard level dataset. Air dropped and submarine discharged XCTDs have no known drop-rate problems. Data from XCTD instruments are included in the CTD dataset (see Chapter 5). XCTD casts make up about 1.1% of the CTD dataset (12,741 of 1,132,680 profiles).

### **XBT Dataset – low and high-resolution Expendable Bathythermographs**

Expendable Bathythermograph ([XBT](#)) probes are torpedo-shaped devices attached to a spool of copper wire. The instrument is launched over the side of a moving ship, from an airplane, or from a submarine. Temperature is estimated by measurements of the resistance in a semiconductor (called a thermistor). For recording the information is sent back to the command unit over the copper wire. Depth is calculated as a function of time since launch using a manufacturer-supplied equation. When the wire has unspooled, the copper wire breaks. There are two major manufacturers of XBTs, Sippican in the United States (original developer) acquired by Lockheed Martin, U.S.A., and [Tsurumi Seiki Co., Ltd](#) in Japan. A third manufacturer, Sparton, is no longer in business. XBTs have been deployed since 1966. XBTs were the major ocean observing system for sea water temperature measurements from 1967 to 2001 (advent of the Argo program). Many researchers (e.g. Heinmiller *et al.*, 1983) reported a systematic error in the recorded depths for XBT drops. Hanawa *et al.* (1995) published depth corrections for XBT types T-4, T-6, and T-7. More recently, there has been a great deal of research into time (year) dependent fall-rate and temperature biases, spurred by a paper by Gourestki and Koltermann (2007) which showed that there was a year to year difference in the XBT bias. For more information on these studies and more on how the XBT bias is handled in the WOD, please see Chapter 4 and the NCEI [XBT bias webpage](#).

### **PFL Dataset – Profiling floats**

Profiling floats are platforms drifting at a predetermined subsurface pressure level in the water column, rising to the surface at set time intervals. Pressure, temperature, salinity, and sometimes

dissolved oxygen and other biogeochemical measurements (BGC-Argo) taken on the ascent or previous descent are transmitted to the designated satellite. Most profiling floats are now operated as part of the Argo project.

Argo profiling float data are obtained from the Argo U.S. Global Data Assembly Center with smaller non-Argo contributions from the World Ocean Circulation Experiment ([WOCE](#)) and the Global Temperature and Salinity Profile Program ([GTSP](#)) Each quarter all new Argo cycles are downloaded from the US-GODAE GDAC and we added them to the PFL file. At the same time, we also get all Argo cycles that already exist in the PFL file but that have been updated since the previous quarter. The latest includes all cycles for which the data mode for one or more variables and the associated data, have been updated,

### **MBT Dataset – Mechanical Bathythermographs, Digital Bathythermographs (DBT), and Micro-bathythermographs ( $\mu$ BT).**

These types of the instruments are no longer in use by oceanographers, so no new data were incorporated into the World Ocean Database since the WOD18 release. Several MBT profiles (56, 0.002% of the total) were removed from the database based on quality check.

#### *i.) Mechanical Bathythermographs*

Mechanical Bathythermographs (MBT) were developed in their modern form around 1938 (Spilhaus, 1938). The instrument provides estimates of temperature as a function of pressure, (which is a proxy for depth) in the upper ocean. Earlier versions of the instrument were limited to making measurements in the upper 140 m of the water column. The last version of this instrument reached a maximum depth of 295 m. Initially, MBTs recorded temperature as a function of pressure (depth) by scratching a line on a smoked glass plate with a stylus, later recording was on paper. Pressure was determined from a pressure-sensitive tube known as a Bourdon tube. MBTs could be dropped from a ship moving at low speed. The accuracy of an MBT is about 0.3°C (see Chapter 7).

#### *ii.) Digital Bathythermographs*

A bathythermograph (developed in Japan) digitally records depth-temperature pairs as it is lowered in the water column. These instruments were used mostly by the Japanese in the mid-1970s and the 1980s in the Pacific Ocean, and less extensively by the Canadians in the North Pacific and North Atlantic (see Chapter 8).

#### *iii.) Micro-Bathythermograph*

Bathythermographs designed to record depth-temperature pairs at high vertical or temporal resolution (see Chapter 13).

### **MRB Dataset – Moored buoys**

Moored buoys are platforms, which are anchored or otherwise stabilized to measure oceanographic and atmospheric data in a small area around a fixed geographic location. Measurement devices are suspended at subsurface levels from a chain attached to the buoy. Temperature is measured using thermistors. Salinity is measured using conductivity sensors similar to those used in standard CTDs.

The moored buoy dataset include data from the Tropical Atmosphere-Ocean ([TAO](#)) buoy array

(in the tropical Pacific), the [TRITON](#) buoy array (in the western tropical Pacific), the [PIRATA](#) buoy array (in the tropical Atlantic), the [RAMA](#) buoy array (in the tropical Indian Ocean), [MARNET](#) buoys and light-ships (in the North Sea and the Baltic Sea). The data in the WOD from the TAO, PIRATA, RAMA and most of the TRITON buoys are daily averages (which is considered sufficient sampling frequency for climatological applications) derived from the original 15-minutes averages submitted by NOAA Pacific Marine Environmental Laboratory ([PMEL](#)). The remainder of the TRITON buoys, the MARNET buoys and light-ships data were acquired via [GTSP](#) (Chapter 9). It should be noted that MRB data submitted under Accession 0063240 in WOD were recorded at the 15-min interval; in 2021 those data were re-converted as daily averages. The daily average profiles matched with the first unique station number each day, and the extra stations were removed, hence total of MRB record counts in WOD23 decreased compared to WOD18.

### **DRB Dataset – Drifting buoys**

Drifting buoys are platforms which are advected by ocean currents, either at the surface, or at predetermined (usually shallow) depths. Drifting buoy data included in WOD23 were acquired from GTSP database, from the Japanese Arctic Buoy program archive, and from the Woods Hole Ice-Tethered Profiler Program (Toole *et al*, 2011). The GTSP data are from the subset of oceanic drifting buoys, which have multiple subsurface temperature measurement devices (thermistors) suspended from a chain, the others are ice drifters with profilers attached (see Chapter 10).

### **UOR Dataset – Undulating Oceanographic Recorders (Towed CTDs)**

Undulating Oceanographic Recorders are specific types of oceanographic vehicle which are towed behind a vessel while ascending and descending in the water column, recording temperature, salinity, and other variables at high vertical and horizontal resolution (see Chapter 11).

### **APB Dataset – Autonomous Pinniped Bathythermographs**

The earliest data in this WOD dataset (circa 1997) consist of bathythermographs attached to sea elephants. Later, CTD-instruments were used. Temperature and salinity information are recorded during dives taken while feeding and transmitted to satellite upon surfacing (see Chapter 12).

### **SUR Dataset – Surface-only data**

Surface-only data are either data taken using some type of bucket, or data from [thermosalinographs](#) connected to the water intake of the ship. These data are not the focus of the WOD. Only selected surface datasets which contained data from specific time periods and ocean areas which were not otherwise well covered by profile data are included in WOD. Note that a “cast” here refers to an entire cruise of surface-only measurements (see Chapter 14).

### **GLD Dataset - Gliders**

GLD dataset contains data collected from reusable autonomous underwater vehicles (AUV) designed to glide from the ocean surface to a programmed depth and back while measuring temperature, salinity, depth-averaged current, and other variables depending on the instrument’s suite (i.e. oxygen, chlorophyll, *etc.*) along a saw-toothed trajectory through the water (see Chapter 15).

## **Plankton Dataset**

The plankton subset in WOD23 includes the content of the previously released WOD05, WOD09, WOD13, WOD18 (Baranova *et al.*, 2013, 2009, 2005), WOD01 (O'Brien *et al.*, 2001), and WOD98 (Conkright *et al.*, 1998). The WOD23 plankton data subset is a collection of measurements from serial bottle and plankton net-tow. The plankton measurements are represented in WOD23 as quantitative and qualitative abundance, and biomass data. The plankton measurements are stored in the OSD dataset (see Chapter 2). Because plankton data now being collected in the specialized databases (i.e. [COPEPOD DB](#) or [others](#)), the content of the WOD23 plankton subset has not been updated since the WOD18 release.

## **Meteorological and Sea state measurements data**

Ship-based research cruises, and some operational oceanographic cruises, recorded atmospheric measurements (air temperature, barometric pressure, wind speed and direction, humidity, visibility, weather conditions, cloud cover and type) and sea state measurements (sea state condition, wave height and period, transparency). Additionally, many of the Equatorial moored buoys record air temperature, wind speed and direction. These measurements augment and complement the subsurface oceanographic profile data. Counts of each meteorological/sea state variable are shown in Table 1.2. In addition to auxiliary information for understanding ocean profile data, these measurements are important in their own right. In addition to retention in the WOD, all atmospheric and sea state measurements are available through the International Comprehensive Atmospheric and Oceanographic Data Set (ICOADS; Freeman *et al.*, 2016).



**Table 1.2. Meteorological and Sea-state parameters stored in the WOD23**

<b>Variables (2<sup>nd</sup> header code#*, units)</b>	<b>OSD</b>	<b>MBT</b>	<b>XBT</b>	<b>CTD</b>	<b>MRB</b>	<b>UOR</b>	<b>DRB</b>	<b>Total</b>
Bottom depth (10, m)	1,823,773	616,061	514,981	606,207	37,547	38,848	-	<b>3,637,424</b>
Water color (14, Forel-Ule color scale number)	282,347	12,438	476	11,315	-	-	-	<b>306,576</b>
Secchi disk visibility depth (15, m)	446,821	12,175	452	16,506	-	-	-	<b>475,954</b>
Wave direction (16, WMO 0877)	361,803	30,033	33,774	8,575	1,952	-	-	<b>436,137</b>
Wave height (17, WMO 1555)	229,309	114,357	51,196	27,455	1,956	32,441	-	<b>456,714</b>
Sea state (18, WMO 3700)	571,051	478,536	57,163	34,250	-	-	-	<b>1,141,000</b>
Wind force (19, Beaufort Scale)	604,603	14,444	3,279	5,206	-	16,667	-	<b>644,199</b>
Wave period (20, WMO 3155 or NODC 0378)	134,413	34,771	44,507	18,284	1,922	32,478		<b>266,375</b>
Wind direction (21, WMO 0877)	1,246,007	653,770	165,285	88,845	742,588	32,857	67,382	<b>2,996,734</b>
Wind speed (22, knots)	610,967	673,474	170,063	69,858	748,788	16,321	67,464	<b>2,356,935</b>
Barometric pressure (23, millibar)	765,439	338,253	33,353	83,016	22,837	33,020	45,187	<b>1,321,105</b>
Dry bulb temperature (24, °C)	1,152,900	622,998	148,389	73,912	797,735	27,727	47,443	<b>2,871,104</b>
Wet bulb temperature (25, °C)	233,569	495,851	52,403	43,691	-	27,031	-	<b>852,545</b>
Weather condition (26, WMO 4501 and WMO 4677)	656,346	514,994	49,760	45,004	-	33,016	-	<b>1,299,120</b>
Cloud type (27, WMO 0500)	364,186	25,589	17,512	26,026	-	-	-	<b>433,313</b>
Cloud cover (28, WMO 2700)	707,891	524,098	32,361	47,299	-	28,705	-	<b>1,340,354</b>
Horizontal visibility (41, WMO 4300)	104,089	185,594	4,160	24,981	-	-	-	<b>318,824</b>
Reference/Sea surface temperature (46, °C)	23,889	1,171,328	117,317	570	6,129	-	68844	<b>1,388,077</b>
Absolute air humidity (45, g m-3)	95,718	1,768	-	677	-	-	-	<b>98,163</b>
Sea surface salinity (47)	2	2,615	12,380		5,571			<b>20,568</b>

\* <https://www.ncei.noaa.gov/access/world-ocean-database/wod-codes.html#second>

### 1.1.6. Economic and scientific justification for maintaining archives of historical oceanographic data: the value of stewardship

Oceanography is an observational science, and it is impossible to replace historical data that have been lost. From this point of view, historical measurements of the ocean are priceless. However, in order to provide input to a “cost-benefit” analysis of the activities of oceanographic data centers and specialized data rescue projects, we can estimate the costs incurred if we wanted to resurvey the world ocean today, in the same manner as represented by the WOD Ocean Station Data (OSD) dataset.

The computation we describe was first performed in 1982 by Mr. Rene Cuzon du Rest, of the former NODC. Here we use an updated average operating cost estimate of \$30,000-\$35,000 per day for a regional class U.S. research ship ([2023 NSF budget to Congress](#)) with a capability to make two ‘deep’ or ten ‘shallow’ casts per day. We define a ‘deep’ cast as extending to a depth of more than 1000 m and a ‘shallow’ cast as extending to less than 1000 m depth. This is an arbitrary definition, but we are only trying to provide a coarse estimate of replacement costs for the WOD23 database. Using this definition, the WOD23 OSD dataset contains approximately 2.31 million ‘shallow’ casts (deepest depth between 2 and 1000 m depth) so that the ship time needed to perform these ‘shallow’ measurements is ~231,000 ship-days or ~633 ship-years, and the cost of the ship time is approximately *\$6.93 billion*. In addition, the WOD23 contains 0.43 million profiles deeper than 1000 m depth, which is estimated as ~217,000 ship-days or about 595 ship-years and the cost of the ship time to make these deep measurements is approximately *\$6.51 billion*. Thus, the total replacement cost of the OSD dataset in WOD23 is about *\$13 billion*. This amount of money is based only on ship-time operating costs, and does not include wages for scientists, technicians, samples/data processing expenses, or any other overhead costs. More recent platforms and instruments, such as XBTs dropped from volunteer merchant ships or autonomous platforms such as Argo floats or gliders, have much lower costs than traditional OSD/CTD and in some cases are publicly available in near-real time. Still, the long-term preservation and maintenance of these data also has a significant economic value for the future studies of historic oceanographic conditions.

### 1.1.7. Data assembly

It is not uncommon in oceanography that measurements of different variables made from the same sea water samples are often maintained as separate databases by different Institutions or principal investigators. In fact, data from the same oceanographic cast may be located at different organizations in different countries. From its inception, NODC/NCEI recognized the importance of building oceanographic databases in which as much data from each station and each cruise as possible are placed together into common standard formats, accompanied by appropriate metadata that make the data useful to future generations of scientists. It was the existence of such databases that allowed the *International Indian Ocean Expedition Atlas* (Wyrтки, 1971) and *Climatological Atlas of the World Ocean* (Levitus, 1982) to be produced without the time-consuming, laborious task of gathering data from many different sources. Part of the development of the WOD has been to expand this data assembly activity by increasing the number of variables as part of standardized databases – even if the variables do not have corresponding WOA fields.

### 1.1.8. Distribution media

WOD is being distributed through the NCEI archive (<https://catalog.data.gov/dataset/ncei-standard-product-world-ocean-database-wod>). For consistency with earlier releases, the WOD23



is also available in a native ASCII format which makes the most efficient use of space on storage media needed to transfer data to users (see <https://www.ncei.noaa.gov/products/world-ocean-database>, yearly and geographically sorted). To further minimize storage space requirements for the ASCII format, the data have been compressed with the **gzip** utility. The archived version of the WOD is in netCDF format, which follows the Climate-Forecast ragged-array convention (<http://cfconventions.org/>). Due to the importance of keeping all oceanographic variables, metadata, meteorological data, and sea state information for a cast together, the CF convention, which require to have all variables in a file use the same array dimension is not entirely followed. For more information on data format see [WOD23 Manual](#) (Garcia *et al.*, 2024).

### 1.1.9. Application software interfaces

Understanding that not all users are comfortable with netCDF files, nor with WOD native ASCII, and to increase the re-usability of WOD data, we have available a suite of programs (in C and FORTRAN) which demonstrate how to convert the ASCII data files to forms, such as comma-separated value (csv) which can be used in standard software packages. A program for converting and displaying data from the netCDF files is also available (see ‘WOD Programs’ section on the [WOD portal](#)). Python routines for reading the WOD native ASCII format have been provided by the International Quality Controlled Database ([IQuOD](#)) project.

### 1.1.10. Units

WOD23 uses *per kilogram* units, i.e.  $\mu\text{mol}\cdot\text{kg}^{-1}$  for oxygen, phosphate, total phosphorus, silicate, nitrite, nitrate,  $\text{nanomoles}\cdot\text{kg}^{-1}$  for helium, argon, and neon, and  $\text{picomoles}\cdot\text{kg}^{-1}$  for CFC-11, CFC-12, CFC-12. While titrated measurements of molecules in sea water are performed on water volumes, it is common practice for oceanographers to use mass-based values as these are directly comparable regardless of a volume of waters dependence on temperature and salinity. Note that the conversion to *per kg* from *per liter* was executed using a constant density ( $1025 \text{ kg}\cdot\text{m}^{-3}$ ) to ensure a simple reconversion to original units if desired.

## 1.2. COMPARISON OF WOD23 WITH PREVIOUS RELEASES

Table 1.3 and Figure 1.3 show the amount of data available from different dataset types that were used in earlier global oceanographic analyses and growth of the WOD data holding. During the past five years, the archives of historical oceanographic data have grown due to special data management and data observation projects that we discuss in section 1.3 of this document, regular routine submissions by scientists and operational ocean monitoring programs as well as due to increasing use of the autonomous data collecting platforms like Argo floats and gliders. With the distribution of the WOD there are now approximately *18.3 million* temperature profiles and *11.0 million* salinity profiles (as well as other profile data and plankton data) available to the international research community in a common format with associated metadata and quality control flags. There has been a net increase of about *1.8 million* oceanographic profiles since publication of *World Ocean Database 2018* (Boyer *et al.* 2018).

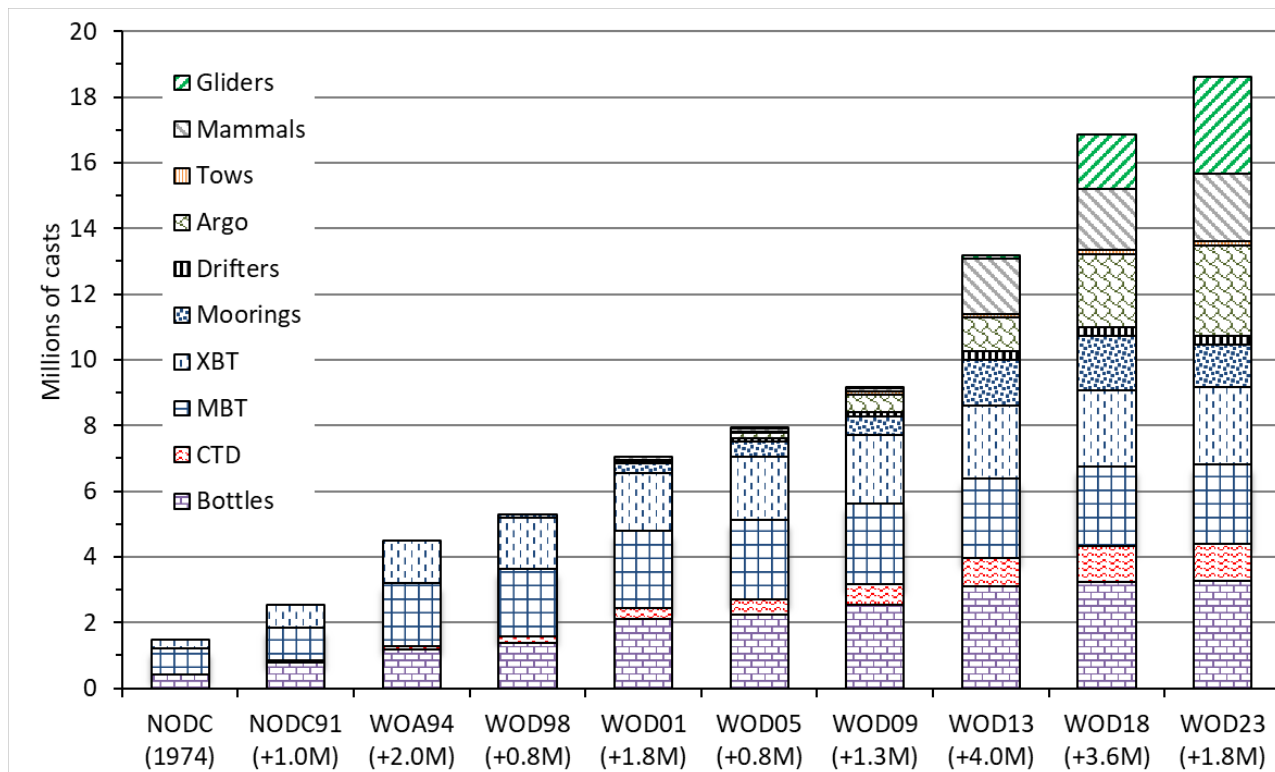


Figure 1.3. Data holding growth over the years of the WOD development.

Figure 1.4. illustrate the percentage of data collected from different instrumental platform in each dataset of WOD23.

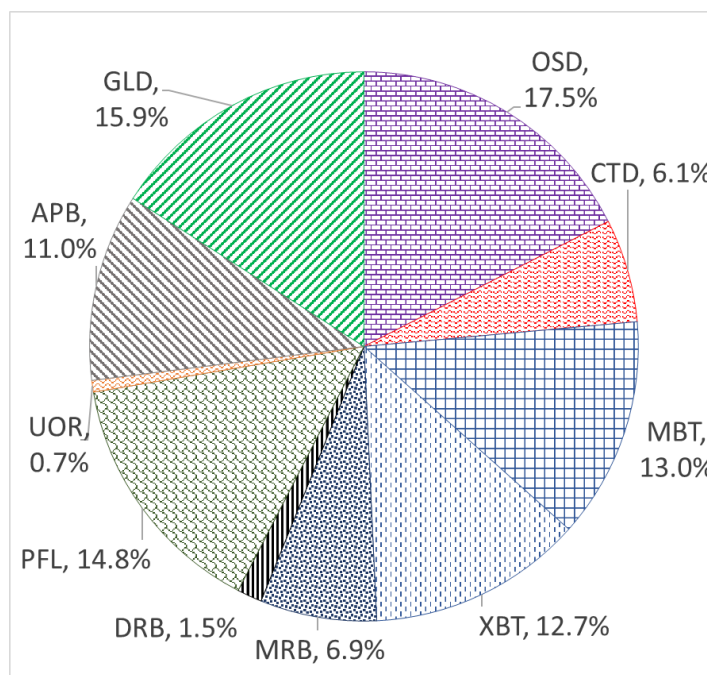


Figure 1.4. Percentage of the data collected by different instruments in WOD23.

### **1.3. DATA SOURCES**

The oceanographic data that comprise the WOD have been acquired through many sources and projects as well as from individual scientists. In addition, many international organizations such as the International Oceanographic Data and Information Exchange and World Data System have facilitated data exchanges, which have provided many data to the WOD.

#### **1.3.1. IODE**

International Oceanographic Data and Information Exchange ([IODE](#)) activities of the IOC have been responsible for the development of a network of National Oceanographic Data Centers in many countries. This network greatly facilitates international ocean data exchange. The IOC was established to support international oceanographic scientific needs including data exchange on an intergovernmental basis (UNESCO, 1979). The WOD became an IODE project in 2001 and has received logistical and planning support since that time. IODE has also been instrumental in facilitating international data flow to the WOD and in promoting and enhancing the use of WOD internationally.

#### **1.3.2. The World Data System**

The World Data System ([WDS](#)) was set up during the International Geophysical Year under the auspices of the International Council of Scientific Unions (ICSU, 1996; Rishbeth, 1991; Ruttenberg and Rishbeth, 1994). Contributions of data from scientists, oceanographic institutions, and countries have also been sent to the World Data Service for Oceanography ([WDS for Oceanography](#)), collocated with NCEI (formerly, World Data Center for Oceanography, Silver Spring), since its inception. There are two other oceanographic centers in the WDS: World Data Center (WDC) for Oceanography, Obninsk, Russia (formerly [WDC-B for Oceanography](#)) and [WDC for Oceanography](#), Tianjin, China. Additional information about the WDS can be found on the [WDS web page](#).

#### **1.3.3. IOC Global Oceanographic Data Archaeology and Rescue Project**

NCEI (then NODC) and several other major oceanographic data centers initiated “data archaeology and rescue” projects around 1991. Based on the success of these projects, the IOC initiated a world-wide project in 1993 known as the Global Oceanographic Data Archaeology and Rescue ([GODAR](#)) project with the goal of locating and rescuing oceanographic data that are stored in manuscript and/or digital form, that are at risk of being lost due to media decay. The international scientific and data management communities have strongly supported this project. Levitus *et al.* (1994) described results from the first phase of this project. With the publication and distribution of the WOD, approximately 3.7 million temperature profiles have been added to the historical archives of oceanographic data since inception of various national data archaeology and rescue projects and the IOC/GODAR project in 1991, and the NCEI/WDS ‘Global Ocean Database Project’ in 1996.

#### **1.3.4. Near-real time data sources**

The Global Temperature and Salinity Profile Program ([GTSPP](#), Searle, 1992; IOC, 1998) is a project sponsored by the IOC/IODE and the Joint Committee for Oceanography and Marine Meteorology ([JCOMM](#)) to develop databases of temperature-salinity profiles reported in “real-

time”. The “real-time” here means that data are transmitted from the measuring platform to data assembling/processing center(s) as soon as they are recorded or with minimal delay, usually within few hours from the moment of data acquisition. WOD incorporates XBT, XCTD, CTD, glider, and pinniped data from GTSP with the expectation that delayed-mode (received some time after 48 hours with full resolution, calibration, and quality assurance) data will be received and incorporated later. Often delayed mode data are never received, leaving the near-real time version the sole record in the WOD for many casts.

Users wanting GTSP real-time data directly can acquire the data via the [NCEI GTSP website](#).

Tropical moored buoy data from the TAO/TRITON array (McPhaden *et al.*, 1998) and the PIRATA and RAMA arrays were obtained from NOAA Pacific Marine Environmental Laboratory (PMEL). Users wanting the complete TAO/TRITON/PIRATA/RAMA buoy database comprised of data that have had the benefit of additional PMEL processing and quality control, can find the data at <https://www.pmel.noaa.gov/tao/drupal/disdel/>. The buoy data stored in WOD23 as a daily averaged profile. Original 15-minutes averages can be accessed at the NCEI archives if desired.

Argo profiling float data are obtained from the Argo U.S. Global Data Assembly Center. Users wanting the most up to date Argo data and quality control should obtain their data via [Coriolis web site](#) (<http://www.coriolis.eu.org/Observing-the-Ocean/ARGO>) or the Global Argo Data Repository site (<https://www.ncei.noaa.gov/products/global-argo-data-repository>).

### **1.3.5. International Research Projects Data**

Data from WOCE, the Climate Variability (CLIVAR) program (WCRP, 1995), and GO-SHIP are maintained at the CLIVAR and Carbon Hydrographic Data Office (CCHDO) and updated to the WOD on a quarterly basis.

### **1.3.6. ICES Contribution**

The International Council for Exploration of the Sea (ICES) has collected data from participating countries since its inception in 1902. ICES has been an important provider of data to the WOD and continues to make available their latest updates, including historical data, on a quarterly basis.

### **1.3.7. Declassified Naval Data Sets**

As a result of the end of the Cold War, the navies of several countries have declassified substantial amounts of oceanographic data that were formerly classified, in some cases at the request of the Intergovernmental Oceanographic Commission (IOC). It should be recognized that some navies have policies of declassifying substantial amounts of data in real-time or with relatively short time delays. For example, the U.S. Navy has contributed approximately 435,000 mechanical bathythermograph (MBT) profiles and the U.S. Coast Guard approximately 217,000 MBT profiles to the NODC(NCEI)/WDC databases. Recent U.S. Navy data have been acquired from the U.S. Navy MOODS database ([NCEI Accession 0001255](#)). In addition, the Australian Navy reports profile data in real-time including data from their Exclusive Economic Zone (EEZ).

**Table 1.3. Comparison of the number of oceanographic casts in WOD products series**

<b>Dataset</b>	<b>NCEI (1974)<sup>1</sup></b>	<b>NCEI (1991)<sup>2</sup></b>	<b>WOA94</b>	<b>WOD98</b>	<b>WOD01</b>	<b>WOD05</b>	<b>WOD09</b>	<b>WOD13</b>	<b>WOD18</b>	<b>WOD23</b>
OSD <sup>3</sup>	425,000	783,912	1,194,407	1,373,440	2,121,042	2,258,437	2,541,298	3,115,552	3,233,155	3,256,037
CTD <sup>4</sup>	n/a	66,450	89,000	189,555	311,943	443,953	641,845	848,911	1,089,421	1,132,680
MBT <sup>5</sup>	775,000	980,377	1,922,170	2,077,200	2,376,206	2,421,940	2,426,749	2,425,607	2,426,301	2,426,245
XBT	290,000	704,424	1,281,942	1,537,203	1,743,590	1,930,413	2,104,490	2,211,689	2,334,267	2,360,444
MRB <sup>6</sup>	n/a	n/a	n/a	107,715	297,936	445,371	566,544	1,411,762	1,656,204	1,277,591
DRB	n/a	n/a	n/a	n/a	50,549	108,564	121,828	251,712	245,592	272,872
PFL	n/a	n/a	n/a	n/a	22,637	168,988	547,985	1,020,216	2,215,341	2,748,011
UOR	n/a	n/a	n/a	n/a	37,645	46,699	88,190	88,190	127,544	127,574
APB	n/a	n/a	n/a	n/a	75,665	75,665	88,583	1,713,132	1,871,303	2,056,367
GLD	n/a	n/a	n/a	n/a	n/a	338	5,857	103,798	1,665,453	2,968,167
<b>Total casts / growth</b>	<b>1,490,000</b>	<b>2,535,163 1,045,163</b>	<b>4,487,519 1,952,356</b>	<b>5,285,113 797,594</b>	<b>7,037,213 1,752,100</b>	<b>7,900,368 863,136</b>	<b>9,155,099 1,254,750</b>	<b>13,190,569 4,035,470</b>	<b>16,864,581 3,674,012</b>	<b>18,625,988 1,761,407</b>
Plankton	n/a	n/a	n/a	83,650	142,900	150,250	218,695	242,727	245,059	245,059
SUR <sup>7</sup>	n/a	n/a	n/a	n/a	4,743	9,178	9,178	9,289	9,289	9,289

<sup>1</sup> Based on statistics from *Climatological Atlas of the World Ocean* (1982).

<sup>2</sup> Based on NODC Standard Product: Global ocean temperature and salinity profiles (2 disc set) (NCEI Accession 0098058).

<sup>3</sup> WOD23 OSD dataset includes data from 179,616 low-resolution CTD and 1,708 low-resolution XCTD casts.

<sup>4</sup> WOD23 CTD dataset includes data from 12,741 high-resolution XCTD casts.

<sup>5</sup> WOD23 MBT dataset includes data from: 2,340,323 MBT, 80,200 DBT and 5,659 Micro-BT casts.

<sup>6</sup> MRB data submitted under Accession 0063240 in WOD18 were recorded at the 15-min interval; in 2021 those data were re-converted as daily averages, the daily average profiles matched with the first unique station number each day, and the extra stations were removed, hence total of MRB record counts in WOD23 decreased.

<sup>7</sup> Surface data are represented differently from cast (profile) data in the database – all observations in a single cruise have been combined into one “cast” with zero depth, value(s) of variable(s) measured, latitude, longitude, and Julian year-day to identify data and position of individual observations.

### 1.3.8. Integrated Global Ocean Service - Volunteer Observing Ship programs

Since the pioneering work of Mathew Maury beginning in 1854, there have been programs in existence to gather meteorological and oceanographic data from merchant ships. These ships are sometimes referred to as Voluntary Observing Ships (VOS) and the programs called Ship-of-Opportunity Programs (SOOP). During the 1970's, the U.S. (Scripps Institute of Oceanography, CA) and France (ORSTOM, New Caledonia) began a SOOP program that focused on the deployment of XBT instruments from VOS platforms in the Pacific Ocean (White, 1995). This program expanded to include the Atlantic and Pacific Oceans and is now supported by [NOAA Ship-of-Opportunity Program](#). Several countries are conducting SOOPs or have conducted them. These programs are coordinated internationally by the World Meteorological Organization ([WMO](#)) and [IOC](#) and are the main provider of XBT data to the WOD through the GTSP.

## 1.4. QUALITY CONTROL FLAGS

Each individual data value and each profile in WOD23 has quality control flags associated with it. A description of these flags and general documentation describing software for reading and using the WOD23 database are inherited from the previous release and can be found in [WOD23 Manual](#) (Garcia *et al.* 2024). WOD also includes Quality Control Flags assigned by data submitters, which could be assigned by either an automatic or manual QC or its combination. It is clear that there are both Type I and Type II statistical errors (for normal distributions) associated with these flags. This means that there is some data that has been flagged as being questionable or unrepresentative when in fact they are not. In the same tile, there are some data that have been flagged as being “acceptable” based on our tests, which in fact may not be the case. In addition, the scarcity of data, non-normal frequency distributions, and presence of different water masses in close proximity to each other results in incorrect assignment of flags. Oguma *et al.* (2003; 2004) discuss skewness of oceanographic data. The WOD flags represent data values used or not used in the calculation of the WOA climatological mean fields, and therefore WOD QC is thus tuned to qualify data representative of the large-scale permanent or semi-permanent features.

The obvious advantage of flagging data is that users can choose to accept or ignore all or part of the flags assigned to data values. The most important flags we set are based on unusual features produced during objective analyses of the data at standard levels. This is because standard statistical tests may be biased for the reasons described above. Data from small-scale ocean features such as eddies and/or lenses are not representative of the large-scale permanent or semi-permanent features we attempt to reproduce with our analyses and will cause unrealistic features such as bullseyes to appear. Hence, we flag these data, and other data that cause such features, as being unrealistic or as questionable data values. It is important to note that an investigator studying the distribution of mesoscale features in the ocean will find data from such features to be the signal they are looking for. As noted by Levitus (1982), it is not possible to produce one set of data analyses to serve the requirements of all possible users. A corollary is that it is also impossible to produce one set of quality control flags for a database that serve the exact requirements of all investigators. As data are added to a database, investigators must realize that flags set for certain criteria being violated in an earlier version of the database may be reset solely due to the addition of new data which may change the statistics of the region being considered. Even data that have



produced unrealistic features may turn out to be realistic when additional data are added to a region of sparse data.

### 1.4.1. Levels of Quality Control

Different oceanographic variables in the WOD datasets have various levels of quality control performed on them. Those oceanographic variables in datasets used for calculating climatological means have the highest level of quality control. This includes all preliminary and automatic quality control checks and subjective checks performed in evaluating the quality of the resultant climatological fields. The automatic checks include minimum/maximum range assessment for 31 ocean areas at 102 standard levels.

Values of temperature in all datasets except APB received the highest level of quality control. Likewise, values of salinity received the highest level of quality control for all datasets except APB. That is due to some peculiarities of the sensors used for collection this data. More details can be found in (Siegelman *et. al*).

Measurements of oxygen, phosphate, silicate, and nitrate content in the OSD dataset received the highest quality control. Oxygen data in the SUR, DRB, UOR, and GLD datasets received a lower level of quality control. Since these data were not used to calculate climatologies, subjective checks were not performed on them. After calculation of climatologies using oxygen data from the OSD, CTD, and PFL datasets, the newly calculated five-degree statistics (mean and standard deviation) were used iteratively to perform additional standard deviation quality control check on oxygen in the OSD, CTD, and PFL datasets. The reason for not using the oxygen data from the DRB, UOR, and GLD datasets is that many of these oxygen data are not calibrated.

Chlorophyll, pH, and alkalinity values in the OSD dataset received a lower level of quality control than oxygen and nutrients. There are no chlorophyll, pH, or alkalinity climatologies calculated for WOA23, so no standard deviation checks were performed. All other checks were done for data in the CTD, PFL, DRB, UOR and GLD datasets.

A lower level of quality control was done on pCO<sub>2</sub>, Dissolved Inorganic Carbon (DIC), Tritium, Helium,  $\Delta^3$ Helium,  $\Delta^{14}$ Carbon,  $\Delta^{13}$ Carbon, Argon, Neon, CFC-11, CFC-12, CFC-113, and  $\Delta^{18}$ Oxygen concentrations in the OSD dataset. Only initial range checks were applied to these variables. These ranges, a single minimum and maximum for all oceans were taken from the WOCE Data Reporting Requirements (WOCE Publication 90-1 *Revision 2*).

BAC data in the CTD and PFL datasets were subject to this lowest level of quality control as well. A. Mishonov set the minimum and maximum values based on the world-wide beam attenuation records reported in Global Transmissometer Database (Gardner *et al.*, 2020).

For more information about the quality control procedures, see [WOD23 Manual](#) (Garcia *et al.*, 2024).

Plankton data have a different set of quality control detailed in Chapter 16 of this document as well as in [WOD23 Manual](#) (Garcia *et al.*, 2024).

## **1.5. OUTLOOK FOR FUTURE ACQUISITIONS OF HISTORICAL OCEAN PROFILE AND PLANKTON DATA AND INTERNATIONAL COOPERATION IN THE “WORLD OCEAN DATABASE PROJECT”**

Substantial amounts of historical ocean data continue to be transferred to NCEI/WDS for archiving and inclusion into databases. The outlook for our ability to continue increasing the amount of such data available to the scientific community is excellent. Based on the positive results of the IOC/GODAR project and the World Ocean Database Project, we have requested the continued cooperation of the international scientific and data management communities in building the historical ocean data archives. There is a particular need for high-resolution CTD data to resolve smaller scale features in the vertical and thus provide objective analyses of variables at greater vertical resolution than present (e.g. Helber *et al.*, 2012 documents the lack of such data for global scale analyses). There is a need for additional historical chlorophyll, nutrient, oxygen, and plankton data so we can improve understanding of ocean biogeochemical cycles.

Improving the quality of historical data and their associated metadata is an important task. Corrections to possible errors in data and metadata is best done with the expertise of the principal investigators who made the original observations, the data center or group that prepared the data, or be based on historical documents such as cruise and data reports (however, one has to also consider that these documents may contain errors). The continuing response of the international oceanographic community to the IODE GODAR and World Ocean Database projects has been excellent. This response has resulted in global ocean databases that can be used internationally without any restriction for studying a wide variety of environmental problems i.e. efforts to develop the International Quality-controlled Ocean Database ([IQuoD](#)), *etc.*

As the amount of historical oceanographic data continues to increase because of international cooperation, the scientific community will be able to make more and more realistic estimates of variability and to place confidence intervals of the more frequently sampled variables such as temperature. It should also be noted that from 2021-2030, the [UN Decade of Ocean Science](#) for Sustainable Development will focus scientific initiatives across the globe on the science needed to ensure sustainable use of ocean resources and long-term ocean health. The U.S. National Committee for the Ocean Decade will encourage participation and serve as a communication channel for the U.S. ocean science community throughout this international effort.

However, it should be noted that the global COVID-19 pandemic, which took hold in 2020, added strain to the maintenance of the observing system and incoming data flow. A survey of the contributing components of the observing system illustrates the impacts of the pandemic from January 2020 through December 2021. The pandemic did not reduce the short-term geographic coverage (days to months) capabilities mainly due to the continuation of autonomous platform observations. In contrast, the pandemic caused critical loss to longer-term (years to decades) observations, greatly impairing the monitoring of such crucial variables as ocean carbon and the state of the deep ocean. So, while the observing system has held under the stress of the pandemic, work must be done to restore the interrupted replenishment of the autonomous components and plan for more resilient methods to support components of the system that rely on cruise-based measurements (Boyer *et al.*, 2023).



## 1.6. LAYOUT OF THE REST OF THIS DOCUMENT

The rest of this document, Chapters 2-16 describe in more detail the oceanographic instrumentation used to collect the data, which are contained in WOD23 and the nature of the measurements themselves, as well as data accuracy, spatial and temporal data distribution and other details.

Chapter 2 describes the OSD dataset, with an emphasis on Ocean Station Data. However, not all chapters neatly fit into one dataset. For instance, Chapter 5 is about the XCTD data, which are spread over the OSD and CTD datasets. Chapters 7, 8, and 13 all detail the data, which are collected by different generations of the bathythermographs instruments, and are stored in the MBT dataset.

## 1.7. REFERENCES AND BIBLIOGRAPHY

- Boyer, T., H.-M. Zhang, K. O'Brien, J. Reagan, S. Diggs, E. Freeman, *et. al.* (2023). Effects of the Pandemic on Observing the Global Ocean. *Bulletin of the American Meteorological Society*, (104)-2, pp. E389-E410. DOI: <https://doi.org/10.1175/BAMS-D-21-0210.1>
- Boyer, T.P., O.K. Baranova, C. Coleman, H.E. Garcia, A. Grodsky, R.A. Locarnini, A.V. Mishonov, C.R. Paver, J.R. Reagan, D. Seidov, I.V. Smolyar, K. Weathers, M.M. Zweng, (2018). World Ocean Database 2018. A.V. Mishonov, Technical Ed., *NOAA Atlas NESDIS 87*.
- Forster, G., A Voyage Round the World in His Britannic Majesty's Sloop Resolution, commanded by Capt. James Cook, during the Years 1772, 3, 4, and 5. (1777), Vol. 1. London, Great Britain.
- Freeman, E., S.D. Woodruff, S.J. Worley, S.J. Lubker, E.C. Kent, W.E. Angel, D.I. *et. al.*, (2017). ICOADS Release 3.0: a major update to the historical marine climate record. *Int. Journal of Climatology*, 37: 2211-2232. <https://doi.org/10.1002/joc.4775>
- Garcia, H. E., T. P. Boyer, R. A. Locarnini, J. R. Reagan, A.V. Mishonov, O. K. Baranova, C. R. Paver, Z. Wang, C. Bouchard, S. Scott, D. Seidov, and D. Dukhovskoy (2024). World Ocean Database 2023: User's Manual. A.V. Mishonov, Tech. Ed., *NOAA Atlas NESDIS 98*, pp 131. <https://doi.org/10.25923/j8gq-ee82>
- Gardner, W.D., A.V. Mishonov, M.J. Richardson, 2020. Global Transmissometer Database V3. <https://odv.awi.de/data/ocean/global-transmissometer-database/>
- Gouretski, V., and K.P. Koltermann (2007). How much is the ocean really warming? *Geophysical Research Letters*, vol. 34, L01610, doi: 10.1029/2006GL027834
- Hanawa, K., P. Rual, R. Bailey, A. Sy, and M. Szabados (1995). A new depth-time equation for Sippican or TSK T-7, T-6 and T-4 expendable bathythermographs (XBT). *Deep Sea Research I*, 42, 1423-1452.
- Heinmiller, R.H., C.C. Ebbesmeyer, B.A. Taft, D.B. Olson, O.P. Nikitin (1983). Systematic errors in expendable bathythermograph (XBT) profiles. *Deep Sea Research I*, 30, 1185-1196. doi:10.1016/0198-0149(83)90096-1
- Helber, R.W.A., Birol Kara, J.G. Richman, M.R. Carnes, C.N. Barron, H.E. Hurlburt, and T. Boyer (2012). Temperature versus salinity gradients below the ocean mixed layer *Journal of Geophysical Research*, 117, C05006, doi:10.1029/2011JC007382
- ICSU (1996). Guide to the World Data Center System, produced by World Data Center-A, *NOAA NGDC*, Boulder, CO, 109 pp.

- IOC (1989). Integrated Global Ocean Services System (IGOSS) – Summary of Ship-of-Opportunity programmes and technical reports. *IOC/INF-804*, 192 pages.
- IOC (1998). Global Temperature-Salinity Profile Programme (GTSP) - Overview and Future. Intergovernmental Oceanographic Commission, Paris, *IOC Technical Series 49*, 12 pp.
- Johnson, G.C. (1995). Revised XCTD Fall-Rate Equation Coefficients from CTD Data. *Journal of Atmospheric & Oceanic Technol.*, **12**, 1367–1373, [doi:10.1175/1520-0426\(1995\)012<1367:RXFREC>2.0.CO;2](https://doi.org/10.1175/1520-0426(1995)012<1367:RXFREC>2.0.CO;2)
- Kizu, S. and K. Hanawa (2002). Start-up transients of XBT measurement. *Deep-Sea Research I*, **49**, 935-940.
- Kizu, S., H. Onoshi, T. Suga, K. Hanawa, T. Watanabe, and H. Iwamiya (2008), Evaluation of the fall rates of the present and developmental XCTDs. *Deep-Sea Research I*, **55**, 571-586
- Kizu, S., H. Yoritaka, and K. Hanawa (2005). A new fall-rate equation for T-5 Expendable Bathythermograph (XBT) by TSK. *Journal of Oceanography*, **61**, 115-121.
- Levitus, S. (1982). Climatological Atlas of the World Ocean, U.S. Gov. Printing Office, Wash., D.C., 173 pp.
- Levitus, S., R. Gelfeld, T. Boyer, and D. Johnson (1994). Results of the NODC and IOC Data Archaeology and Rescue projects. *Key to Oceanographic Records Documentation No. 19*, National Oceanographic Data Center, Wash., D.C., 67 pp.
- Levitus, S., S. Sato, C. Maillard, N. Mikhailov, P. Caldwell, and H. Dooley (2005). Building Ocean Profile-Plankton Databases for Climate and Ecosystem Research, *NOAA Technical Report NESDIS 117*, U.S. Gov. Printing Office, Wash., D.C., 29 pp.
- McConnell, A. (1982). No Sea Too Deep: The History of Oceanographic Instruments. Bristol, *Adam Hilger*, 162 pp.
- McPhaden, M.J., A.J. Busalacchi, R. Cheney, J.-R. Donguy, K.S. Gage, D. Halpern, M. *et. al.* (1998). The Tropical Ocean-Global Atmosphere observing system: A decade of progress. *Journal of Geophysical Research*, **103 (C7)**, 14,169-14,240.
- Mizuno, K. and T.J. Watanabe (1998). Preliminary results of in-situ XCTD/CTD comparison test. *Journal of Oceanography*, **54**, 373, doi:10.1007/BF02742621
- Oguma, S. and Y. Nagata (2002). Skewed water temperature occurrence frequency in the seas off Sanriku, Japan, and intrusions of the pure Kuroshio Water. *Journal of Oceanography*, **58**, 789-796.
- Oguma, S., T. Suzuki, S. Levitus, and Y. Nagata (2003). Skewed occurrence frequency of water temperature and salinity in the subarctic regions. *Journal of Oceanography*, **59**, 921-929.
- Prestwich, J., On Submarine Temperatures (1875). In: *Philosophical Transaction of the Royal Society of London*, Vol. 165, part II, pp. 587-674 Taylor and Francis, London, Great Britain.
- Rishbeth, H. (1991). History and evolution of the World Data Center System. *Journal of Geomagnetism and Geoelectricity*, **43 (Supplement)**, 921-929.
- Ruttenberg, S. and H. Rishbeth (1994). World Data Centers – Past Present and Future. *Journal of Atmospheric and Terrestrial Physics*, **56**, 865-870.
- Searle, B. (1992). Global Ocean Temperature-Salinity Pilot Project. In "*Proceedings of the Ocean Climate Data Workshop*" sponsored by NOAA and NASA, Available from NODC, Silver Spring, MD, pp. 97-108.
- Siegelman, L., F. Roquet, V. Mensah, P. Rivière, E. Pauthenet, B. Picard, and C. Guinet, 2019:

- Correction and Accuracy of High- and Low-Resolution CTD Data from Animal-Borne Instruments. *J. Atmos. Oceanic Technol.*, 36, 745–760, <https://doi.org/10.1175/JTECH-D-18-0170.1>.
- Spilhaus, A.F. (1938). A bathythermograph. *Journal of Marine Research*, 1, 95-100.
- Toole, J.M., R.A. Krishfield, M.-L. Timmermans, and A. Proshutinsky. 2011. The Ice-Tethered Profiler: Argo of the Arctic. *Oceanography*, 24(3):126–135, <https://doi.org/10.5670/oceanog.2011.64>
- UNESCO (1979). A focus for ocean research-Intergovernmental Oceanographic Commission, History, Functions, Achievements. *IOC Technical Series No. 20*, Paris, 64 pp.
- UNESCO (Hanawa, K., P. Rual, R. Bailey, A. Sy, and M. Szabados) (1994). Calculation of New Depth Equations for Expendable Bathythermographs Using a Temperature-Error-Free Methods (Application to Sippican/TSK T-7, T-6 and T-4 XBTs), *IOC Technical Series No. 42*, 46 pp.
- White, W. (1995), Design of a global observing system for gyre-scale upper ocean temperature variability. *Progress in Oceanography*, 36, 169-217.
- WOCE Publication 90-1 Revision 2: *Requirements for WOCE Hydrographic Programme Data Reporting*, T. Joyce and C. Corry editors, unpublished manuscript.
- World Climate Research Program (WCRP) (1995). CLIVAR: A study of climate variability and predictability- Science Plan. *WCRP-89*, Geneva, 157 pp.
- Wyrtki, K. (1971). Oceanographic Atlas of the International Indian Ocean Expedition. *National Science Foundation*, Wash., D.C., 531 pp.
- Wilkinson, M., M. Dumontier, I. Aalbersberg, I. *et al.* (2016). The FAIR Guiding Principles for scientific data management and stewardship. *Sci Data* 3, 160018. <https://doi.org/10.1038/sdata.2016.18>

# CHAPTER 2: OCEAN STATION DATA (OSD), LOW-RESOLUTION CTD, LOW-RESOLUTION EXPENDABLE XCTD, AND PLANKTON

*Hernan E. Garcia, James R. Reagan, Olga K. Baranova, Tim P. Boyer, Ricardo A. Locarnini,  
Alexey V. Mishonov, Christopher R. Paver, Zhankun Wang, Courtney Bouchard, Scott Cross,  
Dan Seidov, and Dmitry Dukhovskoy*

*Ocean Climate Laboratory  
NOAA National Centers for Environmental Information  
Silver Spring, Maryland, USA*

## 2.1. INTRODUCTION

Oceanographic physical, chemical, and biological measurements at geographic locations and depths of interest in the water column (*i.e.*, profiles) collected from sea-going research ships have historically been referred to as Oceanographic Station Data (OSD). Such discrete measurements are frequently referred to as “bottle data” and the entire OSD collection may be alternatively referred to as the “Bottle Dataset”. Here we adopt the general term OSD to refer collectively to low vertical resolution spacing between profile seawater samples, serial (discrete) water column measurements (bottles, buckets), plankton (bottles, net-tows), relatively low vertical (depth or pressure) resolution Expendable Conductivity-Temperature-Depth (XCTD), and relatively low vertical resolution Conductivity-Temperature-Depth (CTD) data in the *World Ocean Database 2023* (WOD23). [WOD23 Manual](#) (Garcia *et al.* 2024) provides a description of the data available in the WOD23.

Historical OSD observational data archived at the National Centers for Environmental Information ([NCEI](#)) have been received as a result of several efforts including the [World Data Service for Oceanography](#) of the World Data System (WDS), international open data sharing, and projects such as the Intergovernmental Oceanographic Commission ([IOC](#)) International Oceanographic Data and Information Exchange ([IODE](#)) Global Oceanographic Data Archaeology and Rescue project ([GODAR](#)) (Levitus *et al.*, 1994, Levitus *et al.*, 2012), IOC IODE World Ocean Database project (Levitus, 2012) and Global Temperature-Salinity Profile Program ([GTSP](#)). GODAR consolidates and digitizes oceanographic data in paper form against degradation or loss of open accessibility. When the project was initially proposed in 1992, most historical oceanographic data was dispersed globally in a variety of repositories in different, incompatible non-digital formats.

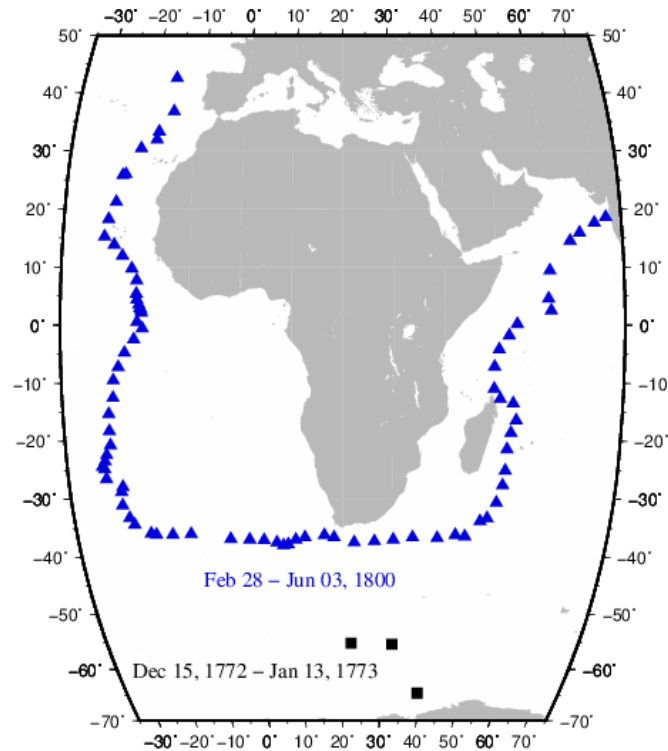
The OSD dataset includes the most frequently measured *in situ* (*i.e.*, collected in the original depth and geographic location) physical, chemical, and biological oceanographic variables as a function of depth or pressure. These variables are also Essential Ocean Variables ([EOV](#)) for ocean physics, biogeochemistry, and biology and ecosystems. The OSD dataset provides the most

comprehensive and representative collection of historical and recent discrete oceanographic observations available without restriction of use to date totaling 3,256,037 casts covering the time period between 1772 and 2022. This is an increase of 22,882 casts (0.7%) when compared to WOD18 (Boyer *et al.*, 2018, Garcia *et al.* 2019). The number of casts in the OSD dataset represent about 18% of the total number of casts available in the WOD23 (18,629,354, see Fig. 1.3 and Tab. 1.3). Not counted in this total of OSD casts is the 245,059 Plankton tows. The description that follows is an overview of the data variables in the OSD dataset.

## **2.2. COMMONLY USED LOW AND LARGE VOLUME WATER COLUMN SAMPLERS**

Most of the historical seawater samples of the ocean's water column in the OSD dataset were obtained from oceanographic research cruises occupying a number of selected oceanographic stations (sometimes-called hydrographic stations) at geographic locations along generally pre-selected cruise tracks. For each station, discrete water samples from the ocean surface to some selected depth of the water column were obtained by means of a variety of specially designed sampling water bottles of different volumes depending on the target measurements. Some of the very early historical oceanographic measurements of the water column were collected by means of wood or metal buckets.

The WOD Team aims at providing access to all oceanographic profile data in the instrumental record irrespective of year of collection. For example, based on the GODAR effort, the earliest *in situ* measurements in the OSD dataset are of sea surface temperature collected on December 15, 23 1772 and January 13, 1773 onboard the United Kingdom (UK) "Britannic Majesty's Sloop" HMS Resolution, commanded by Captain James Cook (Figure 2.1). In an account of the voyage, the entry for December 23, 1772 indicates "We seized this opportunity to hoist out a boat, and continue the experiments on the current, and on the temperature of the sea" (Page 102, Forster, 1777). The data are included in the WOD as cruise GB11003 and the digitized data available in the NCEI archive under [Accession 0000571](#). The OSD dataset includes 80 stations with sea surface temperature and meteorological measurements (dry bulb temperature) collected between February 28 and June 3, 1800 on board the UK sailing ship *Skelton Castle* on a voyage from England to India while in route from England to Bombay as part of the East India Company (Figure 2.1). The data are included in the WOD as cruise GB12994 and the digitized data available in the NCEI archive under [Accession 0095925](#). Figure 2.2 shows a picture of the original data collected in paper form and manually digitized. Anecdotally, we note the entry on April 12 in Figure 2.2 which denotes "shark caught". The [GODAR](#) Project continues to locate and rescue historical oceanographic profile and plankton data that are at risk of being lost due to media decay and/or neglect. Coincidentally, the most recent cast in the WOD23 OSD dataset was collected July 27, 2022; on board of the *RSS James Cook*. Figure 2.3 shows that most of the casts in the OSD were collected on or after the International Geophysical Year (1958-1959).



**Figure 2.1 Geographic positions of *in situ* sea surface temperature measurements collected in the 1772-1800 time period available in the WOD (black squares – locations of the seawater temperature measurements performed during Cpt. Cook’s 1772-1775 voyage on *HMS Resolution*).**

Historically oceanographic water sampling collection and analysis is a labor-intensive process and many types of water sampling devices have been invented since the early days of oceanographic research. The Nansen and Niskin bottles are probably the most commonly used water samplers to date for the serial collection of relatively small volumes of seawater. Nansen bottles, commonly used prior to the late 1960s, were invented by Fridtjof Nansen in 1910. These are cylindrical pressure-resistant metal containers (usually made of brass) with plug valves at each end that allow the collection of small volumes of seawater (about < 1.5 liters) at selected depths in the water column (Sverdrup *et al.*, 1942). The Nansen bottles often included two or more specially designed protected and unprotected mercury-filled glass reversing thermometers inside a small metal case exposed to the water column attached to the outside of each bottle. These thermometers allowed the estimation of the *in-situ* temperature at which each bottle closed in the water column. The Nansen bottles were generally replaced by the Niskin bottles in the late 1960s. Niskin bottles helped minimize some of the problems associated with the collection of Nansen bottle samples (Worthington, 1982). Niskin bottles are cylindrical pressure-resistant plastic containers (to minimize contamination between the bottle and the water sample) with rubber spring-loaded end-caps that allow the collection of a variety of volumes of seawater (about 1.2 to 10 liters). Niskin bottles are frequently mounted around a circular rosette sampler metal frame with the capacity to hold as many as 36 bottles. The bottles can then be closed at any depth or pressure by an electrical command from deck or from preset depth (pressure) values. When the closed Niskin bottles are brought back on deck, water samples can be collected from each bottle and then analyzed for different seawater constituents. The rosette frame may include a CTD and other automated sampling sensor instruments (*e.g.*, fluorometers, transmissometers, pH, O<sub>2</sub>).



1800.	Atmo- sphere.	Sea-Water.	Lat. N.	Long. W.
Feb. 28	•	52°	42° 34'	15° 25'
March 2	•			
4	•	57	36 47	15
7	•	62	33 19	17 34
9	•	64	31 38	50
11	•	66	30 26	20 49
13	•	67	No observ.	59
15	71°	+68	26	23 24
16	72	69½	25 49	24
17	72	70	21 14	55
18	74	72	18 16	26 12
19	74	72	15 17	47
20	75	72	13 52	24 49
21	76	74	-11 58	23 6
22	78	+76½	9 46	21 12
23	80	78	7 43	20 11
24	82	80	5 23	20 13
25	82	80	4 24	20 6
26	84	82	3 33	19 41
27	84	82	2 58	19 27
28	85	82	2 33	19 17
29	86	83	2 7	19 10
30	86	83	35	20 1
31	85	82	33 S.	19
April 1	84	81	2 27	20 41
2	83	80	4 44	22 31
3	82	80	7 12	23 47
4	83	80	9 30	24 50
5	82	79	12 30	25 59
6	80	76	15 17	23 59
7	78	76	18 13	26 5
8	78	76	20 41	26 3
9	80	78	22 22	56
10	80	78	23 28	27 20
11	80	78	24 23	37
12	78	76 Shark 88½	24 48	32
13	76	74	26 29	22
14	74	72	27 30	24 14
15	72	70	28 42	35
16	72	70	31	45

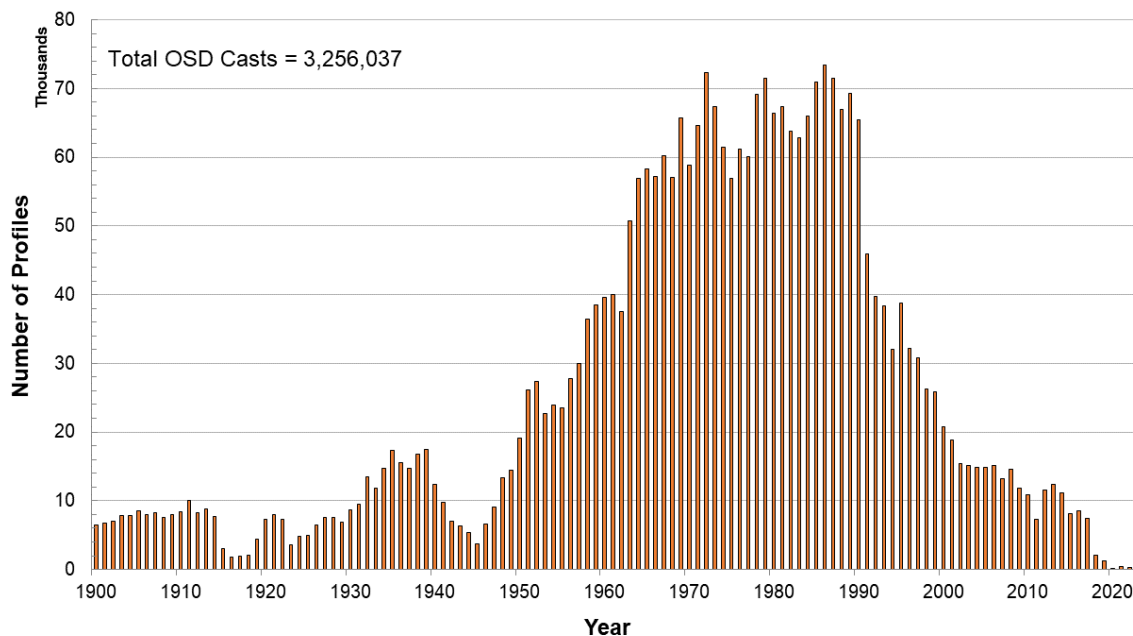
  

1800.	Atmo- sphere.	Sea-Water.	Lat. S.	Long. W.
April 17	70°	68°	33° 10'	23° 37'
18	70	66	34 20	22 46
19	70	66	35 51	19 38
20	68	66	36 2	18 33
21	68	66	36 3	15 11
22	68	64	35 56	11 22
23	68	64		6 47
24	66	64	36 47	3 13
25	64	62	36 34	41 Ex
26	64	62		3 40
27	62	58	37 23	6 20
28	61	58	53	7 40
29	60	58	43	8 40
30	60	58	56 50	12 11
May 1	60	58	7	16 13
2	60	58	30	18 13
3	60	60	34	10 16
4	60	60	37 22	22 23
5	60	62	10	26 38
6	58	64	36 54	50 51
7	62	66	30	34 51
8	64	66	37	39 39
9	62	64	5	43 21
10	62	64	20	45 21
11	62	64	53 43	48 14
12	64	66	33 13	49 38
13	66	66	30 55	51 18
14	68	68	27 33	52 12
15	72	70	23 2	31
16	76	74	21 21	39
17	78	76	18 52	33 16
18	78	76	16 23	34 12
19	80	78	13 27	33 29
20	78	78	12 41	30 55
21	80	78	10 56	49 41
22	80	78	7 8	49 41
23	80	80	4 9	50 32
24	81½	80	1 46	52 27
25	84	81	13 N.	54 2
26	86	82	2 32	50 39
27	84	82	4 55	56 10
28	85	82	9 26	60 31
29	85	82	9 26	60 31
30	85	82	14 31	64 41
June 1	84	82	15 55	66 39
2	86	82	17 39	69 40
3	85½	84	18 37	72
4	Arrived in Bombay Harbour.			

Figure 2.2. Picture of original *in situ* sea surface temperature and meteorological measurements collected between February 28 and June 3, 1800 digitized as part of the GODAR project.

The majority of the analyzed chemical constituents dissolved in seawater in the OSD dataset were obtained from the analysis of a relatively small sample volume of seawater with a few exceptions. The most commonly analyzed chemical constituents in seawater have historically been salinity, dissolved oxygen, and the major dissolved inorganic nutrients: nitrate, silicate, and phosphate. Additional chemical constituents (*e.g.*, trace metals, isotopes, dissolved inorganic and organic carbon, and transient tracers) have been measured with the emergence of more precise and clean chemical measurements and sampling techniques. On the other hand, relatively large volumes of seawater are needed for the analysis of chemical dissolved constituents in trace concentration (*e.g.*, moles per liter of solution) or content (*e.g.*, moles per kilogram) present in seawater such as isotopes (*e.g.*, Argon-39, Krypton-85, Carbon-14, delta N-15). Because of their small content, measuring some isotopes requires a large sampling volume of seawater. For

example, Gerard-Ewing samplers were used during the Geochemical Ocean Sections Study, GEOSECS in the early 1970s and World Ocean Circulation Experiment (WOCE) in the 1990s. In time additional analytical methods have emerged and been used (*e.g.*, Global Ocean Ship-based Hydrographic Investigations Program, [GO-SHIP](#) and [GEOTRACES](#), an International Study of the Marine Biogeochemical Cycles of Trace Elements and Isotopes).



**Figure 2.3. Time series of the number of available OSD casts in WOD23 (1900-2022). There are 3,256,037 casts (1900-2022) and 105,973 casts (1772 to 1900, not shown).**

Salinity-temperature-depth (STDs) and CTDs were introduced around the mid-1960s. Many oceanographic data from the mid-1960s and even from later years were archived at relatively low vertical resolution. These low depth resolution data are stored in the OSD dataset as opposed to the high-resolution CTD dataset. Low-resolution here refers to a limited number or a subset of measurements as a function of depth or pressure. At a minimum, low-resolution CTD and STD measurements are recorded at the depths at which water samples have been collected and usually some data at additional depths.

### **2.3. VARIABLES AND METADATA INCLUDED IN THE OSD DATASET**

The OSD dataset includes *in situ* temperature, salinity (conductivity), dissolved oxygen, dissolved inorganic macro-nutrients (phosphate, nitrate, nitrate + nitrite, nitrite, silicate), chlorophyll, alkalinity (often referred as total alkalinity), pH, partial pressure of carbon dioxide (pCO<sub>2</sub>), dissolved inorganic carbon, Tritium,  $\Delta^{13}\text{Carbon}$  (delta carbon-13),  $\Delta^{14}\text{Carbon}$  (delta carbon-14),  $\Delta^{18}\text{Oxygen}$  (delta oxygen-18), chlorofluorocarbons (CFC-11, CFC-12, and CFC-113), Helium,  $\Delta^3\text{Helium}$  (delta Helium-3), Neon, as well as plankton parameter data (Table 2.1). Each cast may contain simultaneous profiles of one or more variables as a function of depth or pressure. The [GO-SHIP Repeat Hydrography Manual](#) (version 2019) provides examples of sampling and measuring protocols for many of these variables.



The observed level measurements are the same as the data originally submitted by the data originator that have been converted to the WOD data format and standardized units (Table 2.1). The WOD maintains metadata regarding the originator's units. The profiles at standard depth levels are the measurements submitted by the data originator vertically interpolated (following Reiniger and Ross, 1968) to 102 depth levels (0-5500 m depth) if the observations did not occur at the desired standard depths. The profiles include quality flags for observed and standard depth level data. Garcia *et al.* (2024) provides a [WOD23 Manual](#) with more in-depth information about QC and quality flags. The most frequently sampled variable as a function of depth is temperature (Figure 2.4). The number of global temperature observations declines from > 2.5 Million near the surface to < 0.5 Million below about 500 m depth.

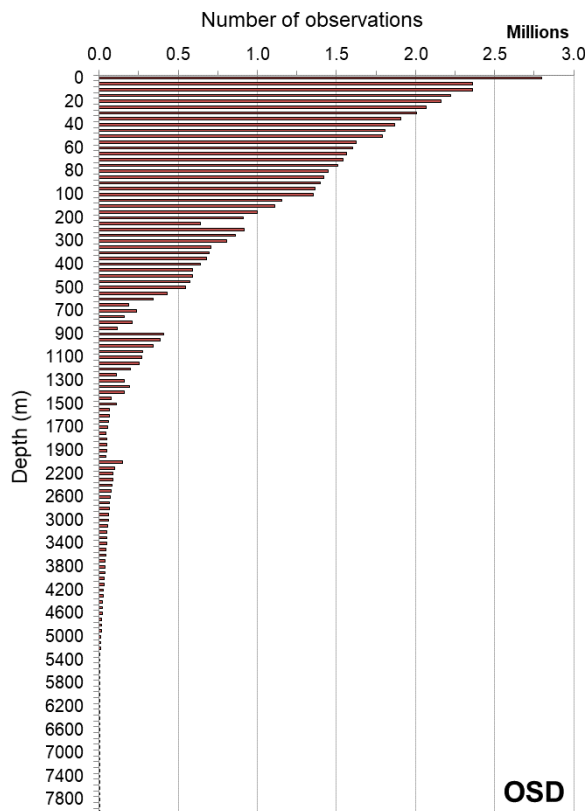


Figure 2.4. Number of temperature observations in WOD23 at standard depth levels.

Physical variables such as temperature, salinity, and hydrostatic pressure are conservative parameters which define the equation of state of seawater (e.g., Millero and Poisson, 1981; Thermodynamic Equation of Sea Water 2010, TEOS-10). Millero (2010) provides a historical summary of the equation of state of seawater. By conservative variables we mean measurements which are not affected directly by biochemical processes.

**Temperature** measurements have been obtained by means of manual (i.e., visual readings of temperature from reversing thermometers in the earliest data) and automated sensors (i.e., digital recordings of temperature from STDs and CTDs) following several International Temperature Scales (ITS) definitions dating back to the early 1900s (e.g., ITS 1927; ITS 1948, ITS 1968, ITS 1990, TEOS-10). Temperature data in the WOD23 include the scale of the measurements if reported by the originator of the data.

**Table 2.1 Depth dependent *in situ* measured variables**

Code	Variable	Standard unit (abbreviation)	Dataset(s) where variable(s) is/are stored
1	Temperature	Degrees Celsius (°C)	OSD, CTD, MBT, XBT, SUR, APB, MRB, PFL, UOR, DRB, GLD
2	Salinity	Dimensionless	OSD, CTD, SUR, MRB, PFL, UOR, DRB, GLD
3	Oxygen	Micro-mole kilogram <sup>-1</sup> (μmol·kg <sup>-1</sup> )	OSD, CTD, PFL, UOR, DRB
4	Phosphate	Micro-mole kilogram <sup>-1</sup> (μmol·kg <sup>-1</sup> )	OSD
6	Silicate	Micro-mole kilogram <sup>-1</sup> (μmol·kg <sup>-1</sup> )	OSD
8	Nitrate and Nitrate + Nitrite	Micro-mole kilogram <sup>-1</sup> (μmol·kg <sup>-1</sup> )	OSD, PFL
9	pH	Dimensionless	OSD, SUR
11	Chlorophyll	Micro-gram liter <sup>-1</sup> (μg·l <sup>-1</sup> )	OSD, CTD, SUR, UOR, DRB
17	Total Alkalinity	Milli-mole liter <sup>-1</sup> (mmol l <sup>-1</sup> )	OSD, SUR
20	Partial pressure of carbon dioxide	Micro-atmosphere (μatm)	OSD, SUR
21	Dissolved inorganic carbon	Milli-mole liter <sup>-1</sup> (mmol·l <sup>-1</sup> )	OSD
24	Transmissivity (Beam Attenuation Coefficient)	Per meter (m <sup>-1</sup> )	CTD
25	Pressure	Decibar	OSD, CTD, UOR, GLD, PFL, DRB
26	Air temperature	Degree Celsius (°C)	SUR
29	Air pressure	Millibar (mbar)	SUR
30	Latitude	Degrees	SUR, APB, UOR
31	Longitude	Degrees	SUR, APB, UOR
32	Julian year-day-1 <sup>1</sup>	Day	SUR, APB, UOR
33	Tritium	Tritium Unit	OSD
34	Helium	Nano-mole kilogram <sup>-1</sup> (nmol·kg <sup>-1</sup> )	OSD
35	Delta Helium-3	Percent (%)	OSD
36	Delta Carbon-14	Per mille (‰); parts per thousand	OSD
37	Delta Carbon-13	Per mille (‰); parts per thousand	OSD
38	Argon	Nano-mole kilogram <sup>-1</sup> (nmol·kg <sup>-1</sup> )	OSD
39	Neon	Nano-mole kilogram <sup>-1</sup> (nmol·kg <sup>-1</sup> )	OSD
40	Chlorofluorocarbon 11	Pico-mole kilogram <sup>-1</sup> (pmol·kg <sup>-1</sup> )	OSD
41	Chlorofluorocarbon 12	Pico-mole kilogram <sup>-1</sup> (pmol·kg <sup>-1</sup> )	OSD
42	Chlorofluorocarbon 113	Pico-mole kilogram <sup>-1</sup> (pmol·kg <sup>-1</sup> )	OSD
43	Delta Oxygen-18	Per mille (‰); parts per thousand	OSD

**Salinity** measurements have been obtained by manual and automated methods (*e.g.*, chemical titrations, chlorinity to salinity formulae, refractometers, salinographs, inductive salinometers, sensors). For the past few decades, bottle salinity sampling and analyses have normally conducted to calibrate the conductivity to salinity measurements of CTDs. Salinity measurements have been obtained using reference standard seawater samples of known salinity (within uncertainty). In 1978 the practical salinity scale (PSS-1978) was adopted defining salinity in terms of an electrical conductivity ratio (UNESCO, 1981; Lewis and Perkins, 1981; Culkin and Ridout, 1998). Under the PSS-1978 definition, salinity values are unitless (Millero, 1993). Seawater standards provide a means to facilitate the inter-comparison of ocean salinity measurements against samples of known electrical conductivity ratio (*e.g.*, Mantyla, 1980; 1987; 1994; Aoyama *et al.*, 2002; Kawano *et al.*, 2005). More recently, a new seawater standard under [TEOS-10](#) is “absolute salinity” ( $S_A$ ; Millero *et al.*, 2008). Unless otherwise indicated, it is likely that the older salinity data collected have not been corrected for “standard sea water” changes (Mantyla, 1994) or converted to a salinity scale other than the scale the measurements were reported in (*e.g.*, practical salinity scale, PSS or  $S_A$ ). Salinity data in the WOD23 include the scale of the measurements and IAPSO standard seawater batch number if reported by the originator of the data.

As it was already mentioned before, relatively low depth resolution CTD profiles present in the OSD dataset may be associated with high-resolution CTD profiles in the CTD dataset. This is done so that users of the OSD dataset have access to CTD values collected at the same time and depth/pressure that water samples are collected and to maintain a more or less concise size for the OSD dataset. Similarly, users of the CTD dataset may have access to low vertical resolution profiles for other physical and chemical variables collected during a cast (Table 2.1).

OSD chemical data received at NCEI are reported by originators of the data in a variety of units that may differ from the WOD standard units (Table 2.1) and the international system of units in oceanography (UNESCO, 1985). When originator’s units differ from the WOD common units, the data are converted from the originator’s units to a common set of units to facilitate the use of the data. For example, originator’s chemical content units reported on a per-volume units (*e.g.*, moles per liter of seawater) were converted to per-mass units (*e.g.*, moles per kilogram of seawater) using a constant density of seawater equal to  $1025 \text{ kg}\cdot\text{m}^{-3}$ . Conversion of chemical units were necessary use standard element atomic weights (CRC, 1993).

Biogeochemical variables such as dissolved oxygen ( $\text{O}_2$ ), dissolved inorganic macronutrients (reactive phosphate, nitrate, and silicate or silicic acid), inorganic carbon species (alkalinity, dissolved inorganic carbon, partial pressure of carbon dioxide) and pH are nominally non-conservative variables. Their contents result from diffusion and advection of waters with varied preformed contents, by biogeochemical processes, and by atmospheric inputs (Redfield *et al.*, 1963).

The WOD23 includes **nitrate plus nitrite** (N+N) and **nitrate** data. The nitrate, nitrite, and N+N content is often estimated by photometric analyses of seawater samples where in one case nitrate is measured indirectly by effectively reducing nitrate to nitrite while in the other only nitrite is measured (*e.g.*, Strickland and Parsons, 1972; Gordon *et al.*, 1993; Becker *et al.*, 2020). The content of nitrate is then obtained by the difference between the estimated contents of N+N and nitrite. It is difficult to verify in some cases whether the older nitrate data collected are N+N or nitrate. Except for low oxygen zones, the nominal nitrite contents in the open ocean are generally low. When reported by the originator of the data, WOD23 includes metadata information about

whether the labeled nitrate measurement is reported as N+N data.

Historical dissolved inorganic carbon, alkalinity, partial pressure of carbon dioxide and pH originator data do not always include information about the methods, instruments, and scales used (Millero *et al.*, 1993; DOE, 1994). When reported, modern chemical data include additional metadata including any use of consensus or certified reference materials (CRM) and measurement scales depending on variables.

The **dissolved O<sub>2</sub>** content is often analyzed following various modifications of the classical “Winkler titration” followed by end-detections by visual (e.g., starch indicator), amperometric, or photometric methods (Winkler, 1888; Carpenter, 1965; Dickson, 1994; Langdon, 2010). Carpenter (1965) outlined a whole bottle titration method that helped minimize the amount of error that was introduced during the O<sub>2</sub> titration. Most modern O<sub>2</sub> chemical titration measurements nominally use modifications of Carpenter’s whole bottle titration method and an amperometric or photometric end-detection with an estimated uncertainty of about 0.5 to 1  $\mu\text{mol kg}^{-1}$ .

It is worth noting that the CTD dataset contains high-resolution O<sub>2</sub> data obtained from electronic sensors mounted on the CTD rosette frame. For example, polarographic O<sub>2</sub> electronic sensors estimate seawater O<sub>2</sub> content by estimating the flux of oxygen molecules per unit time that diffuse through a permeable membrane. The PFL dataset *also* contains a number of relatively high vertical resolution profiles using O<sub>2</sub> optodes (Bittig *et al.*, 2018). These high-resolution O<sub>2</sub> profiles obtained by electronic sensors can be subject to sensor drift problems resulting in relatively lower data quality than O<sub>2</sub> profiles obtained by chemical analysis of discrete water samples Winkler measurements. In recent years, the quality of sensor-based O<sub>2</sub> data have improved dramatically. The CTD O<sub>2</sub> data are often calibrated using discrete O<sub>2</sub> measurements of the water column (Owens and Millard, 1985; Uchida *et al.*, 2010). The O<sub>2</sub> profiles in the OSD, CTD and PFL datasets are kept separate so that users can decide which data to use.

Dissolved gases and tracers help in the interpretation of how ocean surface properties are transmitted into the ocean’s interior, the dynamics of ocean circulation, biochemical cycles, ocean-atmosphere interactions, and to help infer paleotemperatures (Broecker and Peng, 1982).

The OSD dataset includes **noble gases** such as neon, argon, and helium. The distributions of these gases are useful, for example, to further our understanding of the ocean circulation and air-sea gas flux interactions (Schlosser, 1986). The distributions of transient tracers provide estimates of oceanic ventilation rates; a measure of water mass spreading rates from the surface to the ocean interior. Specifically, transient tracers such as bomb-fallout radionuclides and natural isotopes function as “clocks” recording the elapsed time since a parcel of water was last in contact with the oceanic surface layer (e.g., Schlosser *et al.*, 1991; Jenkins, 1982; 1987; Jenkins and Rhines, 1980; Östlund and Rooth, 1990). For example, tritium was delivered to the atmosphere as a result of the atmospheric thermonuclear weapon tests in the late 1950s and early 1960s. Chlorofluorocarbons are man-made gases with high greenhouse potential. Their time history within the water column provides important clues regarding the oceanic uptake of atmospheric gases (Bullister and Weiss, 1988; Smethie, 1993; on Haine *et al.*, 1995). There are a number of man-made freons produced and dissolved in the ocean. The most commonly sampled freons (chlorofluorocarbon, CFC) in the ocean are CFC-11, CFC-12, and CFC-113. CFCs were used worldwide as refrigerants, propellants, and cleaning solvents. The temporal evolution of the CFC contents in oceanic waters is essentially controlled by the atmospheric record. Most of the high-

quality transient tracer data in the OSD dataset were collected starting with the GEOSECS program in the early 1970s, and later as part of WOCE, CLIVAR, GO-SHIP, GEOTRACES.

In addition to the observed data (profiles as a function of depth of each sampled variable), OSD casts include additional information (commonly referred to as “station header information”) such as **ocean surface conditions** (*e.g.*, wave direction and height, sea state), **meteorological observations** (*e.g.*, cloud cover and type, visibility, wind speed and direction, barometric pressure, dry and wet bulb temperature), water color and transparency (*e.g.*, Secchi disk depth).

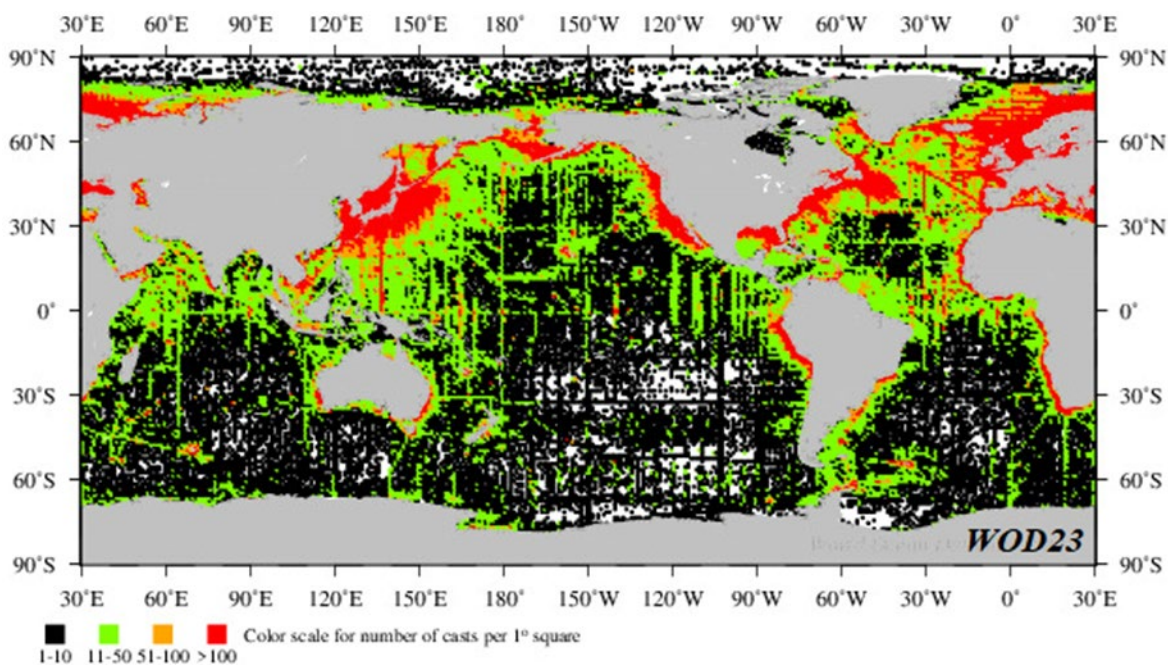
[WOD23 Manual](#) (Garcia *et al.* 2024) describes the WOD23 cast header information, originator’s information about the data collected (instrumentation, methods, units, quality flags, stations and cruise labels, institutions, platforms, principal investigators, *etc.*) and data format. We refer collectively to this information as station metadata. The cast metadata included in the OSD dataset are not meant to substitute any originator’s information included with any oceanographic cruise data reports or scientific manuscripts which may be associated with any particular OSD data values. Metadata are included in the WOD dataset as a means to facilitate identifying information about the measurements that may be available with each cast if the information is available. Metadata are included with each cast in the form of header information when metadata were included with data received at NCEI.

The biogeochemical data in the OSD dataset have been measured using a variety of manual and automated analytical methods. It is difficult to estimate the measurement precision, reproducibility, uncertainty of the historical chemical data in part because (1) there has not been a generally accepted set of standard international analytical oceanographic methods; (2) there has been a continuous availability over time of new or improved analytical techniques for the sampling and determination of the content of dissolved and particulate constituents in seawater; (3) there is the practical difficulty of conducting periodic comparison of the reproducibility oceanographic data collected by oceanographic institutions worldwide. Some major international oceanographic sampling programs have adopted sample and measurement protocols such as for example, WOCE and GO-SHIP. These protocols provide relatively consistent high-quality measurements. In the past few years Consensus or Certified Reference Materials (CRMs) of known chemical quality have been used for the analysis, for example, Dissolved Inorganic Carbon and Alkalinity (DOE, 1994), dissolved inorganic nutrients (*e.g.*, International Nutrient Inter-comparison Voyage, [INIV](#)), Dissolved Organic Carbon (*e.g.*, Dennis Hansell, DOC [Consensus Reference Material](#)). The adoption of CRMs facilitates the interlaboratory comparison of measurements collected by different ocean observing systems. For historical context, Farrington (2000) provides a summary of advances in chemical oceanography for the 1950 to 2000 time period.

## 2.4. OSD DATA COVERAGE

The sampling coverage of the OSD variables is worldwide and for some variables spans several decades (Tables 2.2 and 2.3). Table 2.3 shows that > 87% of all the data in the OSD are from global data sources indicating the substantial global effort to observe the ocean at a great public cost. The number of OSD casts added to WOD has increased greatly since 1974 (Figures 2.3 and 2.6, Table 2.3), however, the coverage for each variable is not uniform in space (Figure 2.5) or time as displayed by the temporal data distributions per each variable contained in the OSD dataset (Figures 2.7-2.29). The largest numbers of oceanographic profiles present in the OSD

dataset are temperature, salinity, dissolved oxygen, and dissolved nutrient measurements. OSD data represents 17.5% of the WOD23 data collection (Figure 1.4).



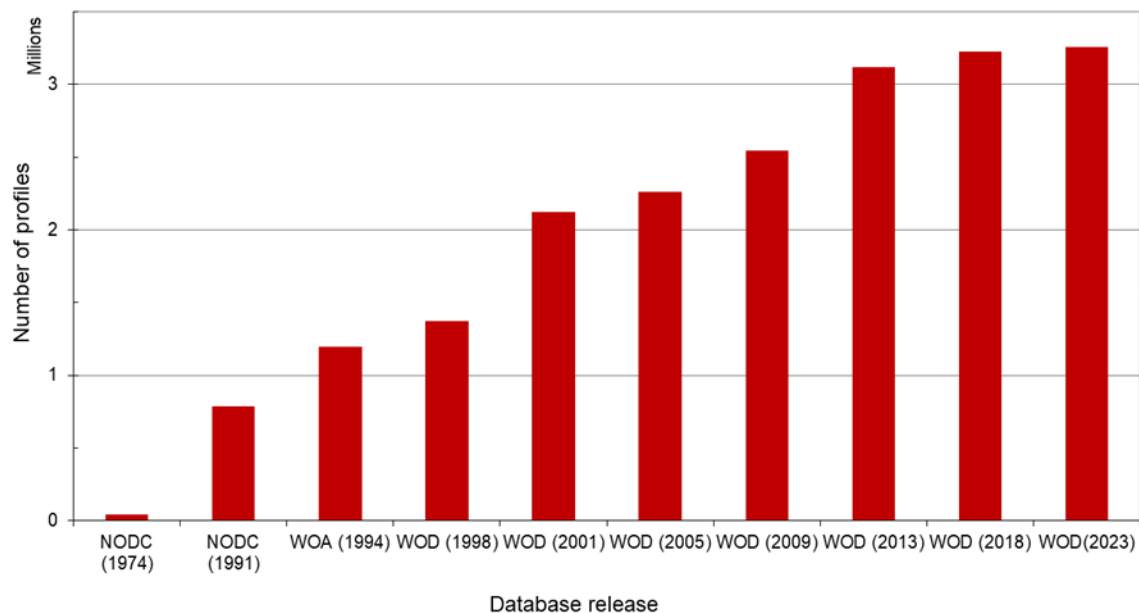
**Figure 2.5. Number and distribution of OSD casts in WOD23 by one-degree squares**

This global non-uniformity of the number of chemical profiles can be attributed to different reasons. First, historical oceanographic cruises typically sampled individual or a limited suite of tracers to deduce aspects of the ocean. In other words, oceanographic cruises have a specific research goal that may require sampling of a limited number of variables. Second, the sampling and analysis of biochemical variables is more labor intensive when compared to temperature or conductivity measurements obtained by CTD instruments.

## **2.5. PARAMETERS AND METADATA NOT INCLUDED IN THE OSD DATASET**

The WOD includes data for other biochemical variables not available as part of the WOD23 release because they did not meet the minimum quality control (QC) that is necessary to be included. This QC comprise of data format, measurements position/date/time, gradients, ranges, and other checks (see details in [WOD23 Manual](#), Garcia *et al.* 2024). The variables not present in the WOD23 release include dissolved and particulate organic carbon, nitrite, total phosphorus, ammonia, various chlorophyll pigments, and primary production. However, the [NCEI archive](#) contains an exact record of the data provided. Users of the WOD can retrieve the original data as sent to NCEI.





**Figure 2.6 Number of OSD casts in NCEI (previously NODC)/WDS databases as a function of time.** The number of casts available at NODC prior to 1994 after Levitus (1982) and NODC Temperature-Profile CD-ROM (1991). The number of casts after 1994 based on the WOA and WOD series.

## 2.6. PROSPECTS FOR THE FUTURE

It is expected that relatively large amounts of historical chemical and biological data still exist in non-digital and digital form at data centers, research institutions, universities, and libraries worldwide. Biogeochemical data is also expected to become available from ongoing and future international oceanographic field programs such as GO-SHIP. We note that the number of OSD casts in the WOD have declined considerably over time, particularly after 2001. For example, the number of O<sub>2</sub> profiles have declined from 322,521 in the 1981 to 2000 time period to 137,726 between 2001 and 2022, a reduction of about 43%. The decline after about 2020 is in part due to the impact of the Coronavirus disease (COVID) pandemic on the observing of the global ocean (Boyer *et al.*, 2023). The proliferation of autonomous Argo profilers provides much greater spatial and temporal coverage than ship board observations using a CTD. The reduction in the number of hydrographic casts to sample *in situ* biochemical variables is concerning. These measurements are considered of higher science-quality (*e.g.*, precision, reproducibility, uncertainty) than the ones gathered recently thought chemical sensors mounted on CTDs and BGC-Argo floats, for example. We note that BGC Argo floats equipped with physical and chemical sensors such as O<sub>2</sub>, Nitrate, pH can contribute much additional spatial and temporal coverage to the WOD.

While the WOD Team aims to ingest all possible open access data from worldwide sources, we are encouraged by the availability and sharing of oceanographic data with NCEI. For example, NCEI operates the Ocean Carbon and Acidification Data System ([OCADS](#)) specializing in the management of ocean carbon and ocean acidification data. We also maintain data sharing cooperation with the IODE network of National Oceanographic Data Centers and international projects and organizations such as the International Council for the Exploration of the Sea ([ICES](#)). NCEI has several operational automated ocean data archival process that seek to maintain a



representative global ocean data archival. For example, NCEI has developed automated ocean profile data archival from the U.S. National Science Foundation (NSF) Arctic Data Center ([ADC](#)), [long-term archive](#) of [Argo core](#) and [BGC Argo](#) data, and CLIVAR and Carbon Hydrographic Data Office ([CCHDO](#)), Ice Tethered Profilers ([ITP](#)) - Woods Hole Oceanographic Institution, Global Temperature and Salinity Profile Programme ([GTSP](#)). At the time of this writing, we are exploring open data sharing with the European Marine Observation and Data Network ([EMODnet](#)). We seek to expand the number of chemical variables in the WOD to accommodate additional tracer elements and isotope data such as from the [GEOTRACES](#) international program.

The WOD is a worldwide source of unrestricted access to the most comprehensive and representative *in situ* observation based global database of the past 250 years (1772 to present). For example, the WOD serves as a source of data to global efforts such as the IOC Ocean Data and Information System ([ODIS](#)) and the [UN Decade of Ocean Science for Sustainable Development \(2021-2030\)](#).

Future releases of the WOD will be enhanced by the addition of more data and metadata. It is hoped that users of the WOD inform us of sources of historical data not present in the database as well as any data or metadata errors that might be present in the database at NCEI. Identification of new sources of chemical oceanographic data to the WOD is beneficial for improving mechanisms for long-term data archival, data management, and data distribution into national and international data archives. Addition of new data will help improve the release of an improved high-quality global, scientifically quality-controlled ocean profile-plankton database, ocean data products, and diagnostic studies as well as help to provide observational constraints on oceanic variability studies. Newly acquired data could be submitted using [Send2NCEI](#) interface. Users feedback is always welcome and should be send to our [help desk](#).

**Table 2.2. The number of Ocean Station Data (OSD) casts as a function of year**

<b>YEAR</b>	<b>COUNT</b>	<b>YEAR</b>	<b>COUNT</b>	<b>YEAR</b>	<b>COUNT</b>	<b>YEAR</b>	<b>COUNT</b>
1772	2	1835	-	1898	6,369	1961	40,023
1773	1	1836	8	1899	6,471	1962	37,539
1774	-	1837	17	1900	6,463	1963	50,694
1775	-	1838	11	1901	6,748	1964	56,941
1776	-	1839	15	1902	7,020	1965	58,234
1777	-	1840	9	1903	7,902	1966	57,189
1778	-	1841	21	1904	7,869	1967	60,209
1779	-	1842	8	1905	8,551	1968	57,094
1780	-	1843	-	1906	8,037	1969	65,663
1781	-	1844	-	1907	8,211	1970	58,817
1782	-	1845	-	1908	7,551	1971	64,612
1783	-	1846	3	1909	7,963	1972	72,319
1784	-	1847	28	1910	8,435	1973	67,301
1785	-	1848	-	1911	10,025	1974	61,437
1786	-	1849	1	1912	8,248	1975	56,940
1787	-	1850	3	1913	8,799	1976	61,238
1788	-	1851	1	1914	7,655	1977	60,087
1789	-	1852	-	1915	3,041	1978	69,199
1790	-	1853	-	1916	1,845	1979	71,488
1791	-	1854	-	1917	1,941	1980	66,332
1792	-	1855	4	1918	2,128	1981	67,406
1793	-	1856	-	1919	4,471	1982	63,719
1794	-	1857	6	1920	7,311	1983	62,784
1795	-	1858	23	1921	7,930	1984	65,958
1796	-	1859	5	1922	7,247	1985	70,973
1797	-	1860	1,053	1923	3,538	1986	73,393
1798	-	1861	1,351	1924	4,817	1987	71,452
1799	-	1862	1,602	1925	4,928	1988	66,885
1800	84	1863	606	1926	6,506	1989	69,245
1801	-	1864	2,329	1927	7,509	1990	65,387
1802	-	1865	352	1928	7,525	1991	45,855
1803	-	1866	1,494	1929	6,930	1992	39,666
1804	10	1867	78	1930	8,667	1993	38,372
1805	1	1868	793	1931	9,557	1994	32,031
1806	-	1869	567	1932	13,460	1995	38,828
1807	-	1870	737	1933	11,786	1996	32,185
1808	-	1871	50	1934	14,686	1997	30,860

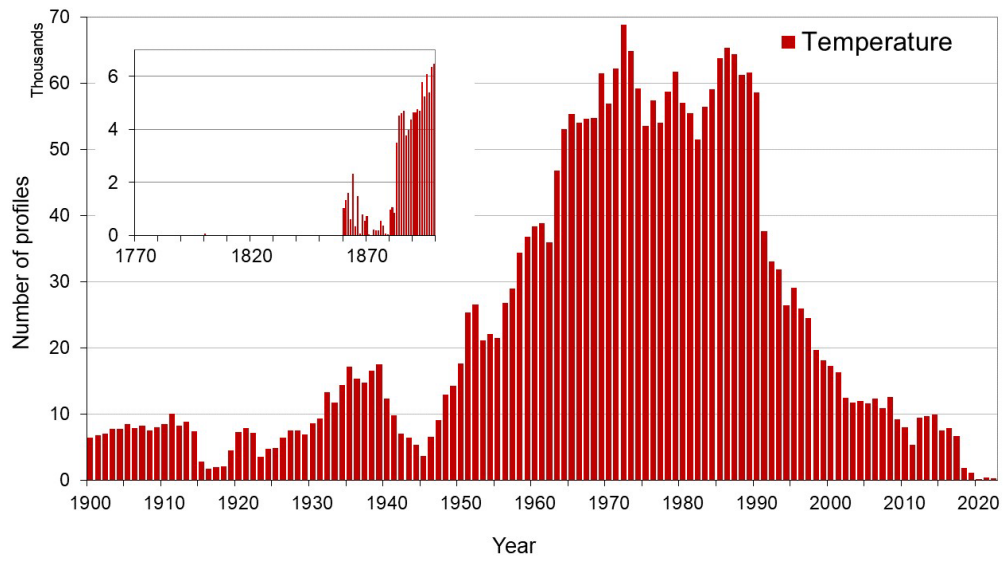
<b>YEAR</b>	<b>COUNT</b>	<b>YEAR</b>	<b>COUNT</b>	<b>YEAR</b>	<b>COUNT</b>	<b>YEAR</b>	<b>COUNT</b>
<b>1809</b>	-	1872	7	1935	17,380	1998	26,263
<b>1810</b>	-	1873	226	1936	15,518	1999	25,798
<b>1811</b>	-	1874	194	1937	14,765	2000	20,717
<b>1812</b>	-	1875	195	1938	16,774	2001	18,785
<b>1813</b>	-	1876	573	1939	17,523	2002	15,414
<b>1814</b>	-	1877	385	1940	12,378	2003	15,069
<b>1815</b>	-	1878	90	1941	9,756	2004	14,825
<b>1816</b>	5	1879	57	1942	7,025	2005	14,845
<b>1817</b>	27	1880	976	1943	6,388	2006	15,070
<b>1818</b>	3	1881	1,071	1944	5,389	2007	13,159
<b>1819</b>	-	1882	873	1945	3,693	2008	14,612
<b>1820</b>	2	1883	3,493	1946	6,595	2009	11,785
<b>1821</b>	-	1884	4,510	1947	9,131	2010	10,865
<b>1822</b>	-	1885	4,599	1948	13,392	2011	7,272
<b>1823</b>	-	1886	4,704	1949	14,425	2012	11,556
<b>1824</b>	2	1887	3,776	1950	19,159	2013	12,446
<b>1825</b>	10	1888	3,971	1951	26,130	2014	11,081
<b>1826</b>	18	1889	4,363	1952	27,393	2015	8,063
<b>1827</b>	30	1890	4,632	1953	22,697	2016	8,495
<b>1828</b>	13	1891	4,655	1954	23,982	2017	7,371
<b>1829</b>	-	1892	4,756	1955	23,486	2018	2,061
<b>1830</b>	-	1893	4,692	1956	27,828	2019	1,267
<b>1831</b>	-	1894	5,788	1957	29,984	2020	37
<b>1832</b>	-	1895	5,233	1958	36,429	2021	424
<b>1833</b>	-	1896	6,087	1959	38,474	2022	319
<b>1834</b>	-	1897	5,381	1960	39,579		

**Table 2.3. Number of OSD casts provided by each country**

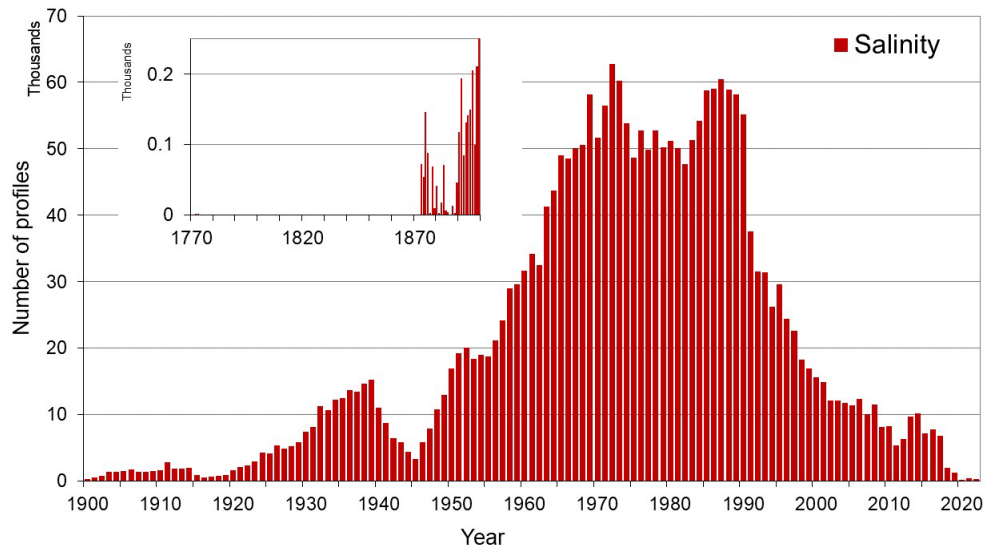
<b>ISO<sup>1</sup> Country Codes</b>	<b>ISO Country Name</b>	<b>OSD Casts</b>	<b>% of Total</b>
<b>SU</b>	Union of Soviet Socialist Republics	745,396	22.89%
<b>JP</b>	Japan	598,761	18.39%
<b>US</b>	United States	415,263	12.75%
<b>SE</b>	Sweden	289,561	8.89%
<b>GB</b>	Great Britain	144,576	4.44%
<b>CA</b>	Canada	124,411	3.82%
<b>NO</b>	Norway	114,125	3.51%
<b>99</b>	Unknown / International	99,560	3.06%
<b>DE</b>	Germany	93,827	2.88%
<b>FI</b>	Finland	65,195	2.00%
<b>DK</b>	Denmark	58,020	1.78%
<b>KR</b>	Korea, Republic of	51,638	1.59%
<b>FR</b>	France	49,971	1.53%
<b>PE</b>	Peru	43,525	1.34%
<b>AU</b>	Australia	39,567	1.22%
<b>NL</b>	Netherlands	35,146	1.08%
<b>ZA</b>	South Africa	28,985	0.89%
<b>RU</b>	Russian Federation	26,945	0.83%
<b>PL</b>	Poland	26,689	0.82%
<b>IS</b>	Iceland	20,876	0.64%
<b>UA</b>	Ukraine	15,917	0.49%
<b>DU</b>	East Germany	15,608	0.48%
<b>IT</b>	Italy	12,074	0.37%
<b>IE</b>	Ireland	11,242	0.35%
<b>BE</b>	Belgium	10,801	0.33%
<b>BR</b>	Brazil	9,572	0.29%
<b>ES</b>	Spain	7,284	0.22%
<b>PT</b>	Portugal	6,539	0.20%
<b>CN</b>	China, The Peoples Republic of	5,509	0.17%
<b>YU</b>	Yugoslavia	5,455	0.17%
<b>EE</b>	Estonia	5,170	0.16%
<b>AR</b>	Argentina	5,046	0.15%
<b>CL</b>	Chile	4,914	0.15%
<b>IN</b>	India	4,488	0.14%
<b>ID</b>	Indonesia	4,397	0.14%
<b>TW</b>	Taiwan	4,062	0.12%
<b>TR</b>	Turkey	3,996	0.12%

<b>ISO<sup>1</sup> Country Codes</b>	<b>ISO Country Name</b>	<b>OSD Casts</b>	<b>% of Total</b>
<b>RO</b>	Romania	3,639	0.11%
<b>VE</b>	Venezuela	3,590	0.11%
<b>EC</b>	Ecuador	3,498	0.11%
<b>GR</b>	Greece	3,489	0.11%
<b>IL</b>	Israel	3,463	0.11%
<b>CI</b>	Cote D'Ivoire (Ivory Coast)	3,185	0.10%
<b>TH</b>	Thailand	2,801	0.09%
<b>GH</b>	Ghana	2,670	0.08%
<b>MG</b>	Malagasy Republic	2,523	0.08%
<b>LV</b>	Latvia	2,266	0.07%
<b>MC</b>	Monaco	2,054	0.06%
<b>SN</b>	Senegal	1,975	0.06%
<b>NZ</b>	New Zealand	1,942	0.06%
<b>CD</b>	Congo	1,864	0.06%
<b>LT</b>	Lithuania	1,626	0.05%
<b>MX</b>	Mexico	1,457	0.04%
<b>NC</b>	New Caledonia	1,344	0.04%
<b>CO</b>	Colombia	1,338	0.04%
<b>MR</b>	Mauritania	1,217	0.04%
<b>NG</b>	Nigeria	980	0.03%
<b>CU</b>	Cuba	976	0.03%
<b>AT</b>	Austria	773	0.02%
<b>AO</b>	Angola t	621	0.02%
<b>EG</b>	Arab Republic of Egypt	544	0.02%
<b>SG</b>	Singapore	412	0.01%
<b>TN</b>	Tunisia	280	0.01%
<b>PH</b>	Philippines	235	0.01%
<b>MA</b>	Morocco	199	0.01%
<b>LB</b>	Lebanon	187	0.01%
<b>PK</b>	Pakistan	167	0.01%
<b>DZ</b>	Algeria	166	0.01%
<b>MY</b>	Malaysia	154	0.00%
<b>PA</b>	Panama	139	0.00%
<b>YE</b>	Yemen	85	0.00%
<b>MT</b>	Malta	66	0.00%
<b>ZZ</b>	Miscellaneous organization	1	0.00%

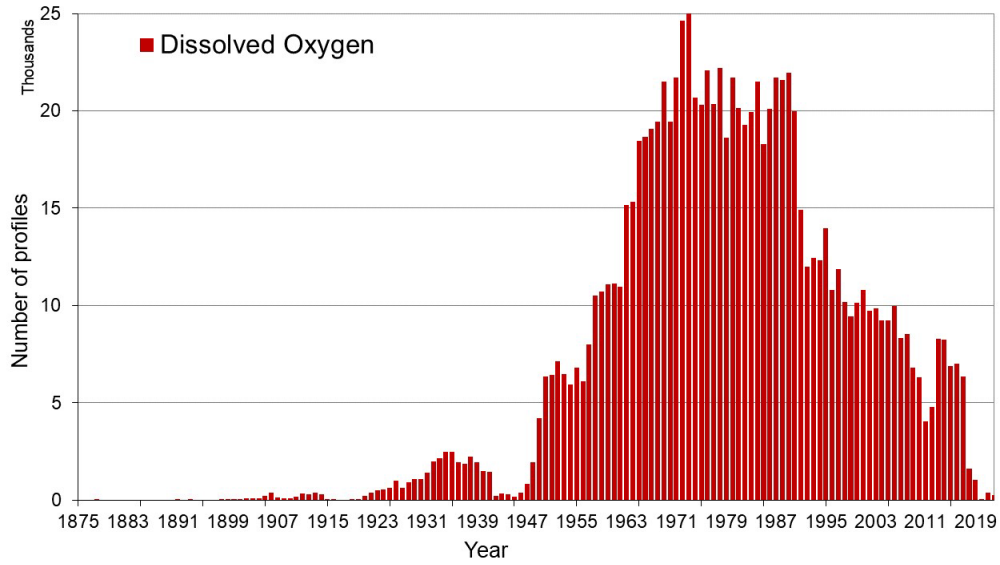
<sup>1</sup> ISO, International Organization for Standardization)



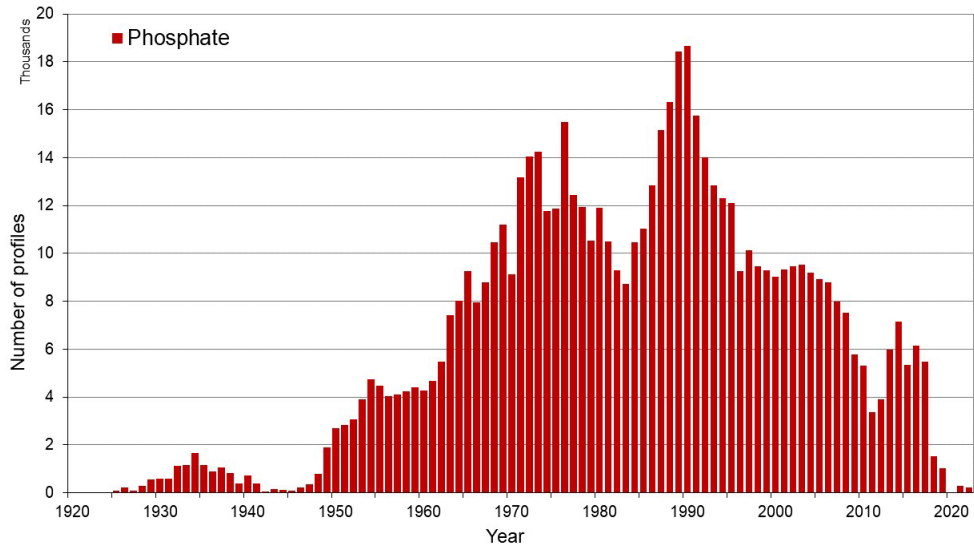
**Figure 2.7. Time series of the number of temperature profiles in WOD23 OSD dataset**



**Figure 2.8. Time series of the number of salinity profiles in WOD23 OSD dataset**

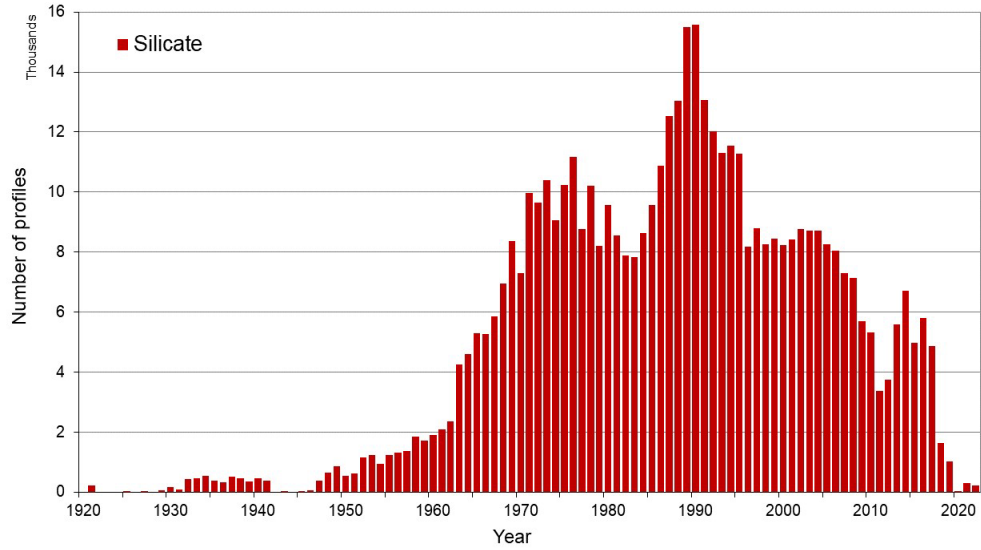


**Figure 2.9. Time series of the number of dissolved oxygen profiles in WOD23 OSD dataset**

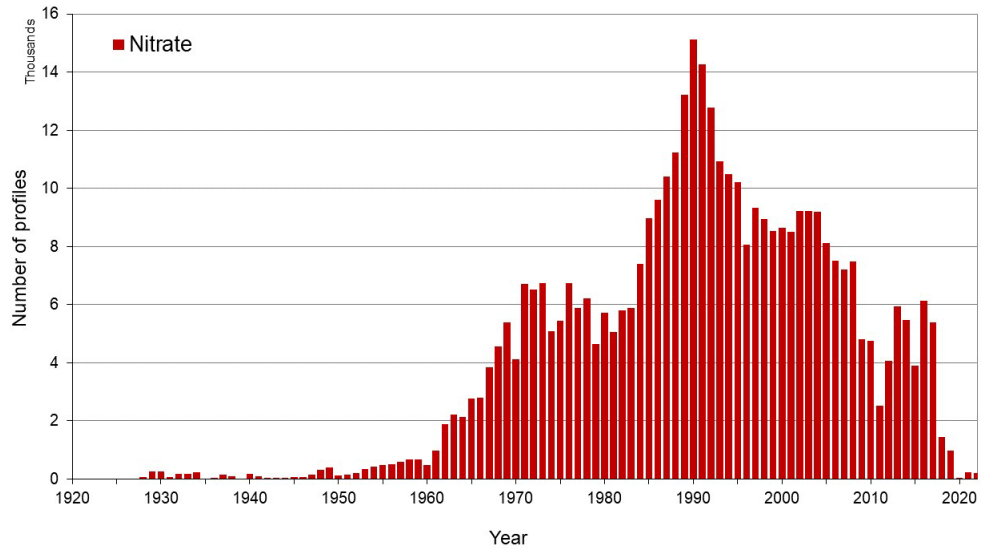


**Figure 2.10. Time series of the number of Phosphate profiles in WOD23 OSD dataset**

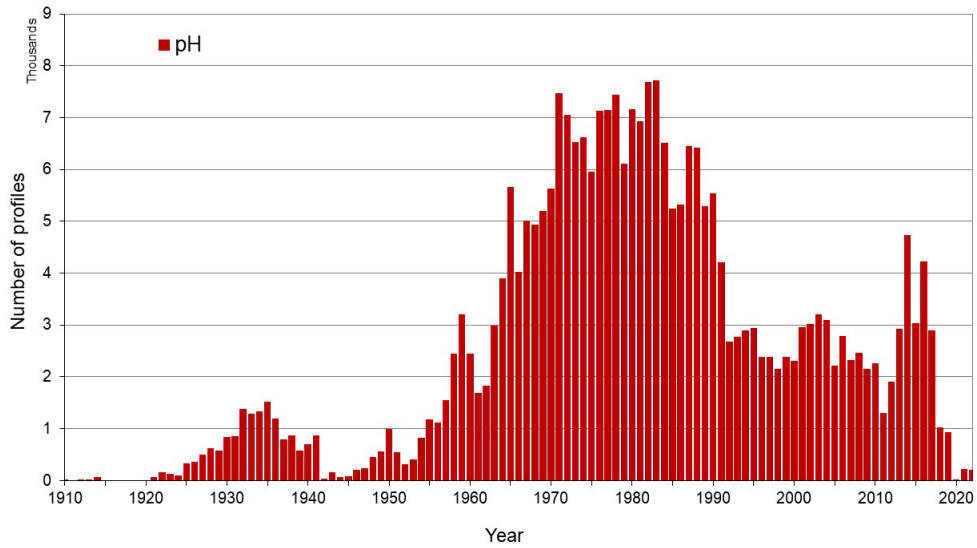




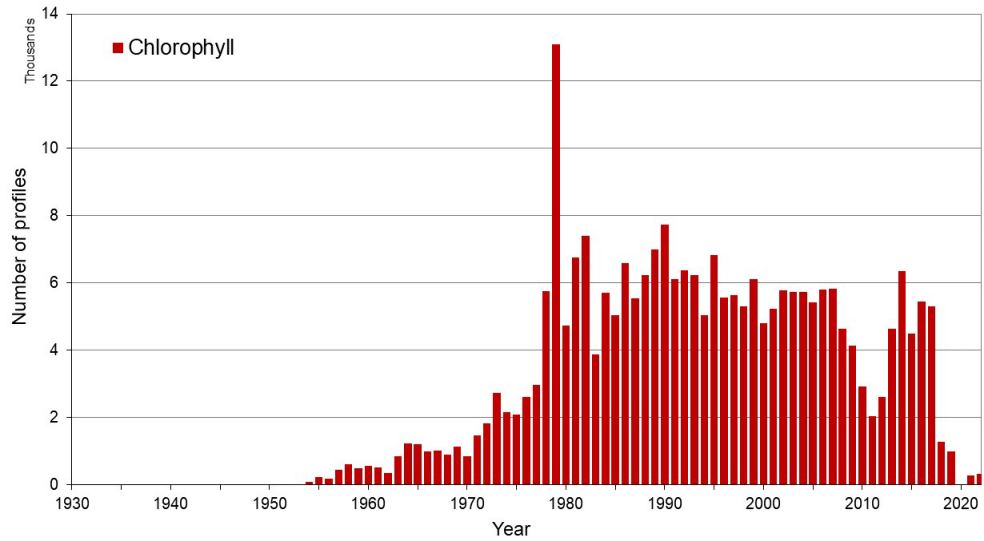
**Figure 2.11. Time series of the number of silicate profiles in WOD23 OSD dataset**



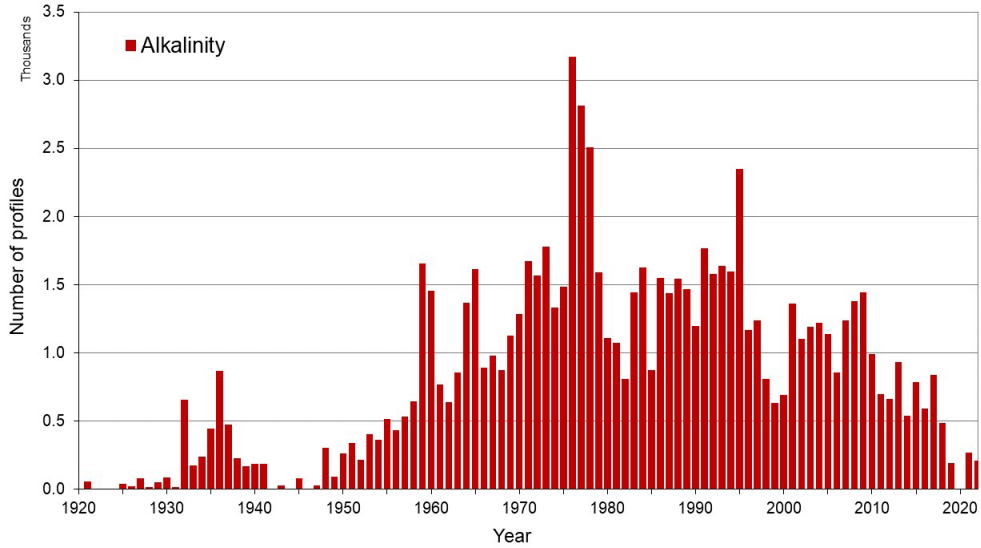
**Figure 2.12. Time series of the number of Nitrate profiles in WOD23 OSD dataset**



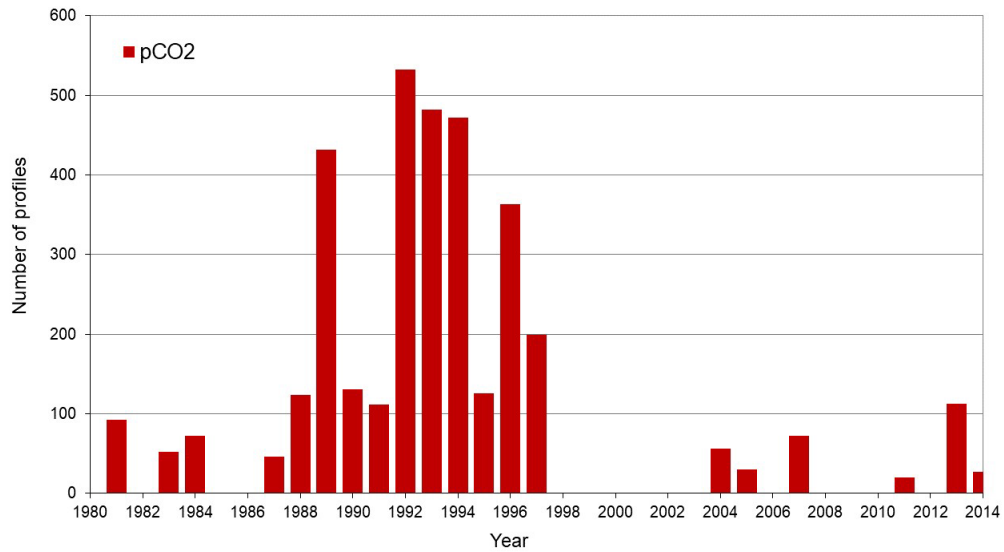
**Figure 2.13. Time series of the number of pH profiles in WOD23 OSD dataset**



**Figure 2.14. Time series of the number of Chlorophyll profiles in WOD23 OSD dataset**



**Figure 2.15. Time series of the number of alkalinity profiles in WOD23 OSD dataset**



**Figure 2.16. Time series of the number of partial pressure of carbon dioxide profiles in WOD23 OSD dataset**

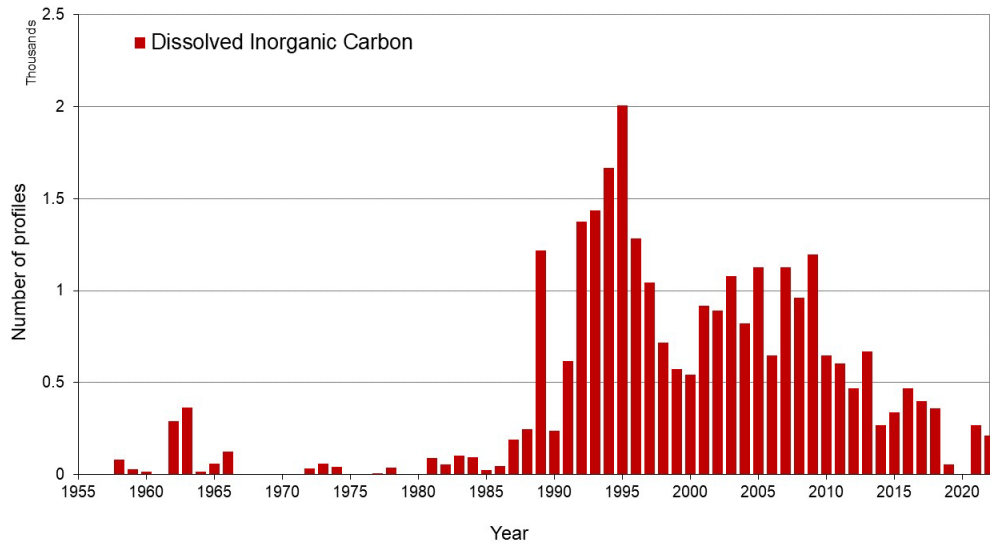


Figure 2.17. Time series of the number of dissolved inorganic carbon profiles in WOD23 OSD dataset

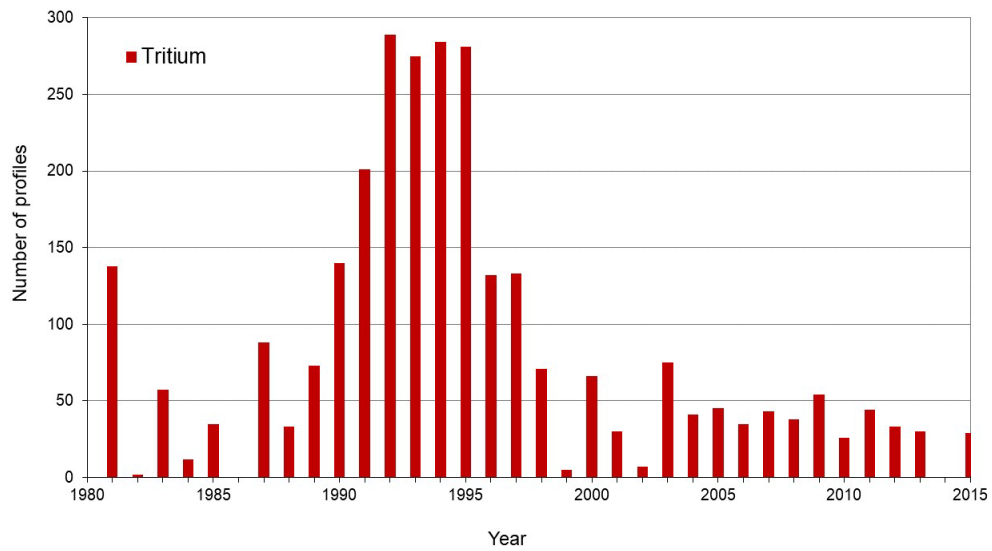
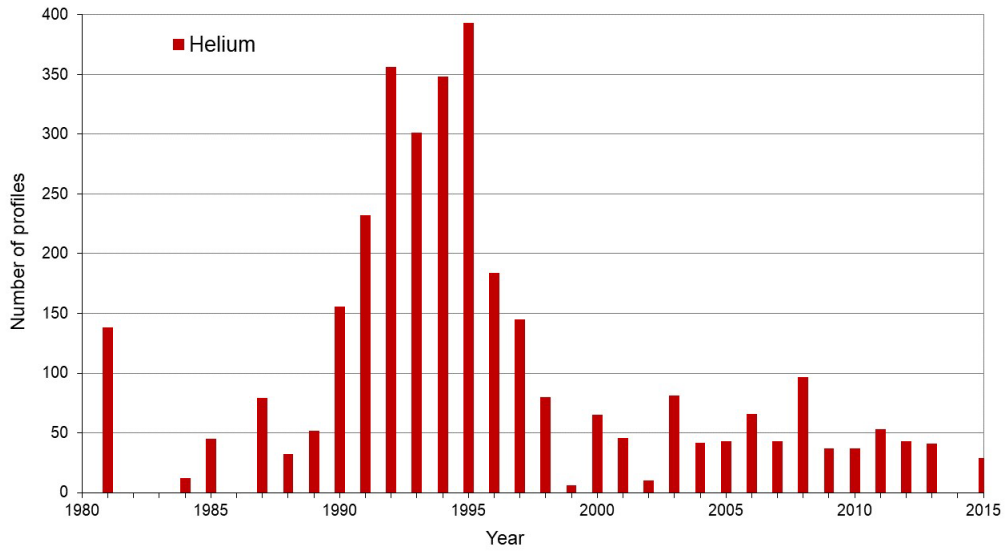
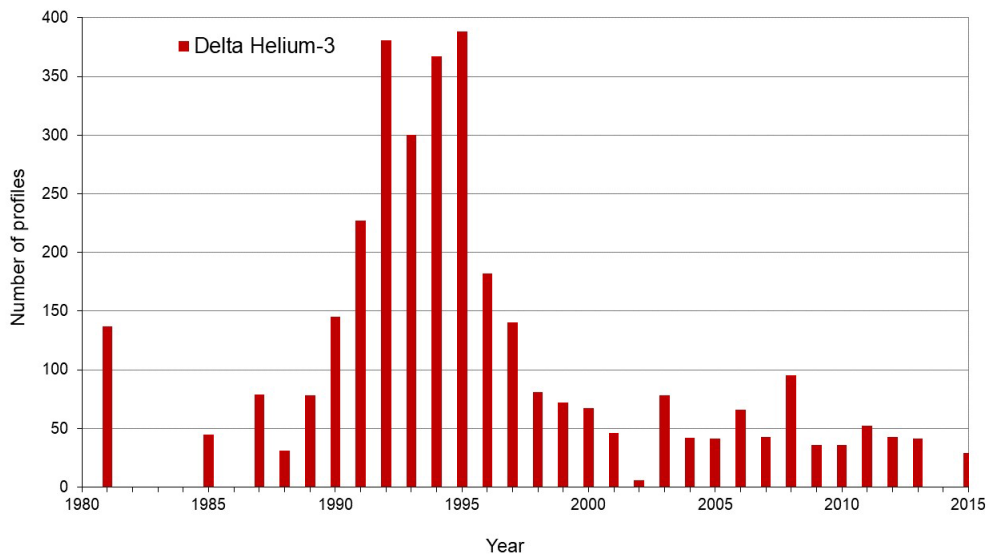


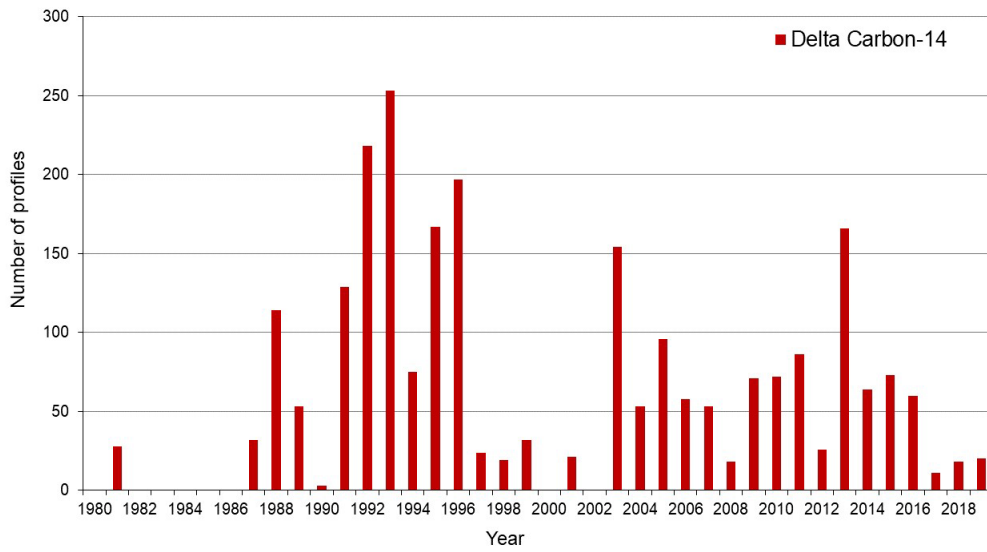
Figure 2.18. Time series of the number of tritium profiles in WOD23 OSD dataset



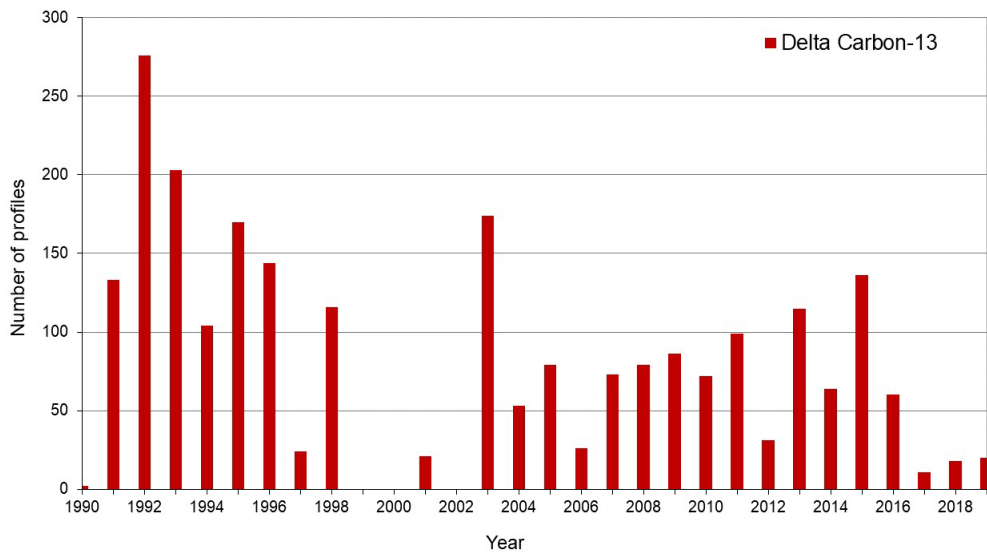
**Figure 2.19. Time series of the number of helium profiles in WOD23 OSD dataset**



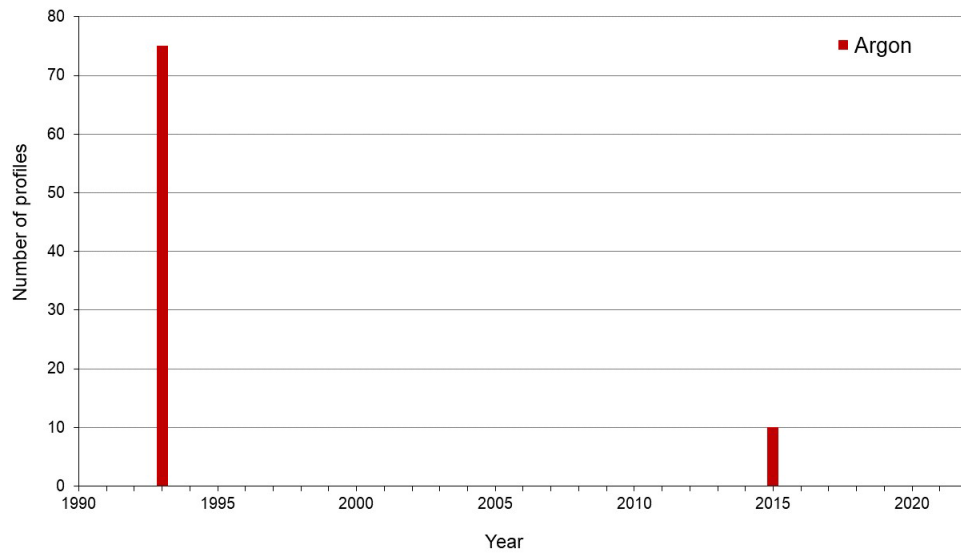
**Figure 2.20. Time series of the number of  $\Delta$ -helium-3 profiles in WOD23 OSD dataset**



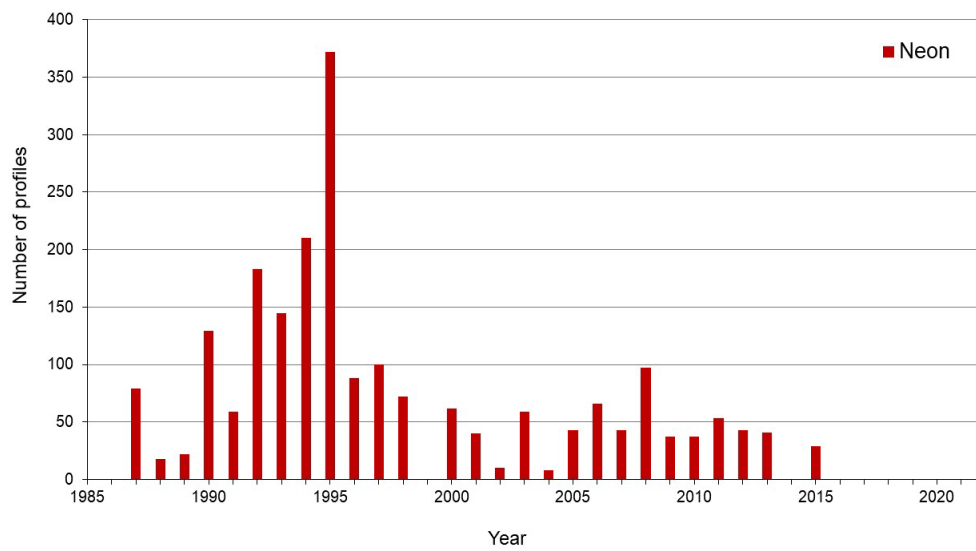
**Figure 2.21. Time series of the number of  $\Delta$ -carbon-14 profiles in WOD23 OSD dataset**



**Figure 2.22. Time series of the number of  $\Delta$ -carbon-13 profiles in WOD23 OSD dataset**

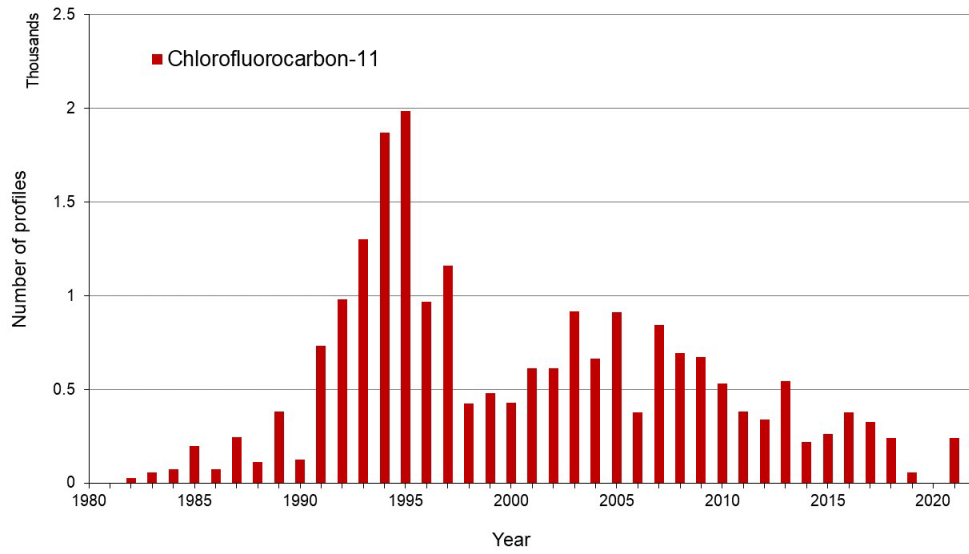


**Figure 2.23. Time series of the number of argon profiles in WOD23 OSD dataset**

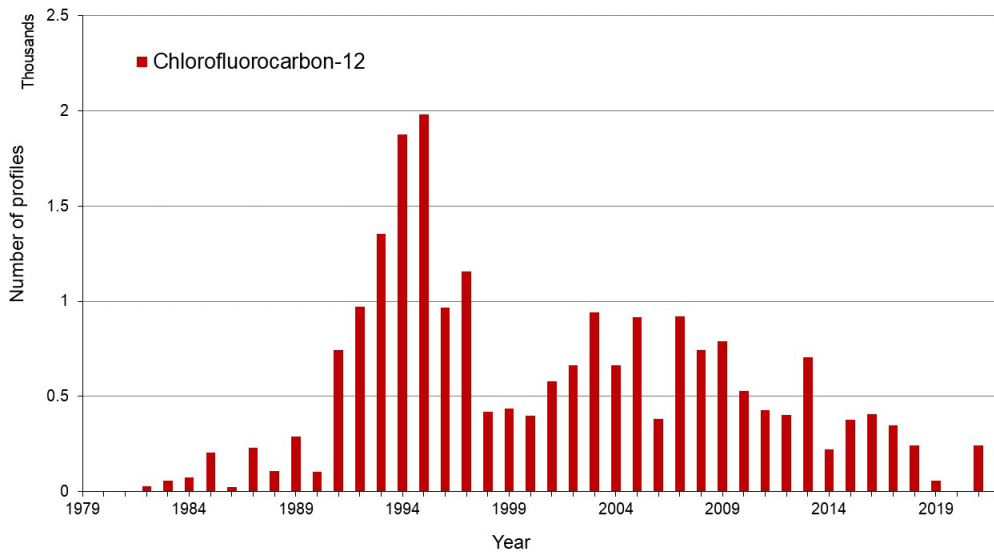


**Figure 2.24. Time series of the number of neon profiles in WOD23 OSD dataset**

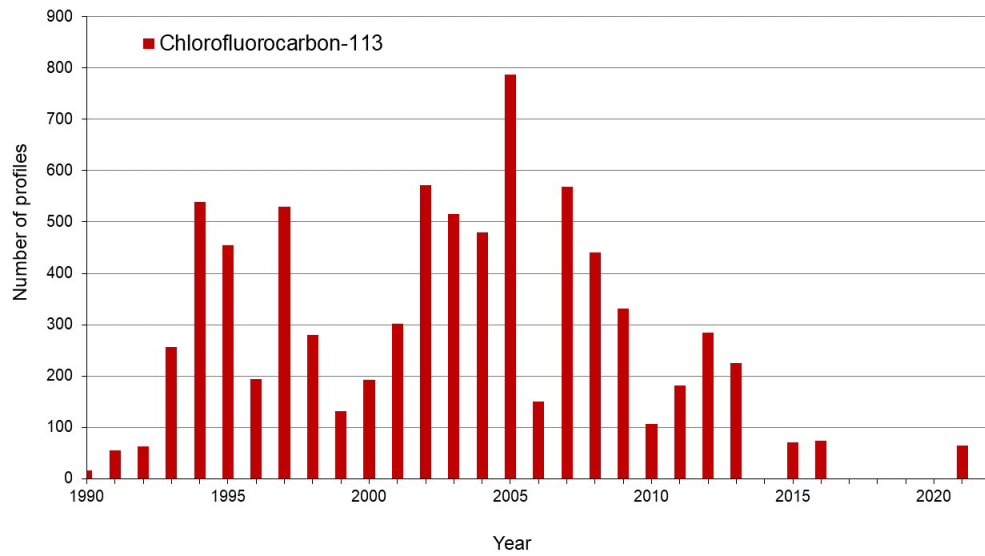




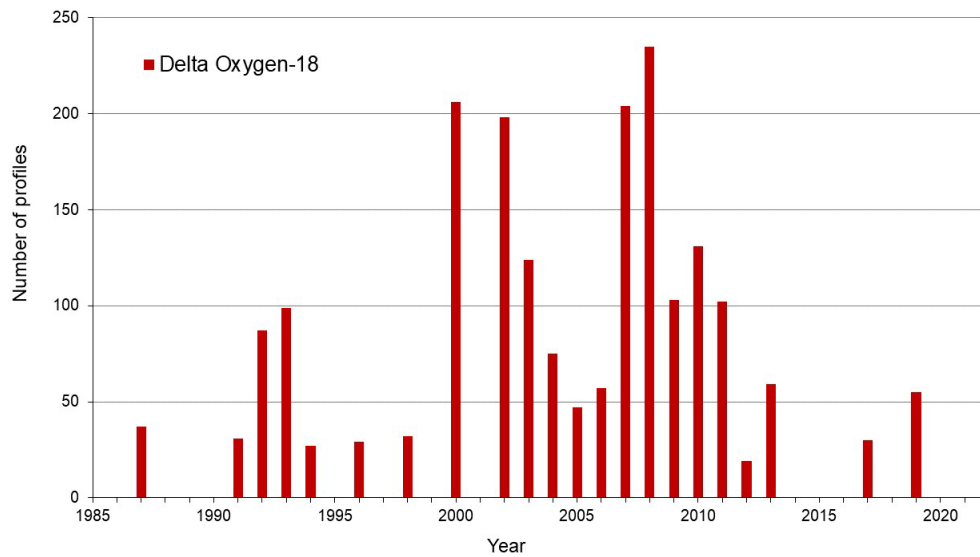
**Figure 2.25. Time series of the number of chlorofluorocarbon-11 profiles in WOD23 OSD dataset**



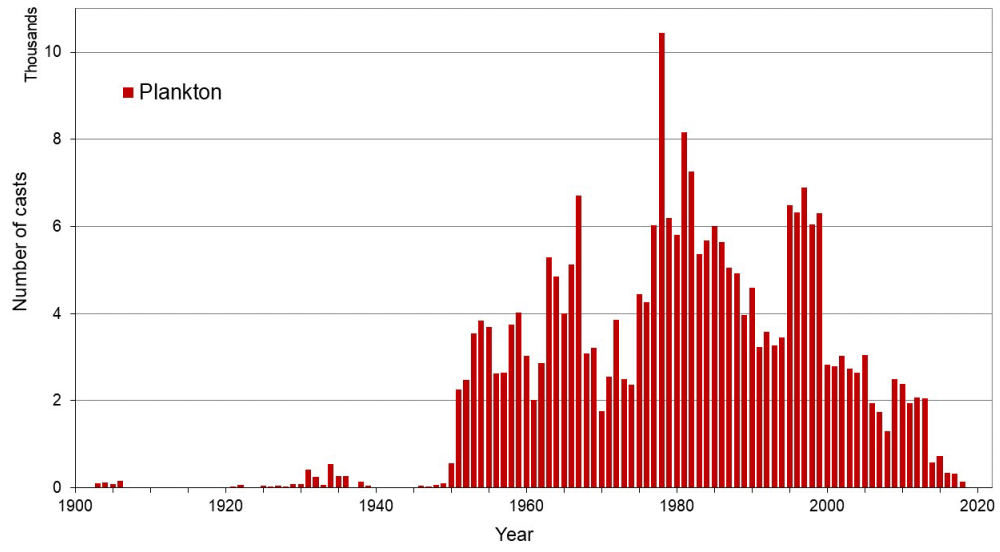
**Figure 2.26. Time series of the number of chlorofluorocarbon-12 profiles in WOD23 OSD dataset**



**Figure 2.27. Time series of the number of chlorofluorocarbon-13 profiles in WOD23 OSD dataset**



**Figure 2.28. Time series of the number of  $\Delta$ -oxygen-18 profiles in WOD23 OSD dataset**



**Figure 2.29. Time series of the number of plankton tows in WOD23 OSD dataset**

## 2.7. REFERENCES AND BIBLIOGRAPHY

- Aoyama, M., M. Joyce, T. Kawano, and Y. Takatsuki (2002). Standard seawater comparison up to P129. *Deep-Sea Res. I*, 49(6), 1103-1114.
- Becker *et al.* (2020). GO-SHIP Repeat Hydrography Nutrient Manual: The Precise and Accurate Determination of Dissolved Inorganic Nutrients in Seawater, Using Continuous Flow Analysis Methods. *Front. Mar. Sci.*, <https://doi.org/10.3389/fmars.2020.581790>
- Bittig *et al.* (2018). Oxygen Optode Sensors: Principle, Characterization, Calibration, and Application in the Ocean. *Front. Mar. Sci.* <https://doi.org/10.3389/fmars.2017.00429>
- Boyer, T.P., O.K. Baranova, C. Coleman, H.E. Garcia, A. Grodsky, R.A. Locarnini, A.V. Mishonov, C.R. Paver, J.R. Reagan, D. Seidov, I.V. Smolyar, K. Weathers, M.M. Zweng (2018): World Ocean Database 2018. A.V. Mishonov, Tech. Ed., *NOAA Atlas NESDIS 87* [PDF]
- Boyer *et al.*, (2023). Effects of the Pandemic on Observing the Global Ocean. *BAMS* <https://doi.org/10.1175/BAMS-D-21-0210.1>.
- Broecker, W.S. and T.H. Peng (1982). *Tracers in the Sea*. Lamont Doherty Geological Observatory, Columbia University, Palisades, New York, 690 pp.
- Bullister, J.L. and R.F. Weiss (1988). Determination of CCl<sub>3</sub>F and CCl<sub>2</sub>F<sub>2</sub> in seawater and air. *Deep-Sea Res.*, 35(5), 839-853.
- Carpenter, J.H (1965). The Chesapeake Bay Institute technique for the Winkler dissolved oxygen titration. *Limn. and Oceanogr.*, 10, 141-143.
- Culkin, F. and J. Smed (1979). The history of standard seawater. *Oceanology Acta*, 2: 355–364.
- CRC (1993). *CRC Handbook of Chemistry and Physics*, D.R. Lide (Ed.), 73<sup>rd</sup> edition (1992-1993), CRC press.
- DOE (U.S. Department of Energy) (1994). *Handbook of methods for the analysis of the various parameters of the carbon dioxide system in sea water*. Department of Energy (DOE), Version 2, A.G. Dickson & C. Goyet (eds.), ORNL/CDIAC-74.
- Dickson, A.G (1994). *Determination of dissolved oxygen in sea water by Winkler titration*. WOCE Hydrographic Program, Operations and Methods Manual, Woods Hole, Mass., U.S.A., Unpublished manuscript.
- Farrington, J.W. (2000). Achievements in Chemical Oceanography. *In: 50 Years of Ocean Discovery, National Science Foundation (1950-2000)*, Ocean Studies Board, National Research Council, National Academy Press, Wash., D.C.
- Forster G. (1777). A Voyage Round The World in His Britannic Majesty's Sloop, Revolution, commanded by Capt. James Cook, during the Years 1772, 3, 4, and 5. (George Forters, F.R.S.) Vol. 1. London, Printed for B. White, Fleet-Str.; J. Robson, Bond Str.; P. Elmsly, Strand; and G. Robinson, Pater-noster-Row.
- Garcia, H.E., T.P. Boyer, R.A. Locarnini, O.K. Baranova, M.M. Zweng (2019). World Ocean Database 2018: User's Manual (prerelease). A.V. Mishonov, Tech. Ed., NOAA, Silver Spring, MD. [PDF]
- Garcia, H.E., T.P. Boyer, R.A. Locarnini, J.R. Reagan, A.V. Mishonov, O.K. Baranova, C.R. Paver, Z. Wang, C. Bouchard, S. Scott, D. Seidov, and D. Dukhovskoy (2024). World Ocean Database 2023: User's Manual. A.V. Mishonov, Tech. Ed., *NOAA Atlas NESDIS 98*, pp 127. <https://doi.org/10.25923/j8gq ee82>

- Gordon, L.I., J.C. Jennings, A.A. Ross, and J.M. Krest (1993). A suggested protocol for continuous flow automated analysis of seawater nutrients (phosphate, nitrate, nitrite, and silicic acid) in the WOCE hydrographic program and the Joint Global Ocean Fluxes Study, WOCE Hydrographic Program Office, Operations manual 91-1, WOCE report 68/91
- Haine, T.W.N., A.J. Watson, and M.I. Liddicoat (1995). Chlorofluorocarbon 113 in the northeast Atlantic. *J. Geophys. Res.*, 100, 10745-10753.
- Jenkins, W.J. (1982). Oxygen utilization rates in the North Atlantic subtropical gyre and primary production in oligotrophic systems. *Nature*, 300, 246-248.
- Jenkins, W.J. (1987).  $^3\text{H}$  and  $^3\text{He}$  in the Beta Triangle: Observations of gyre ventilation and oxygen utilization rates. *J. Phys. Oceanogr.*, 17, 763-783.
- Jenkins, W.J. and P.B. Rhines (1980). Tritium in the deep North Atlantic Ocean. *Nature*, 286, 877-880.
- Kawano, T., M. Aoyama, and Y. Tasatsuki (2005). Inconsistency in the conductivity of standard potassium chloride solutions made from different high-quality reagents. *Deep-Sea Res. I*, 52, 389-396.
- Langdon, C. (2010) Determination of Dissolved Oxygen in Seawater by Winkler Titration using Amperometric Technique, in: The GO-SHIP Repeat Hydrography Manual: A Collection of Expert Reports and Guidelines. Version 1, edited by: Hood, E. M., Sabine, C. L., and Sloyan, B. M., *IOCCP Report no. 14, ICPO Publication Series no. 134*, 18 pp., [[PDF, https://www.go-ship.org/HydroMan.html](https://www.go-ship.org/HydroMan.html)]
- Levitus, S., R. Gelfeld, T. Boyer, and D. Johnson (1994). Results of the NCEI Oceanographic Data and Archaeology and Rescue Project, 73 pp., Key to Oceanographic Records Documentation 19. U.S. Government Printing Office, Washington, D.C.
- Levitus, S. (2012). The UNESCO IOC IODE “Global Oceanographic Data Archeology and Rescue” (GODAR) and “World Ocean Database” projects. *Data Science Journal*, Vol. 11. <https://datascience.codata.org/articles/10.2481/dsj.012-014>
- Lewis, E.L. and R.G. Perkins (1981). The practical salinity scale 1978: conversion of existing data. *Deep-Sea Res.*, 28, 307-328.
- Mantyla, A.W. (1980). Electric conductivity comparisons of standard seawater batches P29 to P84. *Deep-Sea Res.*, 27A: 837-846.
- Mantyla, A.W. (1987). Standard seawater comparison updated. *J. Phys. Oceanogr.*, 17: 543-548.
- Mantyla, A.W. (1994). The treatment of inconsistencies in Atlantic deep-water salinity data, *Deep-Sea Res. I*, 41, 1387-1405.
- Millero, F.S. and A. Poisson (1981). International one atmosphere equation of state of seawater. *Deep-Sea Res.*, 28, 625-629.
- Millero, F.J. (1993). What is PSU? *Oceanography*, 6(3), 67.
- Millero, F.J., R. Feistel, D.G. Wright, T.J. McDougall (2008). The composition of standard seawater and the definition of the reference-composition salinity scale. *Deep-Sea Res. I*, 55 (2008), pp. 50-72
- Millero, F.J. (2010). History of the equation of state of seawater. *Oceanography* 23(3):18–33, doi:10.5670/oceanog.2010.21
- Owens, W.B., Millard, R.C.J. (1985). A new algorithm for CTD oxygen calibrations, *J. Phys. Oceanogr.*, 15, 621-631.

- Östlund, H.G. and C.G.H. Rooth (1990). The North Atlantic tritium and radiocarbon transients 1972-1983. *J. Geophys. Res.*, 95: 20147-20165.
- Reagan, J.R.; T.P. Boyer, H.E. García, R.A. Locarnini, O.K. Baranova, C.N. Bouchard, S.L. Cross, A.V. Mishonov, C.R. Paver, D. Seidov, Z. Wang, D. Dukhovskoy. (2024). World Ocean Atlas 2023. NOAA National Centers for Environmental Information. *Dataset*: [NCEI Accession 0270533](#)
- Reiniger, R.F. and C.K. Ross (1968). A method for interpolation with application to oceanographic data, *Deep Sea Research*, 15: 185-193
- Redfield, A., B. Ketchum, and F. Richards (1963). The influence of organisms on the composition of sea water. Hill, N., editor, in: *The Sea*, vol. 2, p. 224-228, Inter-science, New York.
- Schlosser, P. (1986). Helium: A new tracer in Antarctic oceanography. *Nature*, 321, 233-235.
- Schlosser, P., G. Bönnisch, M. Ehein, and R. Bayer (1991). Reduction of deep-water formation in the Greenland Sea during the 1980's: Evidence from tracer data. *Science*, 251, 1054-1056.
- Smethie, W.M. (1993). Tracing the thermohaline circulation in the western North Atlantic using chlorofluorocarbons. *Progress in Oceanography*, 31, 51-99.
- Strickland, J.D.H. and T.R. Parsons (1972). *A Practical Handbook of Seawater Analysis*. Fisheries Research Board of Canada, Bulletin 169, 2<sup>nd</sup> edition, Ottawa, Canada.
- Uchida H., G.C. Johnson, and K.E. McTaggart (2010). GO-SHIP CTD Oxygen Sensor Calibration Procedures. in: *The GO-SHIP Repeat Hydrography Manual: A Collection of Expert Reports and Guidelines*. Version 1, edited by: Hood, E. M., Sabine, C. L., and Sloyan, B. M., IOCCP Report no. 14, ICPO Publication Series no. 134 [[PDF](#), <https://www.go-ship.org/HydroMan.html>]
- UNESCO (United Nations Educational, Scientific and Cultural Organization) (1985). *The International System of Units (SI) in Oceanography*. Report of IAPSO working group on symbols, units and nomenclature in physical oceanography (SUN), IAPSO Publication Scientifique, No. 32, UNESCO *Tech. Pap. in Marine Sci.*, No. 45.
- Worthington, L.V. (1982). The loss of dissolved oxygen in Nansen bottle samples from deep Atlantic Ocean. *Deep-Sea Res.*, 29(10A), 1259-1266.
- Winkler, L.W. (1888). *Die Bestimmung des in Wasser gelösten Sauerstoffes*. Berichte der Deutschen Chemischen Gesellschaft, 21: 2843-2855.

# CHAPTER 3: CONDUCTIVITY-TEMPERATURE-DEPTH DATA (CTD)

*Christopher R. Paver, Alexey V. Mishonov, Tim P. Boyer, Olga K. Baranova,  
Hernán E. García, Ricardo A. Locarnini, James R. Reagan, Dan Seidov*

*Ocean Climate Laboratory  
National Center for Environmental Information  
Silver Spring, MD*

## 3.1. INTRODUCTION

The Conductivity-Temperature-Depth (CTD) profiling instrument measures temperature, salinity, and pressure - among other variables - with high vertical resolution up to depths of 10,000 m. In practice, most CTD casts sample to considerably shallower depths.

Fundamental physical relationships between temperature (salinity, *etc.*) and electromagnetic properties of sea water are used to develop CTD sensors and appropriate conversion algorithms (Wallace, 1974; Prien, 2001). The sampling rate of CTD sensors – up to 24 Hz for the SBE 911plus (Sea-Bird Scientific; SBE 911plus CTD) - is an important factor that determines the ability of the CTD to make “continuous” measurements. For instance, lowering a CTD at speeds of  $1 \text{ m}\cdot\text{s}^{-1}$  with a typical range of response times for the temperature sensors can provide vertical profiling at resolutions of 0.05 m to 0.3 m.

CTD data submitted to NCEI for archive and inclusion into WOD are stored in their original vertical resolution. In instances where there are more than 6,000 depth-variable measurements in a single profile, those values are generally binned to a lower resolution in WOD. In the past, electronic storage limitations resulted in only selected levels being archived. While processing the original data for inclusion into WOD, reported pressure values (if any) take preference over reported depth values for reporting the depth-measurement values. The pressure values are converted to depths using Saunders and Fofonoff equation (1976).

An earlier version of the CTD instrument was the STD (salinity-temperature-depth) which computed salinity from a conductivity sensor as the instrument was moving vertically through the water column. Because of instrument problems that led to erroneous data values (spikes), this method was replaced by the CTD method for which conductivity measurements are recorded from the instrument and then salinity computed from the conductivity measurement with appropriate calibration information.

Newer sensors have been developed to make continuous measurements of other variables such as dissolved oxygen content, beam attenuation coefficient (BAC), and chlorophyll concentration. All sensor and derived variables collected via the CTD are stored in the NCEI archive, however, only select variables are maintained in WOD. These variables are identified in Table 3.1. It should be noted that amount of the depth profiles shown in Table 3.1 is different from



temperature or salinity profile counts because there are temperature-only or salinity-only profiles exists.

**Table 3.1. List of all variables and profile counts in the WOD23 CTD dataset.**

<b>Variables</b>	<b>Profiles</b>
Depth (reported and calculated)	1,132,680
Pressure	656,483
Temperature	1,131,340
Conductivity	70,152
Salinity	1,099,528
Oxygen	214,674
Chlorophyll (all types: -a, -b, -c)	93,280
Transmissivity (Beam attenuation)	35,553

### **3.2. CTD ACCURACY**

The cited accuracy of CTD measurements represents the results of calibration of CTD sensors by comparison with established standards. This initial accuracy varies with instrument design typically from 0.001°C to 0.005°C for temperature, 0.0003 S·m<sup>-1</sup> to 0.002 S·m<sup>-1</sup> for conductivity (approximately 0.003 to 0.02 equivalent salinity), and 0.015% to 0.08% for pressure. These accuracies are subject to change after prolonged use of the CTD (known as a calibration drift).

The overall quality of CTD measurements does not depend solely on the accuracy of CTD sensors. Other factors such as the difference in response time of temperature and conductivity sensors, varying speeds of the CTD, along with rapid changes in ocean environment can be important sources of erroneous CTD data (Lawson and Larson, 2001).

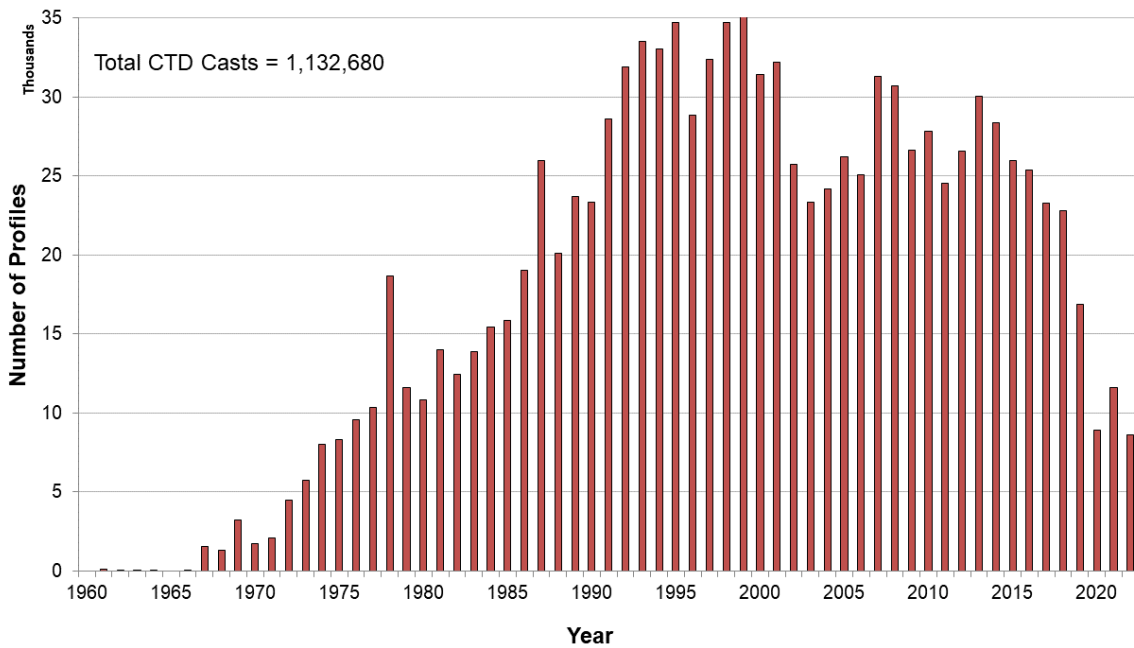
### **3.3. CTD CAST DISTRIBUTIONS**

In WOD23 there are a total of 1,132,680 CTD casts collected over the entire World Ocean, which represents 6.1% of all WOD data. The earliest reported CTD casts available in WOD are from 1961, however increased reporting didn't occur until 1967 during the International Cooperative Effort Toward Understanding of the Oceanography of the Eastern Tropical Pacific (EASTROPAC; Vilchis and Ballance, 2005), thereafter the number of reported casts increased annually (Table 3.2, Figure 3.1). Then in 1978, a jump in reported casts occurred as a result of sampling contributed by the Institute for Marine Research at Kiel University mainly aboard the *R/V Meteor*. Almost a decade later in 1987, reported casts increased to over 20,000 in part to relatively greater contributions from Canada, Great Britain, the Soviet Union, and the United States. Peak reported CTD observations occurred in 1999 at over 35,000 casts. From about 2002 to 2016, total reported casts were averaging around 25,000 per year. The three nations contributing the most CTD casts are the United States, Canada, and Japan, reporting over 52% of the total in WOD (Table 3.3). The Atlantic Ocean is the most reportedly sampled basin in WOD, with roughly 62% of the total (Figure 3.2). The coastal regions of the northern hemisphere are the most densely sampled areas,

mainly around the continent of Europe, the western North Atlantic boundary with North America, and the western North Pacific boundary around central Asia. The least sampled areas for CTD continue to be the Arctic Ocean and the Southern Ocean, more so for the open waters of the South Pacific. Distribution of the high-resolution CTD observations for temperature and salinity at standard depth levels are shown in Figure 3.3.

**Table 3.2. The number of high-resolution CTD casts in WOD23 as a function of year.**

YEAR	CASTS	YEAR	CASTS	YEAR	CASTS	YEAR	CASTS
1961	97	1977	10,334	1993	33,490	2008	30,694
1962	42	1978	18,447	1994	33,037	2009	26,632
1963	71	1979	11,597	1995	34,688	2010	27,820
1964	47	1980	10,858	1996	28,843	2011	24,528
1965	-	1981	13,988	1997	32,277	2012	26,565
1966	12	1982	12,475	1998	34,725	2013	30,059
1967	1,531	1983	13,857	1999	35,812	2014	28,324
1968	1,286	1984	15,441	2000	31,438	2015	25,986
1969	3,229	1985	15,878	2001	32,133	2016	25,395
1970	1,740	1986	18,857	2002	25,645	2017	23,280
1971	2,074	1987	25,911	2003	23,320	2018	22,805
1972	4,451	1988	20,134	2004	24,156	2019	16,853
1973	5,739	1989	23,620	2005	26,194	2020	8,939
1974	7,996	1990	23,317	2006	25,075	2021	11,584
1975	8,257	1991	28,587	2007	31,265	2022	8,608
1976	9,448	1992	31,849				



**Figure 3.1. Temporal distribution of high-resolution CTD casts in WOD23.**

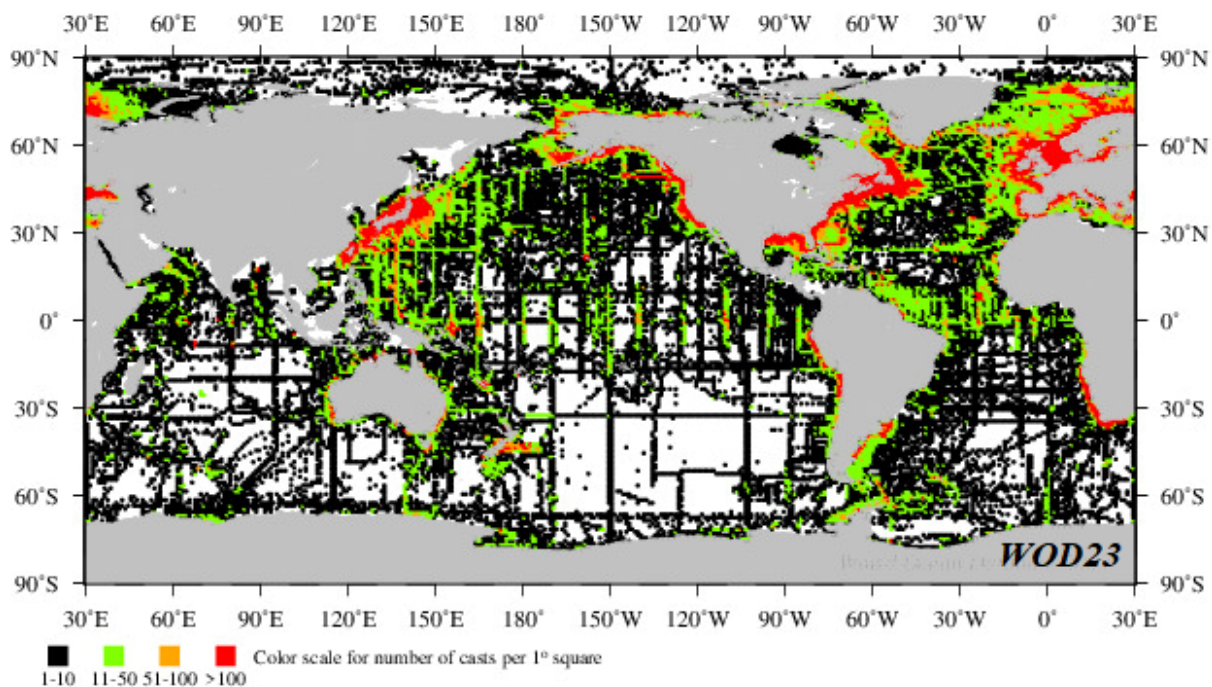


Figure 3.2. Geographic distribution of high-resolution CTD casts in WOD23.

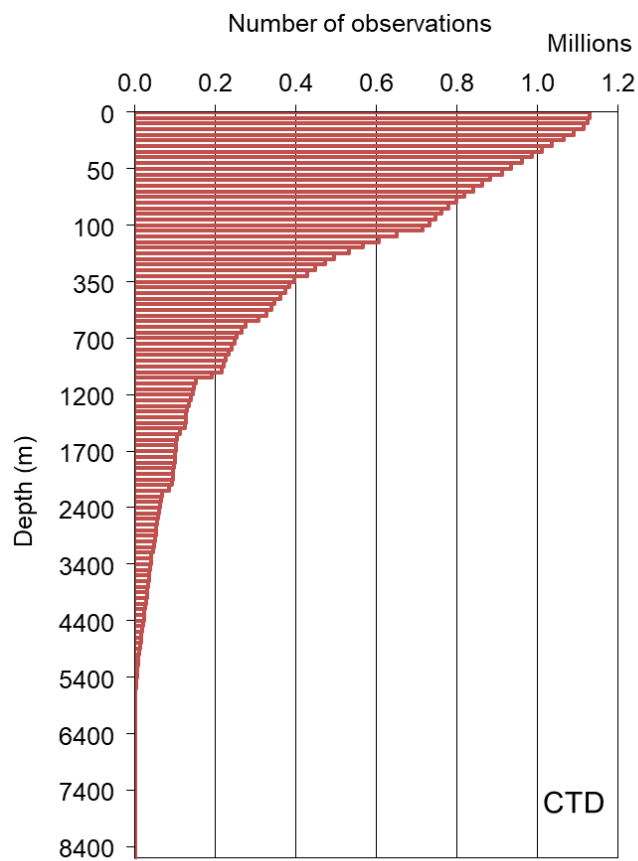


Figure 3.3. Distribution of high-resolution CTD casts at standard depth levels in WOD23.

**Table 3.3. National contributions of high-resolution CTD casts in WOD23.**

<b>ISO<sup>1</sup> 3166-1 Country Codes</b>	<b>Country Name</b>	<b>CTD Casts</b>	<b>% of Total</b>
US	United States	246,710	21.781%
CA	Canada	186,888	16.500%
JP	Japan	165,223	14.587%
NO	Norway	105,126	9.281%
DE	Germany	64,789	5.720%
GB	Great Britain	54,115	4.778%
FR	France	47,683	4.210%
99 <sup>2</sup>	Unknown	29,732	2.625%
SU	Soviet Union	29,514	2.606%
DK	Denmark	24,610	2.173%
AU	Australia	23,863	2.107%
TW	Taiwan	23,744	2.096%
IT	Italy	12,919	1.141%
ZA	South Africa	12,110	1.069%
IE	Ireland	9,734	0.859%
UA	Ukraine	9,187	0.811%
ES	Spain	8,532	0.753%
NZ	New Zealand	8,202	0.724%
AR	Argentina	7,917	0.699%
CL	Chile	6,106	0.539%
GR	Greece	5,914	0.522%
NL	Netherlands	5,489	0.485%
NA	Namibia	5,044	0.445%
IS	Iceland	4,989	0.440%
FI	Finland	3,875	0.342%
PL	Poland	3,862	0.341%
RU	Russian Federation	3,692	0.326%
CN	China	3,259	0.288%
PE	Peru	3,236	0.286%
PT	Portugal	2,843	0.251%
TR	Turkey	2,705	0.239%
EE	Estonia	2,374	0.210%
DU	East Germany	1,692	0.149%
BR	Brazil	1,256	0.111%
ZZ <sup>2</sup>	Miscellaneous organization	1,174	0.104%
IL	Israel	914	0.081%
IN	India	752	0.066%
SE	Sweden	501	0.044%
BE	Belgium	440	0.039%
VE	Venezuela	389	0.034%
PA	Panama	337	0.030%
CY	Cyprus	235	0.021%
EC	Ecuador	217	0.019%
ID	Indonesia	213	0.019%
LT	Lithuania	149	0.013%
BG	Bulgaria	90	0.008%
MX	Mexico	82	0.007%
TN	Tunisia	73	0.006%
EG	Egypt	69	0.006%
LB	Lebanon	42	0.004%
KR	Korea; Republic of	28	0.002%
RO	Romania	27	0.002%
DZ	Algeria	13	0.001%
HR	Croatia	1	0.000%

<sup>1</sup> ISO = International Organization for Standardization.

<sup>2</sup> Codes '99' and 'ZZ' are not official ISO 3166-1 codes. These are supplemental codes for WOD.

### **3.5. REFERENCES AND BIBLIOGRAPHY**

- Lawson, K. and N.G. Larson (2001). CTD, pp. 579-588, doi:10.1006/rwos.2001.0324 in *Encyclopedia of Ocean Sciences* (Eds. J. H. Steele, K. K. Turekian, S. A. Thorpe), Academic Press.
- Mantyla, A. (1987). Standard Seawater comparisons updated. *J. Phys. Oceanogr.*, 17, 543-548.
- Millero, F.J. (1993). What is PSU? *Oceanogr.*, 6(3), 67.
- NOIC (1970). *Calibration procedure for deep sea reversing thermometers*. National Oceanographic Instrumentation Center; Rockville, MD.
- NOIC (1970). *Calibration procedure for STD*. National Oceanographic Instrumentation Center; Rockville, MD
- Park, K. (1964). Reliability of Standard Sea water as a conductivity standard. *Deep-Sea Res*, 11, 85-87.
- Prien, R.D. (2001). Electrical properties of sea water, 832-839, doi:10.1006/rwos. 2001.0328, in *Encyclopedia of Ocean Sciences* (Eds. J. H. Steele, K. K. Turekian, S. A. Thorpe), Academic Press.
- Saunders, P. M. and N. P. Fofonoff (1976). Conversion of pressure to depth in the ocean. *Deep Sea Res.*, 23, 109-111.
- Vilchis, L. and L.T. Balance (2005). A complete listing of expeditions and data collected for the EASROPAC cruises in the Eastern Tropical Pacific, 1967-1968. *NOAA Technical Memorandum NOAA-TM-NMFS-SWFSC-374*. La Jolla, CA.
- UNESCO (1981). Background papers and supporting data on the Practical Salinity Scale. *UNESCO Technical Series, Marine Science*, 37, Paris, 144 pp.
- UNESCO (1987). International Oceanographic Tables. Paris, *Technical Rapport. Marine Sciences*, 195 pp.
- Wallace, W.J. (1974). The Development of the Chlorinity / Salinity Concept in Oceanography, *Elsevier*, New York.
- Wooster, W.S. and B.A. Taft (1958). On the reliability of field measurements of temperature and salinity. *J. Mar. Res.*, 17, 552-566.

# CHAPTER 4: EXPENDABLE BATHY THERMOGRAPH DATA (XBT)

*Ricardo A. Locarnini, Tim P. Boyer, Courtney N. Bouchard, Olga K. Baranova,  
Hernán E. García, Alexey V. Mishonov, Christopher R. Paver, James R. Reagan,*

*Ocean Climate Laboratory  
National Centers for Environmental Information / NOAA  
Silver Spring, MD*

## 4.1. INTRODUCTION

The Expendable Bathythermograph (XBT) was deployed beginning in 1966 and replaced the Mechanical Bathythermograph (MBT) in most measurement programs. The XBT allows the measurement of the upper ocean's temperature profile when launched from underway surface ships, submarines, and aircraft. The system consists of three main components: an expendable measuring probe, a launcher, and an electronic data acquisition unit. The expendable probe includes a thermistor and a spool of copper wire that unwinds as the probe falls through the water column. The launcher holds a second copper wire spool that unwinds as the platform continues its underway trajectory. The launcher and its cable are also a part of the bridge circuit for temperature measurement. The resistance of the wires is also measured though they are thought to be canceled by assuming their equal (or balanced) resistance (Shoichi Kizu, personal communication). Finally, the temperature signal is sent from the launcher through a cable to the data acquisition system, where the data are recorded. While current acquisition systems store the data in digital form, initially the data were recorded in paper strip charts.

The system has different details when the expendable probes are launched from a submarine or from an aircraft. From a submarine, a float carries the expendable probe to the sea surface. Upon reaching the sea surface, the probe detaches from the float and falls through the water column. From an aircraft, the expendable probe and a floating surface unit are deployed with a parachute. After reaching the sea surface, the probe detaches from the floating unit and falls through the water column. The temperature information from the thermistor is transmitted through the copper wire to the floating surface unit, which transmits the data to the acquisition system in the aircraft via a radio signal.

Of all the XBT profiles in *World Ocean Database 2023* (WOD23), 49.3% are known to have been obtained with probes manufactured by Lockheed Martin Sippican (formerly known as Sippican), 2.3% to have been obtained with probes manufactured by Tsurumi Seiki Co. LTD (TSK), and 0.3% to have been obtained with probes manufactured by Sparton or Hermes. There is no manufacturer information for the probe used for just less than half, 48.1%, of the XBT profiles. Each manufacturer has several models of XBT probes which have different maximum sampling depths with the associated launching platform moving at or below the allowed maximum speed. As an example, Table 4.1 below shows the characteristics for some expendable probes

produced by Lockheed Martin Sippican, their WOD variable-dependent instrument code, amount and percentage of the profiles collected by those probes in WOD23.

**Table 4.1. Characteristics of expendable probe models produced by Lockheed Martin Sippican.**

<b>Model</b> <i>(WOD var. dependable instrument code, 5)</i>	<b>Max. Depth</b>	<b>Rated Ship Speed</b>	<b># of profiles in WOD23 XBT subset</b>	<b>% of data # in WOD23 XBT subset</b>
T-4 (208)	460 m	30 kts	453,175	19.20%
Deep Blue™ (214)	760 m	20 kts	423,042	17.92%
T-7 (207)	760 m	15 kts	213,750	9.06%
T-5 (210)	1830 m	6 kts	17,069	0.72%
T-6 (209)	460 m	15 kts	8,083	0.34%
Fast Deep™ (213)	1000 m	20 kts	5,633	0.24%
T-10 (211)	200 m	10 kts	43,273	1.83%
T-11 (212)	460 m	6 kts	525	0.02%
<i>Total for Sippican probes</i>				<i>1,164,550 / 49.3%</i>
<i>Total XBT profiles in WOD23</i>				<i>2,360,444 / 100%</i>

Corresponding models from different manufacturers have comparable characteristics but with some differences as explained below. Sippican's XBT and its correspondent model made by TSK are very different in structure and show systematically different fall-rates, occasionally reaching 3-4% relative difference, according to the side-by-side comparisons. In addition, TSK's T-10 is rated to 300m depth range but Sippican's is to 200m. According to the measurement, a TSK T-10 probe is heavier than a Sippican's T10 probe by more than 5% in the air (Shoichi Kizu, personal communication).

There is probe model information for about 55% of the XBT profiles in WOD23. Overall, the most popular probe model is the T-4, with 19.2% of the XBT profiles in WOD23 known to be obtained with such a probe, while Deep Blue and T-7 probes are known to account for 17.9% and 9.1%, respectively, of the XBT profiles in WOD23 (see Table 4.1). It should be noted, however, that while more than 47,000 XBT profiles obtained with Deep Blue probes have been added to WOD since the release of WOD18, only about 600 new XBT profiles in WOD23 were obtained with T-4 probes. The majority of modern XBT profiles recently added to WOD do have manufacturer and model reported.

The XBT system does not directly measure depth. The depth of each temperature measurement obtained by the expendable probe is estimated using a time-depth equation. This equation converts the time elapsed from the moment the probe enters the water, in seconds, to depth, in meters.

## **4.2. XBT ACCURACY**

Lockheed Martin Sippican reports temperature accuracy of  $\pm 0.10^{\circ}\text{C}$  for their surface ship expendable probes and  $\pm 0.15^{\circ}\text{C}$  for their submarine expendable probes, with a depth accuracy of  $\pm 2\%$  for all probes. Tsurumi Seiki Co. LTD reports temperature accuracy of  $\pm 0.1^{\circ}\text{C}$  and depth accuracy of  $\pm 2\%$  or 5 m, whichever is larger. The accuracy of data from paper strip charts is  $\pm 0.15^{\circ}\text{C}$ .

### 4.3. XBT TIME-DEPTH EQUATION ERROR

Since the XBT system does not measure depth directly, the accuracy of the depth associated with each temperature measurement is dependent on the equation which converts the time elapsed since the probe entered the water to depth. Unfortunately, problems have been found in various time-depth equations (fall-rate equations, FRE) used since the introduction of the XBT system.

The original FRE developed by Sippican for their T-4, T-6, T-7, and Deep Blue models underestimates the probes' fall rate. At a given elapsed time, the falling probe is actually deeper than indicated by the original equation. Thus, the water temperatures are associated by the original FRE with depths that are shallower than the actual depths at which they are measured. The error, first documented by Flierl and Robinson (1977), increases with increasing elapsed time reaching 21 meters, or about a 2.5% error, for depths around 800 meters. Sippican's original FRE was used by TSK for their T-4, T-6, T-7, and Deep Blue models, and by Sparton for their XBT-4, XBT-6, XBT-7, XBT-7DB, XBT-20, and XBT-20DB models.

In 1994, Hanawa *et al.* published an International Oceanographic Commission (IOC, 1994) report detailing a study of XBT fall rates using different probes manufactured by Sippican and TSK and dropped in different geographic locations. A new time-depth equation, the Hanawa *et al.* (1995) FRE, was given, as well as an algorithm for correcting depths for existing data collected using the original FRE. The report emphasized the need to continue to archive existing data with the original FRE only, applying the correction when necessary for scientific research.

Rual *et al.* (1995) and Rual *et al.* (1996) studied Sparton XBT-7 probes. It was determined that the Hanawa *et al.* (1995) FRE was suitable for use with these probes.

Thadathil *et al.* (2002), however, suggest that the Hanawa *et al.* (1995) FRE is not valid for measurements in high-latitude low temperature waters.

Following the report of Hanawa *et al.* (1995) and IOC (1994), TSK altered their software between January and March 1996 to make the Hanawa *et al.* (1995) equation the default equation (Greg Ferguson, personal communication). Sippican did the same around August 1996, (James Hannon, personal communication). However, use of the default is not universal. Depths from XBT drops are recorded using both the original and Hanawa *et al.* (1995) FREs.

Kizu *et al.* (2005) published a new time-depth equation for the TSK T-5 probes, but no manufacturer software has been released with their equation.

Corrections to the time-depth equations for air-dropped XBT probes (AXBT) manufactured by Sippican and Sparton were calculated by Boyd (1987) and Boyd and Linzell (1993b) respectively.

In spite of the fact that the XBT depth bias problem has been known since Flierl and Robinson (1977) and the new FREs was suggested later by Hanawa *et al.* (1994) the impact of the bias on the estimation of the ocean heat content was not realized until the study by Gouretski and Koltermann (2007). They found that the XBT temperature bias is time-dependent and showed that the prevalence of XBT temperature profiles over profiles from other instruments lead to the warm anomaly of the global temperature time series centered around 1980 and first reported by Levitus *et al.* (2005). The Gouretski and Koltermann (2007) study triggered a series of papers dedicated to



the estimation of XBT biases and to the development of the respective bias correction schemes. The idea of time-varying FRE was implemented by Wijffels *et al.* (2008). Wijffels *et al.* (2008), Ishii and Kimoto (2009), and Levitus *et al.* (2009) also developed corrections. In particular, Levitus *et al.* (2009) compared their own corrections with the corrections of Wijffels *et al.* (2008) and Ishii and Kimoto (2009). In their comprehensive study, Gouretski and Reseghetti (2010) demonstrated that the total XBT temperature bias can be decomposed into two components: translated temperature bias due to the systematic error in depth and pure temperature bias which is not related to the FRE used. Both biases were shown to be time-varying.

New time-dependent temperature bias adjustments, as well as time-depth equation corrections, are actively studied and calculated as new analysis and more temperature data from XBT, MBT, and CTD instruments become available (Good, 2011; Hamon *et al.*, 2012; Gourestki, 2012; Cowley *et al.*, 2013; Cheng *et al.*, 2014). In a summary of instrumental biases and errors, Cheng *et al.* (2016) report that XBT data need four different related corrections: time-dependent depth correction, time-dependent temperature bias, temperature-dependent depth and temperature corrections, and surface offset. More details can be found on [WOD XBT correction](#) page.

#### **4.4. CORRECTIONS TO XBT TIME-DEPTH EQUATION ERRORS**

Before the various time-depth equations errors were widely known, a significant amount of data was recorded and archived without notation of what model of expendable probe was used. About 45%, or 1.06 million, of the total 2.36 million XBT temperature profiles in WOD23 have “unknown” model of XBT instrument and discovering this information is a challenge, which we have limited resources to address. Of these, about 0.76 million are positively identified as coming from shipboard drops, about 0.26 million have no associated platform information, and the rest are known to have been dropped from aircrafts or fixed stations. The missing ancillary metadata make it difficult to know whether the reported depths for a particular XBT profile were obtained with an incorrect time-depth equation.

Presently, some XBT data are still recorded and archived with no indication of the FRE used. This is particularly critical since there is more than one FRE in use for many XBT models.

The XBT data in the WOD23 on observed levels report the same data as submitted to NCEI/WDC by the originators. Secondary header 33 indicates reported information on the time-depth equation used by the originator – see [WOD23 Manual](#) (Garcia *et al.*, 2024) for more information on WOD23 format and code descriptions. Secondary header 33 is set to 0 if the original FRE was used, and it is set to 1 if the Hanawa *et al.* (1995) or another amended FRE was used. Secondary header 33 is absent if the FRE used is unreported. Data taken before the introduction of corrected FRE (January 1996) usually have unknown time-depth equation, and it is assumed the original FRE was used unless otherwise noted. Indeed, only 3,330 pre-1996 XBT drops include depths that are known to have been corrected by the originator before being submitted to NCEI/WDC.

The XBT data in the WOD23 interpolated to standard levels uses the appropriate corrected depth when possible using the corrections of Levitus *et al.* (2009) updated through 2022. Since close to half of all XBT profiles are of unknown model, a test was applied to these data to see if a depth correction was necessary. It was assumed that, following the IOC recommendation, data available in the WOD23 were received at NCEI with depths calculated using the original equations

unless otherwise noted. This assumption is not always valid for data collected since new time-depth equations became available on recording software released by each XBT manufacturer. For data collected since January 1996, if the time-depth equation used was not noted, the data were not corrected when interpolating to standard levels and were marked so as not to be used for depth sensitive calculations. Of a total of 709,794 XBT drops during the relevant time period (1996-2022), there are 99,880 drops without time-depth equation information, with the vast majority of them belonging to the period 1996-2000: only 67% of XBT drops from 1996 to 2000 include the information on the time-depth equation used, in contrast to 97% of XBT drops from 2001 to 2022 including that information.

An attempt to ascertain the missing FRE information was made by contacting the data originators. Most of the data originators are large data centers and the information could not be recovered. The actual values of the reported depths can be used to recognize the FRE used, when the full depth trace is reported (Donald Scott, personal communication). Although most data received at NCEI comes with only selected depth levels, when possible, this technique was used.

Secondary header 54 contains information on our decision on whether the depths need correction (i.e. adjusted in order to make data from different XBTs comparable) for each XBT given the criteria listed above. It is clear from Gouretski and Koltermann (2007) that the XBT correction detailed in Hanawa *et al.* (1995) is not the full answer to the drop rate problem, but it is a reference. In order to properly use XBT data collected by different probes, it should all at least have the same meaning for depth. As noted above, the manufacturer provides two definitions of depth in their software since 1996, original FRE and Hanawa FRE. Either all data need to be set to one or the other depth continuum to use the data together. Hanawa *et al.* (1994), the community white paper, details the community agreement to convert original FRE to Hanawa FRE when using (but storing with original FRE). Second header 54 provides this guidance for observed level data. For standard level data, second header 54 gives the actual XBT correction used as per user request in [WODselect](#), or in the default, Levitus *et al.* (2009) correction. If the data user knows that the data from a particular XBT profile most probably correspond, for instance, to the original Sippican FRE, then the data user can introduce and implement the depth correction using the scheme he/she prefers or leave the data as it is.

**Important: the observed level XBT data in wod23 are the same data as submitted by the originators. If you are using observed level XBT data from wod23, please use secondary header 54 to see whether a depth correction is necessary.**

**The standard level XBT data in wod23 were prepared, when needed and possible, using a corrected time-depth equation and the XBT bias adjustment following Levitus *et al.* (2009).**

**XBT bias adjustments were recalculated using an updated data set for world ocean database 2013 and world ocean atlas 2013. Years 2008 to 2012 have their own temperature bias adjustments in the updated Levitus scheme. Values for year 2012 were used for 2013 to 2022.**

**If you are using standard level XBT data from wod23, please use secondary header 54 to see whether a corrected time-depth equation was used, a correction was not needed, or a correction could be needed but there was not enough information.**

The observed level XBT data in wod23 are also offered online through *WODselect* and, in addition to the Levitus scheme adjustment, can be prepared with more than ten different time-dependent temperature adjustments or depth corrections. Please use secondary header 54 to identify the correction used to prepare the downloaded data.

#### **4.5. SURFACE DATA ACQUIRED CONCURRENTLY WITH XBT CASTS**

On a surface ship sometimes, a sea-surface water sample is obtained at the time of the XBT launch. Temperature and salinity obtained from such water sample are usually recorded as ancillary information of the XBT launch. Meteorological conditions at the time of the XBT launch could also be recorded, e.g. air temperature, wind speed and direction, cloud type and cover, barometric atmospheric pressure, as well as sea conditions: wave height and direction, sea state. When available, these data are included in WOD23 as secondary header information for the corresponding XBT drop, from where this data could be ingested by surface-related programs like ICOADS and other.

#### **4.6. XBT PROFILE DISTRIBUTIONS**

There is a total of 2,360,444 XBT profiles for the entire World Ocean in WOD23, which represents 12.7% of all data in WOD23 (see Fig. 1.3), with 517,393 profiles (21.9%) measured in the southern hemisphere and 1,843,051 profiles (78.1%) measured in the northern hemisphere (Figure 4.1). Table 4.2 gives the yearly counts of XBT profiles in WOD23 for the World Ocean. Figure 4.2 shows the time series of the yearly totals of Expendable Bathythermograph profiles for the World Ocean. After its introduction, XBT yearly totals increase dramatically to about, or above, 50,000 for most years from 1970 to 1999, with peaks in the mid-1980's and early to mid-1990's. The decrease in yearly totals in recent years is partially due to the delay between the time of the observations and their reporting by the originators to the NCEI/WDC as well as reducing its use due to Argo floats implementation.

Figure 4.3 shows the distribution of XBT data at standard depth levels. There are significant decreases in the number of observations below the maximum depths sampled by the most popular probe models (see Table 4.1): 460 m (T-4) and 760 m (Deep Blue and T-7).

Although sixty-six known countries contribute XBT data to WOD23, 79% of the profiles are contributed by just seven major contributors: United States, Japan, Great Britain, Australia, Canada, Germany, and France (see Figure 4.4 and Table 4.3). Some country contributions merely reflect the flag of merchant ships in the Ship of Opportunity Program (SOOP), and they do not represent active national scientific programs, e.g. Liberia, Panama, Singapore, and Antigua. Table 4.3 gives detailed information about national contributions of XBT sorted by contribution from each country.

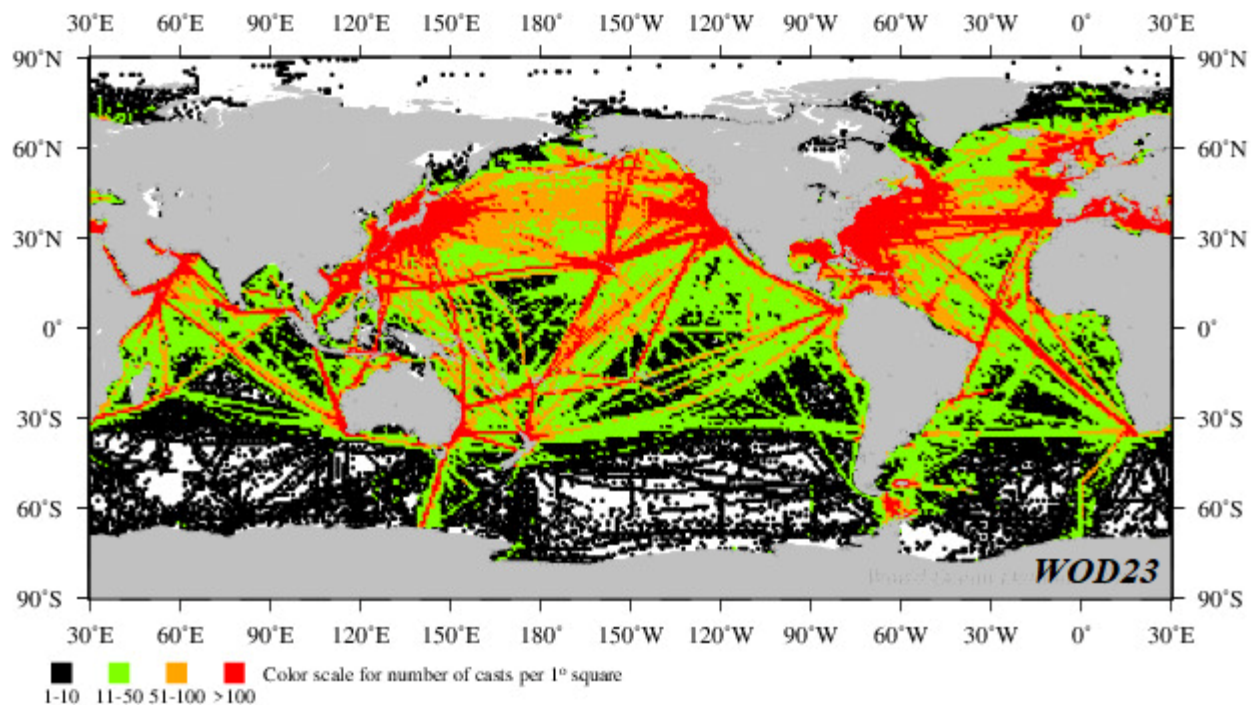


Figure 4.1. Geographic distribution of XBT profiles in WOD23: number of profiles by one-degree squares.

Table 4.2. The number of all XBT profiles as a function of year in WOD23.

Total Number of Profiles = 2,360,444

YEAR	PROFILES	YEAR	PROFILES	YEAR	PROFILES	YEAR	PROFILES
1966	1,750	1981	55,034	1995	78,691	2009	23,176
1967	9,390	1982	56,002	1996	63,728	2010	21,173
1968	26,684	1983	58,917	1997	53,095	2011	19,676
1969	34,321	1984	56,235	1998	50,044	2012	18,056
1970	45,701	1985	68,825	1999	56,034	2013	18,487
1971	57,628	1986	75,291	2000	39,854	2014	19,657
1972	53,173	1987	72,040	2001	31,029	2015	17,603
1973	54,949	1988	62,474	2002	27,809	2016	18,326
1974	54,888	1989	45,255	2003	27,402	2017	16,324
1975	54,430	1990	82,926	2004	32,481	2018	16,038
1976	48,480	1991	72,049	2005	29,618	2019	13,983
1977	54,408	1992	66,159	2006	26,485	2020	5,496
1978	53,387	1993	70,960	2007	23,197	2021	8,203
1979	56,122	1994	69,159	2008	23,746	2022	9,074

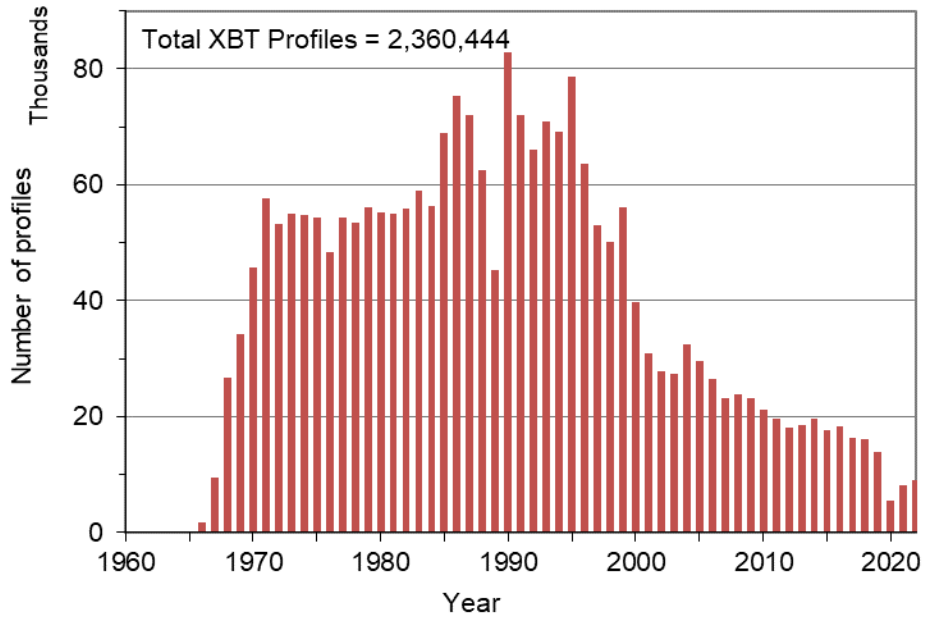


Figure 4.2. Temporal distribution of Expendable Bathythermograph (XBT) profiles in WOD23.

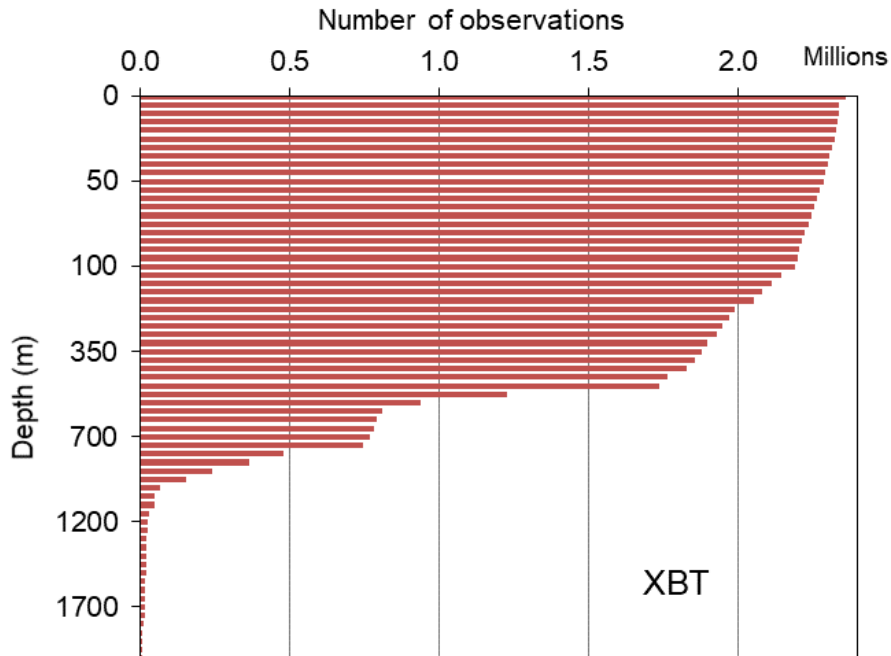
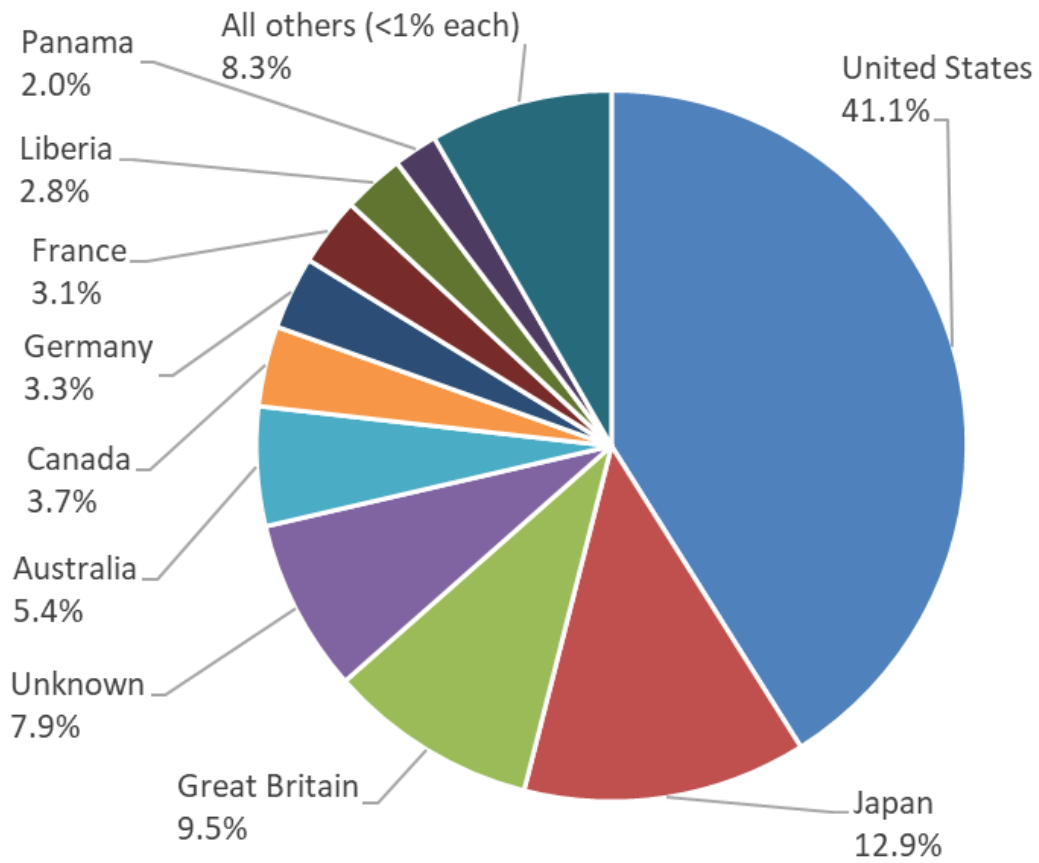


Figure 4.3. Distribution of Expendable Bathythermograph (XBT) data at standard depth levels in WOD23.



**Figure 4.4. XBT data contribution by countries in WOD23. Totals for Panama and Liberia include data obtained from merchant ships in the Ship of Opportunity Program (SOOP).**

**Table 4.3. National contribution of XBT profiles in WOD23.**

<b>ISO<sup>1</sup> Country Code</b>	<b>Country Name</b>	<b>XBT Casts</b>	<b>% of Total</b>
US	United States	969,145	40.74
JP	Japan	304,708	13.03
GB	Great Britain	224,966	9.75
99	Unknown	185,862	7.76
AU	Australia	127,948	5.15
CA	Canada	86,234	3.72
DE	Germany	78,212	3.37
FR	France	73,629	3.08
LR	Liberia	66,048	2.87
PA	Panama	47,391	2.06
SG	Singapore	19,002	0.83
NL	Netherlands	15,799	0.69
SU	Union of Soviet Socialist Republics	13,681	0.62
DK	Denmark	13,162	0.57
AG	Antigua	12,543	0.53
ZA	South Africa	12,131	0.53
BS	Bahamas	11,978	0.52
NO	Norway	8,448	0.37
NZ	New Zealand	6,300	0.27
VC	St. Vincent and Grenadines	6,094	0.26
CY	Cyprus	5,994	0.25
CN	China, The People's Republic of	5,719	0.24
BR	Brazil	5,610	0.23
IS	Iceland	4,574	0.2
SE	Sweden	4,551	0.2
BB	Barbados	4,376	0.19
HK	Hong Kong	4,308	0.19
IT	Italy	3,227	0.14
IN	India	3,083	0.13
ES	Spain	3,001	0.13
TO	Tonga	2,989	0.13
WS	Western Samoa	2,769	0.12
AR	Argentina	2,535	0.11
CL	Chile	2,434	0.11
PH	Philippines	2,302	0.1
MX	Mexico	2,238	0.1
TH	Thailand	1,901	0.08
KW	Kuwait	1,876	0.08
MT	Malta	1,452	0.06

<b>ISO<sup>1</sup> Country Code</b>	<b>Country Name</b>	<b>XBT Casts</b>	<b>% of Total</b>
PL	Poland	1,320	0.06
ID	Indonesia	1,241	0.05
TW	Taiwan	1,082	0.05
BE	Belgium	1,028	0.04
MH	Marshall Islands	936	0.04
FJ	Fiji	866	0.04
YU	Yugoslavia	797	0.03
PT	Portugal	732	0.03
PE	Peru	714	0.03
GR	Greece	658	0.03
RU	Russian Federation	545	0.02
EC	Ecuador	492	0.02
MY	Malaysia	460	0.02
TR	Turkey	307	0.01
SA	Saudi Arabia	197	<0.01
ZZ	Miscellaneous Organization	195	<0.01
UY	Uruguay	146	<0.01
HR	Croatia	82	<0.01
MU	Mauritius	77	<0.01
DU	East Germany	67	<0.01
MG	Madagascar	62	<0.01
KR	Korea, Republic of	53	<0.01
CI	Cote D'Ivoire	43	<0.01
UA	Ukraine	33	<0.01
CO	Colombia	32	<0.01
CR	Costa Rica	29	<0.01
HN	Honduras	13	<0.01
SC	Seychelles	11	<0.01
TT	Trinidad and Tobago	6	<0.01
	<b>Total:</b>	<b>2,360,444</b>	<b>100.00</b>

<sup>1</sup> ISO = International Organization for Standardization: [http://www.iso.org/iso/country\\_codes.htm](http://www.iso.org/iso/country_codes.htm)



## 4.7. REFERENCES AND BIBLIOGRAPHY

- Bailey, R.J., H.E. Phillips, and G. Meyers (1989). Relevance to TOGA of systematic XBT errors, in *Proceedings of the western Pacific International meeting and workshop on TOGA-COARE*, Eds. J. Picaut, R. Lukas, and T. Delcroix, pp. 775-784.
- Bailey, R.J. and A. Gronell (undated). *Scientific quality control at the WOCE Indian Ocean Thermal Data Assembly Centre (WOCE UOT/DAC)*. CSIRO Division of Oceanography, Hobart.
- Bailey, R.J. and A. Gronell (1994). *Quality control cookbook for XBT data*. CSIRO Marine Laboratories report No. 221, Hobart.
- Bane, J.M. Jr., and M.H. Sessions (1984). A field performance test of the Sippican deep aircraft deployed expendable bathythermograph. *J. Geophys. Res.*, 89 3615-3621.
- Boyd, J.D. (1987). Improved depth and temperature conversion equations for Sippican AXBTs. *J. Atmos. Oceanic Technol.*, 4, 545-551.
- Boyd, J.D. and R.S. Linzell (1993a). The temperature and depth accuracy of Sippican T-5 XBTs. *J. Atmos. Oceanic Technol.*, 10, 128-136.
- Boyd, J.D. and R.S. Linzell (1993b). Evaluation of the Sparton tight-tolerance AXBT. *J. Atmos. Oceanic Technol.*, 10, 892-899.
- Boyer, T.P., J.I. Antonov, H.E. Garcia, D.R. Johnson, R.A. Locarnini, A.V. Mishonov, M.T. Pitcher, O.K. Baranova, I.V. Smolyar (2006). World Ocean Database 2005. S. Levitus, Ed., *NOAA Atlas NESDIS 60*, U.S. Gov. Printing Office, Wash., D.C., 190 pp., DVDs.
- Budeus, G. and G. Krause (1993). On-cruise calibration of XBT probes. *Deep-Sea Res.*, 40, 1359-1363.
- Cheng, L., J. Zhu, R. Cowley, T. Boyer, and S. Wijffels (2014). Time, probe type, and temperature variable bias corrections to historical expendable bathythermograph observations. *J. Atmos. Oceanic Technol.*, 31, 1793-1825, doi:10.1175/JTECH-D-13-00197.1.
- Cheng, L., J. Abraham, G. Goni, T. Boyer, S. Wijffels, R. Cowley, V. Gouretski, F. Reseghetti, S. Kizu, S. Dong, F. Bringas, M. Goes, L. Houpert, J. Sprintall, and J. Zhu (2014). XBT science: Assessment of instrumental biases and errors. *Bull. Amer. Meteor. Soc.*, 97, 924-933, doi:10.1175/BAMS-D-15-00031.1.
- Conkright, M., S. Levitus, and T. Boyer (1994). Quality control and processing of historical oceanographic and nutrient data. *NOAA NESDIS Technical Report 79*, Wash., D.C.
- Cowley, R., S. Wijffels, L. Cheng, T. Boyer, and S. Kizu, (2013). Biases in expendable bathythermograph data: A new view based on historical side-by-side comparisons. *J. Atmos. Oceanic Technol.*, 30, 1195-1225, doi:10.1175/JTECH-D-12-00127.1.
- Demeo, R.P. (1969). The validity of expendable bathythermograph measurements. *Trans. Of the Marine Temperature Measurements Symposium. Mar. Tech. Soc.*, 155-179.
- Flierl, G. and A.R. Robinson (1977). XBT measurements of the thermal gradient in the MODE eddy. *J. Phys. Oceanogr.*, 7, 300-302.
- Garcia, H.E., T.P. Boyer, R.A. Locarnini, O.K. Baranova, J.R. Reagan, and M.M. Zweng (2018). World Ocean Database 2018 Documentation. Ed. T. Boyer, *NODC Internal Report*, NOAA Printing Office, Wash., D.C.
- Good, S.A. (2011). Depth biases in XBT data diagnosed using Bathymetry data. *J. Atmos. Oceanic Technol.*, 28, 287-300, doi:10.1175/2010JTECHO773.1.

- Gouretski, V. and K.P. Koltermann (2007). How much is the ocean really warming? *Geophys. Res. Lett.*, 34, L01610, doi:10.1029/2006GL027834.
- Gouretski, V. and F. Reseghetti (2010). On depth and temperature biases in bathythermograph data: Development of a new correction scheme based on analysis of a global database. *Deep-Sea Res. I*, 57, 812-834, doi: 10.1016/j.dsr.2010.03.011.
- Gouretski, V. (2012). Using GEBCO digital bathymetry to infer depth biases in the XBT data. *Deep-Sea Res. I*, 62, 40-52, doi:10.1016/j.dsr.2011.12.012.
- Green, A. W. (1984). Bulk dynamics of the expendable bathythermograph (XBT). *Deep-Sea Res.*, 31, 415-426.
- Hallock, Z.R. and W.J. Teague (1992). The fall rate of the T-7 XBT. *J. Atmos. Oceanic Technol.*, 9, 470-483.
- Hamon, M., G. Reverdin, and P.-Y. Le Traon (2012). Empirical correction of XBT data. *J. Atmos. Oceanic Technol.*, 29, 960-973, doi:10.1175/JTECH-D-11-00129.1.
- Hanawa, K. and H. Yoritaka (1987). Detection of systematic errors in XBT data and their correction. *J. Oceanogr. Soc. of Japan*, 43, 68-76.
- Hanawa, K. and T. Yasuda (1991). Re-examination of depth errors in XBT data and their correction. *J. Atmos. Oceanic Technol.*, 8, 422-429.
- Hanawa, K., P. Rual, R. Bailey, A. Sy, and M. Szabados (1995). A new depth-time equation for Sippican or TSK T-7, T-6 and T-4 expendable bathythermographs (XBT). *Deep-Sea Res.*, 42, 1423-1452.
- Hayes, S.P., L.J. Mangum, J. Picaut, A. Sumi, and K. Takeuchi (1991). TOGA-TAO: a moored array for real-time measurements in the tropical Pacific Ocean. *Bull. Amer. Meteorol. Soc.*, 339-347.
- Heinmiller, R.H., C.C. Ebbesmeyer, B.A. Taft, D.B. Olson, and O.P. Nikitin (1983). Systematic errors in expendable bathythermographs (XBT) profiles. *Deep-Sea Res.*, 30, 1185-1196.
- IOC (1992a). Summary report of the IGOSS task team on quality control for automated systems and addendum to the summary report. *IOC/INF-888*.
- IOC (1992b). Summary report of the IGOSS task team on quality control for automated systems and addendum to the summary report. *IOC/INF-888-append*.
- IOC (1994). Calculation of new depth equations for expendable bathythermographs using a temperature-error-free method (Application to Sippican/TSK T-7, T-6 and T-4 XBTs). *IOC Technical Series No. 42*, 46 pp.
- Ishii, M. and M. Kimoto (2009). Reevaluation of historical ocean heat content variations with time-varying XBT and MBT depth bias corrections. *J. Oceanogr.*, 65, 287-299.
- Kizu, S., H. Yoritaka, and K. Hanawa (2005). A new fall-rate equation for T-5 Expendable Bathythermograph (XBT) by TSK. *J. Oceanogr.*, 61, 115-121.
- Levitus, S. and T. Boyer (1994). World Ocean Atlas 1994, Vol. 5: Interannual variability of upper ocean thermal structure. *NOAA Atlas NESDIS 5*. U.S. Gov. Printing Office, Wash., D.C., 150 pp.
- Levitus, S., R. Gelfeld, T. Boyer, and D. Johnson (1994). Results of the NODC and IOC Data Archaeology and Rescue projects. *Key to Oceanographic Records Documentation No. 19*, National Oceanographic Data Center, Wash., D.C., 67 pp.

- Levitus, S., M. Conkright, T.P. Boyer, R. Gelfeld, D. Johnson, I. Smolyar, C. Stephens, G. Trammell, R. Moffatt, T. O'Brien, and L. Stathoplos (1998). Results of the IOC Global Oceanographic Data Archaeology and Rescue (GODAR) project. *NOAA NESDIS Technical Report*.
- Levitus, S., S. Sato, C. Maillard, N. Mikhailov, P. Caldwell, and H. Dooley (2005). Building Ocean Profile-Plankton Databases for Climate and Ecosystem Research. *NOAA Technical Report NESDIS 117*, U.S. Gov. Printing Office, Wash., D.C., 29 pp.
- Levitus, S., J.I. Antonov, T.P. Boyer, R.A. Locarnini, H.E. Garcia, and A.V. Mishonov (2009). Global ocean heat content 1955-2008 in light of recently revealed instrumentation problems. *Geophys. Res. Lett.*, 36, L07608, doi: 10.1029/2008GL037155.
- McDowell, S. (1977). A note on XBT accuracy. *POLYMODE News*, 29.
- McPhaden, M.J. (1993). TOGA-TAO and the 1991-93 El Niño-Southern Oscillation Event. *Oceanogr.*, 6, 36-44.
- Narayanan, S. and G.R. Lilly (1993). On the accuracy of XBT temperature profiles. *Deep-Sea Res.*, 40, 2105-2113.
- Rual, P., A. Dessier, and J. P. Rebert (1995), New depth equation for 'old' Sparton XBT-7 expendable bathythermographs. *International WOCE newsletter*, 19, 33-34.
- Rual, P., A. Dessier, J.P. Rebert, A. Sy, and K. Hanawa (1996). New depth equation for Sparton XBT-7 expendable bathythermographs, preliminary results. *International WOCE newsletter*, 24, 39-40.
- Singer, J.J. (1990). On the error observed in electronically digitized T-7 XBT data. *J. Atmos. Oceanic Technol.*, 7, 603-611.
- Sy, A. (1991). XBT measurements. *WOCE reports*, 67/91.
- Thadathil, P., A.K. Saran, V.V. Gopalakrishna, P. Vethamony, N. Araligidat, and R. Bailey (2002). XBT fall rate in waters of extreme temperature: A case study in the Antarctic Ocean. *J. Atmos. Oceanic Technol.*, 19, 391-396.
- Wijffels, S.E., J. Willis, C.M. Domingues, P. Barker, N.J. White, A. Gronell, K. Ridgway, and J.A. Church (2008). Changing Expendable Bathythermograph Fall Rates and Their Impact on Estimates of Thermosteric Sea Level Rise. *J. Clim.*, 21, 5657-5672.
- Willis, J.K., D. Roemmich, and B. Cornuell (2004). Interannual variability in upper ocean heat content, temperature, and thermosteric expansion on global scales. *J. Geophys. Res.*, 109, C12036, doi: 10.1029/2003JC002260.
- Wright, D. and M. Szabados (1989). Field evaluation of real-time XBT systems. *Oceans 89 Proceedings*, 5, 1621-1626.
- Wright, D. (1989). Field evaluation of the XBT bowing problem. *NOS OOD Data Report 91-2*. National Ocean Service, NOAA, Rockville, Maryland, U.S.A.

# CHAPTER 5: EXPENDABLE CONDUCTIVITY-TEMPERATURE-DEPTH DATA (XCTD)

*Christopher R. Paver, Ricardo A. Locarnini, Tim P. Boyer, Hernán E. García,  
Alexey V. Mishonov, Dan Seidov, Olga K. Baranova,*

*Ocean Climate Laboratory  
National Center for Environmental Information  
Silver Spring, MD*

## 5.1. INTRODUCTION

An Expendable Conductivity, Temperature and Depth (XCTD) is an ocean profiling instrument, which usually consist of a data acquisition system and launcher onboard the vessel, and a digital XCTD consisting of a thermistor and conductivity sensor. Probes can be launched from ships, submarines, and airborne platforms. Prior to 1999, Tsurumi Seiki Company, Ltd. (TSK) and Sippican Inc. (now Lockheed Martin Corp.; LHM) manufactured XCTDs. In 1999, TSK became the sole manufacturer of XCTDs with LHM becoming a distributor (LHM, XCTD profiling system website; Gille *et al.*, 2009).

The XCTD is a free-falling probe, which is linked to the acquisition system through a thin insulated conductive wire that is used to transmit the temperature and conductivity data back to the acquisition system in real time. Depth is estimated from the elapsed time between when the probe enters the water and the time each temperature-conductivity measurement is made using a fall-rate equation supplied by the vendor. The XCTD probes manufactured by two major companies: Sippican and TSK are very different in workings, conductivity sensors, and data transmission. Sippican-made probes have a 4 Hz sample rate and roughly  $3.2 \text{ m}\cdot\text{s}^{-1}$  fall velocity and can record data every 0.8 m of depth along the profile (Johnson, 1995). TSK-made probes, however, are able to sample at 25 Hz, i.e. every  $\sim 420$  ms, which approximately equal to 17 cm depth interval.

Over the years of collection, XCTD data were submitted in both high and low vertical resolution formats, therefore these data are stored in two WOD23 datasets: high resolution data resides in the CTD dataset (12,740 XCTD casts), and low-resolution data resides in the OSD dataset (1,708 XCTD casts). The earliest XCTD data in WOD23 are 22 casts collected in the Coral Sea in December of 1991.

## 5.2. XCTD PRECISION AND ACCURACY

The precision and accuracy of XCTD data depends on the manufacturer and model. Current specifications from TSK identify similar accuracy and precision for temperature and conductivity. Temperature accuracy and precision are  $\pm 0.02$  °C and 0.01 °C, respectively. Conductivity accuracy and precision are  $\pm 0.03 \text{ mS}\cdot\text{cm}^{-1}$  and  $\pm 0.017 \text{ mS}\cdot\text{cm}^{-1}$ , respectively. Depth accuracy is 4.6 m or 2% of depth, whichever is greater. System response time is 40 ms for conductivity and 100 ms for

temperature (TSK). If these errors are correlated, the salinity error could be as high as  $\pm 0.08$ , otherwise a salinity accuracy of  $\pm 0.05$  is expected (Johnson, 1995). Mizuno and Watanabe (1998) also reported similar numbers.

Despite the XCTD instrument being in use for some time now, some problems with data accuracy may still exist. Early comparison of the XCTD data with CTD performed by Hallock and Teague (1990) concluded that “Examination of temperature and conductivity shows a significant systematic offset of the XCTDs relative to the CTD, suggesting a calibration error”. Later, Sy (1993) revealed that “test results conclusively show that XCTD probes do not meet the manufacturer’s specification”. A test of modified probes indicated: a) “that the XCTD sensor accuracies are better than  $\pm 0.02^\circ\text{C}$  and  $\pm 0.04 \text{ mS}\cdot\text{cm}^{-1}$  without any correction for the conductivity offset” (Alberola *et al.*, 1996); b) that “the system is close to the point of meeting the claimed specification” (Sy, 1996); and c) that “the system is close to providing the performance required by the oceanographic community for upper ocean thermal and salinity investigation” (Sy, 1998). Large amounts of high frequency noise or spiking reported in both XCTD temperature (Gille *et al.* 2009) and salinity (Yuan *et al.* 2004) profiles, required additional data treatment. Nevertheless, XCTD instruments are able to provide data in a more convenient way than traditional CTDs, which encourage data collection in under-sampled regions like the Arctic or the Southern Ocean (Yuan *et al.* 2004, Gille, *et al.* 2009) at higher sampling density. Lancaster and Baron (1984) in Antarctic Surface Waters, and Sprintall and Roemmich (1999) in the Pacific Ocean demonstrate other examples of XCTD deployments.

### 5.3. XCTD FALL-RATE ERROR

The XCTD instrument does not measure pressure or depth directly. The depth of the instrument is computed from the elapsed time from when the probe enters the water through use of a fall-rate equation. Research conducted by Johnson (1995) reveals that the manufacturer-supplied fall-rate coefficients give too slow a descent for some probes. Alberola *et al.*, (1996), showed similar results. Therefore, revised fall-rate equations were introduced (Johnson, 1995; Mizuno and Watanabe, 1998; Koso *et al.*, 2005) and evaluated (Kizu *et al.*, 2008). Uchida *et al.* (2011) developed a data processing method to obtain high-quality XCTD data.

A depth-correction algorithm was applied to XCTD data in WOD09 while computing temperature and salinity values at standard depth levels. For that purpose, depth values were first recalculated back to elapsed time and then two different manufacturer-dependent depth equations were used for adjusted depth calculation. Manufacturer-provided fall rate coefficients for the different XCTD make and models can be found on the Global Temperature and Salinity Profile Programme’s (GTSP) website for the World Meteorological Organization (WMO) 1770 code table.

For data collected by Sippican instruments the equation of Johnson (1995) was used. To indicate that data were subject to such treatment, secondary header code #54 was set to 103. Following procedure and parameters were employed:

$$t = (s_1 \cdot d_x + s_2) - s_3$$

$$d_z = s_a \cdot t + s_b \cdot t^2$$

where:  $s_1 = -1876.17261$ ,  $s_2 = 9317957$ ,  $s_3 = -3052.53296$ ;

$$s_a = 3.227, s_b = -2.17 \cdot 10^{-4};$$

$t$  – time since drop (seconds);

$d_x$  – originally calculated depth (meters);

$d_z$  – new calculated depth (meters).

For data collected by TSK instruments, (Table 5.1) equation of Mizuno and Watanabe (1998) was used. To indicate that data were subject of such treatment, secondary header code #54 was set to 104. The following procedure and parameters were employed:

$$t = (t_1 \cdot d_x + t_2) - t_3$$

$$d_z = t_a \cdot t + t_b \cdot t^2$$

where:  $t_1 = -4672.89697$ ,  $t_2 = 62365712$ ,  $t_3 = -7897.19678$ ;

$$t_a = 3.426, t_b = -4.70 \cdot 10^{-4};$$

$t$  – time since drop (seconds);

$d_x$  – originally calculated depth (meters);

$d_z$  – new calculated depth (meters).

## 5.4. XCTD CAST DISTRIBUTIONS

Table 5.1 gives the yearly counts of XCTD profiles for the World Ocean. Figure 5.1 shows this graphically. There is a total of 14,448 XCTD profiles for the entire World Ocean (12,740 in CTD and 1,708 in OSD) in WOD23.

**Table 5.1. The number of XCTD casts in WOD23 as a function of year**

A total number of casts 14,448 distributed between CTD dataset (12,740) and OSD dataset (1,708).

YEAR	CAST CTD/OSD <sup>1</sup>	YEAR	CASTS CTD/OSD <sup>1</sup>	YEAR	CASTS CTD/OSD <sup>1</sup>	YEAR	CASTS CTD/OSD <sup>1</sup>
1991	22/0	1999	405/182	2007	830/0	2015	436/0
1992	0/0	2000	478/582	2008	997/9	2016	843/0
1993	28/0	2001	327/638	2009	352/0	2017	177/0
1994	0/0	2002	573/126	2010	535/0	2018	278/0
1995	139/0	2003	479/12	2011	305/0	2019	277/0
1996	104/0	2004	551/41	2012	393/0	2020	247/0
1997	131/0	2005	1046/0	2013	595/0	2021	404/0
1998	166/118	2006	1151/0	2014	250/0	2022	221/0

<sup>1</sup> CTD – high-resolution casts; OSD – low-resolution casts

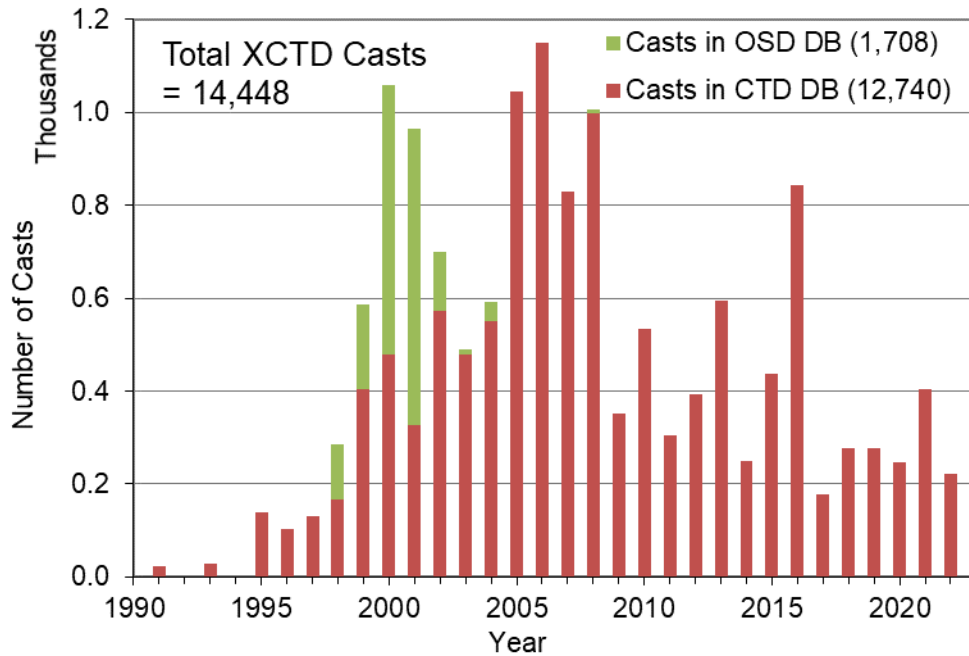


Figure 5.1. Temporal distribution of Expendable Conductivity, Temperature and Depth (XCTD) casts in WOD23 CTD and OSD datasets.

Table 5.2 gives national contributions of XCTD data to WOD23. The geographic distribution of XCTD casts is shown on Figure 5.2.

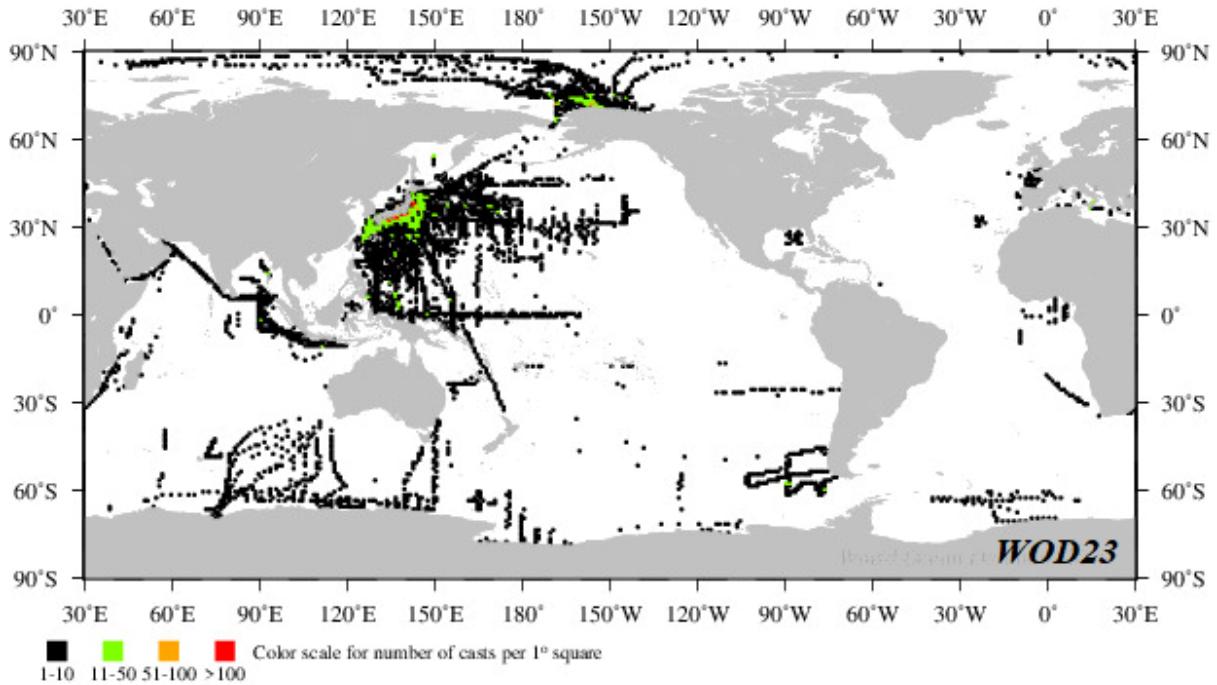
Table 5.2. National contributions of XCTD casts in WOD23.

ISO <sup>1</sup> 3166-1 Country Code	Country Name	XCTD Casts		% of Total
		CTD	OSD <sup>2</sup>	
JP	Japan	10,252	666	75.56%
US	United States	1,319	322	11.36%
FR	France	383	0	2.65%
PA	Panama	337	0	2.33%
CN	China, The People's Republic of	289	0	2.00%
AU	Australia	71	0	0.49%
IN	India	60	0	0.42%
99 <sup>3</sup>	Unknown	29	720	5.18%
<i>Total: 14,448</i>		<i>12,740</i>	<i>1,708</i>	<i>100.0%</i>

<sup>1</sup> ISO = International Organization for Standardization

<sup>2</sup> CTD = high-resolution casts; OSD = low-resolution casts

<sup>3</sup> Code '99' is not from ISO 3166-1. It is a supplementary code used in WOD.



**Figure 5.2. Geographic distribution of Expendable Conductivity, Temperature and Depth (XCTD) casts in WOD23.**

While the majority of XCTD casts (7,628) have no information about the data-collecting organizations, significant amount of XCTD data were collected and submitted by four major institutions. Among them: Japan Oceanographic Data Center (JODC, 2,195 casts), Ocean Research Department of Japan Marine Science and Technology Center (JAMSTEC, 1,376 casts), Arctic Submarine Laboratory (ASL US, 897 Casts), and Japan Meteorological Agency (JMA, 260 casts).

Figure 5.3 illustrates the distribution of the high-resolution and low-resolution XCTD data in CTD (red bars) and OSD (green bars) databases as a function of depth at standard depth levels.



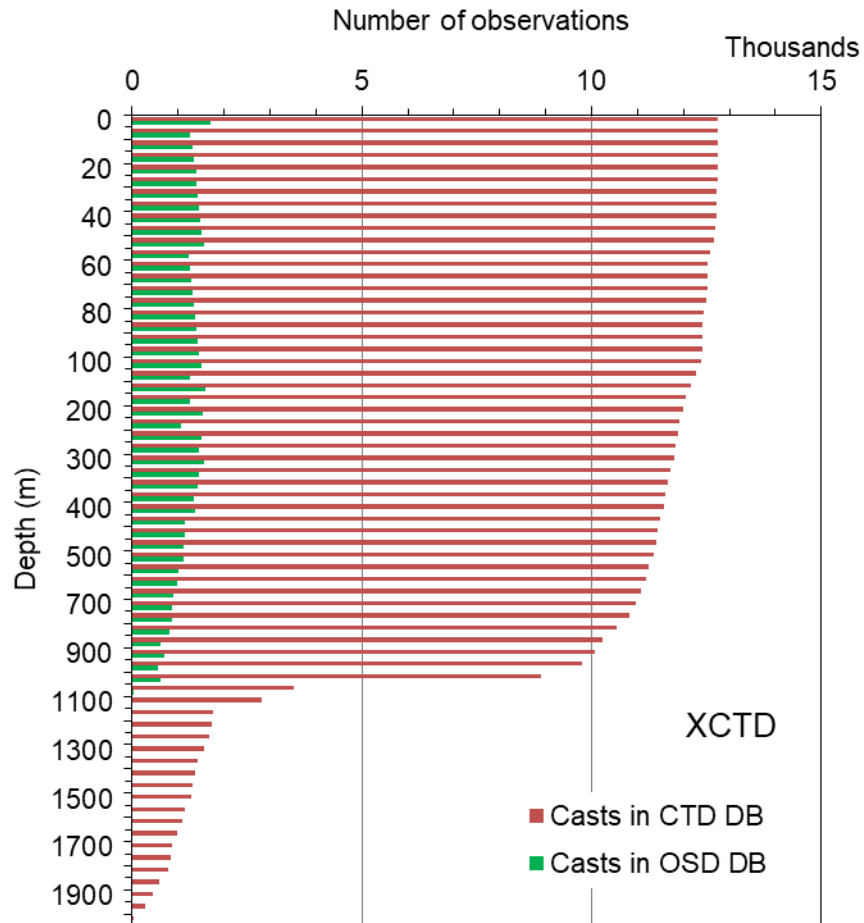


Figure 5.3. Distribution of Expendable Conductivity, Temperature and Depth (XCTD) data at standard depth levels in CTD and OSD datasets of WOD23.

## 5.5. RELEVANT WEB SITES

All websites last accessed: 2024-04-26

Global Temperature and Salinity Profile Programme:

<https://www.ncei.noaa.gov/products/global-temperature-and-salinity-profile-programme>

International Organization for Standardization (ISO): <https://www.iso.org/iso-3166-country-codes.html>

Japan Marine Science & Technology Center (JAMSTEC): <http://www.jamstec.go.jp/e/>

Japan Meteorological Agency (JMA): <http://www.jma.go.jp/jma/indexe.html>

Lockheed Martin, Corp.: <https://www.lockheedmartin.com/en-us/products/oceanographic-instrumentation.html>

Scientific Ice Expeditions Program (SCICEX): <https://nsidc.org/scicex>

Ship of Opportunity Programme (SOOP): <https://community.wmo.int/en/ship-opportunity-programme>

Tsurumi Seiki Co. (TSK <https://tsurumi-seiki.co.jp/en/>)

Tohoku University, Japan: <http://www.tohoku.ac.jp/en/>

U.S. Navy, Arctic Submarine Laboratory (ASL): <https://www.sublant.usff.navy.mil/ASL/>

## **5.6. REFENCES AND BIBLIOGRAPHY**

- Alberola, C., C. Millot, U. Sende, C. Mertens, and J.-L. Fuda (1996). Comparison of XCTD/CTD data. *Deep-Sea Res.*, 43, 859-76.
- Gille, S.T., A. Lombrozo, J. Sprintall, G. Stephenson, and R. Scarlet (2009). Anomalous spiking in spectra of XCTD temperature profiles. *J. Atmospheric Oceanic Tech.*, 26, 1157-1164.
- Hallock, Z.R. and W.J. Teague (1990). XCTD test: reliability and accuracy study (XTRAS) *Tech. note 69*.
- Johnson G.C. (1995). Revised XCTD fall-rate equation coefficients from CTD data. *J. Atmos. Oceanic Technol.*, 12, 1367-73.
- Kizu, S., H. Onoshi, T. Suga, K. Hanawa, T. Watanabe, and H. Iwamiya (2008). Evaluation of the fall rates of the present and developmental XCTDs. *Deep-Sea Res. I*, 55, 571-586.
- Koso, Y., H. Ishii, and M. Fujita (2005). An examination of the depth conversion formula of XCTD-2F. *Technical Bulletin on Hydrography and Oceanography. No. 23*, Japan Coast Guard Hydrographic and Oceanographic Department, Tokyo, Japan, 93-98.
- Lancaster, R.W. and G. Baron (1984). Measuring ASW, oceanographic parameters with XCTD profiling systems. *Sea tech.*, 18-23.
- Mizuno, K. and T. Watanabe (1998). Preliminary results of *in situ* XCTD/CTD comparison test. *J. Oceanogr.*, 54(4), 373-380.
- Morison, J.H., M. Steele, and R. Andersen (1998). Hydrography of the upper Atlantic Ocean measured from the nuclear submarine USS *Pargo*. *Deep-Sea Res. I*, 45(1), 15-38.
- Sprintall, J., and D. Roemmich (1999). Characterizing the structure of the surface layer in the Pacific Ocean. *J. Geophys. Res. – Oceans*, 104, 23297-311.
- Sy, A. (1993). Field evaluation of XCTD performance. *International WOCE Newsletter*, 14, 33-37.
- Sy, A. (1996). Summary of field test of the improved XCTD/MK-12 system. *International WOCE Newsletter*, 22, 11-13.
- Sy, A. (1998). At-sea test of a new XCTD system. *International WOCE Newsletter*, 31, 45-47.
- Yuan, X.J., D.G. Martinson, Z.Q. Dong (2004). Upper ocean thermohaline structure and its temporal variability in the southeast Indian Ocean. *Deep-Sea Res. I*, 51(2), 333-347.

# CHAPTER 6: PROFILING FLOATS DATA (PFL)

*Zhankun Wang, Ricardo A. Locarnini, Tim P. Boyer, Olga K. Baranova, Hernán E. García, Alexey V. Mishonov, James R. Reagan*

*Ocean Climate Laboratory  
National Centers for Environmental Information / NOAA  
Silver Spring, MD*

## 6.1. INTRODUCTION

Profiling floats are autonomous vehicles equipped with oceanographic sensors which measure vertical profiles of oceanographic variables. These autonomous platforms float passively with the ocean currents at a preprogrammed pressure level and then move up and down between the surface and a predetermined mid-water level (Roemmich *et al.*, 2004). They rise to the ocean surface at a predetermined time interval to broadcast collected information to satellites. Satellite telecommunication technology is used to transmit the float position as well as date and time of receipt of the trip to the surface, and in some cases on the preceding dive. Several different sensors may be attached to the profiling float. However, compromises must be made between the weight and power usage of the sensors and the intended lifetime of the profiling float's battery. Most profiling floats are equipped with pressure, temperature, and conductivity sensors (for calculating salinity). Oxygen, nitrate, pH, and chlorophyll-a sensors have also been deployed, as well as transmissometers, optical irradiance sensors, velocity meters, turbulence sensors, and rainfall and wind speed sensing instrumentation. In addition to the measurements of pressure, temperature, salinity, and oxygen present in previous releases, Table 6.1 shows that measurements of nitrate, pH, chlorophyll-a, and transmissivity from Biogeochemical-Argo (Claustre *et al.*, 2010) are included in the PFL dataset of the *World Ocean Database 2023* (WOD23).

**Table 6.1. Variables and profile counts in the PFL WOD23 dataset.**

Variable	Number of Profiles
Temperature	2,747,412
Salinity	2,594,626
Oxygen	241,658
Nitrate	56,379
pH	29,554
Chlorophyll-a	111,113
Transmissivity	4,923
Pressure	2,748,011

The float's active movement is achieved by changes in its buoyancy using external bladders. Oil is pumped from an internal chamber to an external bladder, increasing volume and decreasing density, to force the float to rise to the surface. Oil is then pumped from the external bladder back into the float casing to decrease the volume, increasing the density to the point where the float will sink until it achieves a neutral density commensurate with the pressure level at which it will passively drift with ocean current

Floats are relatively low cost compared with ship-based measurements. Davis *et al.* (2001) calculate that they are equivalent in cost per profile (temperature only) to an XBT. However, their value is much greater since most floats also measure salinity, a significant number of them measure variables such as dissolved oxygen and nitrate, and they are able to measure during any sea or weather condition, with the partial exception of ice cover. Profiling floats are adding measurements in areas and seasons for which little, if any data, were available.

## **6.2. PREDECESSORS OF PROFILING FLOATS**

The precursors of the present profiling floats were neutrally buoyant floats used to track currents at a predetermined level in the ocean. These floats did not measure temperature or conductivity. The first neutrally buoyant floats were designed and deployed by Swallow (1955). These floats sank to their neutrally buoyant level in the water column and were then tracked by a ship nearby. The Swallow floats were used to verify the deep western boundary current predicted by Stommel (1957) (Swallow and Worthington, 1961). In the late 1960s, the SOFAR (SOund Fixing and Ranging) float was developed (Webb and Tucker, 1970; Rossby and Webb, 1970). This was similar to a Swallow float. They differed in that the float was tracked by underwater listening devices which picked up sound emitted by the floats at intervals which allowed geo-location. The listening devices did not have to be in close proximity to the float, eliminating a major limitation of the Swallow float. Further advances led to the RAFOS (the word SOFAR spelled backwards) floats which reversed the geo-location procedure of the SOFAR floats by having the float listen for signals emitted by stationary underwater devices (Rossby *et al.*, 1986). The RAFOS float was smaller than the SOFAR float since it did not need to emit sound, and therefore it was less expensive to deploy. RAFOS floats can carry multiple sensors for measuring ocean properties, including pressure, temperature, and dissolved oxygen. They can be made "isopycnal", or density-following, so that they more accurately follow water parcels in regions where density surfaces are steeply sloped. However, it still required a network of sound sources.

## **6.3. FIRST GENERATION PROFILING FLOATS**

One of the objectives of the World Ocean Circulation Experiment (WOCE, active fieldwork period 1990-1998) was to estimate the mean flow of the World Ocean. To set up a worldwide system of sound sources to achieve this objective using RAFOS floats would have been prohibitively expensive. The Autonomous LAgrangian Circulation Explorer (ALACE) floats (Davis *et al.*, 1992) were the implemented solution. First operationally deployed in the Drake Passage in 1990, these floats eliminated the need for sound sources by surfacing periodically to be geo-located by the Advanced Research and Global Observation Satellites (ARGOS). The tradeoff for manageable costs were small uncertainties introduced in the velocity at depth due to drift while ascending and descending the water column and while broadcasting their signal at the surface. Also, within the framework of the WOCE program, the MARVOR float was created by the Institut Francais de REcherche de la MER (IFREMER) and Tekelec (now Martec), a French engineering firm. MARVOR floats use the same geo-location

principle as RAFOS floats, but they also cycle to the surface to send data to ARGOS satellites. They were first deployed in early 1994 in the Brazil Basin (Ollitrault *et al.*, 1994).

After the success of the new profiling floats, it was a logical step to include in their design oceanographic sensors to record temperature and salinity during the floats' ascent to the surface. In 1991, the first ALACE floats with temperature sensors were deployed, making them Profiling ALACE floats (P-ALACE floats), and in 1994 floats with both temperature and salinity sensors were deployed (Davis *et al.*, 2001). The earliest cast for PFLw23 was collected in 1994.

#### 6.4. PRESENT FLOAT TECHNOLOGY

Further improvements to the P-ALACE float design were made. Float R1, by Webb Research, was introduced at the request of Dr. Steve Riser in 1996 (personal communication, Dan Webb). It was replaced by its successor, the Autonomous Profiling EXplorer (APEX) by Webb Research, which is still in use today. Since 1997, APEX floats have been deployed from merchant vessels moving at speeds up to 25 knots, removing the need to employ research vessels in some areas. Other second-generation floats include the Sounding Oceanographic Lagrangian Observer (SOLO), developed at Scripps Institute of Oceanography. This float replaced the P-ALACE floats' reciprocating high pressure pump with a single stroke hydraulic pump (Davis and Zenk, 2001); the APEX uses a similar pump. This advance allowed the SOLO to more easily reach a desired isobar or isotherm and to cycle between subsurface depths before ascending to the surface. As the P-ALACE was the profiling version of the ALACE float, the PROVOR is the profiling version of the MARVOR float (Loaec *et al.*, 1998), and they have been deployed since 1997.

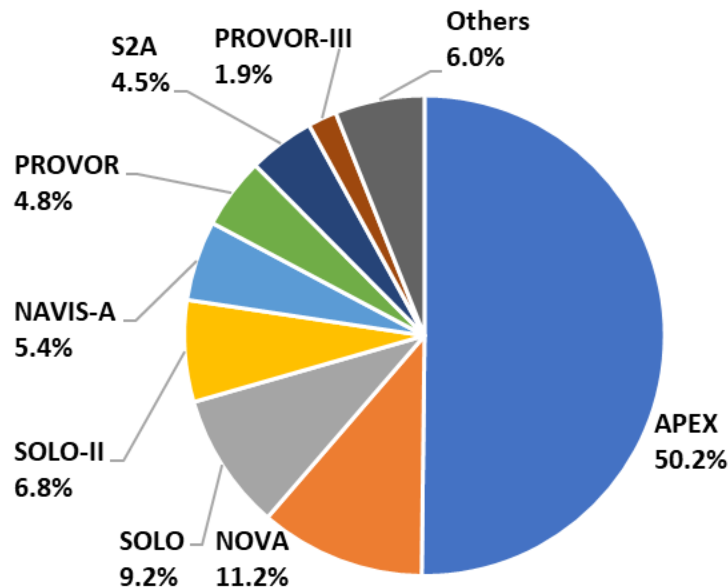


Figure 6.1. Casts from different types of profiling floats (PFL) in WOD23.

Both Martec and MetOcean (Canada) now produce PROVOR floats on the same design. MARVOR and PROVOR floats operate on the same bladder/buoyancy principles as the

ALACE floats. PROVOR floats have the added ability to record and store oceanographic profile data on their descent as well as their ascent. The Japanese Agency for Marine-Earth Science and Technology (JAMSTEC) and Tsurumi Seiki Co. (TSK) have developed and deployed the New profiling floats of Japan (NINJA) (Ando *et al.*, 2003) beginning in 2002. Navigating European Marine Observer (NEMO) floats have been deployed in the Southern Ocean starting in early 2004 by the Alfred Wegener Institute (AWI, Germany). These floats are based on the SOLO design and are equipped with algorithms based on temperature measurements which help them avoid surfacing in ice covered areas. NEMO floats combine this ability with RAFOS positioning, extending the reach of profiling floats to ice-covered regions. New generations of floats, e.g. SOLO-II, S2A by MRV Systems (USA), NOVA (New generation Oceanographic Variable buoyancy Autonomous) by MetOcean, and Navis by Sea-Bird Scientific (USA), include updated software and features, and form part of the Argo array.

Argo is constantly addressing some technological challenges by different innovations. The float lifetime has been doubled with improvement of battery life since the beginning of Argo program. Communications have now improved and most floats use Iridium transmitters, which are high bandwidth and only requires the floats to be on the surface for less than 30 minutes. Newly designed “Deep” types of profiling floats, e.g. Deep SOLO, Deep APEX, and Deep-Arvor (Le Reste *et al.*, 2016), increase their monitoring capabilities beyond the upper 2000 m of the water column, down to the ocean bottom. The deep float models include the Deep SOLO and the Deep APEX capable of reaching 6000m, and the deep ARVOR, the Deep NINJA and the HM4000 capable of reaching 4000m. More than 1000 Deep Argo floats have been deployed to systematically measure full-depth global ocean. BioGeoChemical Argo (BGC-Argo) is the extension of the core Argo by adding additional sensors for dissolved oxygen, pH, nitrate, chlorophyll, suspended particles, and downwelling irradiance measurements. WOD23 also includes data obtained with an air-deployed profiling ocean float, ALAMO (Air Launched Autonomous Micro Observer). Figure 6.1 shows the relative distribution of different types of profiling floats in WOD23. Most of the data, approximately 92%, are known to have been obtained by just seven float types.

#### **6.4.1. The Argo Project**

The Argo project is an umbrella project which coordinates the deployment, quality control, and public access for profiling float data. Argo is not an acronym: it refers to the relationship between the Jason satellite altimeter measuring ocean surface topography and the Argo floats revealing the ocean subsurface structure, evoking the mythical Jason and his ship the Argo (Gould, 2005). Since the year 2000, nearly all data from deployed floats are available through this project. Floats are deployed by individual countries, projects, and institutions, usually with some levels of coordination with Argo project. Float data are captured from the ARGOS and Iridium satellites by the Argo Data Assembly Centers (DACs) and placed on the World Meteorological Organization (WMO) Global Telecommunications System (GTS) within 24 hours. These data are also relayed in near-real-time to the two Argo Global Data Assembly Centers (GDACs): the French Coriolis Center at IFREMER, and the U.S. Global Ocean Data Assimilation Experiment (GODAE) server in Monterey, California hosted by the U.S. Navy. Within 24 hours the data are made available to the public through these sites as well.

Preliminary quality checks are performed at the DACs on the incoming data. These data are the real time data. Further quality control is performed at the DACs, the GDACs, at regional centers, and by the primary investigators responsible for the floats. A delayed mode version of the data is then released and adjustments may be made to the real time data. Each float is assigned a WMO identification number for easy identification. Meetings and workshops on data quality control, data access, and scientific research with floats have been held to keep the scientific community informed and coordinate responses and solutions to quality control and access problems. The goal of Argo is to deploy and maintain a global array of profiling floats to monitor the large-scale circulation of the world ocean, as well as its heat and fresh water content. With this stated goal, pressure, temperature and conductivity/salinity sensors are the core oceanographic sensors, although floats may be equipped with other sensors. Argo has surpassed its goal of 3,000 floats worldwide by 2007, with roughly 3,800 active floats since late 2012 – about 3,900 floats were active in February 2024. The Argo program reached its two-million profile mark by 2018 and is passed three-million profile mark (by the time of this writing, 2.8 million profiles had been incorporated into WOD). The program has deployed more than 15,000 floats and collected over 2.5 million temperature and salinity profiles (Table 6.1). The preference is for the floats to deliver profiles from 2000 decibars, or near the ocean bottom, to the surface every 10 days. Since the floats are deployed for other specific research goals, the parking depth (depth of passive motion) may not be at 2000 decibars. In fact, the recommended parking depth for Argo is 1000 decibars. However, the float should descend to 2000 decibars before beginning to record temperature and salinity. Some floats cycle to the surface at intervals other than 10 days.

The profiling float data in WOD23 consists of data most from the Argo U.S. GODAE server and some from the Global Temperature and Salinity Profile Project (GTSP) and WOCE. The profiling float data in WOD23 for oxygen, nitrate, pH, chlorophyll-a, and transmissivity, are all from the BGC-Argo on Argo U.S. GODAE server. We update the PFL data and metadata quarterly by obtaining all new and updated Argo cycles from the U.S. GODAE G. Any updates on data modes will be captured at the same time. If delayed-mode (DM) data becomes available, real-time data will be replaced with DM data

## **6.5. SENSOR ACCURACY**

The temperature and salinity data from the profiling floats come from various CTD sensors. The P-ALACE floats used an YSI 46016 thermistor, with estimated precision of 0.005°C, and a Falmouth Scientific Inc. (FSI) conductivity sensor with an estimated accuracy of 0.01 mS·cm<sup>-1</sup> (milliSiemens·centimeter<sup>-1</sup>). The pressure sensor used was a Paine strain gauge sensor. The sensor had hysteresis errors on the order of 5 meters initially, which were later reduced by thermally isolating the sensor (Davis *et al.*, 2001). To reduce pressure reading errors, Sea-Bird replaced the Paine strain gauge pressure sensor in their CTDs with a Druck pressure sensor (see Data Problems section, below). Later floats used FSI CTD sensors or CTD sensors from Sea-Bird. The Sea-Bird sensors have 0.002°C temperature accuracy, 0.005 salinity accuracy, and 2.4 db pressure accuracy. All accuracy data are from the product specifications (except the P-ALACE thermistor information from Davis *et al.*, 2001). Sea-Bird specifications are for Sea-Bird 41 CTD for ALACE floats.

For oxygen measurements, the Aanderaa 3835 oxygen optode sensor has an accuracy of 8  $\mu\text{M}$  or 5%, whichever is greater. Accuracy of the Sea-Bird SBE 43 oxygen sensor is 2% of saturation, while that of the SBE 63 is the greater of 3  $\mu\text{mol kg}^{-1}$  or 2%. These values are from the product specifications. Körtzinger *et al.* (2005) discuss oxygen measurements from profiling floats. Recent studies show that measurements with the Aanderaa oxygen optodes 3830 and 4330, or any optodes with reliable in-air measurements, and a proposed in-air oxygen measurement routine will result in accuracies close to 1  $\mu\text{mol kg}^{-1}$  over the entire lifetime of a float (Bittig and Körtzinger, 2015).

The nitrate data from profiling floats obtained using ultraviolet spectrophotometer nitrate sensors have a reported accuracy of 1  $\mu\text{mol kg}^{-1}$  (Johnson *et al.*, 2013). pH data obtained with ion sensitive field effect transistors have an accuracy of 0.01 pH (Johnson *et al.*, 2016). Chlorophyll-a data have an accuracy of the greater of 30% (Fluorescence) and 24% (Radiometer) or 0.03 mg Chl-a  $\text{m}^{-3}$  (Boss *et al.*, 2008; Xing *et al.*, 2011).

## **6.6. DATA PROBLEMS**

Data problems are of two types: 1) Sensor problems, 2) Data stream errors. Each will be examined separately.

### **6.6.1. Sensor problems**

The biggest persisting challenge for profiling float sensors is salinity and dissolved oxygen drift. Conductivity cells are calibrated against samples of standard seawater before deployment of the float. However, even over the course of a short oceanographic cruise, the conductivity sensor on a standard winch-deployed CTD can experience slowly increasing unidirectional errors (drift) due to biofouling and small changes in cell geometry. Profiling floats are designed to be almost constantly immersed in the harsh ocean environment for four years. Therefore, it is to be expected that the conductivity sensor on a float will experience drift. Oka (2005) estimated a salinity drift of  $-0.016 \pm 0.006$  per year from recalibration of three floats recovered after 2-2.5 years of deployment. From examining the extant float data, some floats can experience much larger drifts, or even abrupt deviations from calibration. A number of algorithms for correcting for drift have been proposed (Wong *et al.*, 2003; Böhme and Send, 2005; Durand and Reverdin 2005, Owens and Wong, 2009). Cabanes *et al.* (2016) proposed improvements to those methods to consider interannual variability. Wong *et al.* (2023) evaluated the bias and uncertainty of Argo salinity. They found that there was an increase in salty bias in the raw Argo data beginning 2015 and peaking during 2017-2018 due to sensor problem. However, the salty bias is expected to decrease in the coming years because the manufacturer issue has likely been fixed (Wong *et al.*, 2023). The Argo delayed-mode data are corrected for drift using a combination of algorithms described in Owens and Wong (2009) and Cabanes *et al.* (2016) and described in Wong *et al.* (2023).

For profiling float oxygen, Takeshita *et al.* (2013) showed a significant bias when compared to a climatology oxygen saturation. Bittig *et al.* (2018) found the bias is mainly a result of a reduction in oxygen sensitivity, which is proportional to oxygen values and that decrease in oxygen sensitivity is several percent per year. Similar to profiling float salinity data, adjustments might be made to the profiling float dissolved oxygen observations. Besides non-



adjusted real time data, there are real-time adjusted data and delayed mode data (codes = 1-3, variable specific secondary header 24) in WOD PFL. The data adjustment process for Argo oxygen can be found in Thierry *et al.* (2021). The recommended way for real time adjustment for oxygen is based on previous delayed-mode adjustments (Bittig *et al.*, 2018). If no previous delayed-mode adjustment, the adjustment will be based on the World Ocean Atlas (WOA) surface oxygen saturation field. The purpose of this adjustment is to remove a systematic bias that exists for oxygen optodes (Takeshita *et al.*, 2013, Bittig *et al.*, 2018; Thierry *et al.*, 2021). The delayed mode adjustments for profiling float oxygen are determined using discrete water sample or CTD oxygen profile as reference (Thierry *et al.*, 2021) or using repeated optode in-air data following the method proposed in Bittig *et al.* (2018).

Real time, real time adjusted and delayed-mode data are all available in WOD23. If the pressure, temperature, dissolved oxygen, or salinity adjustment variable is present in a cast (variable specific secondary header 19), the cast has delayed-mode quality control applied by the appropriate DAC. This adjustment variable gives the mean change between delayed-mode and real-time values at the same measurement levels for all levels below 500 meters depth. Salinity and dissolved oxygen drift adjustments and most pressure sensor adjustments are uniform over an entire profile so the adjustment variable is usually a good indicator of the profile change at each level from real-time to delayed-mode. However, there are some cases where a single level or a few levels have their values adjusted. In these cases, the adjustment variable does not represent the change to each level.

Even after delayed-mode adjustment, some residual bias and uncertainty can still remain in Argo salinity (Wong *et al.* 2023) and oxygen data (Garcia *et al.*, 2024b; Gouretski *et al.*, 2024). For example, we did a matchup comparison of oxygen between PFL (delayed mode only) and nearby CTD/OSD casts, an overall negative bias of  $-1.3 \mu\text{mol kg}^{-1}$  was still found between the matched pairs for the global ocean (Garcia *et al.*, 2024b). Gouretski *et al.* (2024) also found a prevailing negative bias for the delayed mode oxygen from Argo varying from  $-4$  to  $-1 \mu\text{mol kg}^{-1}$  among the data from different data assembly centers. Further analysis is needed to evaluate the residual bias in PFL dissolved oxygen.

In 2003, it was found that problems with the Druck Pressure Sensor were causing some floats to stay at the surface for prolonged periods and eventually to become surface drifters. The Druck Pressure Sensor is the successor to the Paine pressure sensor in Sea-Bird CTDs. Even when not severe, the problem may have caused errors in the salinity measurement due to increased biofouling due to prolonged surface exposure. When the problem was found, the CTDs were recalled and the source of the problem was fixed, but this was not possible for floats already deployed. A large number of SOLO floats with FSI CTD packages deployed in the Atlantic Ocean between 2003 and 2006 were found to have a pressure offset problem due to a software error. This error caused pressures to be paired with the temperature measurements from the next lower level, creating the illusion of a cooling ocean. Once the problem was found, a list of such floats was compiled. An effort was made to correct the problem, successful in some floats, not in others. All data from all these problematic floats are included in WOD23. For those data which could not be corrected, all float cycles are flagged. In early 2009, another problem with the Druck pressure sensor was found (J. Willis and D. Roemmich, Argo Steering Team, 2009). This problem causes pressure sensor drift after deployment. Barker *et al.* (2011) reported that about 57% of the profiles from APEX floats, the predominant type of deployed Argo floats – see Figure 6.1, could be immediately corrected for pressure sensor drift, while

only about half of the then uncorrectable APEX profiles could be corrected with the future release of updated metafiles and technical files.

During a normal transmission to the ARGOS satellite, a float needs to stay at the surface between 6 and 12 hours, and it is then when much of the biofouling occurs. This problem is being reduced by the increasing deployment of floats equipped to communicate with two-way communicating Iridium satellites. Two-way communication cuts down on the need for repeated rebroadcasts of the same message, since the broadcasting float can be notified of receipt of the message. This reduces the float's surface time to about 15-30 minutes. While in 2010 only 250 floats had been deployed with Iridium antennas, since 2013 most of the deployed floats use Iridium, which significantly reduced the biofouling issue.

Another identified problem is a thermal lag caused because the thermistor and the conductivity cell are located a small distance from each other. If there is a large vertical gradient in temperature, this can cause erroneous spikes in the salinity field. Work has been done to correct this lag problem and corrections are available in the delayed-mode data. However, the error is quite different between different Sea-Bird sensors found on floats, and not all the necessary metadata is available in all Argo data (G. Johnson, personal communication). Some anomalous spikes in salinity near large temperature gradients, probably caused by the thermal lag error, have been marked by automatic or subjective checks in WOD23.

Biofouling of the optical sensors measuring nitrate, pH, chlorophyll-a, and transmissivity has been reduced or eliminated by exposing the sensors to wave action when the floats surface and removing the sensors from the flow stream of the CTD (Johnson *et al.*, 2017).

### **6.6.2. Data-Stream Errors**

Problems caused by transmission of data from one site to another are always possible. The more data transfers are made, the more possibilities for error. The profiling float data are no exception. The most prevalent error, and one which is not usually recoverable, is errors in transmission of data packages from the float to the ARGOS satellites. Many of these transmission errors result in portions of profiles, or entire profiles containing erroneous information. Most of these errors are of such a nature that they are found and flagged in automatic quality control checks in WOD23 if they have not been removed beforehand. But there may be data with errors of this nature which escaped all quality control steps.

## **6.7. ORIGINATORS FLAGS**

The originators flags from the Argo program are those used by IODE and mostly kept intact in the WOD23 data. The Argo flags are as follows:

- 0 – no quality control (QC) performed
- 1 – good data
- 2 – probably good data
- 3 – bad data that are potentially correctible
- 4 – bad data
- 5 – value changed

8 – estimated value

9 – missing value

(from Argo quality control manual Version 3.8, February 2024, Wong *et al.*, 2024).

Note that not all data marked with originators flag 3 are marked with WOD23 quality control flags. Visual inspection of examples of these data found no reason not to use these data for scientific research. This just means that a quality control test that failed by Argo standards did not fail by WOD23 standards, or that the failing test was not performed for WOD23. The user of WOD23 can choose to use the Argo flags, the WOD23 flags, both, or neither. Note that now Argo replaces the observation with a missing value when the flag is 4, so no observations with Argo flag 4 are found in WOD23. We do not record the data if the Argo QC flag = 4 or 9, flag 8 is not carried in WOD23 either.

Argo also supplies a grey list. This is a list of floats and sensors which have been deemed to have failed at some point. The date of failure is also listed.

**The information on the grey list is used to set a quality control flag for PFL data in WOD23. This grey list is periodically updated. The grey list used to flag data for WOD23 is the version from April, 2023.**

## **6.8. PFL DATA DISTRIBUTIONS**

Figure 6.2 shows the geographic distribution of profiling float casts for the period 1994-2022. It is clear that Argo program has met its goal of full geographic coverage of non-ice-covered ocean: there are a total of 2,748,011 PFL casts for the entire World Ocean, closely divided between the southern hemisphere (1,342,273 casts, or 48.8%) and the northern hemisphere (1,405,738 casts, or 51.2%). PFL data represents 14.8% of all data in WOD23. Table 6.2 and Figure 6.3 shows that about 52.8% of the floats data are of U.S. origin, followed by Japan at about 8.4%. It also shows that many countries around the world contribute profiling float data. The yearly count in Table 6.3 and Figure 6.4 shows the rapid increase with time of recorded profiling float casts, from less than 15,000 a year before 2002, to more than 115,000 a year by 2008, with more than 160,000 casts a year obtained since 2014. The depth distribution (Figure 6.5) shows that many of the surface and near surface values do not exist or are missing: most float sensors are shut down near the surface to avoid biofouling, but it should be noted that surface and near surface data also ingested, not just the primary profile.

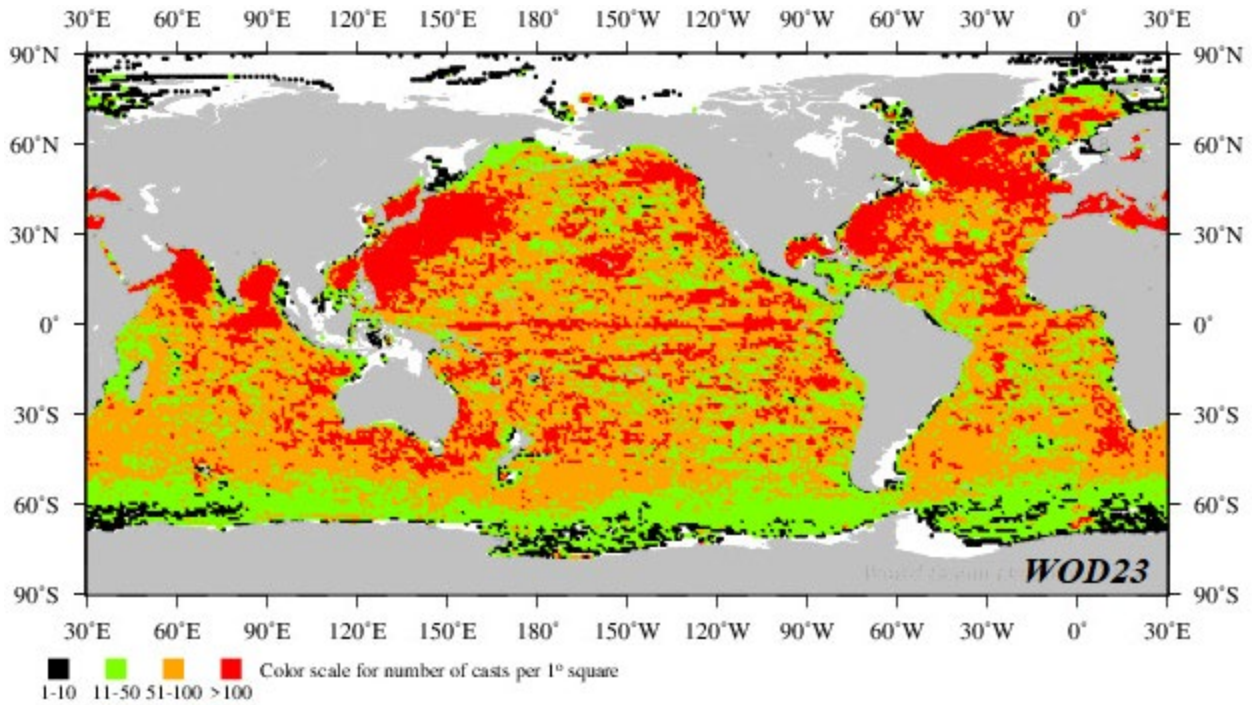


Figure 6.2. Geographic distribution of profiling floats (PFL) casts for the period 1994-2022 in WOD23.

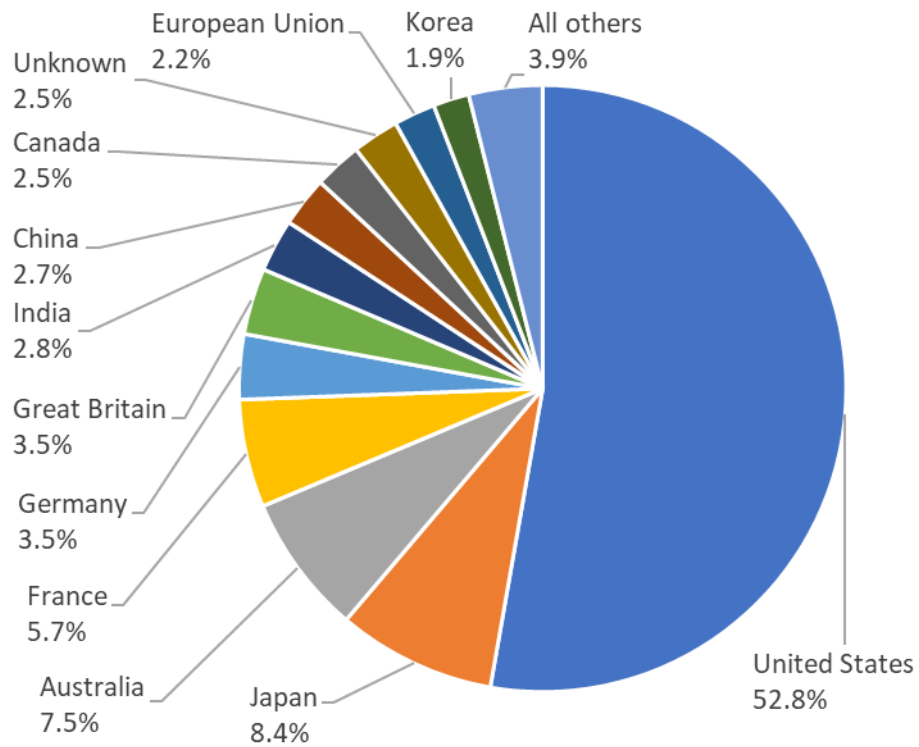


Figure 6.3. Profiling floats (PFL) data contribution by countries in WOD23

**Table 6.2. National contribution of PFL casts in WOD23.**

<b>ISO<sup>1</sup> Country Code</b>	<b>Country Name</b>	<b>PFL Casts</b>	<b>% of Total</b>
US	United States	1,449,968	52.764
JP	Japan	231,883	8.438
AU	Australia	205,105	7.464
FR	France	157,975	5.749
GB	Great Britain	97,258	3.539
DE	Germany	95,417	3.472
IN	India	77,212	2.810
CN	China	73,607	2.679
CA	Canada	69,967	2.546
99	Unknown	68,409	2.489
EU	European Union	60,685	2.208
KR	Korea; Republic of	52,753	1.920
IT	Italy	44,092	1.605
ES	Spain	14,082	0.512
NO	Norway	12,349	0.449
IE	Ireland	6,345	0.231
PL	Poland	5,841	0.213
NL	Netherlands	5,070	0.184
FI	Finland	4,488	0.163
CL	Chile	3,234	0.118
GR	Greece	2,768	0.101
TR	Turkey	2,432	0.089
BG	Bulgaria	2,006	0.073
BR	Brazil	1,163	0.042
MU	Mauritius	1,030	0.037
DK	Denmark	908	0.033
MX	Mexico	761	0.028
NZ	New Zealand	632	0.023
RU	Russian Federation	281	0.010
MA	Morocco	139	0.005
	<i>Total</i>	<i>2,748,011</i>	<i>100.0</i>

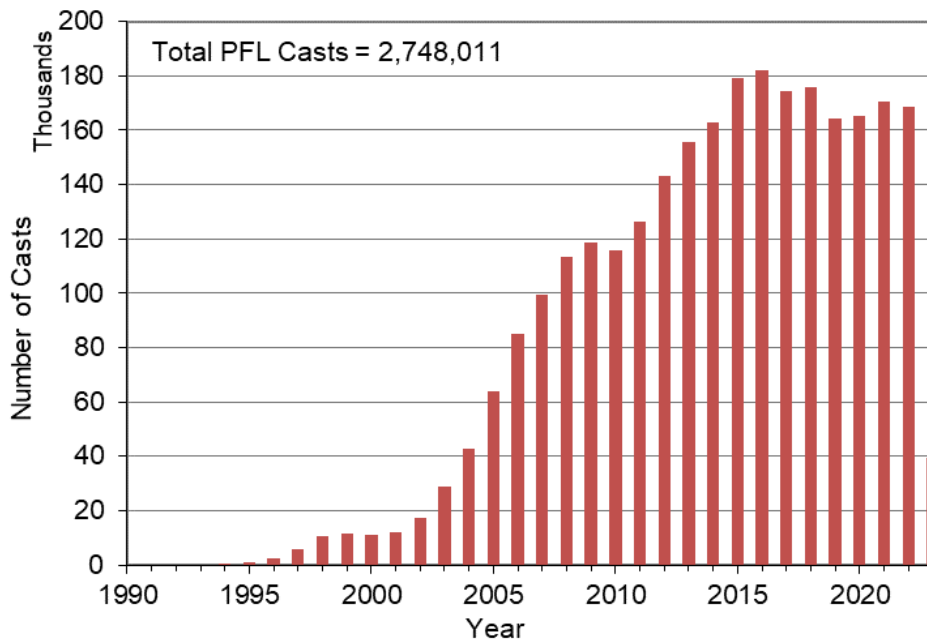
<sup>1</sup> ISO = International Organization for Standardization

[http://www.iso.org/iso/country\\_codes.htm](http://www.iso.org/iso/country_codes.htm)

**Table 6.3. The number of Profiling Float Data (PFL) casts as a function of year in WOD23. Year 2023 is a partial year from January to April.**

Total number of casts = 2,748,011.

YEAR	CASTS	YEAR	CASTS	YEAR	CASTS	YEAR	CASTS
1994	53	2002	17,299	2010	115,676	2017	174,430
1995	1,038	2003	29,054	2011	126,517	2018	175,895
1996	2,549	2004	42,677	2012	142,896	2019	164,432
1997	5,995	2005	64,082	2013	155,724	2020	165,335
1998	10,360	2006	84,969	2014	162,618	2021	170,697
1999	11,653	2007	99,421	2015	179,055	2022	168,540
2000	11,267	2008	113,578	2016	181,814	2023	39,621
2001	11,928	2009	118,838				



**Figure 6.4. Temporal distributions of Profiling Float Data (PFL) casts in WOD23.**

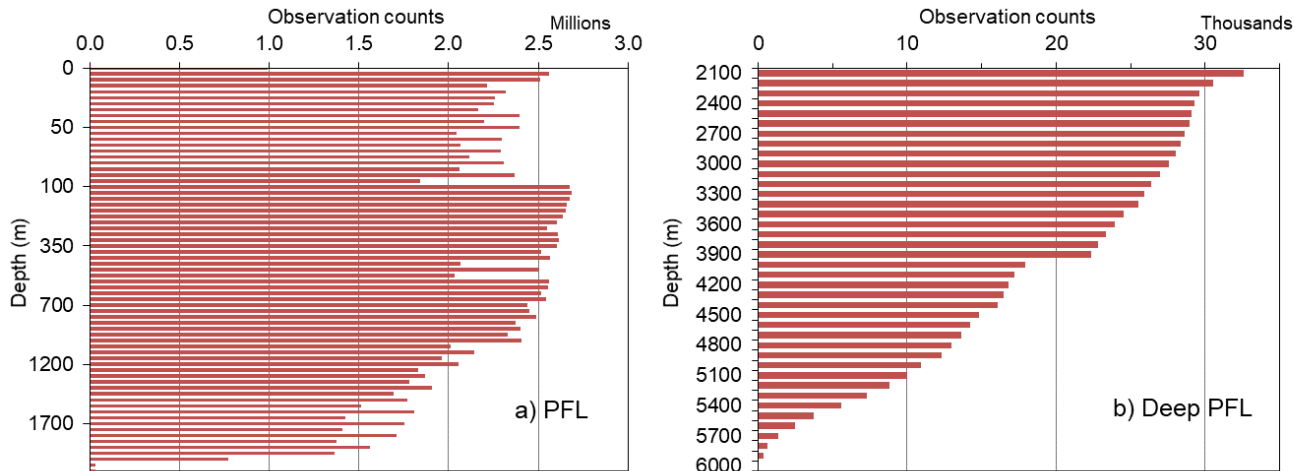


Figure 6.5. Distribution of Profiling Float Data (PFL) data at standard depth levels: a) between 0 and 2000m and b) between 2100 and 7000m depth in WOD23.

## 6.9. RELEVANT WEB SITES

All websites last accessed: 2024-04-26

Aanderaa: <http://www.aanderaa.com/index.php>

Argo: <http://www.argo.ucsd.edu>.

Argo Information Center: <https://www.aoml.noaa.gov/argo/> .

BioGeochemical Argo: <https://biogeochemical-argo.org/>

Deep Argo: <https://argo.ucsd.edu/expansion/deep-argo-mission/>

FSI Scientific, Inc: <http://www.falmouth.com/>

Sea-Bird Scientific: <http://www.seabird.com/>

TSK the Tsurumi-Seiki Co., Ltd.: <https://tsurumi-seiki.co.jp/en/>

## 6.10. REFERENCES AND BIBLIOGRAPHY

- Ando, K., K. Izawa, K. Mizuno, S. Hosoda, A. Inoue, T. Kobayashi, and N. Shikama (2003). *Results of field experiments and laboratory tests of domestic profiling floats (NINJA)*, JAMSTEC, 48, 55-65 (in Japanese).<sup>1</sup>
- Barker, P.M., J.R. Dunn, C.M. Domingues, and S.E. Wijffels (2011). Pressure sensor drifts in Argo and their impacts. *J. Atmos. Oceanic Technol.*, 28, 1036-1049, doi: 10.1175/2011JTECHO831.1.
- Boss, E.D. Swift, L. Taylor, P. Brickley, R. Zaneveld, S. Riser, M.J. Perry, and P.G. Strutton (2008). Observations of Pigment and Particle Distributions in the Western North Atlantic from an Autonomous Float and Ocean Satellite. *Limnology and Oceanography*, 53, 2112-2122, doi:10.4319/lo.2008.53.5\_part\_2.2112.

- Bittig, H. C. and A. Körtzinger (2015). Tackling Oxygen Optode Drift: Near-Surface and In-Air Oxygen Optode Measurements on a Float Provide an Accurate in Situ Reference. *J. Atmos. Oceanic Technol.*, 32, 1536-1543.
- Bittig, H. C., Körtzinger, A., Neill, C., Van Ooijen, E., Plant, J.N., Hahn, J., & Emerson, S.R. (2018). Oxygen optode sensors: principle, characterization, calibration, and application in the ocean. *Frontiers in Marine Science*, 4, 429.
- Böhme, L. and U. Send (2005). Objective Analyses of Hydrographic Data for Referencing Profiling Float Salinities in Highly Variable Environments. *Deep-Sea Res. II*, 52, 651-664.
- Cabanes, C., Thierry, V., & Lagadec, C. (2016). Improvement of bias detection in Argo float conductivity sensors and its application in the North Atlantic. *Deep Sea Research I: Oceanographic Research Papers*, 114, 128-136.
- Claustre, H., J. Bishop, E. Boss, S. Bernand, J.-F. Berthon, C. Coatanoan, K. Johnson, A. Lotiker, O. Ulloa, M.-J. Perry, F. D’Ortenzio, O. Hembise Fanton D’Andon, J. Uitz (2010). Bio-optical profiling floats as new observational tools for biogeochemical and ecosystem studies. *Proceedings of the “OceanObs’09: Sustained Ocean Observations and Information for Society” Conference*, Venice, Italy, 2009, J. Hall, D. E. Harrison, and D. Stammer, Eds., doi:10.5270/OceanObs09.cwp.17.
- Davis, R.E., D.C. Webb, L.A. Reiger, and J. Dufour (1992). The Autonomous Lagrangian Circulation Explorer (ALACE). *J. Atmos. Oceanic Technol.*, 9, 264-285.
- Davis, R. E., & Zenk, W. (2001). Subsurface Lagrangian observations during the 1990s. In *International Geophysics* (Vol. 77, pp. 123-139). Academic Press.
- Davis, R.E., J.T. Sherman, and J. Dufour (2001). Profiling ALACEs and Other Advances in Autonomous Subsurface Floats. *J. Atmos. Oceanic Technol.*, 18, 982-993.
- Durand, F. and G. Reverdin (2005). A Statistical Method for Correcting Salinity Observations from Autonomous Profiling Floats: An ARGO Perspective. *J. Atmos. Oceanic Tech.*, 22, 292-301.
- Garcia, H.E., Z. Wang, C. Bouchard, S.L. Cross, C.R. Paver, J.R. Reagan, T.P. Boyer, R.A. Locarnini, A.V. Mishonov, O. Baranova, D. Seidov, and D. Dukhovskoy (2024b). World Ocean Atlas 2023, Volume 3: Dissolved Oxygen, Apparent Oxygen Utilization, and Oxygen Saturation. A. Mishonov, Tech. Ed. *NOAA Atlas NESDIS 91*.
- Gould, W.J. (2005). From Swallow floats to Argo- the development of neutrally buoyant floats. *Deep-Sea Res. II*, 52, 529-543.
- Gouretski, V., Cheng, L., Du, J., Xing, X., & Chai, F. (2024). A consistent ocean oxygen profile dataset with new quality control and bias assessment. *Earth System Science Data Discussions*, 2024, 1-27 (under review).
- Johnson, K.S., L.J. Coletti, H.W. Jannasch, C.M. Sakamoto, D.D. Swift, and S.C. Riser (2013). Long-Term Nitrate Measurements in the Ocean Using the in situ Ultraviolet Spectrophotometer: Sensor Integration into the APEX Profiling Float. *J. Atmos. Oceanic Technol.*, 30, 1854-1866, doi:10.1175/JTECH-D-12-00221.1.



- Johnson, K.S., H.W. Jannasch, L.J. Coletti, V.A. Elrod, T.R. Martz, Y. Takeshita, R.J. Carlson, and J.G. Connery (2016). Deep-Sea DuraFET: A Pressure Tolerant pH Sensor Designed for Global Sensor Networks. *Analytical Chemistry*, 88, 3249-3256, doi:10.1021/acs.analchem.5b04653.
- Johnson, K.S., J.N. Plant, J.P. Dunne, L.D. Talley, and J.L. Sarmiento (2017). Annual Nitrate Drawdown Observed by SOCCOM Profiling Floats and the Relationship to Annual Net Community Production. *J. Geophys. Res. Oceans*, 122, 6668-6683, doi:10.1002/2017JC012839.
- Körtzinger, A., J. Schimanski, and U. Send (2005). High-Quality Oxygen Measurements from Profiling Floats: A Promising New Technique. *J. Atmos. Oceanic Technol.*, 22, 302-308.
- Le Reste, S., V. Dutreuil, X. Andre, V. Thierry, C. Renaut, P.-Y. Le Traon, and G. Maze (2016). "Deep-Arvor": A New Profiling Float to Extend the Argo Observations Down to 4000-m Depth. *J. Atmos. Oceanic Technol.*, 33, 1039-1055, doi:10.1175/JTECH-D-0214.1.
- Loaec, G., N. Cortes, M. Menzel, and J. Moliera (1998). PROVOR: A Hydrographic Profiler Based on MARVOR Technology. *Proceedings, IEEE-Oceans '98*, Nice, France.
- Oka, E. (2005). Long-term Sensor Drift Found in Recovered Argo Profiling Floats, *J. Oceanogr.*, 61, 775-781.
- Ollitrait, M., N. Cortes, G. Loaec, and J.P. Rannou (1994). MARVOR float present results from the SAMBA experiment. *Proceedings, IEEE-Oceans '94*, Brest, France.
- Owens, W. B. and A. Wong (2009). An improved calibration method for the drift of the conductivity sensor on autonomous CTD profiling floats by  $\theta$ -S climatology. *Deep-Sea Res. I*, 56, 450-457.
- Roemmich, D., S. Riser, R. Davis, and Y. Desaubies (2004). Autonomous Profiling Floats: Workhorse for Broad-scale Ocean Observations. *Mar. Tech. Soc. J.*, 38, 31-39.
- Rossby, T. and D. Webb (1970). Observing abyssal motions by tracking Swallow floats in the SOFAR Channel. *Deep-Sea Res.*, 17, 359-365.
- Rossby, T., D. Dorson, and J. Fontaine (1986). The RAFOS System. *J. Atmos. Oceanic Technol.*, 3, 672-679.
- Stommel, H. (1957). A survey of ocean current theory. *Deep-Sea Res.*, 4, 149-184.
- Swallow, J.C. (1955). A neutral-buoyancy float for measuring deep currents. *Deep-Sea Res.*, 3, 74-81.
- Swallow, J.C. and L.V. Worthington (1961). An observation of a deep countercurrent in the Western North Atlantic. *Deep-Sea Res.*, 8, 1-19.
- Takeshita, Y., Martz, T.R., Johnson, K.S., Plant, J.N., Gilbert, D., Riser, S.C., & Tilbrook, B. (2013). A climatology-based quality control procedure for profiling float oxygen data. *Journal of Geophysical Research: Oceans*, 118(10), 5640-5650.
- Thierry, V. H. Bittig, and the Argo-BGC team (2021). Argo Quality Control Manual for Dissolved Oxygen Concentration, v2.1 <http://dx.doi.org/10.13155/46542> .

- Webb, D.C. and M.J. Tucker (1970). Transmission Characteristics of the SOFAR Channel. *The J. of the Acoustical Soc. of America*, 48, 767-769.
- Wong, A.P.S., G.C. Johnson, and W.B. Owens (2003). Delayed-Mode Calibration of Autonomous CTD Profiling Float Salinity Data by  $\theta$ -S Climatology. *J. Atmos. Oceanic Technol.*, 20, 308-318.
- Wong, A. P., Gilson, J., & Cabanes, C. (2023). Argo salinity: bias and uncertainty evaluation. *Earth System Science Data Discussions*, 15, 383-393.
- Wong, A., R. Keeley, T. Carval, and the Argo Data Management Team (2024). Argo Quality Control Manual for CTD and Trajectory Data. <http://dx.doi.org/10.13155/33951>
- Xing, X.A., A. Morel, H. Claustre, D. Antoine, F. D'Ortenzio, A. Poteau, and A. Mignot (2017). Combined Processing and Mutual Interpretation of Radiometry and Fluorimetry from Autonomous Profiling Bio-Argo Floats: Chlorophyll a Retrieval. *J. Geophys. Res. Oceans*, 116, C06020, doi:10.1029/2010JC006899.

---

<sup>1</sup>English version of Argo Information Center Newsletter available on their website (see above).

# CHAPTER 7: MECHANICAL BATHYTHERMOGRAPH DATA (MBT)

*Courtney N. Bouchard, Alexey V. Mishonov, Tim P. Boyer, Ricardo A. Locarnini*

*Ocean Climate Laboratory  
National Centers for Environmental Information  
Silver Spring, MD*

## **7.1. INTRODUCTION**

The Mechanical Bathythermograph (MBT) is an instrument developed during the late-1930's (Spilhaus, 1938) that can be dropped from either a stationary or moving surface ship to produce an upper ocean temperature profile. This instrument was a substantial improvement of an instrument known as the "oceanograph" which was designed by Dr. Carl Rossby and Dr. Karl Lange (Rossby and Montgomery, 1934) for studying the upper ocean thermal structure. The introduction of the MBT allowed ships to make synoptic surveys of oceanographic regions and for discovery of fine structure of the ocean's thermal structure. Spilhaus (1941) used the instrument to identify "fine" structure (in the horizontal) from temperature profiles near the edge of the Gulf Stream. Pressure is determined from a pressure sensitive tube known as a Bourdon tube. A temperature sensitive element in the nose of the MBT enables the instrument to trace temperature as a function of depth. The wide-spread model of MBT became the one much improved later by Ally Vine and Mike Ewing in WHOI.

Different varieties of the MBT have different maximum depth ranges with 295 m being the deepest depth measured from any version of the instrument. Earlier versions of the instrument were limited to making measurements in the upper 140 m of the water column. Spilhaus (1987) gave a review of the development of the MBT. Couper and LaFond (1970) provide another more comprehensive review.

In most countries and institutions, the use of the MBT has been replaced by the XBT. Only 1.6% of all the MBT profiles in our archives were collected between 1991 and 2000 (Table 7.1). The MBT data in WOD23 does not extend past year 2000 with only a few historical profiles added in the late 1990s and 2000.

## **7.2. MBT ACCURACY**

The accuracy of the MBT has been the subject of several studies. Leipper and Burt (1948) report the results of comparisons between MBT temperature measurements and near simultaneous reversing thermometer measurements, which were made by D. Pritchard of the U.S. Navy Electronics Laboratory in Lake Meade. By comparing the temperature traces on the up and down casts of the MBT it was inferred that there was "an almost complete absence of internal

waves of large amplitude and short period, hysteresis of the instruments, or rapid temperature changes due to advection”. These results are reproduced in Table 7.2 given below. Clearly, there is good agreement between the reversing thermometer measurements (which typically had an accuracy of 0.02°C at this period of time) and the MBT measurements. However, there is a problem with interpreting the results from Table 7.2 because it is not clearly stated in the table, or the text of the technical report of Leipper and Burt, what temperature units were used. Throughout their report, Leipper and Burt use the Fahrenheit scale. If this scale applies to the results in Table 7.2, then the agreement is impressive. If the results are in degrees Celsius, the agreement is less impressive, but the data are still useful for many scientific purposes. We presume that ‘Shallow’ in Table 7.1 is referring to the instrument with 140m depth range and ‘Deep’ refer to the instrument with 295m depth range. Other studies attribute an accuracy of about 0.5°F to the MBT instrument. This figure is comparable to the accuracy of expendable bathythermograph (XBT) probes for which the thermistor sensing element is not calibrated (Tabata, 1978). For example, individual XBT probes are accurate to a few tenths of a degree Celsius. Although both MBT and XBT probes are an order of magnitude less precise than reversing thermometers, the standard error of the mean of any estimate based on these temperature measurements decreases with the increase in number of data used. This applies to random errors. Hence, historical bathythermograph measurements provide valuable information when estimating global-scale features by averaging over many measurements in space and/or time.

**Table 7.1. Comparison of observations taken with Mechanical Bathythermographs and reversing thermometers.**

Reproduced from D. Leipper and W. Burt (1948).

<b>TABLE 2.3. Observations taken with bathythermographs and reversing thermometers</b>			
<b>BT</b>	<b>No. of stations</b>	<b>No. of thermometer observations</b>	<b>Standard Deviation of Temperature Differences*</b>
# 1784A (Shallow)	9	20	0.15
# 1258A (Deep)	10	41	0.19
# 514A (Deep)	12	36	0.10

We reproduce this table as it appeared in the work by Leipper and Burt (1948).

### **7.3. SURFACE DATA ACQUIRED CONCURRENTLY WITH MBT CASTS**

On occasions, a sea-surface water sample is taken at the time of the MBT cast. Temperature and salinity of the water sample are usually measured and recorded as ancillary information of the MBT cast. Meteorological conditions at the time of the MBT cast could also be archived, e.g. air temperature, wind speed and direction, cloud type and cover, barometric atmospheric pressure, as well as sea conditions: wave height and direction, sea state.

A significant amount of ancillary meteorological information was recovered by the NODC/OCL through the digitization of historical MBT cards from the Scripps Institution of Oceanography and the Woods Hole Oceanographic Institution.

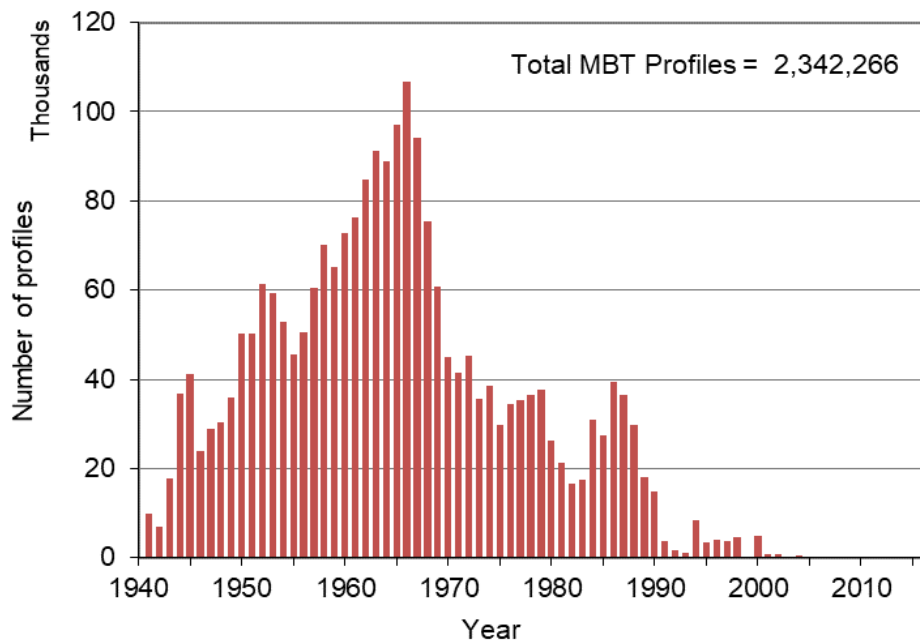


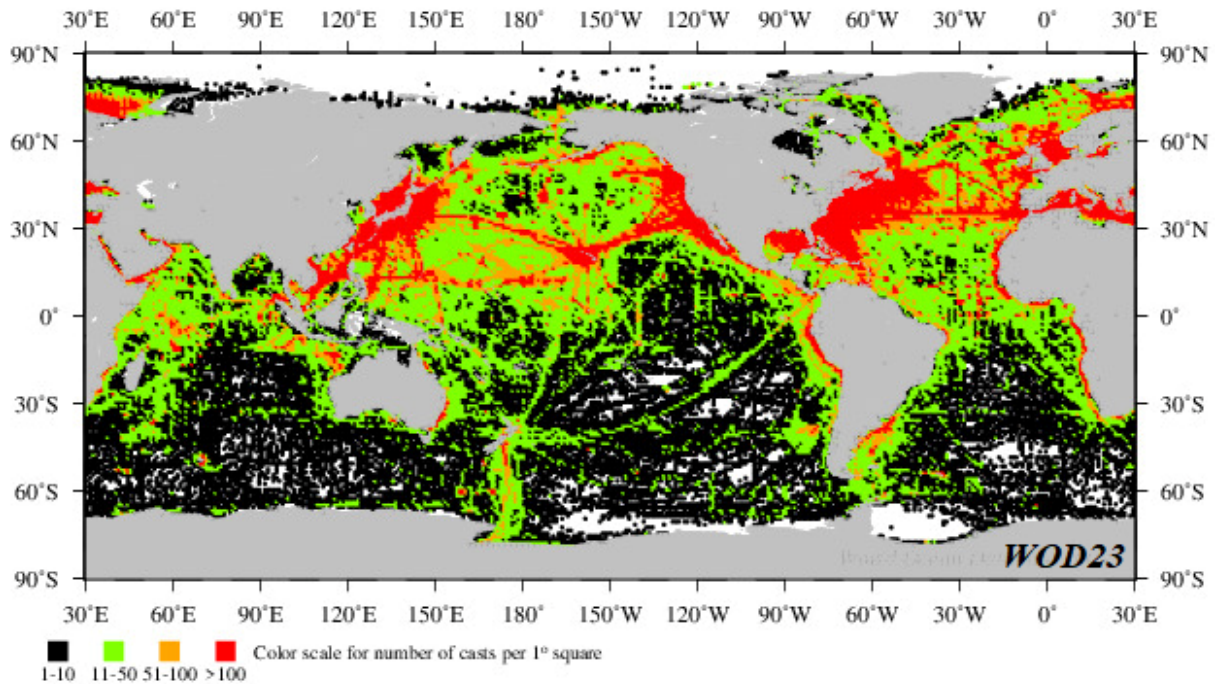
Figure 7.1. Temporal distribution of Mechanical Bathythermograph (MBT) profiles in WOD23.

#### 7.4. MBT PROFILE DISTRIBUTION

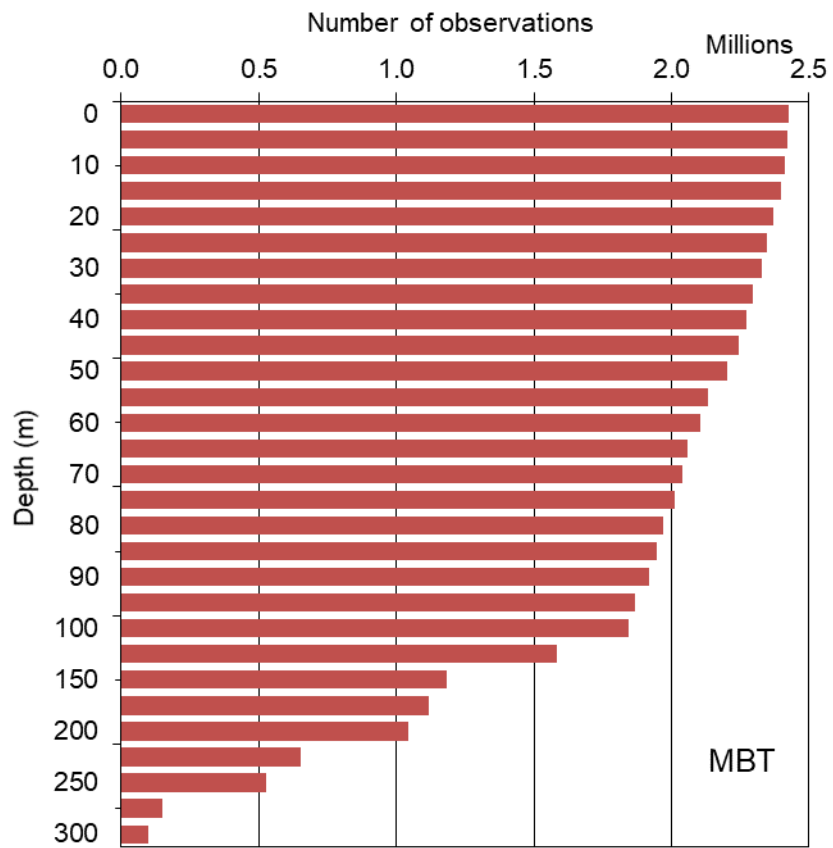
Table 7.1 gives the yearly counts of MBT profiles for the World Ocean and Figure 7.1 shows the time series of those yearly totals. There is a total of 2,339,415 MBT profiles for the entire World Ocean with only about 11% measured in the southern hemisphere and 89% profiles measured in the northern hemisphere (Figure 7.2). MBT data (including DBT and micro-BT data, see Chapters 8 and 13 respectively) represents 13% of entire WOD23 data collection. Figure 7.3 represents distribution of Mechanical Bathythermograph (MBT) data at standard depth levels. Table 7.3 gives national contributions of MBT profiles. As Table 7.2. demonstrate, there is no new MBT data collected after year 2000. There however a possibility exists that some historic MBT data is still not declassified or digitized and resides in different archives. If and when such data will become available to us they will be added to WOD collection.

**Table 7.2. Number of all MBT profiles as a function of year in WOD23.**  
**Total Number of Profiles = 2,339,415**

YEAR	COUNT	YEAR	COUNT	YEAR	COUNT
1941	9,990	1962	84,853	1983	17,481
1942	7,014	1963	91,213	1984	30,910
1943	17,767	1964	88,951	1985	27,285
1944	36,785	1965	97,112	1986	39,504
1945	41,086	1966	106,585	1987	36,525
1946	23,822	1967	94,162	1988	29,684
1947	28,808	1968	75,372	1989	18,024
1948	30,304	1969	60,830	1990	14,921
1949	36,038	1970	44,886	1991	3,829
1950	50,296	1971	41,386	1992	1,554
1951	50,248	1972	45,353	1993	1,169
1952	61,310	1973	35,533	1994	8,417
1953	59,341	1974	38,586	1995	3,453
1954	52,914	1975	29,808	1996	3,905
1955	45,467	1976	34,360	1997	3,859
1956	50,510	1977	35,455	1998	4,552
1957	60,424	1978	36,608	1999	0
1958	70,102	1979	37,806	2000	4,819
1959	65,214	1980	26,362		
1960	72,656	1981	21,182		
1961	76,417	1982	16,608		



**Figure 7.2. Geographic distribution of Mechanical Bathythermograph (MBT) profiles in WOD23.**



**Figure 7.3. Distribution of Mechanical Bathythermograph (MBT) data at standard depth levels in WOD23.**

**Table 7.3. National contributions of Mechanical Bathythermograph profiles in WOD23.**

<b>ISO<sup>1</sup> Country Codes</b>	<b>Country name</b>	<b>MBT Casts</b>	<b>% of Total</b>
US	United States	1,169,301	49.96
SU	Union of Soviet Socialist Republic	449,993	19.2
JP	Japan	296,410	12.7
CA	Canada	184,845	7.9
GB	Great Britain	118,638	5.1
DE	Germany	25,005	1.1
AU	Australia	18,376	0.8
99	Unknown/International	16,386	0.7
FR	France	13,538	0.6
AR	Argentina	10,995	0.5
NL	Netherland	8,088	0.3
IT	Italy	6,267	0.3
PE	Peru	5,212	0.2
CL	Chile	4,161	0.2
PT	Portugal	2,628	0.1
NZ	New Zealand	2,435	0.1
CD	Congo, the Democratic Republic	1,234	0.1
BE	Belgium	1,218	0.1
NO	Norway	913	<0.1
EC	Ecuador	885	<0.1
CO	Colombia	747	<0.1
VE	Uruguay	673	<0.1
IN	India	540	<0.1
MG	Madagascar	405	<0.1
GR	Greece	327	<0.1
SN	Spain	245	<0.1
ES	Estonia	195	<0.1
SL	Russian Federation	187	<0.1
CI	Cote D'Ivoire	99	<0.1
MC	Monaco	97	<0.1
NG	Nigeria	89	<0.1
BR	Brazil	82	<0.1
TH	Thailand	77	<0.1
ZA	South Africa	20	<0.1
GH	Ghana	12	<0.1

<sup>1</sup>ISO = [International Organization for Standardization](http://www.iso.org)



## 7.5. REFERENCES AND BIBLIOGRAPHY

- Bralove, A.L. and E.I. Williams, Jr. (1952). A study of the errors of the bathythermograph. Final Report. National Scientific Laboratories Inc., *Contract No. NObsr 52348*, 49 pp.
- Cascviano, D.L. (1967). Calibration Monitoring of Mechanical Bathythermographs, GMT, Dec / Jan 1966-67, 19-21.
- Couper, B.K. and E.C. LaFond (1970). Mechanical Bathythermograph: An Historical Review. In *Advances in Instrumentation*, Paper 735-70, Instrument Society of America, 25, Part 3, pp 735-70.
- Dinkel, C.R. and M. Stawnychy, (1973). Reliability Study of Mechanical Bathythermographs, *Mar. Tech. Soc. J.*, 7(3), 41-47.
- Gouretski, V. and K.P. Koltermann (2007). How much is the ocean really warming? *Geophys. Res. Lett.*, 34, L01610, 10.1029/2006GL027834.
- Hazelworth, J.B. (1966). Quantitative analysis of some bathythermograph errors. *Technical Report ASWEPS No. 11*, U.S. Naval Oceanogr. Off., pp. 27.
- IOC (1975). Guide to oceanographic and marine meteorological instruments and observing practices. UNESCO, Paris, 5 pp. and 12 chapters.
- Leipper, D.F. and R.M. Adams (1952). Some methods used in representing bathythermograph data. The A&M College of Texas, Dept. of Oceanogr., *Tech. Rep. 1*, 6 pp., 9 figs.
- Leipper, D.F., R.M. Adams, and Project staff (1952). Summary of North Atlantic Weather Station Bathythermograph data 1946-1950. The A&M College of Texas, Dept. of Oceanogr., *Tech. Rep. 3*, 2 pp., 40 figs.
- Leipper, D.F. and Project staff (1954). Summary of North Pacific Weather Station Bathythermograph data 1943-1952. The A&M College of Texas, Dept. of Oceanogr., *Tech. Rep. 7*, 2 pp., 64 figs.
- Leipper, D.F. and W.V. Burt (1948). Annual Report, 1947-48. Bathythermograph Processing Unit. Scripps Inst. of Oceanogr., *Oceanography Rep. No. 15*, Scripps Inst. of Oceanogr., La Jolla, CA, 78 pp.
- Leipper, D.F. and W.V. Burt (1948). Bathythermograph processing unit. Annual report to 1 July 1948. *Oceanographic report No. 15*. 1948. Scripps Inst. of Oceanogr., La Jolla, CA, 59p. SIO Reference 48-14.
- Levitus, S., R. Gelfeld, T. Boyer, and D. Johnson (1994). Results of the NODC and IOC Data Archaeology and Rescue projects. Key to Oceanographic Records Documentation No. 19, National Oceanographic Data Center, Wash., D.C., 67 pp.
- Levitus, S., M. Conkright Gregg, T.P. Boyer, R. Gelfeld, L. Stathoplos, D. Johnson, I. Smolyar, C. Stephens, G. Trammell, R. Moffatt, and T. O'Brien (1998). Results of the IOC Global Oceanographic Data Archaeology and Rescue (GODAR) project. *NOAA NESDIS Technical Report*.
- Levitus, S., S. Sato, C. Maillard, N. Mikhailov, P. Caldwell, and H. Dooley (2005). Building Ocean Profile-Plankton Databases for Climate and Ecosystem Research. *NOAA Technical Report NESDIS 117*, U.S. Gov. Printing Office, Wash., D.C., 29 pp.
- Levitus, S., J.I. Antonov, T.P. Boyer, R. A. Locarnini, H.E. Garcia, and A.V. Mishonov (2009). Global ocean heat content 1955-2008 in light of recently revealed instrumentation problems. *Geophys. Res. Lett.*, 36, L07608, doi: 10.1029/2008GL037155.

- NODC (1966). Atlas of bathythermograph data, Indian Ocean. U.S. Naval Oceanographic Office, *NODC Publication G6*, 129 pp.
- Robinson, M.K. and E.M. Drollinger (1969). Bibliography of reports based on bathythermograph temperature data, *SIO Reference Series 69-16*, pp. 104.
- Rosby, C-G. and R.B. Montgomery (1934). The layer of frictional influence in wind and ocean currents, in "Papers in Physical Oceanography and Meteorology of the Massachusetts Institute of Technology and Woods Hole Oceanographic Institution", Vol. III, No. 3, pp. 73.
- Smed, J. (1978). Inventory of Oceanographic Investigations at North Atlantic Ocean Weather Stations 1947-1962. ICES, Charlottenlund, Denmark, 63 pp.
- Spilhaus, A.F. (1938). A bathythermograph. *J. Mar. Res.*, 1, 95-100.
- Spilhaus, A.F. (1941). Fine structures on the edge of the Gulf Stream. *EOS, Transactions*, Amer. Geophys. Union, 22, 478-484.
- Spilhaus, A.F. (1987). On Reaching 50: [An Early History of the Bathythermograph](#), *Sea Tech.*, 28, 19-28.
- Stewart, R.L. (1963). Test and Evaluation of the Mechanical Bathythermograph, Unpublished manuscript, *Mar. Sci. Dept.*, U.S. Naval Oceanogr. Office, 33 pp.
- Tabata, S. (1978). Comparison of observations of sea surface temperatures at Ocean Weather Station P and NOAA Buoy Stations and those made by merchant ships traveling in their vicinities, in the Northeast Pacific Ocean. *J. Applied Meteorol.* 17, 374-385.
- U.S. Naval Oceanographic Office (1968). Instruction Manual for Obtaining Oceanographic Data, *Publication 607, Sup. of Documents*, Wash., D.C.
- U.S. Weather Bureau (1956). Ocean Station Vessel Meteorological Records Survey: Atlantic and Pacific. U.S. Gov. printing Office, U.S. Gov. Printing Office, Wash., D.C., 106 pp.
- Vine, A.C. (1952). Oceanographic Instruments for Measuring Temperature, in Symposium on Oceanographic Instrumentation, Rancho Santa Fe, California.

# CHAPTER 8: DIGITAL BATHY THERMOGRAPH (DBT) PROFILES

*Courtney N. Bouchard, Alexey V. Mishonov, Tim P. Boyer, Ricardo A. Locarnini*

*Ocean Climate Laboratory  
National Centers for Environmental Information  
Silver Spring, MD*

## **8.1. INTRODUCTION**

The Digital Bathythermograph (DBT) is an instrument developed to record and report temperature profile data electronically. The self-contained underwater instrument includes a thermistor and a strain gauge. Temperature and depth/pressure measurements are automatically recorded in the underwater unit as it is lowered in the water column. Upon recovery, the underwater unit is connected to a computer to retrieve the data.

All DBT profiles are stored in the MBT dataset of WOD23.

It should be noted that the DBT instrument is obsolete now and not in use anymore, therefore, there is no new DBT data in WOD23 compared to WOD18.

## **8.2. DBT ACCURACY**

The DBT has a temperature accuracy of  $\pm 0.05^{\circ}\text{C}$ . However, Pankajakshan *et al.* (2003) report temperature errors of  $-0.3^{\circ}\text{C}$  to  $+1.0^{\circ}\text{C}$  in Indian DBT data from the Indian Ocean. No errors were observed in DBT data collected in the Pacific Ocean by Japanese and United States institutions.

## **8.3. DBT PROFILE DISTRIBUTIONS**

Table 8.1 gives the yearly counts of DBT profiles for the World Ocean. Figure 8.1 shows the time series of the yearly totals of Digital Bathythermograph profiles for the World Ocean. There is a total of 80,200 DBT profiles for the entire World Ocean with about 6.0% measured in the southern hemisphere and 94.0% profiles measured in the northern hemisphere. Table 8.2 gives national contributions of DBT data. Figure 8.3 illustrate distribution of Digital Bathythermograph (DBT) data at standard depth levels in WOD23.

**Table 8.1. The number of Digital Bathythermograph (DBT) profiles as a function of year in WOD23.**

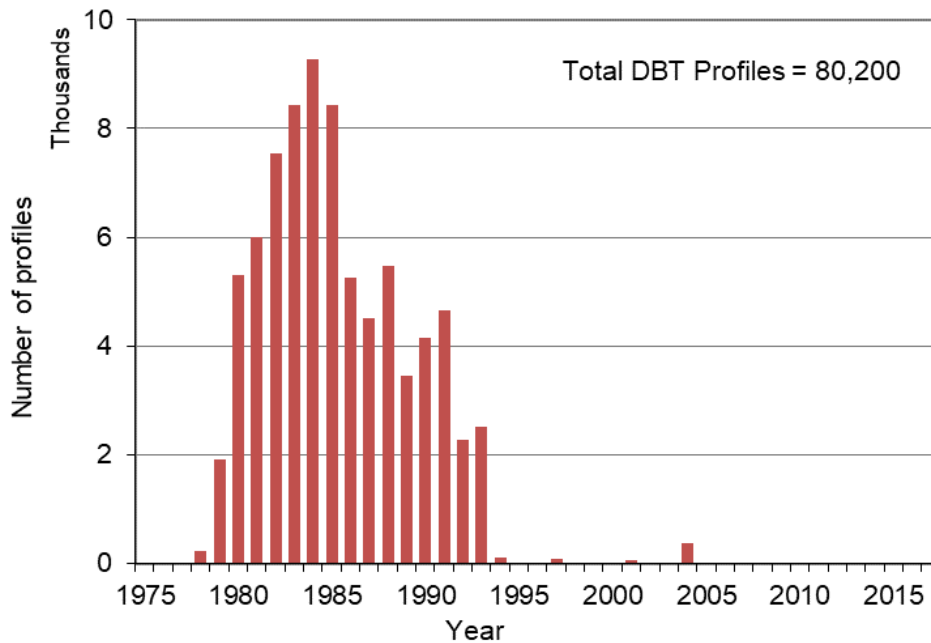
The total number of casts = 80,200.

YEAR	CASTS	YEAR	CASTS	YEAR	CASTS	YEAR	CASTS
1977	27	1984	9,271	1991	4,662	1998	0
1978	234	1985	8,427	1992	2,285	1999	0
1979	1,926	1986	5,255	1993	2,507	2000	1
1980	5,310	1987	4,505	1994	121	2001	73
1981	6,000	1988	5,478	1995	2	2002	19
1982	7,539	1989	3,443	1996	27	2003	23
1983	8,440	1990	4,148	1997	88	2004	389

**Table 8.2. National contributions of Digital Bathythermograph (DBT) profiles in WOD23.**

ISO <sup>1</sup> Country Codes	Country Name	DRB Casts	% of Total
JP	Japan	69,098	86.2
CA	Canada	11,102	13.8
	<i>Total</i>	<i>80,200</i>	<i>100.00</i>

<sup>1</sup>ISO = [International Organization for Standardization](http://www.iso.org)



**Figure 8.1. Temporal distribution of Digital Bathythermograph (DBT) profiles in WOD23**

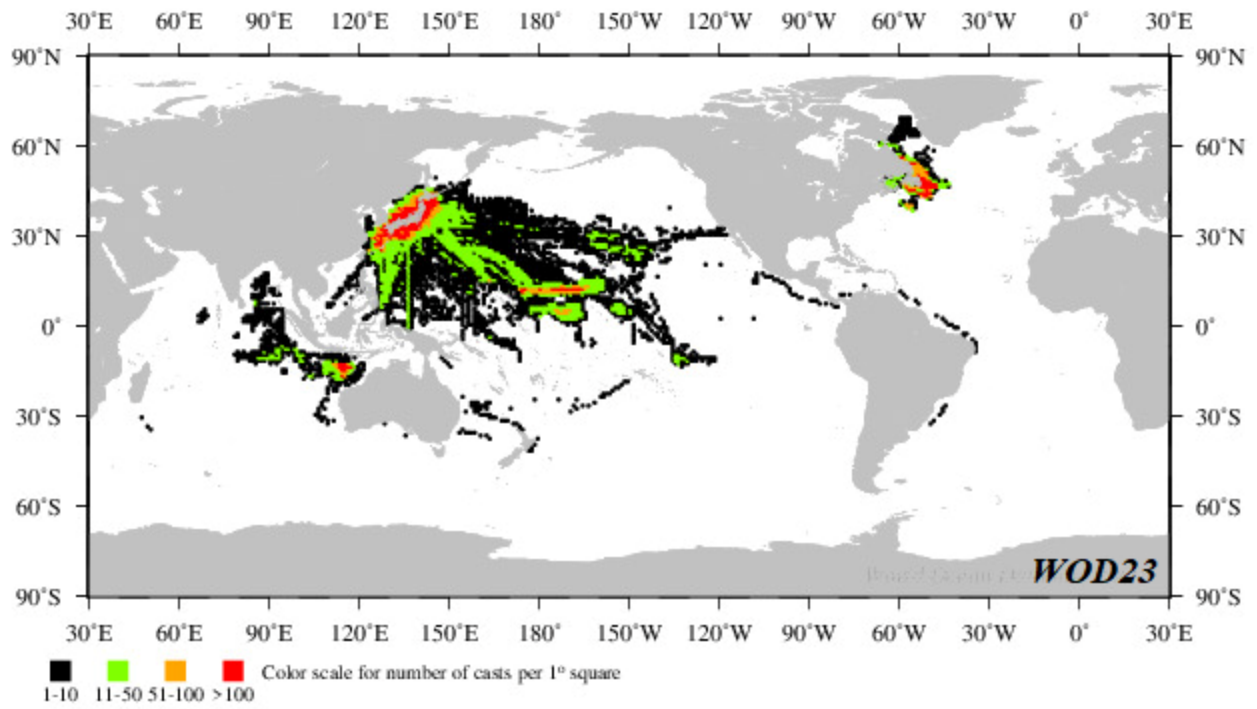


Figure 8.2. Geographic distribution of Digital Bathythermograph profiles in WOD23

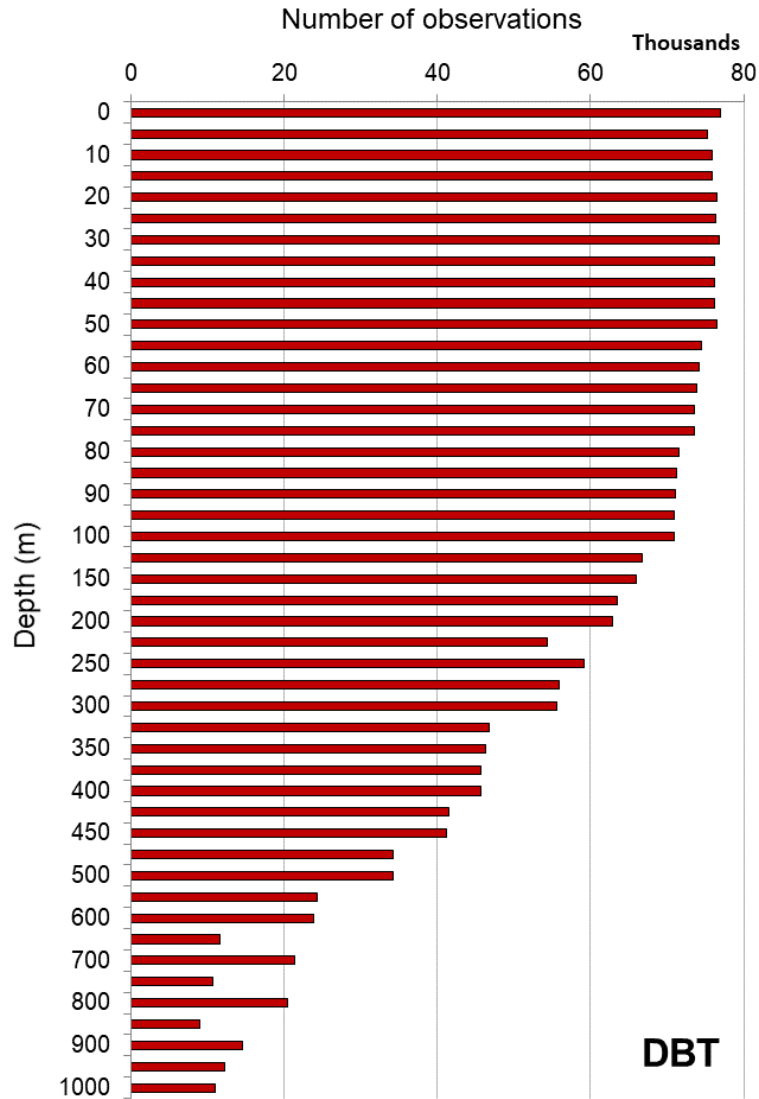


Figure 8.3. Distribution of Digital Bathythermograph data at standard depth levels in WOD23.

#### 8.4. REFERENCES AND BIBLIOGRAPHY

Pankajakshan T., G.V. Reddy, L. Ratnakaran, J.S. Sarupria, and V.R. Babu (2003). Temperature error in digital bathythermograph data. *Indian J. Mar. Sci.*, 32, 234-240

# CHAPTER 9: MOORED BUOY DATA (MRB)

*Alexey V. Mishonov, Scott L. Cross, Tim P. Boyer, Ricardo A. Locarnini, Dan Seidov*

*Ocean Climate Laboratory  
National Centers for Environmental Information  
Silver Spring, MD*

## 9.1. INTRODUCTION

As the National Data Buoy Center website (<https://www.ndbc.noaa.gov/>) reports, “In March 1966, the Panel on Ocean Engineering of the Interagency Committee on Oceanography convened a group of Federal agency representatives to address the problems and possibilities associated with automated data buoy networks. This group recommended a national system of ocean data buoys and the Committee asked the United States Coast Guard to conduct a feasibility study of a consolidated national data buoy system”. After ten months of work, the study report made the following conclusions:

- extensive requirements exist for oceanographic and meteorological information to satisfy both operational and research needs in the oceanic and Great Lakes environments;
- automatic, moored buoys were capable of meeting a significant portion of those needs; and that
- a network of such buoys, would be an essential element of an overall environmental information and prediction system (Shea, 1987).

As further explained in the U.S. Department of Commerce’s publication NDBCM WO547, “The National Data Buoy Project (NDBP) was established in December 1967 for the purpose of developing a national capability to deploy and operate networks of automatic buoys to retrieve useful information describing the marine environment on a reliable, real time basis”. As noted by Shea (1987) in “A History of NOAA” – “By the 1960’s, scientists had recognized the need for more detailed information on environmental conditions over vast marine areas which remained largely uncovered except for occasional observations from ships or aircraft of opportunity, oceanographic research expeditions, or the few existing ocean station vessels. As a result, a number of Federal Agencies and universities began programs to develop and implement networks of buoys which could routinely and automatically report environmental conditions like temperature, wind speed and direction, etc.”

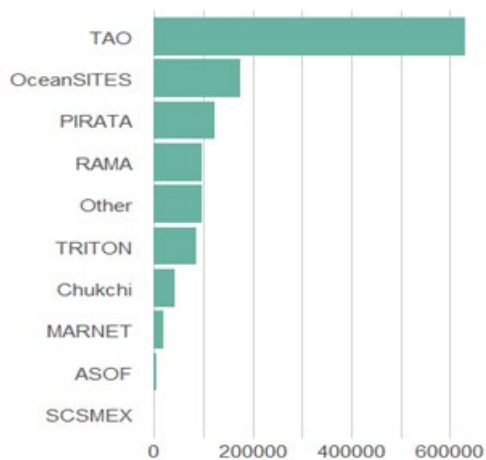
The Data Buoy Cooperation Panel [website](http://www.jcommops.org/dbcp/network/maps.html) describes moored buoys as “normally relatively large and expensive platforms. If a moored buoy goes adrift it represents a potential loss of costly equipment and a possible hazard to navigation. For these reasons the ARGOS system has been used for location determination for moored buoys. In addition, some World Meteorological Organization (WMO) Member countries use the ARGOS system for normal transmission of meteorological observations from moored buoys” (see <http://www.jcommops.org/dbcp/network/maps.html>).

The WOD23 MRB dataset contains data on daily averaged values of water temperature and salinity collected by sensors located on moored buoys (MRB) during the period from November 3, 1977 to December 31, 2022. The dataset contains a total of 1,277,591 profiles, which reflects both additions to, as well as some subtractions from the WOD18.

A significant new source of MRB profiles in WOD23 is the [OceanSITES collection](#). OceanSITES is an international network of deep-ocean observing systems, including many of programs listed above, as well as a number of other, smaller programs. Net of any duplicates downloaded from other sources, OceanSITES contributed 176,456 MRB profiles to WOD23.

Some MRB profiles have been removed since the WOD18 version as well. In particular, data submitted under NCEI Accession 0063240, from Baffin Bay, were originally included in WOD18 at full temporal (i.e. 15-minute) resolution. These data were converted to daily averages in 2021, and it is these averaged data that are included in WOD23. This has resulted in a reduction of 375,247 profiles.

Together, these programs have contributed over 1.1 million profiles, over 80% of the total in the database (See Figure 9.1, and related web-links below for additional information). Please note that sampling frequencies among the observational programs differ; the number of profiles depends not only on the number of instruments and the length of the record but also the rate at which the instrument samples, and if the data is averaged or post-processed.



**Figure 9.1. Numbers of MRB daily profiles by observing program**

**Major moored-buoy programs.** As part of the Tropical Ocean-Global Atmosphere ([TOGA](#)) program, efforts were made to enhance the real-time ocean observing system in the tropical Pacific Ocean. The Tropical Atmosphere Ocean ([TAO](#)) array of moored buoys spans the tropical Pacific from 137°E to 95°W and from 8°S to 8°N. The TAO system began in 1985 as a regional-scale set of meridional arrays on both sides of the Equator at 110°W and 165°E and has expanded over the years to its present size of approximately 49 active moorings.

The buoys in the TAO array are typically separated by 2-3 degrees of latitude and 10-15 degrees of longitude. The array provides surface wind, rain rates, sea surface temperature (SST), upper ocean temperature, as well as subsurface temperatures and salinity down to a depth of 500 meters, and current measurements (Mangum, 1994; Mangum *et al.*, 1994; McPhaden, 1995; McPhaden *et al.*, 1998). The majority of TAO moorings are ATLAS moorings developed at NOAA's Pacific Marine Environmental Laboratory (PMEL) Seattle, WA, in the 1980's (<https://www.pmel.noaa.gov/gtmba/pmel-theme/pacific-ocean-tao>). The ATLAS mooring is a taut wire surface mooring with a toroidal float, deployed in depths of up to 6000 meters (Milburn *et al.*, 1996). The expansion of this array is the result of international collaboration between scientists from France, Japan, Korea and the USA, and its current support is from the US, Japan and France. The first ATLAS mooring was deployed in December 1984. Collected data are transmitted to shore in real time using the ARGOS System (<http://www.argos-system.org/>), processed by Collecte



Localisation Satellites (CLS, <http://www.cls.fr/>) or Service ARGOS Inc., and placed on the Global Telecommunication System (GTS, <https://community.wmo.int/en/activity-areas/global-telecommunication-system-gts>). Post-recovery processing and analysis of the data is performed at PMEL. The TAO array now supports programs like the Global Climate Observing System (GCOS, <https://gcos.wmo.int/en/home>), World Climate Research Programme (WCRP, <https://www.wcrp-climate.org/>), Climate Variability and Predictability Programme (CLIVAR, <http://www.clivar.org/>), and the World Weather Watch Programme (WWW, [Website](#)).

[PIRATA](#) (Prediction and Research Moored Array in the Tropical Atlantic) is a project designed by a group of scientists involved in CLIVAR, and is implemented by the group through multinational cooperation. Contributions are provided by France with the participation of L'Institut de Recherché pour le Développement (IRD) in collaboration with Meteo-France, Centre National de la Recherche Scientifique (CNRS), Universities and French Research Institute for Exploitation of the Sea (IFREMER), by the Brazilian Instituto Nacional de Pesquisas Espaciais (INPE) and Diretoria De Hidrografia E Navegação (DHN), and by the USA (NOAA/PMEL). The purpose of PIRATA is to study ocean-atmosphere interactions in the tropical Atlantic that are relevant to regional climate variability on seasonal, inter-annual and longer time scales (<https://www.pmel.noaa.gov/gtmba/pirata>). As of this writing there are 17 active moorings in this network.

The [RAMA](#) (Research Moored Array for African–Asian–Australian Monsoon Analysis and Prediction) Project - a key element of the Indian Ocean Observing System (IndOOS) is the basin-scale moored-buoy array intended to cover the tropical Indian Ocean. In this respect, RAMA is the Indian Ocean equivalent of the TAO/TRITON array in the Pacific and PIRATA grid in the Atlantic (McPhaden *et al.*, 2009). RAMA started in 2000 as Indian and Japanese national efforts when JAMSTEC deployed two TRITON moorings and NIO (National Institute of Oceanography, India) began subsurface mooring deployment along the equator (McPhaden *et al.*, 2006). As of 2023 RAMA has produced data from 35 mooring sites (<https://www.pmel.noaa.gov/gtmba/pmel-theme/indian-ocean-rama>). The planned array will consist of 38 surface and 8 subsurface moorings. Mooring equipment, ship time, personnel, and/or logistic support has been provided by several nations including Japan, India, the United States, Indonesia, China, France and the African countries participating in the Agulhas and Somali Current Large Marine Ecosystems (ASCLME) project: Comoros, Mozambique, Kenya, Tanzania, Madagascar, South Africa, Seychelles and Mauritius (McPhaden *et al.*, 2009). Data collected by RAMA buoys are distributed by Service Argos via Global Telecommunications System (GTS) as well as via the PMEL, JAMSTEC and NIO websites (see links below).

The [MARNET](#) (Marine Environmental Monitoring Network in the North Sea and Baltic Sea) project has four buoys located in the North Sea and five buoys in the Baltic Sea. The program uses existing platforms as a base for instrument installation; in the North Sea two unmanned lightships and two North Sea Buoys (NSB II and NSB III) are used, and in the Baltic Sea two large discus buoys, a stabilized mast, semi-submersible buoy and a pier/platform near the Kiel lighthouse are used. The main components of the measuring equipment are sensors, data acquisition unit, data storage system, and data collection platform (DCP). Sensors with analog and digital outputs are connected to the data acquisition unit. The raw data are transmitted via DCP and satellite (METEOSAT) to the land-based station at the Bundesamt für Seeschifffahrt und Hydrographie (BSH, <http://www.bsh.de/>). The data storage is a security backup in case the satellite communications system breaks down. The moorings have oceanographic sensors that measure

temperatures at 5 to 8 depth levels (depending on water depth), conductivity at 2 to 4 depth levels, oxygen concentration at 2 depth levels, radioactivity at 1 or 2 depth levels, currents, water levels, nutrient analyzers and samplers for micro-contaminants (accommodated in deck containers), and sea water pumping units. For more information see the MARNET website ([https://www.bsh.de/EN/TOPICS/Monitoring\\_systems/MARNET\\_monitoring\\_network/marnet\\_monitoring\\_network\\_node.html](https://www.bsh.de/EN/TOPICS/Monitoring_systems/MARNET_monitoring_network/marnet_monitoring_network_node.html)).

The Arctic/Subarctic Ocean Fluxes ([ASOF](#)) project is a collection of research products from several countries. The ASOF project was established in 2000, and from 2000 to 2008 measurements at several locations in the Arctic and Subarctic were collected with the aim of estimating a freshwater budget for Arctic inflows and outflows. A second phase of the project, since 2008, combines the ongoing scientific work of ASOF I with application of the results to broader questions of science and society. The ASOF mooring data in WOD23 comes from the Canadian Arctic Throughflow Study ([CATS](#)), which gathered data in Nares Strait from 2003 to 2006 (Rabe *et al.*, 2010). The moorings were instrumented with Sea-Bird Electronics, Inc. model 37-IM temperature and conductivity sensors at up to 4 depth levels (CATS, 2007).

**Table 9.1. National contributions of MRB profiles in WOD23**

**Total number of profiles = 1,277,591**

<b>ISO<sup>1</sup> Country Code</b>	<b>Country Name</b>	<b>MRB Profiles</b>	<b>% of Total</b>
US	United States	755,851	59.2%
JP	Japan	190,126	14.9%
DE	Germany	114,705	9.0%
99	Unknown / International	63,363	5.0%
BR	Brazil	62,017	4.9%
FR	France	56,785	4.4%
NL	Netherlands	16,938	1.3%
CA	Canada	5,490	0.4%
GR	Greece	4,495	0.4%
IT	Italy	4,301	0.3%
GB	United Kingdom	2,615	0.2%
TW	Taiwan	905	0.1%
<i>Total:</i>		<b>1,277,591</b>	<b>100.0%</b>

<sup>1</sup> ISO stands for International Standards Organization

[The Circulation of the North Central Chukchi Shelf](#) project placed 5 moorings on the shelf of the Chukchi Sea, in water 46–54m deep, from 1993 to 1996. The moorings were instrumented with Sea-Bird Electronics, Inc. model SBE 16 temperature/ conductivity sensors at two depths (Weingartner *et al.*, 2005).

Five countries collected most of the moored buoy data in WOD23: USA, Japan, Germany, Brazil, and France. Significant amounts of data have no country information mostly because of the multinational nature of its acquisition and processing; those data were obtained from internet-based web portals of the international research Programs (*i.e.* RAMA, *etc.*) Table 9.1 provides detailed information on each country's contributions.

The Central and Northern California Ocean Observing System ([CeNCOOS](#)) acts as a Data Assembly Center for and Monterey Bay Aquarium Research Institute ([MBARI](#)) moorings M1 and M2 in Monterey Bay. The M1 mooring measures temperature at 0, 60 and 100m along with chemical, biological and meteorological information at the surface. Mooring M2 was replaced by the NDBC mooring 46042 in 2011 (CeNCOOS, 2018)

## **9.2. MRB DATA PRECISION AND ACCURACY**

The nominal accuracy of MRB temperature and salinity data depends on the temperature and conductivity sensors used. For TRITON buoys, for example, sensor range and accuracy are: conductivity 0-70/0.003  $\text{ms}\cdot\text{cm}^{-1}$ ; temperature -3.0 – 33.0/0.002°C; depth 0-1000 pounds per square inch absolute (psia) / 0.15% full scale (Kuroda, 2001; Ando *et al.*, 2005). Data acquired during TAO and PIRATA programs were collected from PROTEUS and ATLAS buoys using Sea-Bird Electronics SEACAT sensors which have sea surface temperature accuracy of 0.01°C for the PROTEUS mooring and 0.03°C for ATLAS moorings; subsurface temperature accuracy is 0.01°C for the PROTEUS mooring and 0.09°C for ATLAS moorings (Freitag *et al.*, 1994; Cronin and McPhaden, 1997). RAMA data are collected mostly from ATLAS and TRITON moorings. In February 2008 JAMSTEC deployed several mini-TRITON buoys with slack-line moorings with all its sensors equipped to measure pressure so data can be interpolated to standard depth (McPhaden *et al.*, 2009).

MARNET data were collected using oceanographic sensors calibrated at the BSH's calibration laboratory by means of triple point thermometer, gallium cells, reference resistors and resistance bridges of the highest available precision, as well as salinometers calibrated with Copenhagen standard sea water. The three seawater baths used for temperature and conductivity calibration reach a temperature stability of  $\pm 1 \cdot 10^{-3}$  °C. After deployment, the sensors are checked and cleaned at monthly intervals. During each monthly check, an *in situ* comparative measurement is carried out using a reference CTD system.

The moorings from the CATS study, part of the ASOF project, were instrumented with SBE 37-IM temperature/conductivity sensors (Rabe *et al.*, 2010). The temperature measured by this sensor is accurate to 0.002°C, and conductivity to 0.0003  $\text{S m}^{-1}$ . The measurements from the Circulation of the North Central Chukchi Shelf project were collected using SBE 16 temperature and conductivity sensors. Using this sensor, temperature was measured to 0.005 °C, and conductivity to 0.0005  $\text{S m}^{-1}$  (CATS, 2007). The sensors were calibrated before deployment and after recovery, and linearly interpolated calibration coefficients were applied to the data during processing (Weingartner, 2007).

## **9.3. MRB PROFILE DISTRIBUTIONS**

Table 9.2 gives the yearly counts of MRB profiles for the World Ocean; this is graphically illustrated in Figure 9.2.

The geographic distribution of the MRB profiles for 1977-2022 is shown in Figure 9.3. There are 1,277,591 MRB profiles for the entire World Ocean, which represents 6.9% of the entire WOD23 data collection. Note that this number is smaller than the total listed for WOD18, reflecting the change of one particular dataset from 15-minute observations averages to daily

observations averages. Majority of the MRB data were collected in the tropical regions of the Pacific, Indian and Atlantic oceans within 15°S – 15°N latitudinal band and came from relatively large-scale and ongoing programs. Some specialized programs placed buoys in different locations of the interests as can be seen in Figure 9.3. Among major research and monitoring projects, out of the total, 629,473 were contributed by Tropical Atmosphere Ocean (TAO) Buoy Array program, 124,502 by PIRATA Buoy Array program, 85,391 by TRITON (Triangle Trans-Ocean Buoy Network) program, and 98,023 profiles were collected by Research Moored Array for African-Asian-Australian Monsoon Analysis and Prediction (RAMA). Shelf Basin Interaction Project (SBI) contributed 43,005 profiles, South China Sea Monsoon Experiment project provided 905 profiles, the Central and Northern California Ocean Observing System (CeNCOOS) project contributed 3,150 profiles. MARNET program contributed 19,444 profiles from the North and Baltic Seas, and Japan Meteorological Agency (JMA) programs have contributed 73,347. The Arctic/Subarctic Ocean Fluxes (ASOF) project contributed 6,618 profiles, Circulation of the North Central Chukchi Shelf project provided 43,005 profiles, and 905 profiles collected during the South China Sea Monsoon Experiment ([SCSMEX](#)). As mentioned before, the [OceanSITES](#) program contributed 176,456 MRB profiles to WOD23. It should be noted that there are about 93K MRB profiles in WOD23 which have no project or program information submitted/collected.

**Table 9.2. The number of MRB profiles in WOD23 as a function of year**

Total number of profiles = 1,277,591

YEAR	PROFILES	YEAR	PROFILES	YEAR	PROFILES	YEAR	PROFILES
1977	173	1989	8,716	2001	36,905	2012	43,621
1978	240	1990	8,881	2002	36,165	2013	38,983
1979	1,184	1991	10,114	2003	39,633	2014	38,316
1980	1,583	1992	24,231	2004	44,390	2015	42,588
1981	1,513	1993	33,545	2005	45,976	2016	42,118
1982	1,294	1994	47,461	2006	45,929	2017	38,628
1983	1,734	1995	44,649	2007	47,371	2018	36,497
1984	1,830	1996	38,202	2008	48,573	2019	29,139
1985	2,316	1997	34,337	2009	48,182	2020	28,886
1986	3,693	1998	40,869	2010	46,388	2021	22,838
1987	4,767	1999	44,171	2011	52,099	2022	19,237
1988	7,332	2000	42,294				

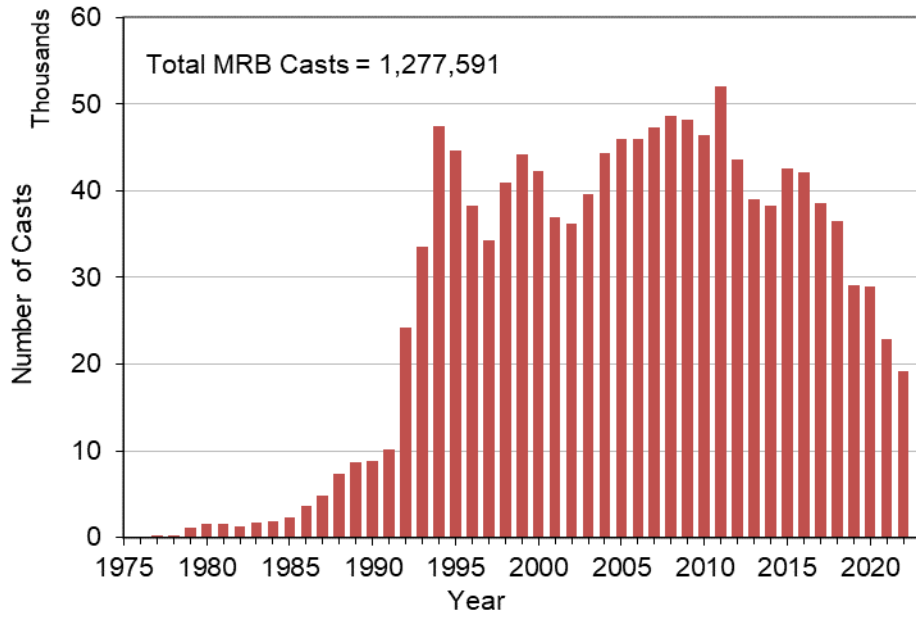


Figure 9.2. Temporal distribution of the Moored Buoy (MRB) profiles in WOD23.

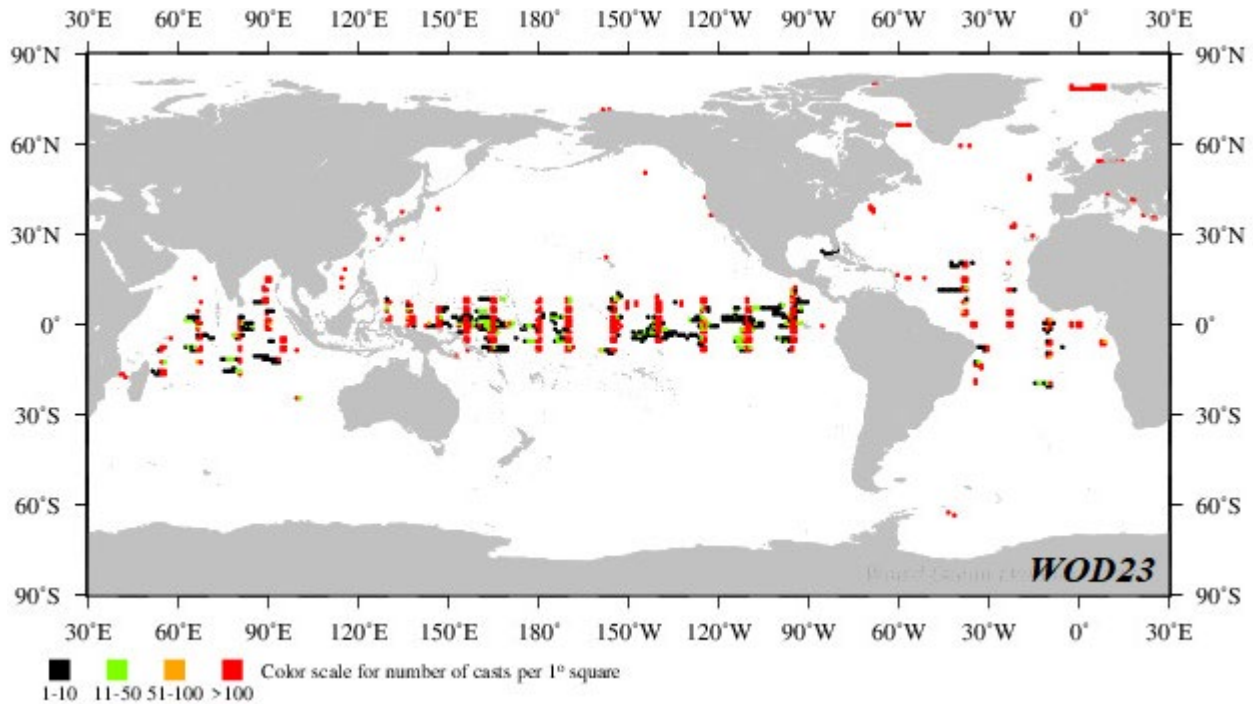


Figure 9.3. Geographic distribution of the Moored Buoy (MRB) profiles collected by major research programs in WOD23 by one-degree squares

Figure 9.4 shows the distribution of the temperature and salinity MRB data as function of depth. The majority of the moored buoys are designed to sample only the upper layer of the ocean, so most of the data were collected within upper 750 meters of the water column.

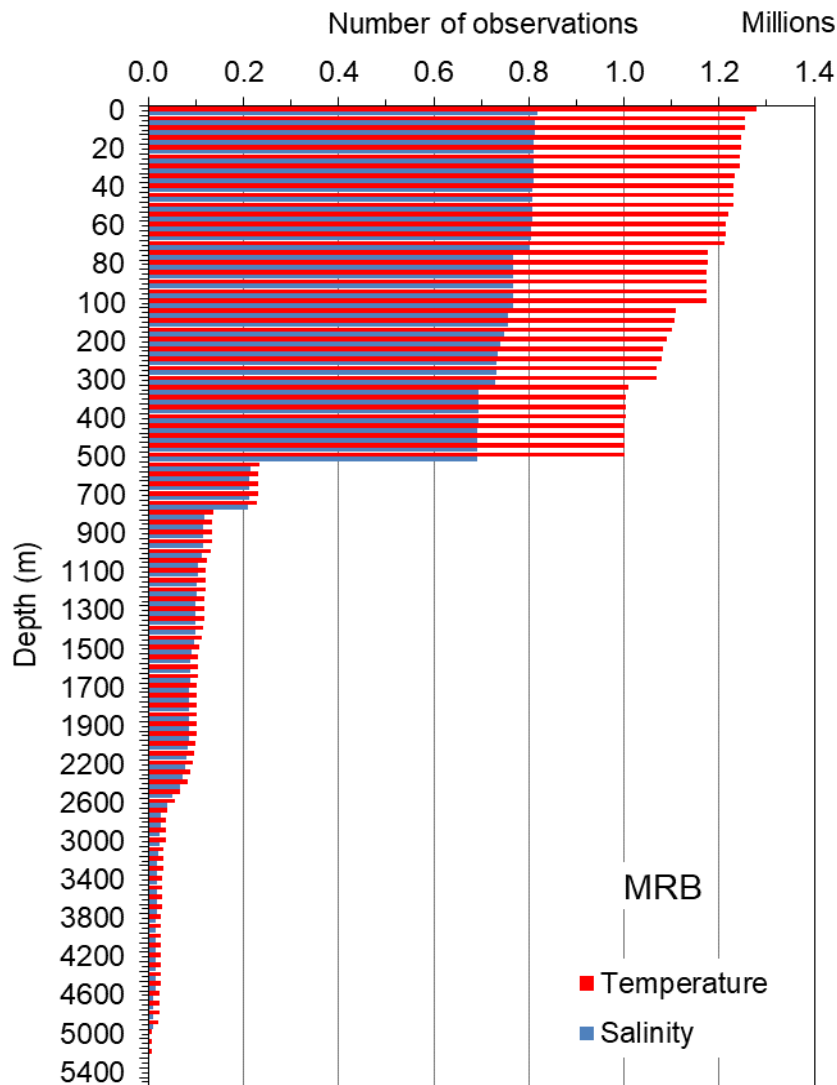


Figure 9.4. Distribution of the Moored Buoys (MRB) data at standard depth levels in WOD23

#### 9.4. RELEVANT WEB SITES

All websites last accessed: 2024-04-26

ARGOS Program: <http://www.argos-system.org/>

Centre national de la recherche scientifique (CNRS): <http://www.cnrs.fr/>

Canadian Arctic Throughflow Study (CATS):

[http://www.udel.edu/CATS/healy\\_2007/expedition/Cruise\\_Report.pdf](http://www.udel.edu/CATS/healy_2007/expedition/Cruise_Report.pdf)

Central and Northern California Observing System (CeNCOOS):



<https://www.mbari.org/data/cencoos>  
Diretoria De Hidrografia E Navegação (DHN), Brazil: <https://www.marinha.mil.br/dhn/>  
L'Institut de recherché pour le développemen (IRD): <http://www.ird.fr/>  
French Research Institute for Exploitation of the Sea (IFREMER): <http://www.ifremer.fr/>  
Instituto Nacional de Pesquisas Espaciais (INPE), Brazil: <http://www.inpe.br/>  
JAMSTEC TRITON Buoy project: [http://www.jamstec.go.jp/jamstec/TRITON/real\\_time/](http://www.jamstec.go.jp/jamstec/TRITON/real_time/)  
MARNET description available at:  
[https://www.bsh.de/EN/TOPICS/Monitoring\\_systems/MARNET\\_monitoring\\_network/marnet\\_monitoring\\_network\\_node.html](https://www.bsh.de/EN/TOPICS/Monitoring_systems/MARNET_monitoring_network/marnet_monitoring_network_node.html) .  
METEOSAT (ESA): <http://www.esa.int/>  
National Data Buoy Center: <https://www.ndbc.noaa.gov/>  
National Data Buoy Center OceanSITES page: <https://dods.ndbc.noaa.gov/oceansites/>  
PIRATA Program: <https://www.pmel.noaa.gov/gtmba/pirata>  
RAMA Program links:  
    PMEL: <https://www.pmel.noaa.gov/tao/drupal/disdel/>  
    JAMSTEC: <http://www.jamstec.go.jp/tropicbuoy/index.html>  
    NIO: <https://www.nio.res.in/>  
South China Sea Monsoon Experiment:  
    [http://www.pmel.noaa.gov/tao/proj\\_over/scsmex/scsmex-display.html](http://www.pmel.noaa.gov/tao/proj_over/scsmex/scsmex-display.html)  
Tropical Atmosphere Ocean Project: <https://www.pmel.noaa.gov/gtmba/>  
WMO-IOC Data Buoy Cooperation Panel: <https://community.wmo.int/en/data-buoy-co-operation-panel>

## **9.5. REFERENCES AND BIBLIOGRAPHY**

- Ando, K., T. Matsumoto, T. Nagahama, I. Ueki, Y. Takatsuki, Y. Kuroda (2005). Drift characteristics of a moored conductivity-temperature-depth sensor and correction of salinity data. *J. Atmos. Oceanic Technol.*, 22, 282-291.
- Cronin, M.F. and M.J. McPhaden (1997). The upper ocean heat balance in the western equatorial Pacific warm pool during September-December 1992. *J. Geophys. Res.*, 102(C4), 8533-8553.
- Data Buoy Cooperation Panel, <http://www.jcommops.org/dbcp/>.
- Freitag, H.P., Y. Feng, L.J. Mangum, M.J. McPhaden, J. Neander, L.D. Stratton (1994). Calibration procedures and instrumental accuracy estimates of TAO temperature, relative humidity and radiation measurements. *NOAA TM ERL PMEL-104*: 32 pp.
- Hayes, S.P., L.J. Mangum, J. Picaut, A. Sumi, and K. Takeuchi (1991). TOGA-TAO: A moored array for real-time measurements in the tropical Pacific Ocean. *Bull. Amer. Meteorol. Soc.*, 72, 339-347.
- Kuroda, Y. (2002). TRITON: Present status and future plan (TOCS No.5 Report). *TRITON Office*, March 2002, Yokosuka Japan, pp 77
- Mangum, L.J. (1994). TOGA-TAO Array Sampling Schemes and Sensor Evaluations, 1994: *Proc. of the Oceans '94 OSATES*, 2, 402-406.

- Mangum, L.J., H.P. Freitag, and M.J. McPhaden (1994). TOGA TAO array sampling schemes and sensor evaluations. *Proc. of the Oceans '94*, 1316. Brest, France.
- McPhaden, M.J. (1995). The Tropical Atmosphere Ocean (TAO) array is completed. *Bull. Amer. Meteorol. Soc.*, 76, 739-741.
- McPhaden, M.J., A.J. Busalacchi, R. Cheney, J.-R. Donguy, K.S. Gage, D. Halpern, M.Ji, P. Julian, G. Meyers, G.T. Mitchum, P.P. Niiler, J. Picaut, R.W. Reynolds, N. Smith, K. Takeuchi (1998). The Tropical Ocean-Global Atmosphere observing system: A decade of progress. *J. Geophys. Res.* 103 (C7), 14,169-14,240.
- McPhaden M.J., Y. Kuroda and V.S.N. Murty (2006). Development of an Indian Ocean Moored Buoy Array for Climate Studies. *CLIVAR Exchanges*, No. 11(4), Int. CLIVAR Project Office, Southampton, UK, 3-5.
- McPhaden, M.J., G. Meyers, K. Ando, Y. Masumoto, V.S.N. Murty, M. Ravichandran, F. Syamsudin, J. Vialard, L. Yu, and W. Yu (2009). RAMA: The Research Moored Array for African-Asian-Australian Monsoon Analysis and Prediction. *Bull. Amer. Meteor. Soc.*, **90**, 459-480.
- Milburn, H.B., P.D. McLain, C. Meinig (1996). ATLAS Buoy – Reengineered for the Next Decade. Ocean'1996. Available from <https://ieeexplore.ieee.org/document/568312>
- National Data Buoy Center - History. Available from: <https://www.ndbc.noaa.gov/ndbc.shtml#History>
- National Data Buoy Center: Development of national data buoy systems (1971). *U.S. DoC/NOAA publication NDBCM WO547*, 39pp.
- Rabe, B., A. Münchow, H.L. Johnson and H. Melling (2010). Nares Strait hydrography and salinity field from a 3-year moored array. *J. Geophys. Res.* 115, C07010, doi:10.1029, 2009JC005966.
- Shea, E.L. (1987). A history of NOAA. Ed. S. Theberge, *NOAA Central Library*.
- Weingartner, T., K. Aagaard, R. Woodgate, S. Danielson, Y. Sasaki, and D. Cavalieri (2005). Circulation on the north central Chukchi Sea shelf, *Deep Sea Research Part II: Topical Studies in Oceanography*, 52(24-26), 3150-3174



# CHAPTER 10: DRIFTING BUOY DATA (DRB)

*Hernan E. Garcia, Alexey V. Mishonov, Tim P. Boyer, Ricardo A. Locarnini, Dan Seidov*

*Ocean Climate Laboratory  
National Oceanographic Data Center  
Silver Spring, MD*

## 10.1. INTRODUCTION

Drifting buoys are a cost-effective means for obtaining meteorological and oceanographic data from remote ocean areas. They form an essential component of the marine observing systems that were established as part of many operational and research programs. Drifting buoys are used as a practical alternative to acquiring data from inaccessible regions as opposed to maintaining costly manned stations (DBCP, 2018; IABP, 2018).

The first drifting buoys, drift bottles, were used in the early 1800s in an effort to map surface currents. The bottles were weighted so that they were almost entirely submerged and usually carried a note that recorded launch location and time. Bottles were used because previous attempts at mapping ocean currents using ship drift measurements proved unreliable due to the added effect of wind on the movement of the ships (Lumpkin and Pazos, 2006). With the advent of radio, the position of the drifters could be transmitted from small, low-drag antennae and triangulated from the shore. In the early 1970s, positions started to be gathered via satellites. As technology improved, drifters started to obtain meteorological measurements, sea surface temperatures, as well as oceanographic measurements (IADP, 2018; Lumpkin and Pazos, 2006).

### 10.1.1. Arctic Ocean Buoy Program

The first sea ice buoys used by the [Arctic Data Buoy Program](#) were deployed in the ice floes of the Arctic Basin in 1979; they recorded meteorological parameters such as surface atmospheric pressure, air temperature, wind speed, as well as geographic position. Data were transmitted and collected via the ARGOS system and then distributed on the Global Telecommunication System (GTS) (IABP, 2018; GTS, 2018).

Between the years 1985 and 1994, the Arctic Data Buoy Program of the Polar Science Center of the [Applied Physics Laboratory at the University of Washington](#) deployed 24 modified data buoys in ice floes on the Arctic Ocean. These were the first buoys, as well as the first sea ice buoys, to be equipped with Seabird CTD sensors for collecting oceanographic data along with the meteorological data. These modified buoys, known as Polar Ocean Profile (POP) buoys, measured subsurface ocean temperature, salinity, and depth. They also measured air temperature and barometric pressure. Measurements were taken at twelve-minute intervals. The direction and velocity of the sea ice floe was interpolated from changes in position from each buoy. Due to being subjected to the stresses and strains of the Arctic pack ice, these buoys varied greatly in their

longevity, though the battery pack was designed to last for approximately three years (Rigor, 2002; IABP, 2018; JAMSTEC, 2018).

The components of a Polar Ocean Profile Buoy start with an ARGOS antenna with air temperature and barometric pressure sensors in a fiberglass shroud that protrudes from the ice floe. This sits on a flotation/ablation skirt that is directly on top of the ice. Within the ice itself are the buoy electronics assembly housing and an alkaline (D-cell) battery pack, all encased in an aluminum hull. Attached to the bottom of the hull, extending into the water column, is a 24-conductor electromagnet cable upon which an SBE-16 SEACAT CTD sensor is attached. The SBE-16 SEACAT has a total of 6 sensors, placed at depths of 10, 40, 70, 120, 200, and 300 meters; a depth sensor is added to the sensors at 40, 120, and 300 meters. At the very end of the electromechanical cable is a 50-pound ballast weight (IABP, 2006).

### **10.1.2. Global Temperature-Salinity Profile Program (GTSP)**

The Marine Environmental Data Service ([MEDS](#), Canada) collects the data from all drifting buoys via the Global Telecommunication System (GTS). MEDS has been a Responsible National Oceanographic Data Center (RNODC) since January 1986 under the auspices of the Intergovernmental Oceanographic Commission (IOC). They acquire, process, quality control, and archive real-time drifting buoy data that is reported over the GTS as well as delayed-mode data that are acquired from other sources. Over 200,000 new records are captured monthly from the GTS by MEDS. The GTSP program through MEDS only includes data from drifting buoys that send subsurface data. This drifting buoy data includes buoy position, date, time, surface and subsurface water temperature, salinity, air pressure, temperature and wind direction (MEDS, 2018). Currently, buoy data from GTSP in the WOD23 database comes mostly from the United States, Japan and France and consists of temperature readings and some have meteorological measurements such as wind speed, wind direction, dry bulb temperature, and barometric pressure.

### **10.1.3. JAMSTEC Buoys**

In the early 1990s, the Japan Marine Science and Technology Center ([JAMSTEC](#)) developed a polar ocean profiler buoy, the Ice Ocean Environmental Buoy ([IOEB](#)), as a joint project with the Woods Hole Oceanographic Institution ([WHOI](#)). This was the first attempt to develop a drifting ice buoy equipped with not only meteorological, sea ice and oceanographic sensors, but also with other sensors, such as optical sensors and time series collection devices, that would determine the activities of marine organisms. The first IOEB was deployed in the Beaufort Sea in April 1992; the second was deployed April 1994 into the Arctic Transpolar Drift. These first buoys lacked mobility and had little consistency in measurements due to the large number of different sensors on them. In addition, the buoys were expensive to assemble and required large camps and lots of equipment and materials to install them in the ice. They also had to be recovered to analyze collected sediment samples (JAMSTEC, 2018).

JAMSTEC and MetOcean Data System Ltd. developed a new drifting buoy in 1999, named [J-CAD](#) (JAMSTEC Compact Arctic Drifter), and its mission was to conduct long-term observations in the Arctic Ocean multi-year ice zones as a participant of the International Arctic Buoy Program ([IABP](#)). Since 2000, the J-CAD has been used to measure the structure of upper

ocean currents and water properties. J-CADs have been installed into the sea ice in various regions of the Arctic Ocean and have been collecting oceanographic and meteorological data. The data J-CAD buoys collect are: air temperature, barometric pressure, wind direction, wind speed, sea surface temperature, platform heading, platform tilt, latitude, longitude, date and time of reading, GPS drift speed, GPS drift direction, CTD sensor depth, pressure, temperature, conductivity, salinity, potential temperature, density, and several ADCP parameters. (The ADCP data are not available through the WOD series.) The sensors measure data at one-hour intervals and the J-CAD deployment location varies by different projects' requirements (JAMSTEC, 2018; Kikuchi *et al.*, 2002).

The total weight of the J-CAD system was designed to be 255 kg or less so it can be deployed using a small, light crane system. The maximum external diameter of the underwater sensors is 28 cm; each sensor can be lowered through a 30 cm hole in the ice that can be drilled with simple equipment. It is equipped with three types of sensors: meteorological, oceanographic, and buoy status sensors. The J-CAD buoys consist of a floatation collar made of foam resin buoyancy material (Surlyn Ionomer resin manufactured by DuPont Co.) enclosed by aluminum. The housing for instruments, also made from aluminum and foam resin, holds the data logger/controller engine (Tattletale model 8) with 48MB flash card memory, a GPS receiver, two satellite communication systems, the GPS interface MetOcean Digital Controller, and two 245 Ah lithium battery packs to supply power. On the top of the aluminum enclosure is an ARGOS antenna mast that includes the air temperature sensor, the barometer port, and two GPS antennas. There is also a PC interface for the physical downloading of data from the flash card memory, to configure the data logger, and to set various sensor operating parameters (JAMSTEC, 2018).

Meteorological sensors equipped on the J-CAD consist of a YSI Inc. model 44032 high-precision thermistor for air temperature, a Paroscientific Inc. model 216B barometer, and a RM Young Co. model 5106-MA anemometer. The outside air or sea ice temperature is measured from the thermistor placed at the top of the ARGOS antenna mast. The barometer port is also at the top of the mast and is covered by a water trap and a Gore-Tex membrane to protect it from moisture. Finally, the wind sensor is vertically mounted on the top of the J-CAD tower; this tower is designed to withstand 120-knot winds (JAMSTEC, 2018).

The ocean temperature and conductivity data are obtained from Sea-Bird SBE37IM CT sensors (CT stands for Conductivity and Temperature), two of which are equipped with pressure sensors that are part of the CT instrument. On a J-CAD buoy, four CT and two CTD sensors can be mounted. The CT sensors are usually attached at 25m, 50m, 80m, and 180m. The two CTD sensors are usually placed at 120m and 250m. These depths can be adjusted to the sea area under observation. There are also two WorkHorse 300 kHz ADCPs from RD Instruments attached at 12m (facing downward) and at 260m (facing upward/downward) depth to measure the underwater currents. These ADCPs also measure the heading, pitch, and roll of the buoy and have a thermistor to measure the water temperature at the ADCPs' depth (JAMSTEC, 2018; Kikuchi *et al.*, 2002).

The J-CAD is equipped with sensors that check the physical status of the buoy. A model TCM2, three-axis magnetometer (Precision Navigation Inc.) measures the platform's orientation. It is mounted inside the hull and provides estimates of platform direction and vertical tilt. There is also a compass that indicates the rotation of the ice base that the J-CAD platform is installed upon. Two GPS receivers are attached to the ARGOS mast. One receiver is a Jupiter model TU30-D140-231 (Conexant Systems Inc.) and is interfaced with the MetOcean Digital Controller. The data from this GPS is used as the J-CAD position reported for the data. The second GPS is an integral

part of the Panasonic KX-G7101 ORBCOMM Subscriber Communicator but is only used as a complement to the ORBCOMM satellite system. Finally, there is a sensor to measure the temperature of the water and/or ice surrounding the J-CAD hull. It is an YSI model 44032 high-precision thermistor that is in constant contact with the inside wall of the platform hull. The instrument is safely inside the J-CAD and, due to the high thermal conductivity of aluminum, the interior wall temperature matches the outside temperature, giving an accurate reading (JAMSTEC, 2018).

In the spring of 2000, an international research team supported by the U.S. National Science Foundation (NSF) was formed to conduct annual expeditions to the North Pole. These expeditions established a group of un-manned platforms, collectively referred to as an observatory, to record as much data as possible. Drifting buoys from the IABP and the JAMSTEC J-CAD are major components of this project, entitled the North Pole Environmental Observatory (NPEO) Project. JAMSTEC continues to deploy buoys at a rate of 2 per year for the NPEO. The Pacific Marine Environmental Laboratory (PMEL) also maintains drifting weather buoys as part of this program (NPEO, 2018; Kikuchi *et al.*, 2002). WOD23 DRB database contains 14,067 profiles collected between 2000-2022 and contributed by NPEO project.

#### **10.1.4. Ice-Tethered Profiling Buoys (ITP)**

In 2006, WHOI developed the [ITP](#) buoy platform to effectively map the water column in ice-covered seas with a profiling drifter. This drifter builds off the technology developed for the moored profiler (MP) instrument, also developed at WHOI, and from the innovations of the ARGO profiling floats. Instead of being instrumented at fixed depths like other ice drifters, the ITP platform profiles at high resolution through the water column as the ice floe moves, and returns measurements along with the instrument's position from GPS. The platform is lightweight and able to be deployed by helicopter or Twin Otter aircraft through a standard 25cm hole augured in the ice. The instruments are relatively inexpensive, so they can be considered expendable and several can be deployed at one time (ITP, 2018; Krishfield *et al.*, 2008).

The drifter is composed of two parts: the surface package that sits on the ice surface, and the underwater profiling package on the tether line. The surface package is made from yellow-painted foam and contains the Iridium modem, GPS receiver, data controller, batteries and an interface to the underwater part of the drifter. On the outside of the surface package, a temperature sensor measures air temperature. The surface package is designed to be expandable and support other instrumentation as well as increased power and battery loads in the future. When deployed, the surface package is placed on a wooden pallet to help avoid melting and deformation of the ice floe (ITP, 2018; Krishfield *et al.*, 2008).

Below the surface, a plastic-jacketed wire rope tether line extends up to 800m into the water column. A ballast weight at the bottom of the line keeps it oriented vertically. The profiling package is ballasted to be neutrally buoyant at mid-profile depth. The profiling package moves along the line using a traction drive similar to that used by the moored profiler (MP) buoys. The profiling package is instrumented with a Sea Bird Electronics, Inc. model 41-CP CTD, the same instrument used by ARGO floats. The package returns data at a sample rate of 1Hz and the surface unit sends it to shore at near-real time over the Iridium link. Two related platforms, the Ice-Tethered Winch (ITW) and Ice-Tethered Micro-mooring (ITM) have also been deployed. The ITW

uses a profiling Arctic winch to sample closer to the sea surface. The ITMs, which have instruments at fixed depths, are instrumented with Sea-Bird SBE37-SI microCATs and Nortec Aquadopp current profilers to measure ocean currents (ocean currents data not available in the WOD series) (ITP, 2018; Krishfield *et al.*, 2008).

As of September 2018, 110 individual ITP, ITM and ITW packages have been launched, supporting the Beaufort Gyre Observing System ([BGOS](#)), Beaufort Gyre Freshwater Experiment (BGFE), North Pole Environmental Observatory (NPEO), European Union DAMOCLES, Nansen and Amundsen Basins Observational System (NABOS), and Hybrid Arctic/Antarctic Float Observation System (HAFOS) projects in the Arctic and the National Institute of Water and Atmospheric Research (NIWA) project in the Antarctic (ITP, 2018).

### **10.1.5. PAICEX Russian Drifting Ice Camp data**

Data rescue coordinated by the Global Oceanographic Data Archaeology and Rescue (GODAR, 2018) project resulted in a digitization and submission by the [P.P Shirshov Institute of Oceanology](#), Russian Academy of Sciences, of Soviet and Russian drifting ice camp data collected between 1950 and 2009. The April 2007 and April 2009 ice camp data, from the [PAICEX](#) project, also featured drifting mooring lines. The camp started at [ice-station Barneo](#), around 89°N, and drifted along with the ice. The moorings measured temperature and salinity at three depths (75m, 100m and 125m) every 3 minutes for approximately 2 weeks. The PAICEX program also collected bottle profiles; these data are part of the WOD23 OSD dataset (see Chapter 2 of this document).

## **10.2. DRB ACCURACY**

The SBE-16 SEACAT that is used in the AOBP's POP buoy is designed to accurately measure and record temperature and conductivity. It is powered by internal batteries that give it a year or more of recording time. The time-base is accurate to within 3 minutes per year. There is also an internal battery back-up to support the memory and the real-time clock. Data from the AOBP's POP buoy's SBE-16 SEACAT consists of temperature and conductivity measurements from pre-determined depths along the cable. It is capable of temperature measurements ranging from -5 to +35°C with an accuracy of 0.01°C and has a resolution of 0.001°C. The conductivity measurement range is from 0 to 7 S m<sup>-1</sup> with an accuracy of 0.001 S m<sup>-1</sup> and resolution of 0.0001 S m<sup>-1</sup> (Sea-Bird Electronics Inc., 2013).

The foremost concern of the POP buoy's accuracy was conductivity sensor drift due to fouling. Over a year, it seemed that the normal instrumental drift that occurs with age and use fell to less than 1% of the original accuracy. Because the buoys were not usually recovered or revisited, their approach to minimize fouling was to use light baffling shrouds coated with anti-fouling paint around the conductivity cell. More recently, Sea-Bird has provided anti-fouling tubes on the ends of the conductivity cells. The Arctic environment, being cold and dark for half of the year, is detrimental to the growth of fouling organisms. The few sensors that were recovered showed no evidence of fouling or fouling drift. Over time, fouling was generally found to not be a serious problem in the Arctic, though there were occasional problems with shallow sensors in the summer (Morrison, Pers. Com.; Rigor, 2002). Another problem with the POP buoys was inaccurate surface air temperatures that were caused by the small size of the buoy. The air temperature sensor was inside a fiberglass shroud that created a microcosm that would heat up in the summer and be drifted

over and insulated by snow in the winter. This difference in internal and external environments rendered the air temperature readings unfit for scientific use (Rigor *et al.*, 2000).

For data transmitted through GTSPP, the MEDS data quality control consists of two main parts: validation and verification. The data validation consists of reformatting the data to the MEDS processing format, this allows the data to be checked for its readability and correct interpretation. When the reformatting is complete, then the data values themselves are quality controlled or verified. This is to ensure that the number and codes represent reasonable physical quantities that exist in the given time and location. There are three parts to the verification process:

- checking the drift track,
- checking the variable values, and
- checking for duplicate profiles.

The track is checked to make sure that the date is valid, not listed as a future date or one that is farther in the past than the buoy was deployed, and to make sure that the position is not over land. The inferred speed between each measurement location is also checked to make sure that it is reasonable. Values of variables are checked against the regional range as well as others for validity and any spikes in gradients or large inversions; any discrepancies are flagged with specific flags. Duplicate checking will identify any data that are versions of the same observation. Exact matches where each version of the same observation is identical usually results in one observation being deleted, unless the data were gathered by two different methods, then both observations are specifically flagged and kept in the database. The results of the quality control procedure are the setting of flags or making corrections where instrument failure or human error is evident on the data that needs it (MEDS, 2018).

J-CAD buoys use six Sea-Bird SBE-37 IM CT sensors, two of which are equipped with pressure sensors. The SBE-37 IM accurately measures conductivity and temperature with optional pressure. It has an internal battery, non-volatile memory and uses an inductive modem to transmit data and receive commands. It is specifically designed for moorings and other long-duration, fixed-site deployments. Over 100,000 measurements can be taken before the battery runs low and its real-time clock is accurate to within 2.6 minutes per year. The range of temperature and conductivity measurements match the IABP's POP's SBE-16 SEACAT (-5 to +35°C and 0 to 7 S m<sup>-1</sup> respectively), but the SBE-37 IM has an initial temperature accuracy of 0.002°C and initial conductivity accuracy of 0.0003 S m<sup>-1</sup>. The pressure sensor used has a range of 0 to 7,000 meters and is accurate to within 1%. Resolution of the temperature, conductivity, and pressure data are 0.0001°C, 0.00001 S m<sup>-1</sup>, and 0.002% respectively (Sea-Bird Electronics Inc., 2018).

The ITP drifters use the Sea-Bird Electronics, Inc. model 41-CP CTD package to measure temperature, conductivity and pressure at 1 Hz resolution as the instrument profiles the water column. Temperature measurements using the 41-CP are accurate to 0.002°C, and the conductivity sensor has an accuracy of 0.002 (equivalent salinity). Pressure is accurate to within 2 dbar. The instrument demonstrates good stability and is suitable for long deployments with little measurement drift. Some ITP drifters also include a dissolved oxygen sensor, the SBE-43, which is accurate to 2% of saturation (Sea-Bird Electronics Inc., 2018). Also included on select ITP drifters is a Seapoint fluorometer to measure chlorophyll (Seapoint Sensors Inc., 2013). Depending on the water depth, the drifters may take up to 6 profiles per day (Krishfield, 2008). ITM instruments are outfitted with Sea-Bird SBE37-SIP microCATs, which have temperature sensors

accurate to  $\pm 0.002$  °C, conductivity sensors accurate to  $\pm 0.0003$  S/m, and pressure sensors accurate to  $\pm 0.1\%$  of the range (Sea-Bird Electronics Inc, 2018).

### 10.3 DRB PROFILE DISTRIBUTIONS

There are data from 227,871 drifting buoy casts in WOD18, which were submitted by five major research programs. DRB data represents 1.5% of WOD23 data collection. The majority of DRB data came from the Ice-Tethered Profiler program, which contributed 93,750 profiles, and surface drifters equipped with thermistor chains via GTSP data system (82,427 casts). JAMSTEC provided 40,450 casts from J-CAD buoys, the Arctic Ocean Buoy Program (AOPB) submitted 8,240 profiles, and the P.P Shirshov Institute of Oceanology, Russian Academy of Sciences submitted 11,243 casts from the [PAICEX](#) project (see Figure 10.1).

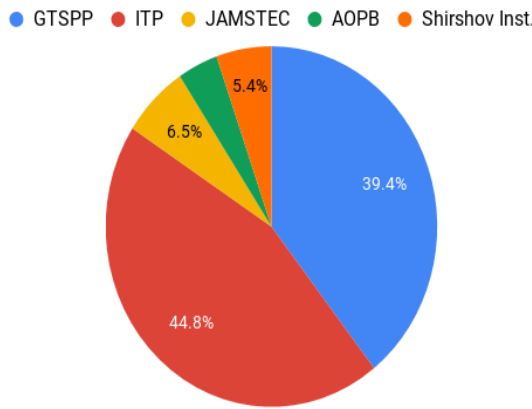


Figure 10.1. Distribution of the Drifting Buoy data in WOD23 among major research programs

The geographic distribution of the DRB casts is illustrated in Figure 10.2 for global ocean. The majority of the DRB casts distributed in the northern hemisphere in Pacific, Atlantic and Arctic oceans, and small amount found in the Southern Ocean. There are also profiles from the Mediterranean Sea, Persian Gulf and Red Sea as well as the northern Indian Ocean, but they are only a minor part of the global DRB profiles distribution.

The temporal distribution of the DRB data is shown in Table 10.1 as well as in Figure 10.3.

Table 10.1. The number of DRB profiles in as a function of year in WOD23. The total number of profiles = 272,872

YEARS	CASTS	YEARS	CASTS	YEARS	CASTS	YEARS	CASTS
1985	217	1995	0	2005	10,334	2014	13,494
1986	482	1996	0	2006	4,655	2015	7,158
1987	447	1997	0	2007	9,231	2016	3,064
1988	1,387	1998	3	2008	7,131	2017	3,493
1989	1,510	1999	4,770	2009	12,897	2018	6,874
1990	1,175	2000	12,611	2010	5,301	2019	21,249
1991	1,422	2001	62,952	2011	6,999	2020	10,891
1992	606	2002	9,249	2012	14,737	2021	4,806
1993	462	2003	7,905	2013	17,841	2022	1,305
1994	532	2004	5,682				

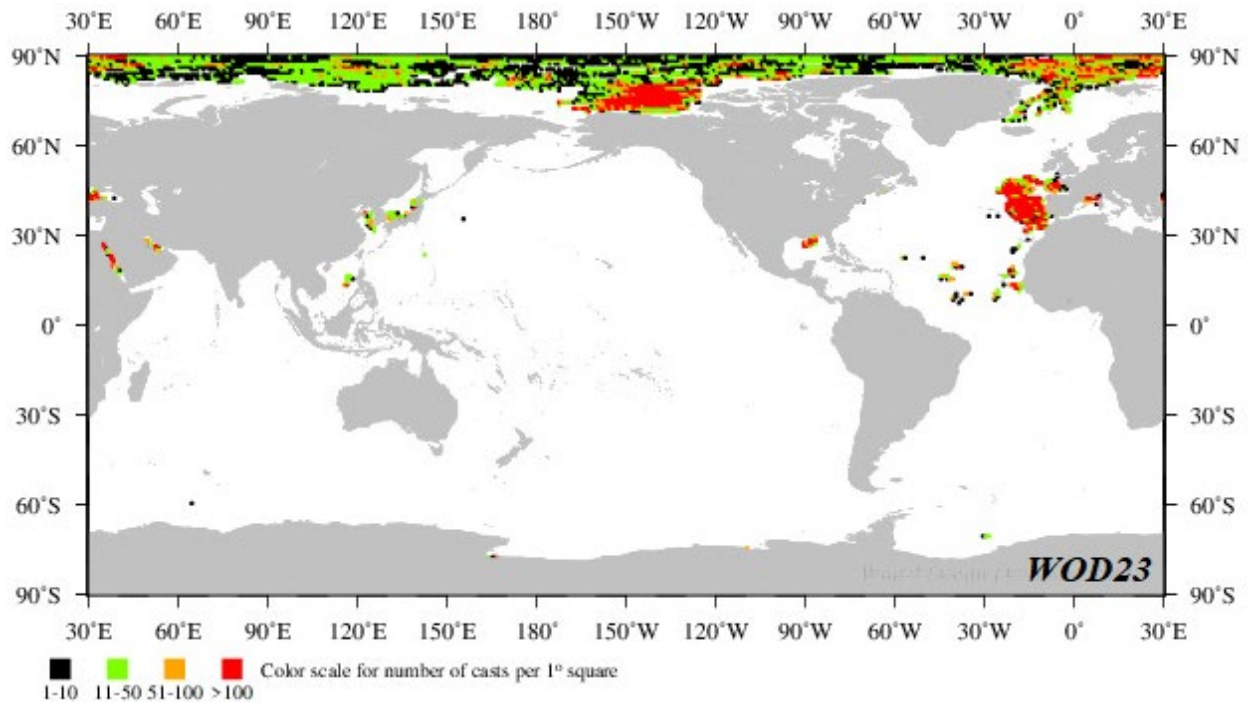


Table 10.2 gives national input to the DRB dataset by each contributing country.

**Table 10.2. National contributions of DRB casts in WOD23.**

ISO <sup>1</sup> Country Code	Country Name	DRB Casts	% of Total
US	United States	161,294	59.11
FR	France	58,100	21.29
JP	Japan	40,453	14.82
RU	Russian Federation	11,243	4.12
99	Unknown / International	1,781	0.65
AU	Australia	1	<0.01
<i>Total:</i>		<i>272,872</i>	<i>100.00</i>

<sup>1</sup> ISO = International Organization for Standardization  
[http://www.iso.org/iso/country\\_codes.htm](http://www.iso.org/iso/country_codes.htm)



**Figure 10.2. Geographic distribution of the Drifting Buoy (DRB) data by one-degree squares in WOD23.**

It should be noted that in year 2001, the prominent "drifter buoy program" referred to as the Global Drifter Program ([GDP](#)), was fully established, providing extensive global data on ocean currents and sea surface temperature through a network of satellite-tracked drifting buoys; this program is primarily managed by NOAA's AOML in Miami. This results in significant amount of the data collected in that year and added to WOD.



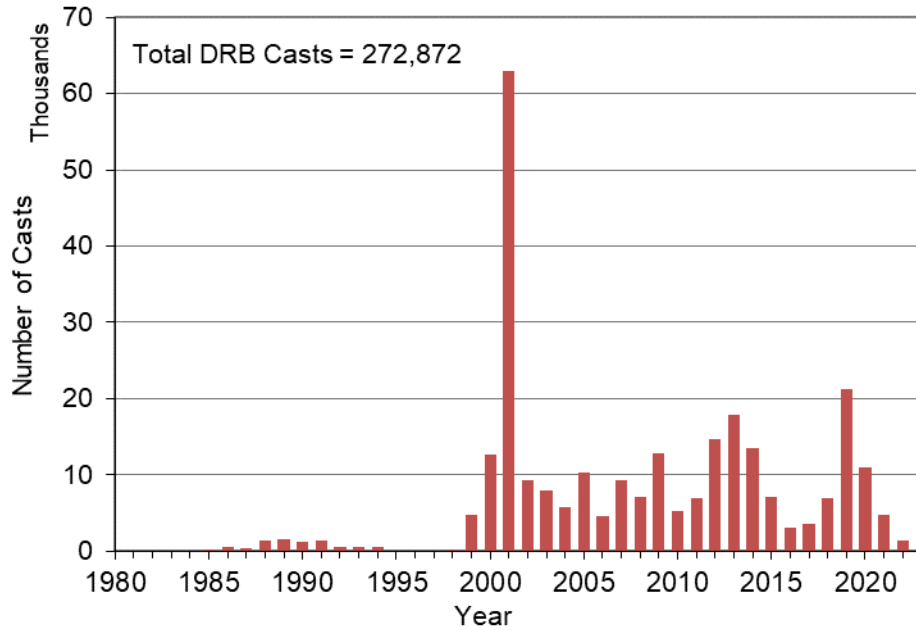


Figure 10.3. Time series of the Drifting Buoy (DRB) casts as a function of year in WOD23.

Distribution of the DRB data as a function of depth at standard depth levels is illustrated in Figure 10.4.

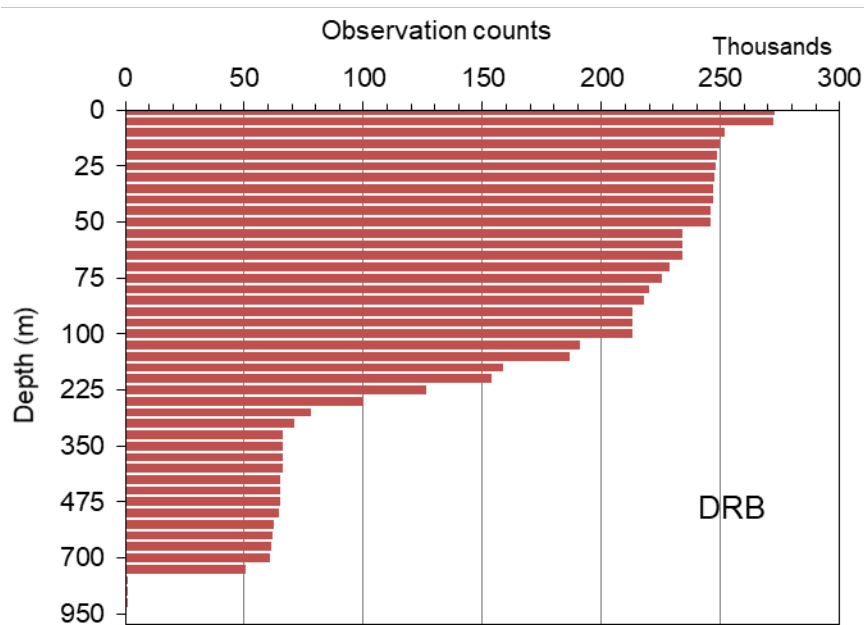


Figure 10.4. Distribution of the Drifting Buoy (DRB) data at standard depth levels in WOD23.

### 10.4. RELEVANT WEB SITES

DBCP, 2018. Data Buoy Cooperation Panel, Joint WMO-IOC Technical Commission for Oceanography and Marine Meteorology, <http://www.jcommops.org/dbcp/>

GODAR, 2018.

[https://www.iode.org/index.php?option=com\\_content&view=article&id=18&Itemid=57](https://www.iode.org/index.php?option=com_content&view=article&id=18&Itemid=57)

GTS, 2018. The Global Telecommunication System, World Meteorological Organization, <https://community.wmo.int/en/activity-areas/global-telecommunication-system-gts>

IABP, 2018. International Arctic Buoy Program, Polar Science Center, Applied Physics Laboratory, University of Washington, Washington, USA, <http://iabp.apl.washington.edu> and [http://iabp.apl.washington.edu/overview\\_contributions.html](http://iabp.apl.washington.edu/overview_contributions.html)

ITP, 2018. Ice-Tethered Profiler, Woods Hole Oceanographic Institution, Massachusetts, USA, <http://www.whoi.edu/page.do?pid=20756> .

JAMSTEC, 2018. JAMSTEC Compact Arctic Drifter (J-CAD), Arctic Ocean Climate System Group, Global Warming Observational Research Program, Institute of Observational Research for Global Change, Japan Agency for Marine-Earth Science and Technology, Kanagawa, Japan, [http://www.jamstec.go.jp/arctic/J-CAD\\_e/jcadindex\\_e.htm](http://www.jamstec.go.jp/arctic/J-CAD_e/jcadindex_e.htm).

MEDS, 2018. Marine Environmental Data Service (MEDS), Department of Fisheries and Oceans, Ontario, Canada, <http://www.meds-sdmm.dfo-mpo.gc.ca/isdm-gdsi/drib-bder/index-eng.html>

NDBC, 2018. National Data Buoy Center (NDBC), National Weather Service, National Oceanic and Atmospheric Administration, U.S. Dept. of Commerce, Mississippi, USA, <http://www.ndbc.noaa.gov>.

NPEO, 2018. The North Pole Environmental Observatory (NPEO) Project, Office of Polar Programs, National Science Foundation, Virginia, USA, <http://psc.apl.washington.edu/northpole/index.html>.

Sea-Bird Electronics Inc., 2018. <http://www.seabird.com>.

Seapoint Sensors Inc., 2018. <http://www.seapoint.com/scf.htm>.

## **10.5 REFERENCES AND BIBLIOGRAPHY**

Kikuchi, T., K. Hatakeyama, K. Shimada, T. Takizawa, and J. Morison (2002). Oceanographic observation under the multi-year ice of the Arctic Ocean using J-CAD (JAMSTEC Compact Arctic Drifter). *Mombetsu-02 Symposium*, Feb. 2002. Mombetsu, Hokkaido, Japan.

Krishfield, R., J. Toole, A. Proshutinsky, and M.-L. Timmermans (2008). Automated Ice-Tethered Profilers for Seawater Observations under Pack Ice in All Seasons. *J. Atmos. and Oceanic Technol.*, **25**, 2091 – 2105.

Lumpkin, R. and M. Pazos (2006). Measuring Surface currents with Surface Velocity Program drifters: the instrument, its data, and some recent results. Chap. 2 of *Lagrangian Analysis and Prediction of Coastal and Ocean Dynamics* (LAPCOD). Eds. A. Griffa, A.D. Kirwan, J.J. Mariano, T. Ozgokmen, and T. Rossby.

Morison, J. (Personal Communication, Feb. 2006). Principal Oceanographer, Polar Science Center, Applied Physics Laboratory, U. Washington, Seattle, WA, USA.

- Rigor, I.G. and A. Heiberg (1997). International Arctic Buoy Program data report 1. January - 31 December 1995. U. of Washington, Seattle. Applied Physics Laboratory. *Technical memorandum, May 1997*. APL-UW TM 4-97, 173p. + append.
- Rigor, I., R. Colony, and S. Martin (2000). Variations in Surface Air Temperature Observations in the Arctic, 1979 - 1997, *J. Clim.*, 13(5): 896-914.
- Rigor, I. (2002). IABP drifting buoy, pressure, temperature, position, and interpolated ice velocity. Compiled by the Polar Science Center, Applied Physics Laboratory, U. of Washington, Seattle, in association with NSIDC. Boulder, CO: *National Snow and Ice Data Center*. Digital media.

# CHAPTER 11: UNDULATING OCEAN RECORDER DATA (UOR)

*James R. Reagan, Christopher R. Paver, Tim P. Boyer, Alexey V. Mishonov*

*Ocean Climate Laboratory  
National Centers for Environmental Information  
Silver Spring, MD*

## 11.1. INTRODUCTION

The first ship-towed ocean recorder was developed by Sir Alister Hardy for underway plankton sampling. As stated at the Sir Alister Hardy Foundation for Ocean Science website (<http://www.sahfos.ac.uk/about-us/history/history-of-the-cpr-survey.aspx>): “Sir Alister Hardy started his career as a fishery biologist in Lowestoft, England. In 1925 he embarked on a two-year voyage to the Antarctic on the ship *Discovery*. He designed the prototype Continuous Plankton Recorder (Mark I) specifically for the expedition. After his return in 1927, Hardy designed a smaller version of the Continuous Plankton Recorder (Mark II) for use on merchant ships. This model is essentially the same as that used routinely today. In September 1931, the *SS Albatross* towed the first Continuous Plankton Recorder (CPR), and the survey was born. The CPR Survey was based in Hull until 1950, when it moved to Edinburgh under the administration of the Scottish Marine Biological Association (SMBA). In 1959, the first transatlantic route was towed from Reykjavik to Newfoundland.”

As ship speeds increased through the 1950's and 1960's, a need for a fast CPR was identified. In addition to this, a need to measure plankton concentrations at more than a single depth level (~10 m) had also been identified. Thus, a fast CPR (FCPR) and a prototype undulator were developed. By the early-1970's, as technology advanced in the form of more environmental data sensors and larger data storage, the Undulating Oceanographic Recorder (UOR) was born (Reid *et al.*, 2003). The UOR Mark I was developed jointly through the Oceanographic Laboratory at Edinburgh and the Plessey Marine Systems Unit. Further development took place at the Plymouth Marine Laboratory (<http://www.pml.ac.uk/>) where UOR Mark 2 was developed (Reid *et al.*, 2003). In the early-1960's the Longhurst Hardy Plankton Recorder (LHPR) was also developed, allowing vertical measurements of plankton to be recorded (Longhurst, 1966). For a more in-depth history of the CPR and UOR, please review Reid *et al.* (2003).

The modern UOR is a self-contained oceanographic sampler which can be towed from research vessels and merchant ships at speeds up to 25 knots. It can be launched and recovered by non-scientist crew members while the vessel is underway. It can be used to carry instrumentation to sample plankton continuously and to measure chlorophyll, radiant energy, temperature, and conductivity, all of which are recorded with depth (Aiken, 1981; Burt, 2000). This technique is often used for large marine ecosystem sampling or frontal zones because of its convenience and uninterrupted data coverage (Williams and Lindley, 1980; 1998; Pollard, 1986) and its ability to sample a large area in a reasonable period of time (Brown *et al.*, 1996). It is also expanding towards

a wider set of sensors used, such as light absorption sensors and attenuation meters (Barth and Bogucki, 2000).

The WOD23 UOR dataset consists of temperature, salinity, conductivity, chlorophyll concentration, pressure, and a small number of oxygen profiles (see Table 11.1 for details) collected by CTD and fluorometer sensors mounted on a SeaSoar-type towed vehicle. The SeaSoar towed vehicle was developed by Chelsea Technologies Group (<http://www.chelsea.co.uk>) from an original design by the Institute of Oceanographic Sciences (now the Southampton Oceanography Centre, UK). “Typical instrumentation includes a SeaBird 911+ CTD with pressure and dual conductivity and temperature sensors, and optical instrumentation such as fluorometer, transmissometer, PAR sensor, an experimental bioluminescence sensor and others.

Under the original open-ocean configuration, SeaSoar undulates between the surface and about 400 meters depth while being towed on faired cable at about eight knots. A typical dive cycle takes about 12 minutes to complete, providing an up- and down profile every 3 km.” (<https://www.whoi.edu/what-we-do/explore/underwater-vehicles/towed-vehicles/seasoar/>).

**Table 11.1. Profile count for major variables in the WOD23 UOR dataset.**

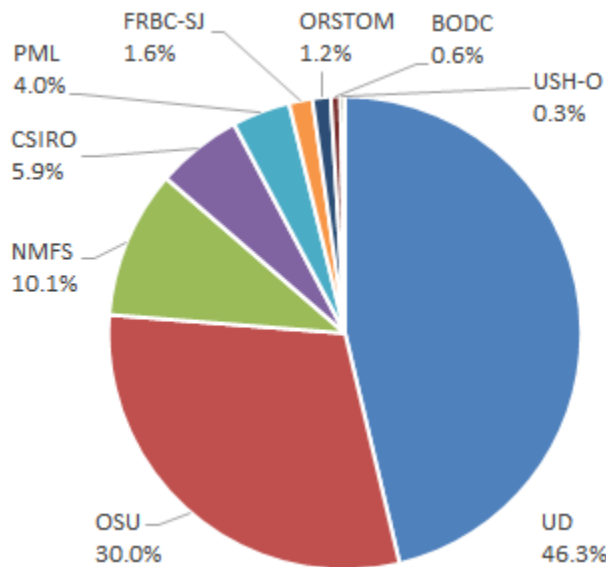
<b>Variables</b>	<b>Profiles</b>
Temperature	127,554
Salinity	125,729
Oxygen	361
Chlorophyll	20,254
Pressure	118,757
Conductivity	41,485

Much of the WOD23 UOR data were submitted to NCEI by nine institutions (see Figure 11.1) and collected in the framework of several major international programs (see Figure 11.2) in the Atlantic, Pacific and Indian oceans from 1992 till 2000.

The majority of data (41,485 casts) were collected by the University of Delaware (UD, see Fig. 11.1) during the Delaware Circulation and Dye Experiment (DECADE, see Fig. 11.2). This experiment studied the mixing and secondary circulation in the Delaware River plume. Data were acquired by means of a Scanfish undulating towed vehicle equipped with a Chelsea Ltd. MKIII Aquatracka fluorometer fitted to a Sea Bird SBE-911 CTD (Houghton *et al.*, 2004).

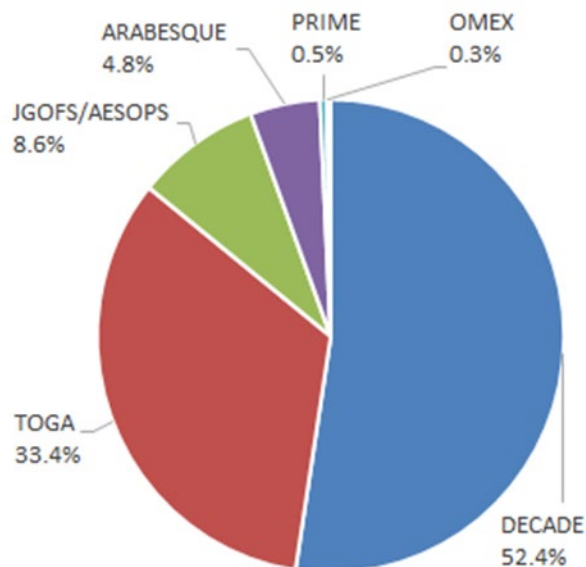
Oregon State University (OSU, Corvallis, OR) has submitted 26,854 UOR casts to NCEI, of which 20,026 casts were taken as part of the international research program “Tropical Ocean Global Atmosphere/ Coupled Ocean Atmosphere Response Experiment” (TOGA/COARE).

The Australian Commonwealth Scientific and Industrial Research Organization (CSIRO) provided 5,269 casts and the Office de la Recherche Scientifique et Technique d'Outre-Mer program (ORSTOM, France) submitted 1,118 casts, all of which were a part of the TOGA/COARE. The TOGA program studied the interaction of the ocean and atmosphere in the western Pacific warm pool region. Field measurements were made along the ~155°E line in 1992 - 1993 (for further details see the TOGA/COARE website at <http://www.soest.hawaii.edu/COARE/index.html>). A substantial amount of data was provided by the NOAA National Marine Fisheries Service (NMFS) in Seattle, WA, which contributed 9,054 casts measured along the Oregon coast.



**Figure 11.1. Distribution of the UOR data in WOD23 among the known contributing institutions**

The Oregon State University (Corvallis, OR) team also submitted data collected in the framework of the Joint Global Ocean Flux Study (JGOFS) – Antarctic Environment and Southern Ocean Process Study (AESOPS) Program - 6,828 casts from the Antarctic Polar Front Zone area.



**Figure 11.2. Distribution of the UOR data in WOD23 among the known contributing projects**

During the U.K. ARABESQUE project in 1994, 3,829 casts were collected in the Indian Ocean by groups from the Plymouth Marine Laboratory (PML, U.K.: 3553 casts) and the

University of Southampton – Department of Oceanography (USH-O, U.K.: 276 casts). The ARABESQUE project aimed to understand the microbial biogeochemistry in the upper ocean of the Arabian Sea. Its focus was to understand how the cycling processes of carbon and nitrogen were linked to climate change. The field program was timed to coincide with the Southwest Monsoon. It also continued from the inter-monsoon period until the onset of the Northeast Monsoon ([https://www.bodc.ac.uk/resources/products/data/bodc\\_products/](https://www.bodc.ac.uk/resources/products/data/bodc_products/)).

The UOR dataset also includes 363 profiles submitted by the British Oceanographic Data Center (BODC) that were collected during the Plankton Reactivity in the Marine Environment (PRIME) program, which was a thematic project funded by the National Environment Research Council of UK (NERC) to study plankton's role in oceanic biogeochemical fluxes. The PRIME data included in WOD23 were collected in the northeast Atlantic in 1996 (<https://www.bodc.ac.uk/resources/inventories/edmed/report/195/>).

BODC also submitted Ocean Margin Exchange (OMEX) project data (218 casts). The aim of the OMEX project was to study, observe, and model the physical, chemical and biological processes and fluxes taking place along the ocean margin, the interface between the open ocean, and the continental shelf. The first phase of the project, OMEX I, concentrated on studying the processes taking place along the northwest European shelf break. (<https://www.bodc.ac.uk/omex/>).

More recently (2012), the Fisheries and Oceans Canada – Marine Environmental Data Section (MEDS) submitted 32,937 UOR casts to NCEI, which were new to WOD18. Much of this data is off of the Northeast coast of North America (see Fig. 11.5). This data complements historical data from the Fisheries Research Board of Canada Biological Station in St. John, Newfoundland (FRBC-SJ: 1,439 casts). Finally, the Woods Hole Oceanographic Institution (WHOI) submitted 30 casts to NCEI as part of the Gulf Integrated Spill Response Consortium (GISR). These data are new to WOD23.

It should be noted that the current UOR data holdings in WOD23 is not exhaustive. There exists a substantial amount of UOR data up to the present day that has not been submitted to the National Centers for Environmental Information and is therefore not included in WOD23. If and when the data is submitted to NCEI, the data will then be added to WOD.

## **11.2. UOR DATA PRECISION AND ACCURACY**

The accuracy of UOR data depends on the performance of the sensors used and the post-processing of the data. A SeaSoar undulating vehicle is capable of carrying various instrumental packages. For the data stored in the WOD23 database, the Sea-Bird Electronics SBE 911*plus* CTD instrument was used most often. Please see section 3.2 for CTD accuracy information.

It is presumed that UOR data submitted into WOD23 were corrected for effects of:

- a) variable flow rate (Huyer *et al.*, 1993),
- b) thermal mass (Lueck, 1990; Morrison *et al.*, 1993), and
- c) offset between temperature and conductivity data (Larson, 1992; Morrison *et al.*, 1993).

### 11.3. UOR PROFILE DISTRIBUTIONS

OUR data represent 0.7% of the entire WOD23 data collection. Table 11.2 gives the yearly counts of UOR casts for the World Ocean and Figure 11.3 illustrates this graphically.

Table 11.2. The number of all UOR casts as a function of year in WOD23.

Total number of casts = 127,574

YEAR	CASTS	YEAR	CASTS	YEAR	CASTS	YEAR	CASTS
1976	1	1986	6	1996	2,286	2006	1,385
1977	0	1987	2	1997	8,254	2007	660
1978	0	1988	35	1998	1,429	2008	1,315
1979	0	1989	257	1999	1,634	2009	1,359
1980	0	1990	1,407	2000	10,344	2010	1,232
1981	0	1991	2,387	2001	1,299	2011	1,229
1982	4	1992	13,909	2002	1,391	2012	1,469
1983	1	1993	18,606	2003	4,043	2013	0
1984	2	1994	8,383	2004	39,961	2014	0
1985	0	1995	1,933	2005	1,331	2015	0

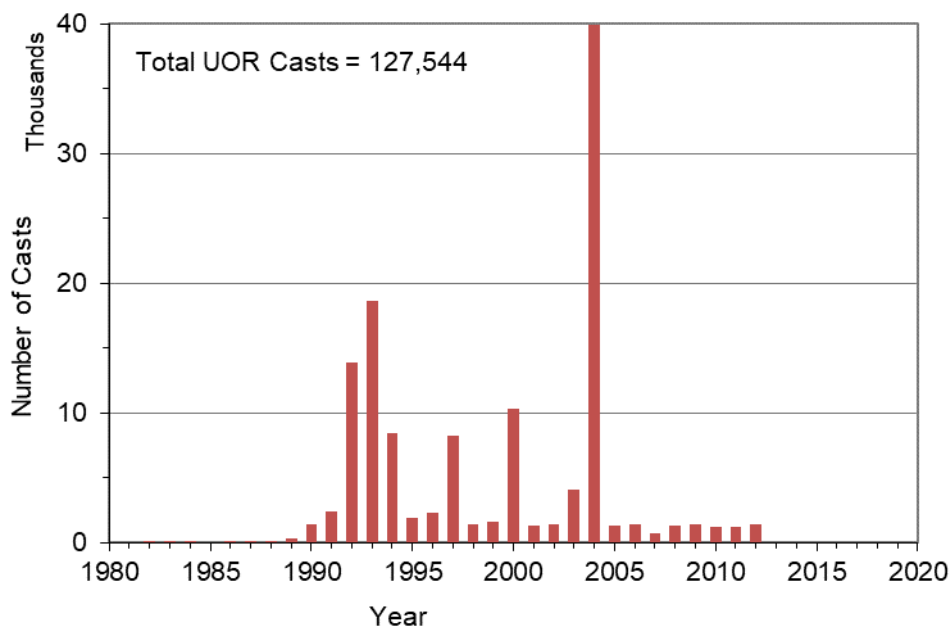


Figure 11.3. Temporal distribution of Undulating Ocean Recorders (UOR) casts in WOD23.

Table 11.3 gives the numerical and percentage contribution from various countries to the UOR dataset (also shown graphically in Figure 11.4). The geographic distribution of UOR casts is shown in Figure 11.5. Much of the data is localized and related to the projects discussed in section 11.1. Figure 11.6 illustrates the distribution of the UOR data as a function of depth at

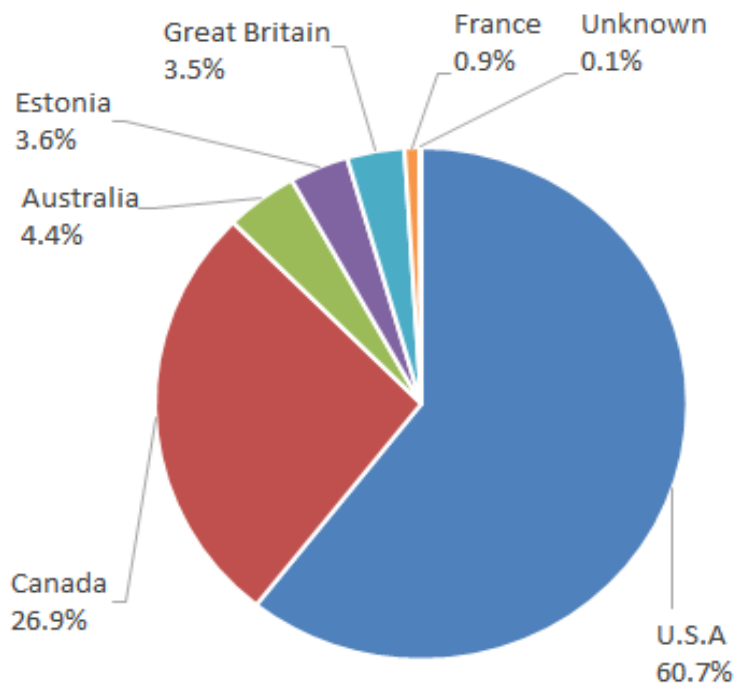


observed depth levels.

**Table 11.3. National contributions of UOR casts in WOD23.**

ISO <sup>1</sup> Country Code	Country Name	UOR Casts	% of Total
US	United States	77,423	60.68
CA	Canada	34,255	26.86
AU	Australia	5,597	4.39
EE	Estonia	4,650	3.64
GB	United Kingdom	4,410	3.46
FR	France	1,118	0.88
99	Unknown	121	0.09
<i>Total</i>		<i>127,574</i>	<i>100.00</i>

<sup>1</sup> ISO = International Organization for Standardization  
[http://www.iso.org/iso/country\\_codes.htm](http://www.iso.org/iso/country_codes.htm)



**Figure 11.4. Distribution of the Undulating Ocean Recorders (UOR) data in WOD23 among the contributing countries.**

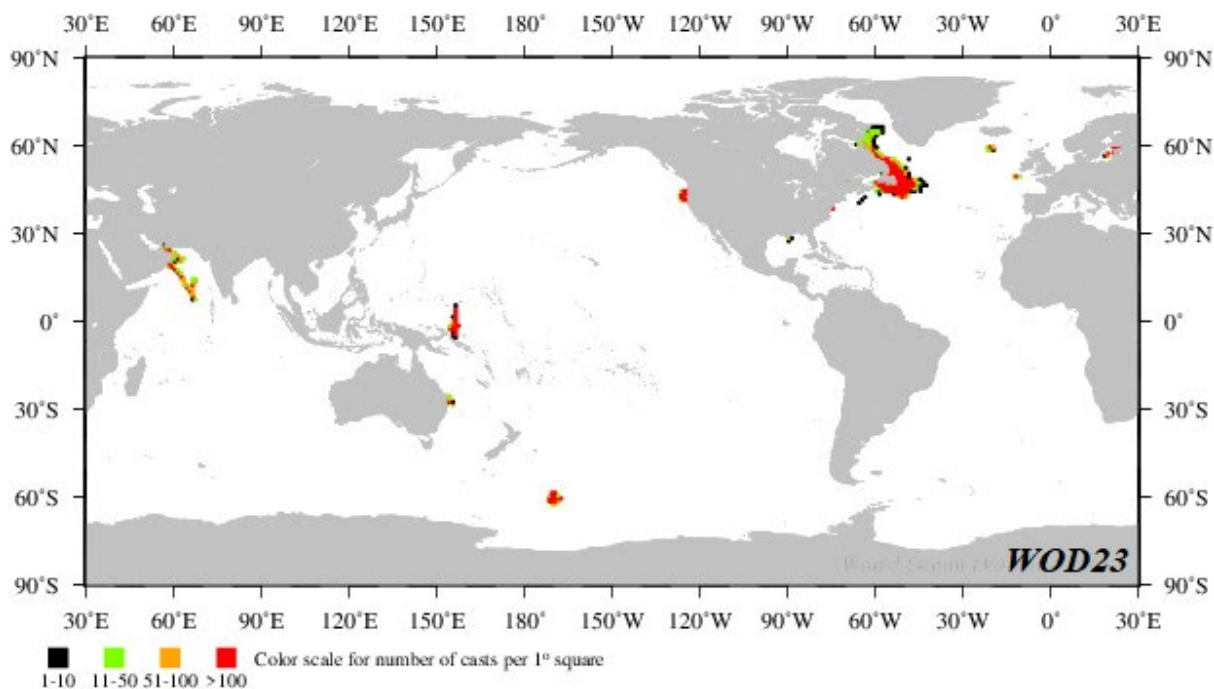


Figure 11.5. Geographic distribution of Undulating Ocean Recorders (UOR) casts in WOD23 by one-degree squares.

#### 11.4. RELEVANT WEB SITES

ARABESQUE Project:

[https://www.bodc.ac.uk/resources/products/data/bodc\\_products](https://www.bodc.ac.uk/resources/products/data/bodc_products) Australian Commonwealth Scientific and Industrial Research Organization (CSIRO): <http://www.csiro.au>

Chelsea Technologies:

<https://www.chelsea.co.uk/products/marine-science/towed-vehicles>

College of Oceanic and Atmospheric Sciences at Oregon State University, Corvallis, OR:

<http://ceos.oregonstate.edu/>

JGOFS-AESOPS: [http://usjgofs.whoi.edu/jg/dir/jgofs/southern/rr-kiwi\\_6/](http://usjgofs.whoi.edu/jg/dir/jgofs/southern/rr-kiwi_6/)

Ocean Margin Exchange (OMEX): <http://www.bodc.ac.uk/omex/>

Plankton Reactivity in the Marine Environment (PRIME):

<https://www.bodc.ac.uk/resources/inventories/edmed/report/195/> Plymouth Marine Laboratory: <http://www.pml.ac.uk/>

Tropical Ocean Global Atmosphere/Coupled Ocean Atmosphere Response Experiment (TOGA-COARE):

<https://www.ncei.noaa.gov/access/metadata/landing-page/bin/iso?id=gov.noaa.ncdc:C00577> ; <http://www.soest.hawaii.edu/COARE/>

WHOI SeaSoar page: <https://www.whoi.edu/what-we-do/explore/underwater-vehicles/towed-vehicles/seasoar/>

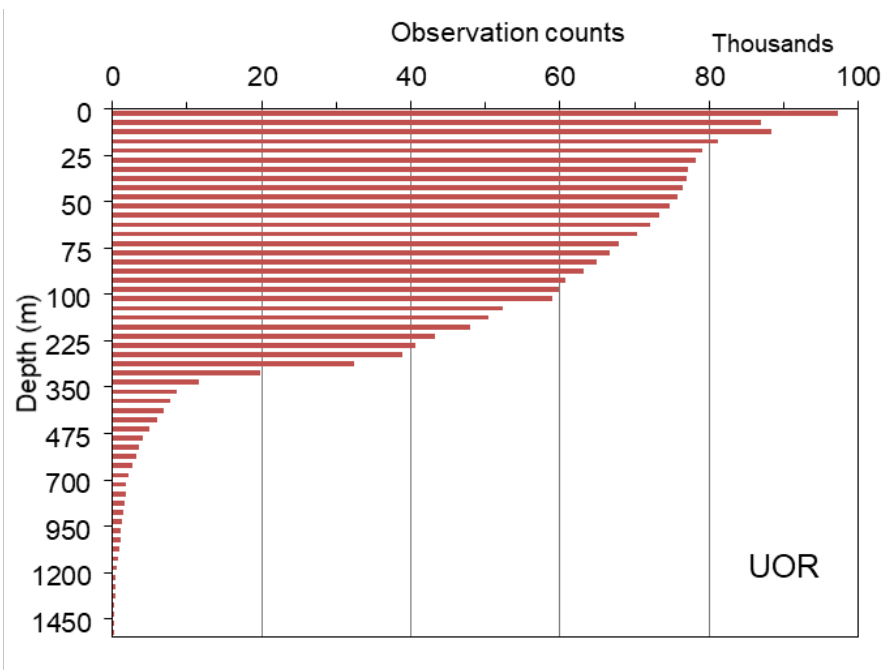


Figure 11.6. Distribution of Undulating Ocean Recorders (UOR) data at standard depth levels in WOD23.

## 11.5. REFERENCES AND BIBLIOGRAPHY

- Aiken, J. (1981). Undulating Oceanographic Recorder Mark 2. *J. Plankton Res.*, 3(4), 551-560.
- Barth, J.A. and D.J. Bogucki (2000). Spectral light absorption and attenuation measurements from a towed undulating vehicle. *Deep-Sea Res.*, 47, 323-342.
- Burt, R. (2000). Undulators come of age. International Ocean System, available at [https://www.researchgate.net/publication/291487460\\_Undulators\\_come\\_of\\_age](https://www.researchgate.net/publication/291487460_Undulators_come_of_age).
- Brown J., K. Brander, A.E. Hill (1996). Scafish: high performance towed undulator. *Sea Tech.*, 9, 23-27.
- Houghton, R.W., C.E. Tilburg, R.W. Garvine and A. Fong (2004). Delaware River plume response to a strong upwelling-favorable wind event. *Geophys. Res. Lett.*, 31(7): doi:10.1029/2003GL018988.
- Huyer, A., P.M. Kosro, R. O'Malley, J. Fleishbein (1993). Seasoar and CTD Observations during a COARE Surveys Cruise, W9211C, 22 Jan to 22 Feb 93, OSU Data Report.
- Larson, N. (1992). *Oceanographic CTD Sensors: Principles of operation, sources of error, and methods for correcting data*. Sea-Bird Electronics, Inc. Bellevue, Washington, USA.
- Longhurst, A.R., A.D. Keith, A.D. Bower and D.L.R. Seibert (1966). A new system for the collection of multiple serial plankton samples. *Deep-Sea Res.* 13, 213-222.
- Lueck, R.G. (1990). Thermal inertia of conductivity cells: Theory. *J. Atmosph. Oceanic Tech.* 7(5), 741-755.
- Morison, J., R. Andersen, N. Larson (1993). The Correction for Thermal-Lag Effects in Sea-Bird CTD Data. *J. Atmos. Oceanic Tech.* 11(4), 1151-1164.

- Pollard, R. (1986). Frontal surveys with a towed profiling conductivity/temperature/depth measurement package (SeaSoar). *Nature*, 323, 433-435.
- Reid, P.C., J.M. Colebrook, J.B.L. Matthews, J. Aiken, Continuous Plankton Recorder Team (2003). The Continuous Plankton Recorder: concepts and history, from Plankton Indicator to undulating recorders. *Progress in Oceanography*, 58, 117-173, doi:10.1016/j.pocean.2003.08.002.
- Williams, R. and J.A. Lindley (1980). Plankton of the Fladen Ground during FLEX 76. I. Spring development of the plankton community. *Mar. Biol.*, 57(2), 73-78.
- Williams, R. and J.A. Lindley (1998). Strategy and application for sampling Large Marine Ecosystems with the Continuous Plankton Recorder and Undulating Oceanographic Recorder/Aquashuttle. *Large marine ecosystems of the Indian Ocean: assessment, sustainability and management*. Ed. By K. Sherman, E.N. Okemwa and M.J. Ntiba. Oxford, Blackwell Science, 45-60.

# CHAPTER 12: AUTONOMOUS PINNIPED BATHYTHERMOGRAPH DATA (APB)

*Alexey V. Mishonov, Ricardo A. Locarnini, Olga K. Baranova, Tim P. Boyer*

*Ocean Climate Laboratory  
National Centers for Environmental Information  
Silver Spring, MD*

## 12.1. INTRODUCTION

The first usage of marine mammals as sampling platforms is credited to Pers Scholander (Scholander, 1940; cited after Fedak, 2004 and Ropert-Coudert and Wilson, 2005). Based on a description of a depth gauge provided by Lord Kelvin in the 19<sup>th</sup> century, Scholander developed depth gauges to record diving depths of whales. Various research groups have used other marine mammals to carry sensors and data loggers followed up this pioneering work.

Data from sensor instruments attached to marine animals (instrumented animals) such as sea turtles, sea birds, sharks, tuna, and marine mammals, were initially collected for the principal purpose of studying animal ecology (Le Bœuf *et al.*, 1988; Block, 2005). In addition to animal ecology studies, scientists can use instrumented animals as autonomous ocean profilers to enhance sparse oceanographic observations in specific oceanic regions (McCafferty *et al.*, 1999; Boehlert *et al.*, 2001; Charrassin *et al.*, 2002; Lydersen *et al.*, 2002; Hooker and Boyd, 2003; Fedak, 2004; Ropert-Coudert and Wilson, 2005; Roquet *et al.*, 2009; Padman *et al.*, 2010). The data supplied by instrumented animals could potentially fill data gaps due to harsh environmental conditions in areas such as the Bering Sea, Gulf of Alaska, and the Southern Ocean, especially in winter and in ice-covered waters. These data can also help fill spatial and temporal gaps due to remoteness of some areas such as the southeast Pacific, and spatial gaps between routes of Ships-of-Opportunity (or Voluntary Observing Ships, [VOS](#)). The use of instrumented marine animal data has been shown to improve state estimates of under-sampled regions like the Southern Ocean (Roquet *et al.*, 2013).

Temperature profiles from instrumented animals are less expensive than those obtained by traditional instruments such as Expendable Bathythermographs (XBT) (Boehlert *et al.*, 2001). After recovering instruments from animals, the equipment can be re-used. The vertical resolution of the available pinniped data is better than the vertical resolution of bottle station data but generally worse than the resolution of XBT and Conductivity-Temperature-Depth (CTD) data.

The Autonomous Pinniped Bathythermograph (APB) dataset presented in the WOD23 contains *in situ* temperature data from temperature-depth records (TDRs) and conductivity-temperature-depth satellite relay data loggers (CTD-SRDLs) attached to pinnipeds (*e.g.*, elephant seals). The instrumented animals include northern elephant seals (*Mirounga angustirostris*), southern elephant seals (*Mirounga leonina*), Weddell seals (*Leptonychotes weddellii*) and narwhals (*Monodon monoceros*).

## 12.2. DATA SOURCES

The APB data that comprise WOD23 have been acquired through different sources and projects. Table 12.1 shows the contributing projects and number of casts submitted. Geographic positions were determined using the ARGOS satellite transmitters. The half-watt satellite platform transmitter terminals (PTT; Model ST-6, Telonics, Mesa, Arizona) were affixed near the animal's head using epoxy. The antenna was oriented to be out of the water when the seal surfaced. The PTT transmitted every 34 seconds while the seals were at the surface (Boehlert *et al.*, 2001).

The Autonomous Pinniped Environmental Samplers Project ([APES](#)), a project under the NOAA/National Marine Fisheries Services (NOAA/NMFS), submitted northern elephant seal data equipped with the [WildLife Computers Mk3](#) TDR. The Tagging of Pacific Predators ([TOPP](#)), a project under the Census of Marine Life, submitted northern elephant seal data equipped with the [WildLife Computers Mk9](#) and [Mk10](#) TDRs. The Sea Mammal Research Unit in Scotland ([SMRU](#)) established in 1978 by the Natural Environment Research Council (NERC, UK) initiated the Southern Elephant Seals as Oceanographic Samplers ([SEaOS](#)) program in 2004, which submitted southern elephant seal data deployed with Autonomous CTD-Satellite Relay Data Loggers (CTD-SRDLs). Data were also received from the Australian Integrated Marine Observing System (IMOS).

The Marine Mammals Exploring the Oceans Pole to Pole ([MEOP](#)) project started in the framework of the International Polar Year (IPY) project in 2008, aiming at coordinating the cooperation between several national programs to provide a comprehensive, synoptic oceanographic coverage in Polar Regions. This consortium, a collaboration of scientists from many countries who work on instrumented marine animals, serves as a Data Assembly Center for instrumented marine animal data. They combine global data from national observing programs (Australia, Brazil, Canada, China, Denmark, France, Germany, Norway, South Africa, Sweden, the United Kingdom, and the USA) and periodically release a quality-controlled database. They also release best practices for collection and data formats and maintain documentation and links to publications that feature the data. More details can be found at the relevant [Copernicus web site](#) and in (Roquet *et al.*, 2013, 2014).

NCEI also received some APB near-real time data from IMOS, [SMRU](#) and [SEaOS](#) through the Global Temperature-Salinity Profile Program ([GTSP](#)) system. The [GTSP](#) system provides NCEI with most near-real time APB profiles. Some programs provide delayed-mode profiles later and WOD replaces the near-real time data when post-processed versions are available. About 342,018 profiles are still near-real time versions from GTSP. At the time of the writing, WOD Team is working on inclusion of the most recent update of the MEOP database in WOD.

## 12.3. INSTRUMENTATION AND ACCURACY

The northern elephant seals were deployed with [WildLife Computers](#) (Redmond, USA) Mk3, [Mk9](#), and [Mk10](#) TDRs.

Earlier submissions of northern elephant seal data used the Mk3 TDR. This is a slower responding internal thermistor using a simple time lag to account for the response of the thermistor

(Simmons *et al.*, 2009; Boehlert *et al.*, 2001). The Mk3 has a temperature resolution of 0.1°C and an accuracy of 0.5°C with a manufacturer’s stated minimum recording temperature of 4.8°C (Boehlert *et al.*, 2001). The pressure transducers on the TDR were calibrated prior to deployment using a pressure station. The Mk 3 TDRs used had two transducer channels. In order to increase the accuracy on shallower dives, TDRs were programmed to use channel #1 for depths <450 m (with accuracy <2 m) and channel #2 for depths >450 m (with accuracy <4 m) (Boehlert *et al.*, 2001).

**Table 12.1. Projects contributing to the WOD23 APB dataset and number of profiles submitted.**

Project or data originator	Number of casts
Autonomous Pinniped Environmental Samplers (APES, WOD project code 371)	75,665
SMRU Southern Elephant Seals Oceanographic Samplers (SEAOS, WOD project code 586)	2,129
Integrated Marine Observing System (IMOS, WOD DB code 18)	140,378
Marine Mammals Exploring the Oceans Pole to Pole (MEOP, WOD project code 705)	323,720
Tagging of Pacific Predators (TOPP, WOD project code 682)	1,171,943
Data submitted via GTSP without project information (WOD DB code 2)	342,023
Data without project/institutions/DB information	18
<i>Total:</i>	<i>2,051,828</i>

The [Mk9](#) and [Mk10](#) are fast-response (*i.e.* fastloc) archival TDRs. Both TDRs are configured with multiple sensors. The depth sensor has a 12-bit analog-to-digital converter; it provides highly accurate measurements from -40 to +1000 m, with 0.5 m resolution and an accuracy of  $\pm 1\%$ . In addition, measurements from 1000 to 1500 m are made with a lesser degree of accuracy. Measurements can be recorded throughout the range at full resolution. The temperature sensor is a 12-bit analog-to-digital converter; it has a range of -40° to +60°C, with 0.05°C resolution and an accuracy of  $\pm 0.1^\circ\text{C}$ . Measurements can be recorded throughout the range at full resolution.

The southern elephant seals were equipped with CTD-SRDs, a specific configuration of [Valeport](#)’s CTD. The CTD-SRDs are designed and manufactured by [SMRU](#). Temperature was measured by the [Valeport](#) fast response Platinum Resistance Thermometer (PRT), with a range of -5°C to +35°C, an accuracy of  $\pm 0.005^\circ\text{C}$ , and a resolution of 0.001°C. Conductivity was measured by the [Valeport](#) inductive coils with a range of 0 to 80mS/cm, an accuracy of  $\pm 0.01\text{mS/cm}$ , and a resolution of 0.002mS/cm. Pressure was measured by the Keller PA-3L sensor, with a range of 2000 dbar, an accuracy of 2 dbar  $\pm (0.3 \text{ to } 0.035\% \cdot \text{reading})/ \text{K}$ , and a resolution of 0.05 dbar. The TDR capability can retain a continuous record of depth readings (4 sec sample rate), which can be retrieved by Bluetooth link if the tag is recovered. During profiling, the TDR records all individual temperature and salinity measurements at 1 Hz during profiling. (Boehme *et al.*, 2009).

The latest version of a miniaturized CTD-Satellite Relay Data Loggers (CTD-SRD) is the ARGOS tag 9000 series CTD-SRD designed and built at the [SMRU](#) (University of St. Andrews, UK) and calibrated at [Valeport](#) Ltd (Devon, UK). It has a 401 MHz RF unit and antenna for data

transfer by way of the ARGOS system, a lithium-thionyl chloride (Li-SOCl<sub>2</sub>) D-cell battery (LSH 201) and a Hitachi H8/3048 microprocessor programmed to act as the data logger, data compression tool and to schedule data transfer (*i.e.* Boehme *et al.*, 2009).

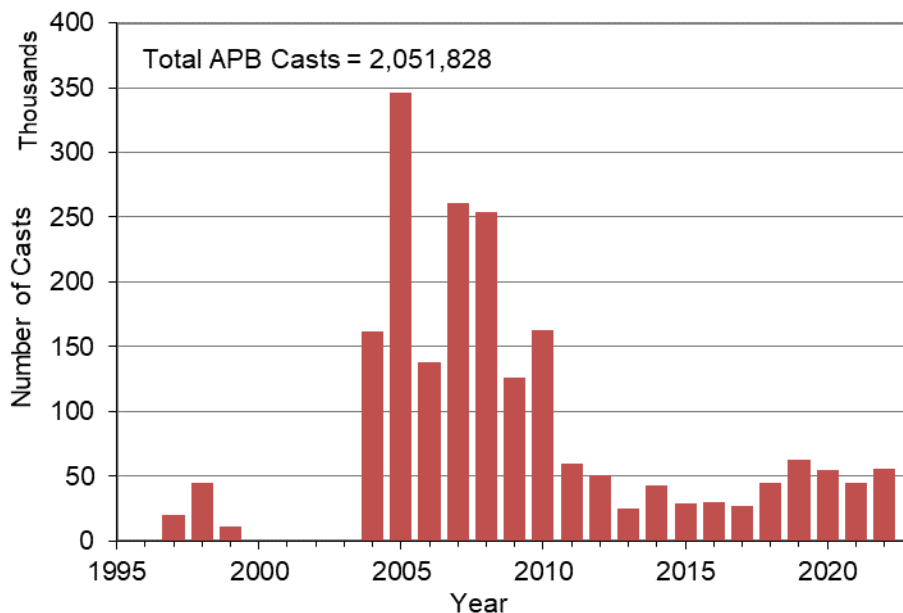
### 12.4. GEOGRAPHICAL, TIME, AND DEPTH DISTRIBUTION OF DATA

The WOD23 has a total of 2,051,828 vertical profiles collected from APB between 1997 and 2022 (Figure 12.1 and Table 12.2). APB data represents 11% of the WOD23 collection. Figure 12.2 shows the geographic distribution of the dataset, and Figure 12.3 shows depth distribution of the data on the standard depth levels of the entire APB data set.

**Table 12.2. The number of all APB casts as a function of year in WOD23.**

Total number of casts =2,051,828

YEAR	CASTS	YEAR	CASTS	YEAR	CASTS	YEAR	CASTS
1997	19,875	2004	162,219	2011	29,250	2017	26,878
1998	44,626	2005	346,045	2012	50,893	2018	44,748
1999	11,164	2006	137,722	2013	25,307	2019	62,405
2000	0	2007	260,772	2014	42,894	2020	54,268
2001	0	2008	254,092	2015	29,072	2021	45,320
2002	0	2009	126,029	2016	30,187	2022	55,713
2003	0	2010	162,349				



**Figure 12.1. Temporal distribution of the Autonomous Pinniped Bathythermograph (APB) casts in WOD23.**



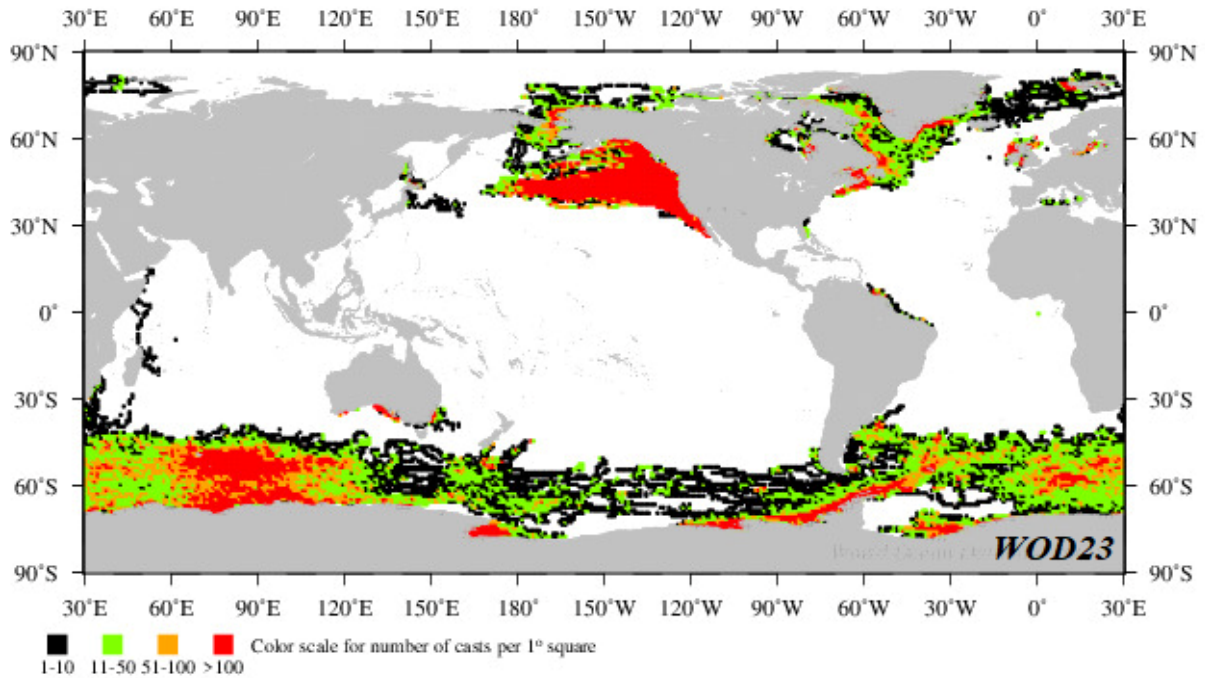


Figure 12.2. Geographical distribution of the Autonomous Pinniped Bathythermograph (APB) data in WOD23 by one-degree squares.

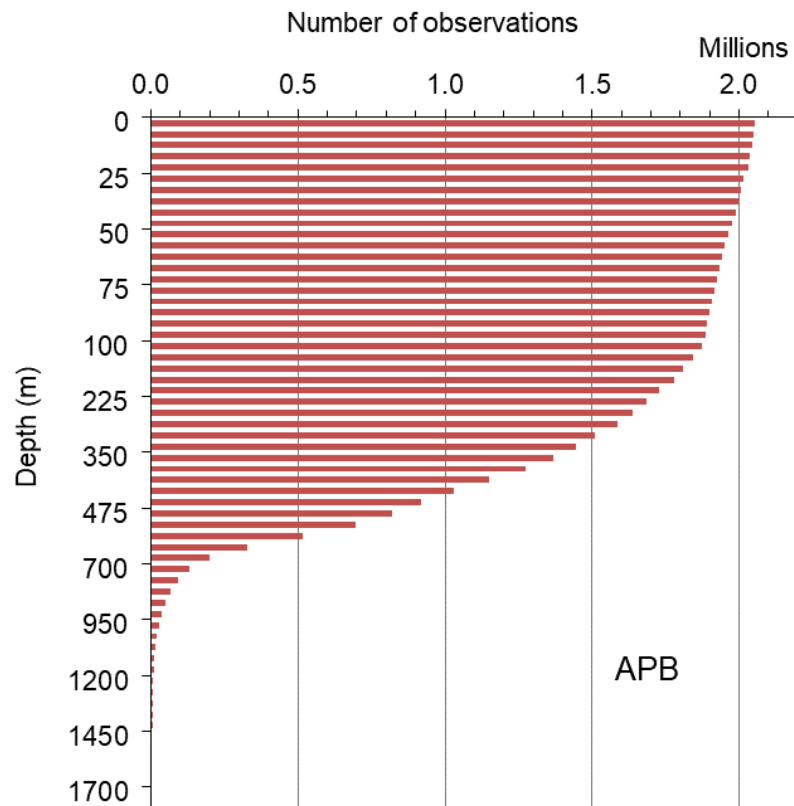


Figure 12.3. Distribution of the Autonomous Pinniped Bathythermograph (APB) data at standard depth levels in WOD23.

## 12.5. REFERENCES AND BIBLIOGRAPHY

- Block, B.A. (2005). Physiological ecology in the 21<sup>st</sup> century: Advancements in biologging science, *Integrative and Comparative Biol.*, 45, 305-320.
- Boehlert, G.W., D.P. Costa, D.E. Crocker, P. Green, T O'Brien, S. Levitus, and B.J. LeBoeuf (2001). Autonomous pinniped environmental samples: using instrumented animals as oceanographic data collectors, *J. Atmos. Oceanic Technol.*, 18: 1882-1893.
- Boehme, L., P. Lovell, M. Biuw, F. Roquet, J. Nicholson, S.E. Thorpe, M.P. Meredith, and M. Fedak (2009). Technical Note: Animal-borne CTD-Satellite Relay Data Loggers for real-time oceanographic data collection, *Ocean Sci.*, 5, 685-695.
- Charrassin, J.B., Y.H. Park, Y.Le Maho, and C.A. Bost (2002). Penguins as oceanographers unravel hidden mechanisms of marine productivity, *Ecol. Lett.*, 5(3), 317-319.
- Fedak, M. (2004). Marine animals as platforms for oceanographic sampling: a “win/win” situation for biology and operational oceanography, *Memoirs of the National Institute of Polar Research (Japan)*, Special Issue, 58, 133-147.
- Hooker, S.K., and I.L. Boyd (2003). Salinity sensors on seals: use of marine predators to carry CTD data loggers, *Deep-Sea Res. I*, 50(7), 927-939.
- Le Boeuf, B.J., D.P. Costa, A.C. Huntley, and S.D. Feldkamp (1988). Continuous deep diving in female northern elephant seals, *Mirounga angustirostris*. *Canadian J. Zoology*, 66, 446-458.
- Lydersen, C., O.A. Nøst, P. Lovell, B.J. McConnell, T. Gammelsrod, C. Hunter, M.A. Fedak, and K.M. Kovacs (2002). Salinity and temperature structure of a freezing Arctic fjord – monitored by white whales (*Delphinapterus leucas*). *Geophys. Res. Lett.*, 29(23), 2119.
- McCafferty, D.J., I.L. Boyd, T.R. Walker, and R.I. Taylor (1999). Can marine mammals be used to monitor oceanographic conditions? *Marine Biol.*, 134, 387-395.
- Padman, L., D.P. Costa, S.T. Bolmer, M.E. Goebel, L.A. Huckstadt, A. Jenkins, B.I. McDonald, and D.R. Shoosmith (2010). Seals map bathymetry of the Antarctic continental shelf. *Geophys. Res. Lett.*, 37, L21601, 1-5.
- Ropert-Coudert, Y., R.P. Wilson (2005). Trends and perspectives in animal-attached remote sensing. *Frontiers in Ecology and the Environment*, 3(8), 14-17.
- Roquet, F., Y.-H. Park, C. Guinet, F. Bailleul, J.-B. Charrassin (2009). Observations of the Fawn Trough Current over the Kerguelen Plateau from instrumented elephant seals. *Journal of Marine Systems*, 78(1-3), 377-393.
- Roquet, F.C., and 28 co-authors (2013). Estimates of the Southern Ocean general circulation improved by animal-borne instruments, *Geophys. Res. Lett.*, 40, doi:10.1002/2013GL058304.
- Roquet F., Williams G., Hindell M. A., Harcourt R., McMahon C. R., Guinet C. *et al.* (2014). A Southern Indian Ocean database of hydrographic profiles obtained with instrumented elephant seals. *Nature Scientific Data*, 1:140028, doi: 10.1038/sdata.2014.28.
- Scholander, P.F. (1940). Experimental investigations on the respiratory function in diving mammals and birds. *Hvalrådets Skrifter* 22:1-131.

- Simmons, S.E. (2009). Pinnipeds as ocean-temperature samplers: calibrations, validations, and data quality. *Limnol. Oceanogr.: Methods* 7, 2009, 648-656.
- Treasure, A.M., and 27 co-authors (2017). Marine Mammals Exploring the Oceans Pole to Pole: A review of the MEOP consortium. *Oceanography* 30(2):132–138, <https://doi.org/10.5670/oceanog.2017.234>.

# CHAPTER 13: MICRO BATHYTHERMOGRAPH DATA (MICRO BT)

*Courtney N. Bouchard, Alexey V. Mishonov, Tim P. Boyer, Ricardo A. Locarnini*

*Ocean Climate Laboratory  
National Centers for Environmental Information  
Silver Spring, MD*

## **13.1. INTRODUCTION**

The Micro Bathythermograph (Micro BT) is a high-accuracy temperature and pressure instrument developed to record and report data electronically. WOD23 includes data collected with micro BT instruments manufactured by RBR Ltd. and Sea-Bird Electronics (SBE). The self-contained underwater instrument includes a rapid response thermistor and a strain gauge pressure sensor. Temperature and depth/pressure measurements are automatically archived in the underwater unit as it is lowered in the water column attached to a net, cable, or towed vehicle. The instrument can be programmed to measure and archive data at desired intervals. Upon retrieval, the underwater unit is connected to a computer and data are retrieved and archived. The micro BT instruments can also provide real time data using an underwater cable.

Micro BT instruments can measure temperatures over a varied range of depths, with RBR Ltd. instruments being able to measure to a maximum depth of 1000 m, and SBE instruments to a maximum depth of 7000 m.

All micro BT profiles are stored in the MBT dataset of WOD23. It should be noted that the micro BT instrument is obsolete now and not in use anymore, therefore, there is no new micro BT data in WOD23 compared to WOD18.

## **13.2. MICRO BT ACCURACY**

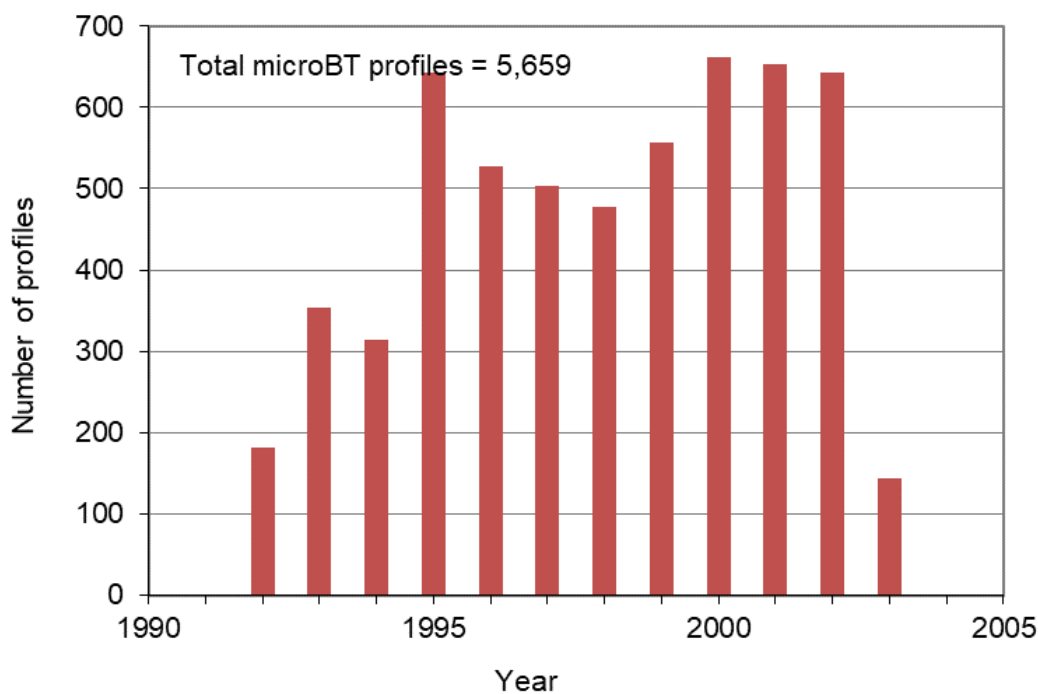
RBR Ltd. reports a temperature resolution of 0.1°C, and SBE reports a temperature accuracy of  $\pm 0.002^\circ\text{C}$ . Both manufacturers report a pressure accuracy of  $\pm 0.1\%$  of full-scale range.

## **13.3. MICRO BT PROFILE DISTRIBUTIONS**

Table 13.1 gives the yearly counts of micro BT profiles for the World Ocean. Fig. 13.1 shows the temporal distribution of Micro Bathythermograph profiles for the World Ocean. There is a total of 5,659 micro BT profiles for the entire World Ocean, all measured in the northern hemisphere (Figure 13.2) by the U.S. (Table 13.2). Distribution of the micro BT data at the WOD23 standard depth levels is shown in Figure 13.3.

**Table 13.1. The number of all Micro BT profiles as a function of year in WOD23.  
Total Number of Profiles = 5,659**

YEAR	CASTS	YEAR	CASTS	YEAR	CASTS	YEAR	CASTS
1992	182	1995	642	1998	478	2001	653
1993	354	1996	528	1999	556	2002	643
1994	314	1997	504	2000	662	2003	143



**Figure 13.1. Temporal distribution of micro Bathythermograph data in WOD23.**

**Table 13.2. National contributions of Micro Bathythermograph profiles in WOD23.**

ISO <sup>1</sup> Country Code	Country Name	Micro BT Count	% of Total
US	United States	5,659	100.0

<sup>1</sup>ISO = [International Organization for Standardization](http://www.iso.org)

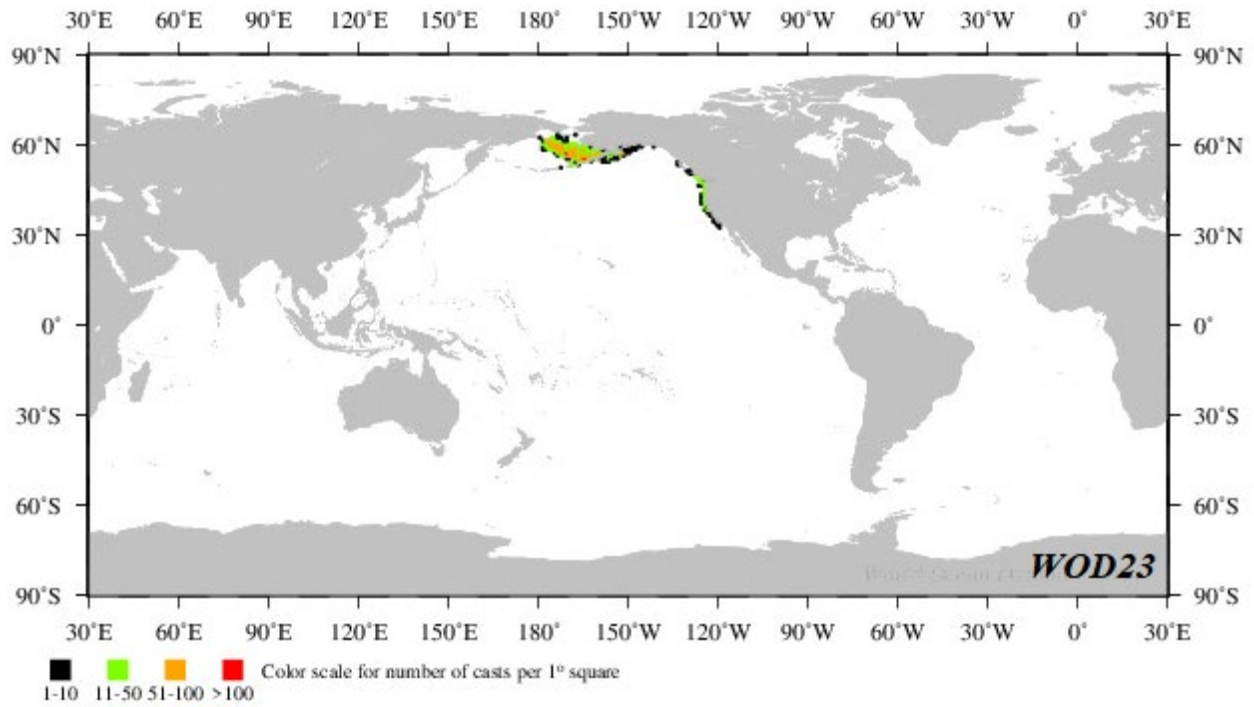


Figure 13.2. Geographic distribution of Micro Bathythermograph data in WOD23.

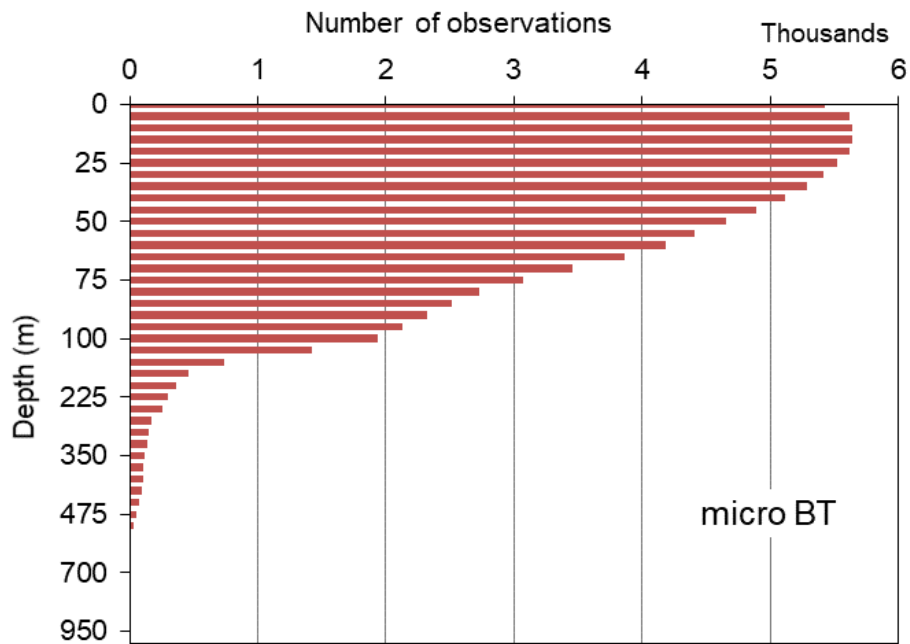


Figure 13.3. Distribution of micro Bathythermograph data at standard depth levels in WOD23.

# CHAPTER 14: SURFACE-ONLY DATA (SUR)

*Alexey V. Mishonov, Zhankun Wang, Tim P. Boyer, Ricardo Locarnini, James R. Reagan*

*Ocean Climate Laboratory  
National Centers for Environmental Information  
Silver Spring, MD*

## 14.1. INTRODUCTION

The major focus of the WOD23 is sub-surface profile data. Therefore, surface data are included in WOD23 only either if/when it has been collected together with measurements of other oceanographic variables of interest (*e.g.*, chlorophyll, CO<sub>2</sub>, pH, etc.; see Table 14.1), or if the data covers under-sampled time periods (*e.g.*, ICES Atlantic data for 1900-1939), or data provided by scientific ship-of-opportunity (SOOP) programs. For example, WOD23 contains many observations from the Institut de Recherche et Développement (IRD), formerly ORSTOM, which provided NCEI with surface salinity data from SOOP for the Tropical Pacific (Henin and Grelet, 1996). It should be noted that surface-only data-oriented projects exist, which hold much more comprehensive surface data collections than WOD23. For example, the International Comprehensive Ocean-Atmosphere Data Set (ICOADS) contains more than 455 million individual marine reports (Freeman *et al.*, 2017) and the NCEI [Surface Underway Marine Database](#) (NCEI-SUMD, Wang, 2018), which is a combination of the Shipboard Automated Meteorological and Oceanographic System (SAMOS, Shawn *et al.*, 2009), the Global Ocean Surface Underway Data (GOSUD, Loic, 2015), and the Atlantic Oceanographic and Meteorological Laboratory Thermosalinograph dataset (AOML-TSG, Goni, 2017). The SUMD contains over 450 million (temperature and/or salinity) observations from more than 490 ships and Un-crewed Surface Vehicles (Wang, 2018). The majority of the SUR data in WOD23 are salinity, temperature, and mole fraction of CO<sub>2</sub> in seawater (Table 14.1). Table 14.2 lists the number of SUR observations as a function of year of collection since 1867.

It should be noted that due to the reasons mentioned above, no new SUR data have been added to WOD collection between WOD18 and WOD23 releases. Therefore, all statistics below are still the same as in the previously released [WOD18 documents](#). In this document the data density bins were updated to refresh the distribution map (Figure 14.2) and data distribution map for the 19<sup>th</sup> century has been added to highlight the early researchers' efforts in Arctic (Figure 14.3) before year of 1900.

## 14.2. SUR ACCURACY

Samples of the seawater may have been collected from the continuous flow of water pumped from subsurface depths (*e.g.*, ship's water intake) or have been drawn from a bucket deployed from the ship deck over the board. A comprehensive review of the sampling techniques and its influence on the collected data precision can be found in Reverdin *et al.* (1994). When data

came from bucket samples, the precision of the sea surface salinity is believed to be about  $\pm 0.1$  (Delcroix and Picaut, 1998; Delcroix *et al.*, 2005). When data were collected by Thermosalinograph (TSG), sea surface salinity and temperature readings were recorded approximately every 10 seconds (Thomas *et al.*, 1999). Data precision of more modern measurements is limited by the characteristics of the instrument (Delcroix *et al.*, 2005). The accuracy of a Sea-Bird SBE 45 MicroTSG sensor (example of a modern day TSG instrument) is  $\pm 0.002^\circ\text{C}$  for temperature and  $\pm 0.005$  for salinity (see: <https://www.seabird.com/sbe45-microtsg-thermosalinograph/product?id=54627900541>).

**Table 14.1. List of parameters and number of observations in the SUR dataset of WOD23.**

<b>Parameter [nominal abbreviation]</b>	<b>Reporting unit (nominal abbreviation)</b>	<b>Number of observations</b>
Temperature [t]	Degree centigrade ( $^\circ\text{C}$ )	506,062
Salinity [S]	Unit less	1,958,361
pH	Unit less	84
Total Chlorophyll [Chl] unless specified	Micro-gram per liter ( $\mu\text{g}\cdot\text{l}^{-1}$ )	44,256
Phaeophytin	Micro-gram per liter ( $\mu\text{g}\cdot\text{l}^{-1}$ )	119
Primary Production	Micro-gram of Carbon per liter-day ( $\mu\text{gC}\cdot\text{l}^{-1}\cdot\text{day}^{-1}$ )	119
Alkalinity [TALK]	Milli-equivalent per liter ( $\text{meq}\cdot\text{l}^{-1}$ )	84
Partial Pressure of Carbon Dioxide [pCO <sub>2</sub> ]	Micro-atmosphere ( $\mu\text{atm}$ )	37,124
Mole fraction of CO <sub>2</sub> in seawater [XCO <sub>2</sub> sea]	Parts per million (ppm)	132,793
Conductivity	Siemens per meter ( $\text{S}\cdot\text{m}^{-1}$ )	52,284
CO <sub>2</sub> warm <sup>1</sup>	Degree centigrade ( $^\circ\text{C}$ )	60,262
Mole fraction of CO <sub>2</sub> in atmosphere [XCO <sub>2</sub> atm]	Parts per million (ppm)	169,469
Barometric pressure	Millibar (mb)	164,881
Latitude	Degrees of latitude	2,098,020
Longitude	Degrees of longitude	2,098,020
Julian Day	Day	2,098,020

<sup>1</sup> CO<sub>2</sub>warm is the temperature change (e.g., warming) for seawater as it transits from the ship's water intake line to the CO<sub>2</sub> analysis instrumentation location.

### **14.3. DATA COVERAGE**

The earliest surface temperature data included in WOD23 were collected in 1867 by Norwegian sailors from the ships *Isbjornen* and *Ishavet* in the North Sea, Norwegian Sea, and in the North Atlantic waters around Iceland (Table 14.2). SUR data were routinely collected in the 19<sup>th</sup> century (insert on Figure 14.1 and Figure 14.3), but most of the SUR data were collected after the late 1990s (Figure 14.1). The WOD23 SUR dataset consists of 9,257 cruises (Figure 14.2). Surface data collected before 1955 were often bucket samples, data acquired after 1957 were, most often, from thermosalinographs (TSG) and other underway systems.



There are noticeable data gaps after the First World War (1914-1918) and during the Second World War (1939-1945) and after it up to 1960. A huge increase in surface data collection (mainly *in situ* sea surface temperature and salinity measurements) occurred in the 1990s. These data were mainly acquired by the TSG instruments mounted on ships-of-opportunity. Data collected over that period contains more than 70% of the entire SUR dataset in WOD23 with almost all data being collected along shipping routes in the Pacific and northern Atlantic Ocean (Figure 14.2). The major country contributions to the SUR dataset are France (41.8% of all SUR data) and Australia (32.5%). Unlike many other datasets in WOD where temperature dominates the data collection/distribution, salinity dominates the data held in SUR (Table 14.1). This is due to many salinity-only data contributions (1,589,463 observations) from the Oceanic Lab of the Institution of French Oceania (Noumea, New Caledonia) during the 1969-1999 time period. Adding salinity observations from SUR improves the surface data coverage in the World Ocean Atlas.

Table 14.3 lists the input of data to the SUR dataset by country of origin. Figure 14.2 shows that the majority of SUR data were collected along the main commercial shipping routes of the Atlantic and Pacific Oceans. In terms of volume of data, about 87% of the observations in SUR were acquired from two main sources: International Council for the Exploration of the Sea (ICES) and the Oceanic Lab of the Institution of French Oceania (Noumea, New Caledonia). The remaining 13% came from the National Institute of Oceanography in India, Scripps Institution of Oceanography, National Institute for Environmental Studies, Institute of Ocean Sciences, Sidney, Australia, the Office of Scientific and Technical Research Overseas (ORSTOM) - New Caledonia (before independence), and several others.

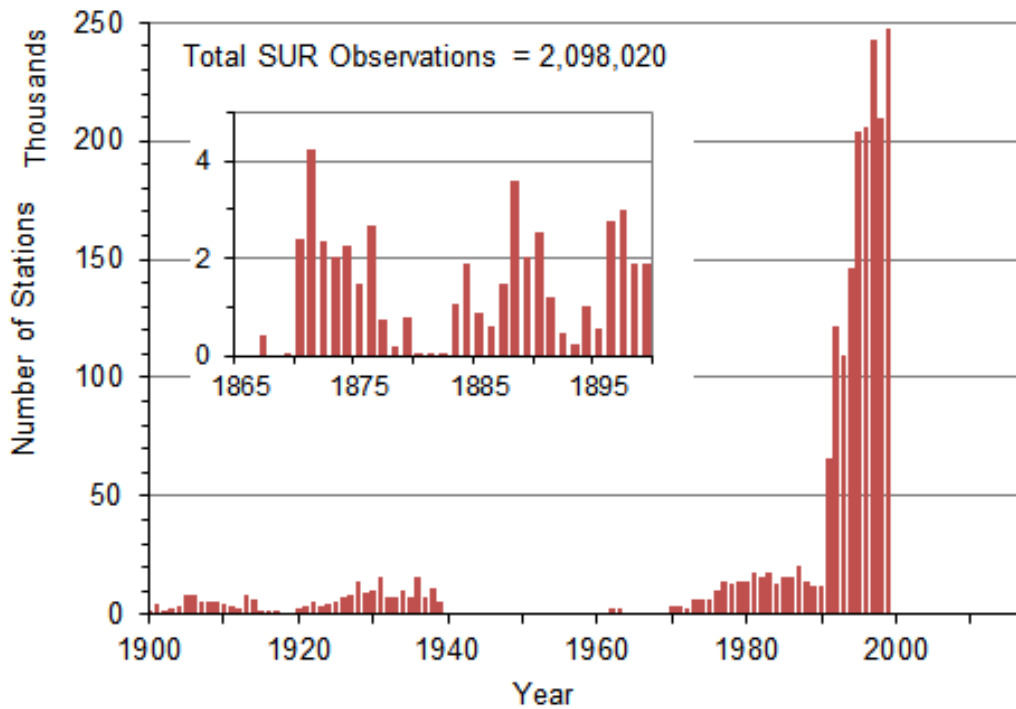


Figure 14.1. Temporal distribution of surface (SUR) observations in WOD23.

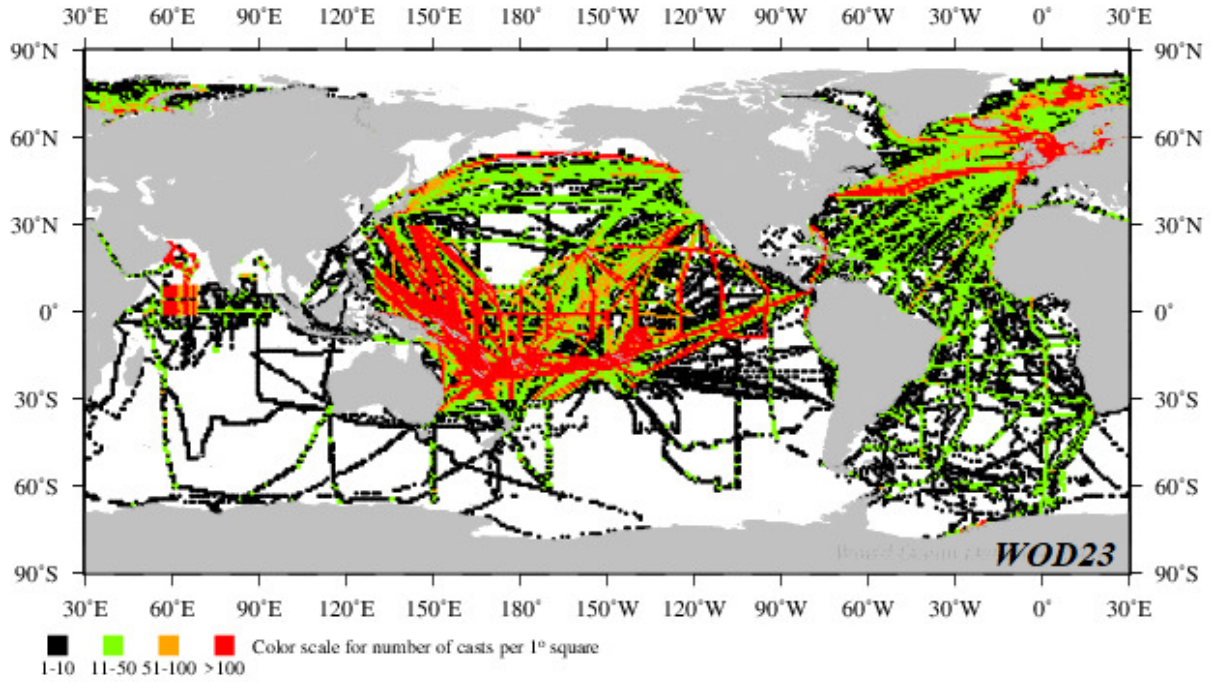


Figure 14.1. Temporal distribution of surface (SUR) observations in WOD23.

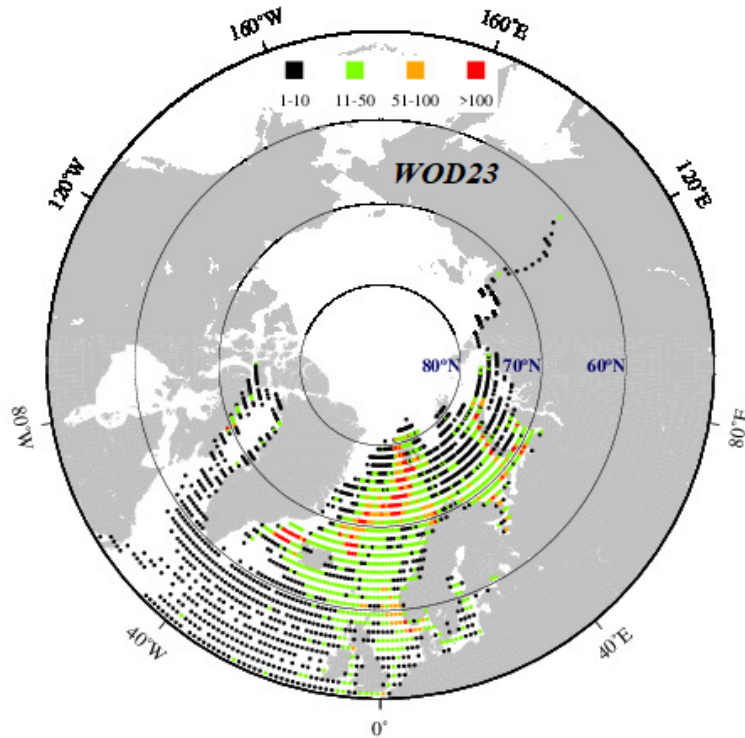


Figure 14.3. Geographic distribution of surface (SUR) observations by one-degree squares collected by European polar researchers in Arctic before 1900 year in WOD23.

**Table 14.2. The number of all SUR observations as a function of year in WOD23.**

Total number of observations (counts) = 2,098,020

Year	Amount	Year	Amount		Year	Amount		Year	Amount
1867	398	1903	2,242		1939	5,740		1975	803
1868	0	1904	3,695		1940	48		1976	2,902
1869	44	1905	8,609		1941	0		1977	2,951
1870	2,421	1906	7,897		1942	0		1978	4,553
1871	4,261	1907	5,781		1943	0		1979	6,790
1872	2,366	1908	5,170		1944	0		1980	6,670
1873	2,029	1909	5,557		1945	0		1981	11,799
1874	2,240	1910	4,502		1946	0		1982	8,847
1875	1,480	1911	3,556		1947	0		1983	10,626
1876	2,691	1912	2,469		1948	0		1984	6,629
1877	725	1913	7,881		1949	0		1985	10,969
1878	187	1914	5,956		1950	0		1986	12,685
1879	780	1915	1,882		1951	0		1987	17,032
1880	68	1916	1,753		1952	26		1988	10,917
1881	41	1917	1,659		1953	22		1989	6,991
1882	15	1918	55		1954	0		1990	3,954
1883	1,075	1919	113		1955	0		1991	2,658
1884	1,884	1920	2,838		1956	0		1992	5,183
1885	861	1921	3,699		1957	839		1993	0
1886	601	1922	5,532		1958	0		1994	0
1887	1,475	1923	3,945		1959	0		1995	43,560
1888	3,589	1924	4,150		1960	0		1996	15,646
1889	2,013	1925	5,666		1961	555		1997	14,831
1890	2,523	1926	7,143		1962	2,961		1998	3,007
1891	1,197	1927	8,633		1963	2,972		1999	0
1892	468	1928	13,579		1964	0		2000	0
1893	214	1929	8,935		1965	0		2001	0
1894	1,003	1930	9,921		1966	0		2002	0
1895	570	1931	15,847		1967	0		2003	0
1896	2,777	1932	6,975		1968	0		2004	0
1897	3,005	1933	7,590		1969	0		2005	0
1898	1,885	1934	10,173		1970	0		2006	79
1899	1,885	1935	7,355		1971	0		2007	199
1900	1,975	1936	16,058		1972	0		2008	145
1901	4,820	1937	7,488		1973	973		2009	697
1902	1,294	1938	10,858		1974	514		2010	267

**Table 14.3. National contributions of observations, and number of cruises by country of origin in the SUR dataset in WOD23.**

ISO <sup>1</sup> Country Codes	Country Name	Number of Cruises	Number of Observations	% of Total
FR	France	3,272	876,382	41.77
AU	Australia	85	681,879	32.50
99	Unknown	3,378	161,543	7.70
US	United States	63	100,492	4.79
DE	Germany	93	63,698	3.04
NO	Norway	245	59,714	2.85
JP	Japan	66	57,406	2.74
NC	New Caledonia	1,229	41,655	1.99
CA	Canada	34	18,682	0.89
GB	United Kingdom	345	16,514	0.79
DK	Denmark	178	8,274	0.39
PL	Poland	23	2,824	0.13
FI	Finland	18	2,593	0.12
IN	India	111	1,537	0.07
NL	Netherlands	21	1,309	0.06
SU	Union of Soviet Socialist Republics	1	1,068	0.05
LV	Latvia	38	1,010	0.05
SE	Sweden	15	710	0.03
BE	Belgium	3	283	0.01
PT	Portugal	9	199	0.01
IE	Ireland	27	164	0.01
EE	Estonia	3	84	0.00
<i>Total:</i>		<i>9,257</i>	<i>2,098,020</i>	<i>100.00</i>

<sup>1</sup> ISO = International Organization for Standardization: [http://www.iso.org/iso/country\\_codes.htm](http://www.iso.org/iso/country_codes.htm)

#### **14.4. REFERENCES AND BIBLIOGRAPHY**

- Delcroix, T., J. Picaut (1998). Zonal displacement of the western equatorial Pacific "fresh pool". *J. Geophys. Res.*, 103(C1), 1087-1098 (97JC01912).
- Delcroix, T., M.J. McPhaden, A. Dessier, Y. Gouriou (2005). Time and space scales for sea surface salinity in the tropical oceans. *Deep-Sea Res. I*, 52(5), 787-813.
- Freeman, E., S.D. Woodruff, S.J. Worley, S. J. Lubker, E.C. Kent, W.E. Angel, D.I. Berry, P. Brohan, R. Eastman, L. Gates, W. Gloeden, Z. Ji, J. Lawrimore, N.A. Rayner, G. Rosenhagen, and S.R. Smith (2017). ICOADS Release 3.0: a major update to the historical marine climate record. *Int. J. Climatol.*, 37: 2211-2232. doi:[10.1002/joc.4775](https://doi.org/10.1002/joc.4775)
- Goni, G.J. (2017). Underway sea surface temperature and salinity data from thermosalinographs collected from multiple platforms assembled by NOAA Atlantic Oceanographic and Meteorological Laboratory (AOML). NOAA National Centers for Environmental Information. Dataset. <https://www.ncei.noaa.gov/archive/accession/AOML-TSG> .

- Henin, C. and J. Grelet (1996). A merchant ship thermo-salinograph network in the Pacific Ocean. *Deep-Sea Res.*, 43, 1833-1855.
- Johnson, D.R., T.P. Boyer, H.E. Garcia, R.A. Locarnini, J.I. Antonov, O.K. Baranova, A.V. Mishonov, D. Seidov, I.V. Smolyar, and M.M. Zweng (2009). World Ocean Database 2009: Documentation. Ed. S. Levitus. *NODC Internal Report 18*, U.S. Gov. Printing Office, Wash., D.C.
- Petit de la Villéon, L., IFREMER (2015). Sea surface temperature and salinity from the Global Ocean Surface Underway Data (GOSUD) from 1980-01-03 to present. NOAA National Centers for Environmental Information. Dataset. <https://www.ncei.noaa.gov/archive/accession/IODE-GOSUD> .
- Reverdin, G., D. Cayan, H.D. Dooley, D.J. Ellett, S. Levitus, Y. du Penhoat, A. Dessier (1994). Surface salinity of the North Atlantic: Can we reconstruct its fluctuations over the last one hundred years? *Prog. Oceanogr.*, 33(4), 249-386.
- Shawn, S., J.J. Rolf, K. Briggs, M.A. Bourassa (2009). Quality-Controlled Underway Oceanographic and Meteorological Data from the Center for Ocean-Atmospheric Predictions Studies (COAPS) - Shipboard Automated Meteorological and Oceanographic System (SAMOS). NOAA National Centers for Environmental Information. Dataset. <https://doi.org/10.7289/v5qj7f8r>
- Thomas, G.G., S. Cook, Y-H. Daneshzadeh, W.S. Krug, R. Benway (1999). Surface salinity and temperature from ships of opportunity. *Sea Tech.*, 2, 77-81.
- Wang, Z. NOAA/NCEI (2017). Quality-controlled sea surface marine physical, meteorological and other *in situ* measurements from the NCEI Surface Underway Marine Database (NCEI-SUMD). *NOAA National Centers for Environmental Information*. Dataset. <https://www.ncei.noaa.gov/archive/accession/NCEI-SUMD>.

## CHAPTER 15: GLIDER DATA (GLD)

*Alexey V. Mishonov, Tim P. Boyer, Ricardo A. Locarnini, James Beauchamp,  
Courtney N. Bouchard*

*Ocean Climate Laboratory  
National Centers for Environmental Information/ NOAA  
Silver Spring, MD*

### 15.1. INTRODUCTION

A glider is an autonomous underwater vehicle (AUV) propelled by buoyancy force that moves from the ocean surface along a slant trajectory through the water column to a programmed depth and back to the surface while measuring oceanographic parameters (Eriksen *et al.*, 2001; Rudnick *et al.*, 2004). Modern gliders carry various sensors to measure oceanographic parameters such as pressure, temperature, conductivity, chlorophyll *a* fluorescence, CDOM (colored dissolved organic matter) fluorescence, nitrate, oxygen, transmissivity, optical backscatter, acoustical backscatter, and downwelling radiance (Davis *et al.*, 2008; Glenn *et al.*, 2008; Niewiadomska *et al.*, 2008; Johnson *et al.*, 2009). Gliders can travel thousands of kilometers while making several hundred descents and ascents underway, thus achieving high vertical and horizontal resolution. Since gliders can be retrieved and reused, they represent one of the most cost-effective tools for oceanographic data collection. The annual operating cost of a glider is equivalent to a fraction of one ship-day (Eriksen *et al.*, 2001).

The original concept of a glider was invented by Douglas Webb in 1986 and was based on the thermal engine intended for global range (D. Webb, personal communication, May 2006). In 1986 Douglas Webb described to Henry Stommel the ideas of a glider with buoyancy engine harvesting propulsion energy from ocean thermal gradients (Stommel, 1989). Stommel later became an enthusiastic supporter and funding for a contract was received through the Office of Naval Technology (D. Webb, personal communication, May 2006). The glider with a battery-powered buoyancy engine was tested in 1991 at Wakulla Springs, FL and in Seneca Lake, NY (Simonetti, 1992; Webb and Simonetti, 1997; Webb *et al.*, 2001). The U.S. patent [US5291847A](#) for the concept of “Autonomous propulsion within a volume of fluid” was filed by Douglas Webb in 1992 and granted in 1994. At this time, the TDY Industries LLC Teledyne Benthos, Inc. holds this patent.

Gliders are equipped with a Global Positioning System (GPS) navigation to locate the vehicle. A satellite data relay is used to send its position and other data to shore-based computers while the operators program the gliders’ depth and mission. Modern gliders can reach a maximum depth of 6000 m with maximum range up to 10,000 km ([Deepglider](#), Osse and Eriksen, 2007; Steinberg and Eriksen, 2022). Their battery lifetime ranges from a few weeks to several months. Gliders’ speed is typically less than  $0.5 \text{ m}\cdot\text{s}^{-1}$  (Eriksen *et al.*, 2001; Davis *et al.*, 2002; Rudnick *et al.*, 2004). Gliders are used to perform diverse scientific missions, each requiring the use of a different instrument configuration.



## 15.2. GLD ACCURACY

The temperature and salinity data collected from the gliders mostly come from CTD sensors manufactured by two major players in the field of the oceanographic instrumentation – the [Sea-Bird Scientific](#) and [RBR Ltd.](#)

Most commonly used Sea-Bird glider instrumental package – the Glider Payload CTD (SB [GPCTD](#)) could be also equipped with SBE43F dissolved oxygen sensor and paired with [ECO FLNTU](#) optical sensor.

Sensor packages produced by RBR Ltd. [RBRlegato<sup>3</sup>](#) and [RBRcentauro<sup>3</sup>](#) could be equipped with variety of sensors including CTD, dissolved oxygen, fluorescence, pH, and others.

Glider data in WOD23 collected from Slocum, Spray, and Seaglider/Deepglider platforms most often were recorded by the SB GPCTD. Data from the SeaExplorer platform are mainly acquired by RBR sensors. There are also other sensors used for data collection, among which are oxygen sensors Aanderaa Optodes 4330, 4831 (mostly used on Slocums), 5013 (mostly used on SeaExplorers), SBE63 (mostly used on Sprays), and CTD packages SBE41CP (mostly used on Seagliders and Sprays).

Brief review of the mostly used sensor's characteristics is provided below extracted from the manufacturers' products specifications and data sheets with more details available from their web-sites.

The SB GPCTD have a measurement range for conductivity from 0 to 9  $\text{S}\cdot\text{m}^{-1}$ , for temperature from  $-5$  to  $+42^\circ\text{C}$  and for pressure from 0m up to 2000m, with initial accuracy of  $\pm 0.0003 \text{ S}\cdot\text{m}^{-1}$  for conductivity, 0-5 VDC for temperature (SBE 43) and frequency (SBD 43F), and  $\pm 0.1\%$  for pressure. Typical stability of the sensors is: conductivity -  $0.0003 \text{ S}\cdot\text{m}^{-1}$  per month, temperature –  $0.0002^\circ\text{C}$  per month, and pressure -  $\pm 0.05\%$  of full-scale range per year. Signal resolution is: for conductivity -  $0.00001 \text{ S}\cdot\text{m}^{-1}$ , for temperature –  $0.001^\circ\text{C}$ , and for pressure -  $0.002\%$  of full-scale range. The SB ECO FLNTU fluorometer and scattering sensor provides fluorescence sensitivity of  $0.025 \mu\text{g l}^{-1}$  Chl within typical range of 0-50  $\mu\text{g l}^{-1}$  Chl with 99% linearity.

The RBS sensors packages [RBRlegato<sup>3</sup>](#) and [RBRcentauro<sup>3</sup>](#) are features pump-free design and shared the CTD sensors configuration. They differ by housing material, depth rating, and physical appearance. Sensor's conductivity measurement range is 0-85  $\text{mS}\cdot\text{cm}^{-1}$  with initial accuracy of  $\pm 0.003 \text{ mS}\cdot\text{cm}^{-1}$ , resolution of  $< 0.0001 \text{ mS}\cdot\text{cm}^{-1}$ , and typical stability of  $0.10 \text{ mS}\cdot\text{cm}^{-1}$  per year. Temperature measurements range is from  $-5^\circ\text{C}$  to  $35^\circ\text{C}$ , with initial accuracy of  $\pm 0.002^\circ$ , at resolution  $< 0.00005^\circ\text{C}$ , and typical stability of  $0.002^\circ\text{C}$  per year. Pressure is measured within 2000dbar range with initial accuracy of  $\pm 0.05\%$  of full range with time constant  $< 10\text{s}$  and typical stability of  $\pm 0.05\%$  full scale per year.

### 15.3. GLIDER DESIGN AND OPERATION

Several types of operational gliders shown in Table 15.1 collected data which are now stored in WOD23. In general, gliders have similar features and functionality which could be illustrated by *Seaglider* (Figure 15.1): this platform is 1.8 m long, has a wing span of 1 m, 1.4 m antenna mast, and weighs 52 kg (Eriksen *et al.*, 2001). It was designed to operate with pitch angles from 10° to 75°. The vehicle alternately dives and climbs to a commanded depth from the surface down to a maximum depth of 1 km and back to the surface every 3 to 9 hours. It remains at the surface for 5 minutes and during that time the Iridium/GPS antenna is raised above the air-sea surface by pitching the vehicle nose down at 45° (Eriksen *et al.*, 2001; Hines, 2005; Rudnick *et al.*, 2004). The glider obtains its GPS fixes, transmits collected data at 180 bytes s<sup>-1</sup>, relays its position, and receives instructions via the Iridium satellite phone network before diving again (Rudnick *et al.*, 2004). It travels at a speed of 0.25 m·s<sup>-1</sup>, driven by buoyancy control - a hydraulic system that moves oil in and out of an external rubber bladder to force the glider to move, respectively, up or down. Shifting its battery pack relative to its body, causes it to pitch its nose up or down or roll its wings to change compass heading (Hines, 2005).



Figure 15.1 Seaglider licensed to Kongsberg. Photo [source](#)

Table 15.1. Different glider platform capabilities<sup>1</sup> and number of platforms & casts in WOD23.

Glider platform & WOD Ocean Vehicle code	Max. depth, m	Typical speed, m·sec <sup>-1</sup>	Max. Range, km	Endurance, days	Platforms & Casts in WOD23	Reference
<b>Seaglider</b> 701	1,000	0.25	4,600	300	<b>109</b> 280,473	<a href="#">Kongsberg</a>
<b>Slocum:</b> 702					<b>306</b> 2,329,910	<a href="#">Teledyne</a>
Alkaline	1,000	0.35	1,200	50		
Rechargeable	1,000	0.35	3,000	120		
Lithium	1,000	0.35	13,000	500		
<b>Spray</b> 703	1,500	0.25	4,700	180	<b>49</b> 124,574	<a href="#">SCRIPPS</a>
<b>SeaExplorer</b> 704	1,000	0.25-0.51	1,700	110	<b>15</b> 29,996	<a href="#">Alseamar</a>
<b>Deepglider</b> 701	6,000	0.225	8,500	380	<b>16</b> 25,257	<a href="#">Osse &amp; Eriksen, (2007)</a>

<sup>1</sup> Capabilities above based on standard load packages

The [Seaglider](#) (Eriksen *et al.*, 2001) was developed at the University of Washington (UW). Currently, the UW manufactures *Seagliders* only for UW employees/students, with iRobot and most recently Kongsberg Underwater Technology Incorporated manufacturing *Seagliders* for those outside of UW (Kongsberg, 2014). The [Deepglider](#) was developed in UW based on the enhanced *Seaglider* platform (Osse and Eriksen, 2007) with updated battery and fortified hull



allowing extended horizontal range of 10,000 km and able to dive down to 6,000 m depth. It carries an instrument suite similar to the *Seaglider*. These platforms are most often equipped with the Sea-Bird Scientific Glider Payload conductivity-temperature-depth instrumental package ([GPCTD](#)). It could be also furnished with an oxygen sensor: [SBE43](#), [Aanderaa optode 4531](#), or other. In addition, the optical sensor can be added ([ECO FLNTU](#)) for measuring backscattering, turbidity, and chlorophyll (Davis *et al.*, 2002; Rudnick *et al.*, 2004). Output of the pressure sensor is used for controlling the vehicle as well as recording the depth at which the measurements are taken (Eriksen *et al.*, 2001). *Seaglider* dynamics and performance are discussed at length by Eriksen *et al.* (2001) and further details can be found on the *Seaglider* [web page](#).

The [Slocum](#) gliders (Webb *et al.*, 2001) are manufactured by Teledyne Webb Research Corporation. This is a very versatile platform, which can support over forty different sensor configurations and can be operated remotely or pre-programmed. Its long endurance makes possible a large data collection range over an extended period of time. This platform is most widely used by oceanographers all over the globe. WOD23 contain more than 2.3 million profiles collected from three different generations of the *Slocum* platform (G1, G2, and G3).

[Spray](#) gliders are capable of diving to 1500 m depth with descent angles exceeding 20° and can conduct over 800 dives along a 4000 km long trajectory path. The *Spray* gliders (Sherman *et al.*, 2001) were developed at the Scripps Institution of Oceanography (Rudnick *et al.*, 2004). [Bluefin Robotics](#) licensed the technology from Scripps in 2004. The next generation of this platform – [Spray 2](#) is available from [MRV systems](#).

The [SeaExplorer](#) glider platform, developed and manufactured by the [Alseamar](#), is multi mission platform popular among the European Institutions. It is a very cost-effective solution for data collection as it reduces reliance on large vessels with high daily running costs: no surface supervision boat is required during the mission. The *SeaExplorer* is easy to operate and can be deployed and recovered by reduced crews in coastal waters using small boats. It is equipped with a wide range of the standard sensors, which includes CTD (pumped or non-pumped, most often the [RBRlegato<sup>3</sup> C.T.D.](#)), Dissolved Oxygen (Optode or Electrochemical) Chlorophyll, Turbidity, CDOM, PAR, ADCP, Nitrates, pCO<sub>2</sub>, Echo sounder, Fluorometers options (Puck).

Detailed information on gliders specifications and functions can be found in Rudnick *et al.* (2004), Eriksen *et al.* (2001), Sherman *et al.* (2001), Webb *et al.* (2001), and via the web links provided in Table 15.1 and in the References section below.

The accuracy of CTD instruments used on gliders varies with the instrument design. Typically, the accuracy of salinity measurement is approximately 0.003 to 0.02 and accuracy of temperature measurement is from 0.001°C to 0.005°C. For detailed information on CTDs and their accuracy, refer to section 3.2 of this document.

## **15.4. GLD PROFILES DISTRIBUTIONS**

GLD data represents 15.9% of the entire WOD23 data collection. Figure 15.2 illustrates the geographical distribution of 2,968,167 glider casts in WOD23 collected from 2002 through 2022. Data density is illustrated by bins of different colors indicating the amount of the profiles found in a single 1°x1° latitude-longitude square.

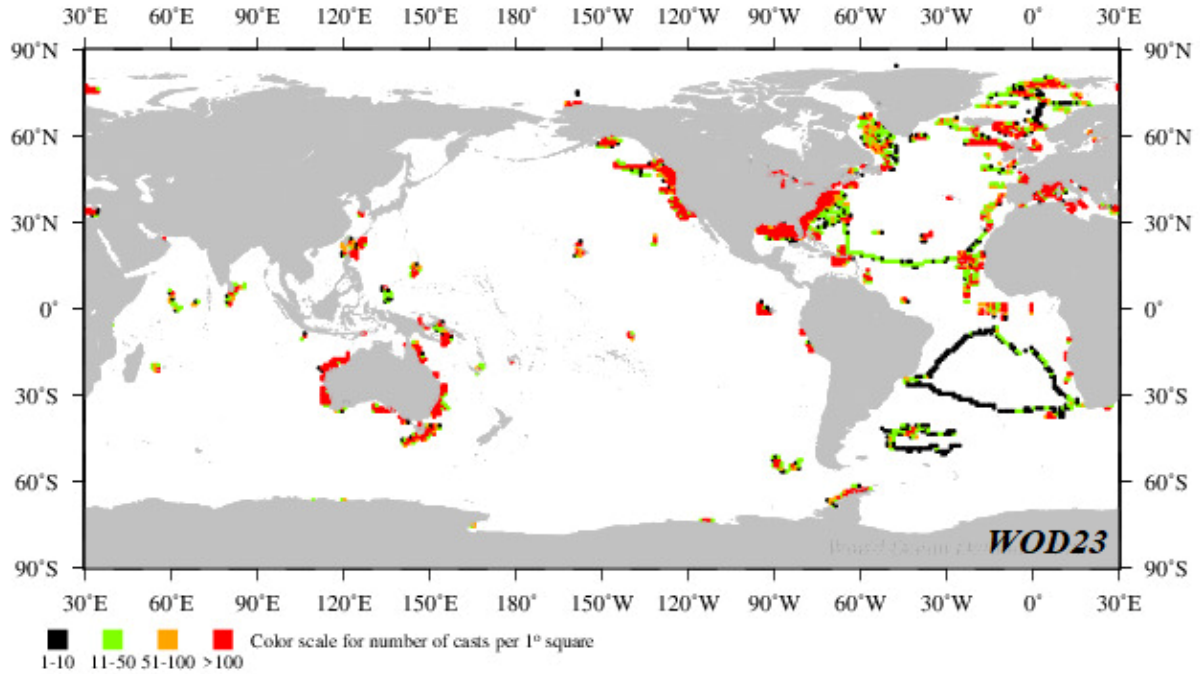


Figure 15.2. Geographical distribution of Glider (GLD) data in WOD23: number of casts per 1°square.

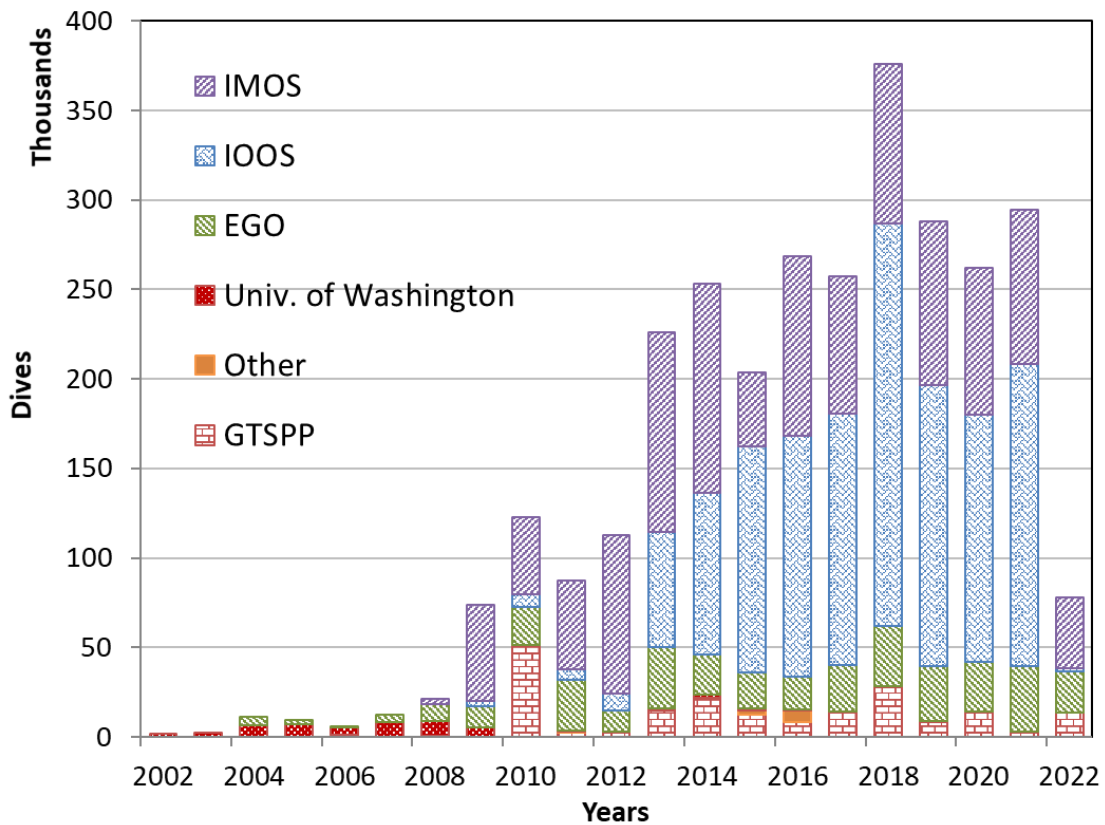


Figure 15.3. Temporal distribution and major sources of Glider (GLD) data in WOD23.

**Table 15.2. The number of all Glider (GLD) casts as a function of year in WOD23.**  
**Total Number of Profiles = 2,968,167**

YEAR	CASTS	YEAR	CASTS	YEAR	CASTS	YEAR	CASTS
2000	12	2006	5,461	2012	112,831	2018	376,081
2001	0	2007	12,460	2013	226,055	2019	288,385
2002	1,492	2008	21,312	2014	253,040	2020	261,878
2003	2,313	2009	73,741	2015	203,683	2021	294,330
2004	11,001	2010	123,539	2016	268,354	2022	78,054
2005	9,161	2011	87,368	2017	257,616		

Figure 15.3 and Table 15.2 show the temporal distribution of glider casts in WOD23 over the period of data collection. It should be noted that after the initial period of the technology development, 2002-2008, the sharp increase in the data collection reflects the growing interest in gliders as a convenient and versatile platform for oceanographic research.

Figure 15.4 shows contribution of the glider data made by different programs as a percentage of the total amount. The major contributor of the glider data in WOD23 is the U.S. Integrated Ocean Observing System ([IOOS](#)) followed by the Australian Integrated Marine Observing System ([IMOS](#)). The data submitted by IOOS (42.9% of all glider data in WOD23) are collected via the [IOOS Underwater Glider Network Map](#) and includes current and historical glider missions dating back to 2005 ranging from Gulf of Mexico ([GCOOS](#)), Southern California ([SCCOOS](#)), Northern Pacific ([NANOOS](#)), Central and Northern California ([CeNCOOS](#)), Great Lakes ([GLOS](#)), Mid-Atlantic ([MARACOOS](#)), and the Atlantic Oceanographic and Meteorological Lab ([AOML](#)). All data collected from U.S.-funded research programs/glider platforms are submitted by IOOS to the NCEI archives and then added to the WOD after established conversion and quality control procedures. Those data mostly cover areas of the US economic zone and North American shelf regions.

Data collected by University of Washington Applied Physics Lab ([APL, UW](#)) from their *Seaglid*ers and *Deepglid*ers were the main source of glider data in WOD for 2002-2009 and provided data from different areas of Atlantic and Pacific oceans, including unique data collected in Puerto Rico Trench zone. Now, UW data submitted to NCEI archives via IOOS channel.

The [Australian Ocean Gliders facility](#) operates a fleet of gliders measuring oceanographic parameters on shelf and boundary currents in Australian waters close to the shores. It operates a number of *Slocum* gliders in the Coral Sea, East Australian Current off New South Wales and Tasmania, Southern Ocean southwest of Tasmania, the Leeuwin and Capes Currents off South Western Australia and the Pilbara and Kimberly regions off North Western Australia. Data are distributed via Australian Integrated Marine Observing System ([IMOS](#)) portal and represent 36.2% of all glider data in WOD23.

Fast growing “Everyone’s Gliding Observatories” ([EGO](#)) program supported by [Copernicus Program](#) is very active in expanding glider use. Substantial amount of the glider data (12.3% of WOD23 GLD data collection) have been collected by various European Institutions mostly around the Europe as well as in Atlantic and Pacific oceans. These data distributed via Everyone’s Gliding Observatories portal ([EGO](#)).

Canada's Department of Fisheries via the Global Temperature-Salinity Profile Program (GTSPP) has submitted a noticeable (6.5%) volume of glider data in 2010-2022.

It should be noted, that because of the global glider community's efforts, the majority of the glider data in WOD23 have improved their associated metadata information leaving only 0.4% of the data in the entire GLD database with no programs/submitters information provided ("Others" on Figure 15.4). Compared to 22.3% of such data in WOD18, this makes possible a significant improvement in assigning proper credit to the collecting and submitting agencies, institutions, and individuals involved in data collection.

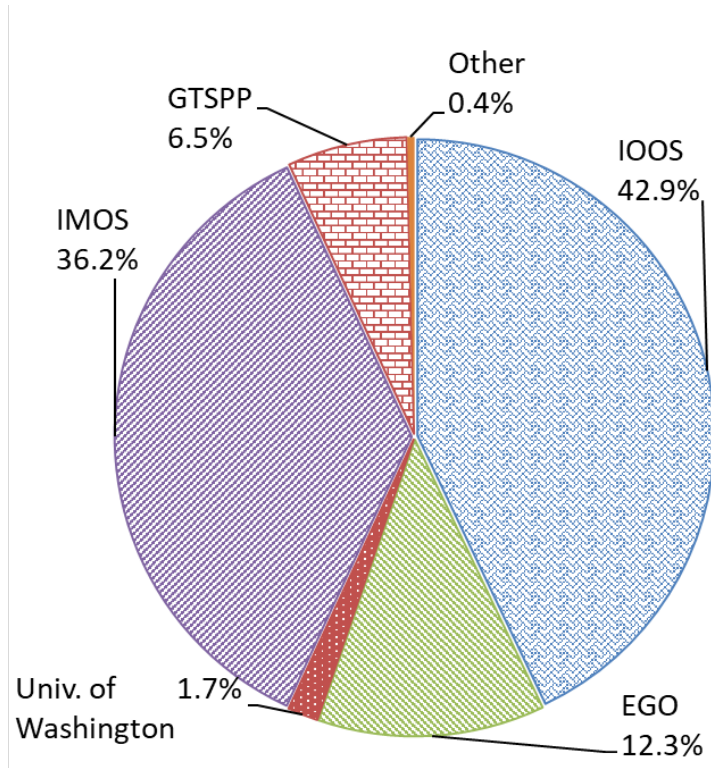


Figure 15.4. Contribution of Glider (GLD) data by different sources in WOD23.

Figure 15.5 shows global geographic distribution of the glider data submitted by different glider data centers/programs mentioned above. This figure clearly illustrates that at the current stage of the glider observation technology, the majority of the data are collected close to the contributing countries' economic zones with the only few examples of the trans-Atlantic deployments. The development of deep-range gliders and advanced battery packs, is expected to provide glider data that cover larger areas in the near future.

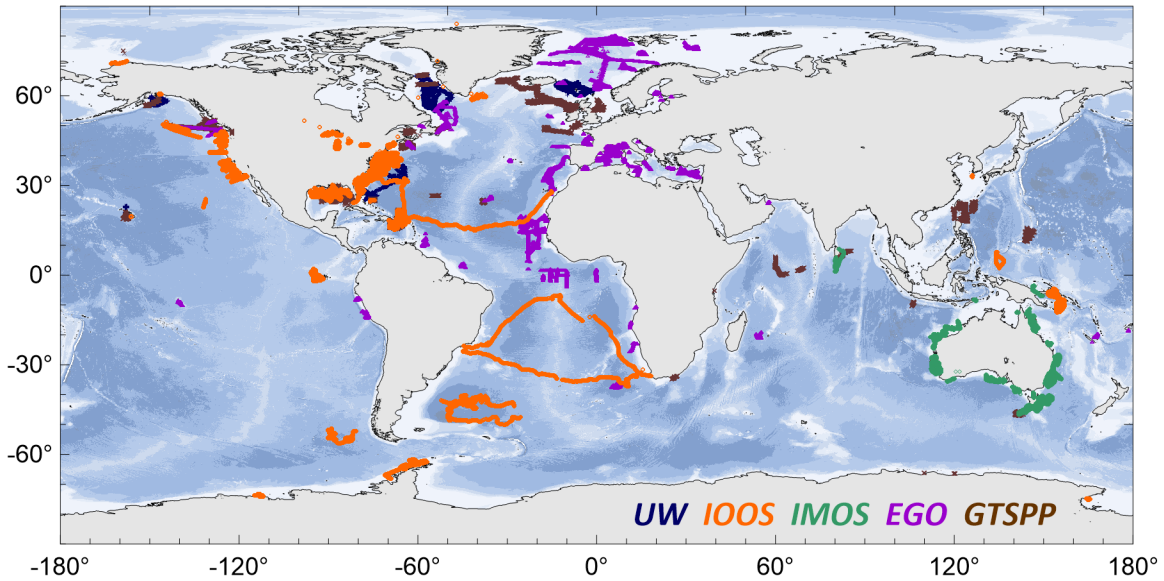


Figure 15.5. Geographical distribution of Glider (GLD) data submitted by different programs in WOD23.

Figure 15.6 shows depth distribution of the glider data at the standard depth levels. Because of the gliders hull depth restraints, the majority of glider data in WOD23 are in the upper 1000m layer of the ocean (Fig. 15.6a). However, data from deep-ocean gliders developed at University of Washington, (which provided 1.7% of the glider data in WOD23, see Fig. 15.4) are present in WOD23, and can be spotted in the deep ocean regions (*e.g.* Puerto Rico Trench on Fig. 15.5). There are 3,940 deep ocean glider profiles going below 1000 m depth and 222 profiles reaching 7200 m depth (Fig. 15.6b).

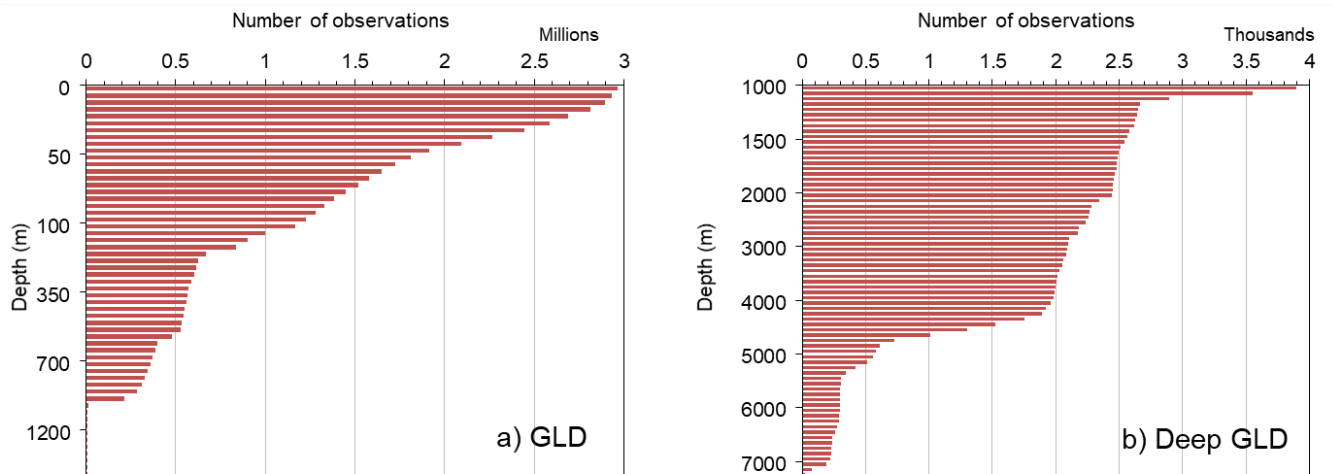


Figure 15.6. Distribution of Glider (GLD) data at standard depth levels in WOD23: a) from surface to 1500m depth, b) Deep Gliders (UW) – from 1000m to 7200m depth.



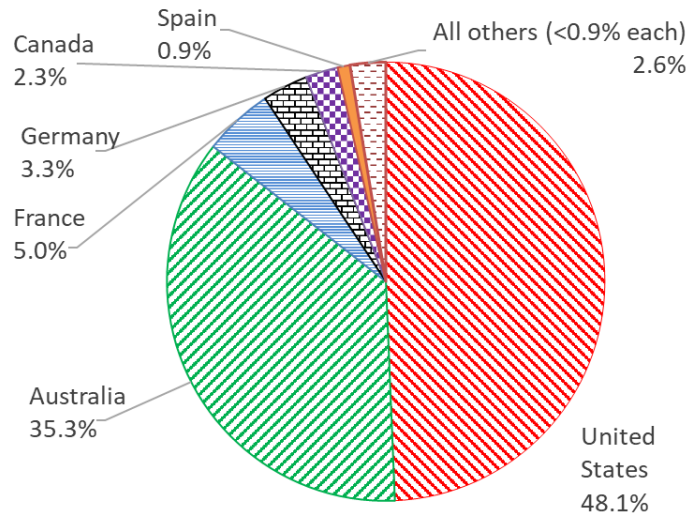
**Table 15.3. National contribution of Glider (GLD) profiles in WOD23.**

Cast amounts from shaded rows are summarized and shown on Figure 15.5 as ‘All others’

ISO <sup>1</sup> Country Code	Country Name	GLD Casts	% of Total
US	United States	1,464,620	49.34%
AU	Australia	1,074,589	36.20%
FR	France	152,713	5.15%
DE	Germany	100,050	3.37%
CA	Canada	71,124	2.40%
ES	Spain	27,051	0.91%
EE	Estonia	21,839	0.74%
IT	Italy	19,605	0.66%
GB	Great Britain	15,591	0.53%
NO	Norway	13,439	0.45%
SE	Sweden	2,591	0.09%
CY	Cyprus	2,114	0.07%
GR	Greece	1,362	0.05%
FI	Finland	677	0.02%
IE	Ireland	622	0.02%
99	Unknown	180	0.01%
<i>Total</i>		<i>2,968,167</i>	<i>100.00%</i>

<sup>1</sup> ISO = International Organization for Standardization: [http://www.iso.org/iso/country\\_codes.htm](http://www.iso.org/iso/country_codes.htm)

There are not too many countries able to manufacture, maintain, support glider operations, and generate constant data stream at this time. Table 15.3 and Figure 15.7 illustrate the glider data contribution sorted by country. Two major contributors of the glider data to WOD23 and to the NCEI archives are the USA (IOOS) – submitted ~49.3%, and Australia (IMOS) - submitted ~36.2% of all glider data stored in WOD23. European Union countries contributed ~14.5% of the data so far, with France and Germany leading the way collecting and submitting 5.1% and 3.4% of all data respectively.



**Figure 15.7. Glider (GLD) data contribution by countries in WOD23.**

## 15.5. RELEVANT WEB SITES

Applied Physics Laboratory - University of Washington (Seaglider):  
<http://www.apl.washington.edu/project/project.php?id=seaglider>

Australian Integrated Marine Observing System (IMOS) Ocean Gliders Facility:  
<http://imos.org.au/facilities/oceangliders/>

Oceanographic systems Laboratory, Woods Hole Oceanographic Institute:  
<https://www.whoi.edu/what-we-do/understand/departments-centers-labs/aope/aope-labs-and-groups/>

AUV Laboratory, Massachusetts Institute of Technology, Sea Grant College Program:  
[http://auvlab.mit.edu/MURI/1997\\_Rprtfinal.html](http://auvlab.mit.edu/MURI/1997_Rprtfinal.html)

Bluefin Robotics (Spray Glider):  
<https://gdmissionsystems.com/products/view-all-products-programs-and-services?tags=SeaCoastal> Ocean Observation Lab – Rutgers University:  
<http://rucool.marine.rutgers.edu/>

CTD Instrument:  
<https://oceanexplorer.noaa.gov/technology/ctd/ctd.html> Everyone's Gliding Observatories:  
[www.ego-network.org](http://www.ego-network.org)

Global Temperature and Salinity Profile Programme: <https://www.nodc.noaa.gov/GTSPP/>

iRobot (Seaglider): <https://pdf.archiexpo.com/pdf/irobot/irobot-1ka-seaglider/51219-114445.html>

Kongsberg Underwater Technology, Inc. (Seaglider): <https://www.km.kongsberg.com/>  
Rutgers: <https://marine.rutgers.edu/main/announcements/the-challenger-glider-mission-south-atlantic-mission-complete>

SBE 911 plus CTD: <https://www.seabird.com/profiling/sbe-911plus-ctd/family?productCategoryId=54627473769>

SCRIPPS Institute of Oceanography (Spray Glider): <http://spray.ucsd.edu/>

Teledyne Marine (Slocum Glider):  
<https://www.teledynemarine.com/brands/webb-research/slocum-glider> U.S. Integrated Ocean Observing System (IOOS): <https://gliders.ioos.us/>

U.S. Patents depository:  
<https://patentimages.storage.googleapis.com/42/e3/b2/c1d4299f929b98/US5291847.pdf>

## 15.6. REFERENCES AND BIBLIOGRAPHY

Davis, R.E., C.C. Eriksen, and C.P. Jones (2002). Autonomous Buoyancy-Driven Underwater Gliders, *Chapter 3 18*: 25-37.

- Davis, R.E., M.D. Ohman, D.L. Rudnick, and J.T. Sherman (2008). Glider surveillance of physics and biology in the southern California Current System, *Limnol. Oceanogr.*, 53(5, part 2) 2151-2168.
- Eriksen, C.C., T.J. Osse, R.D. Light, T. Wen, T.W. Lehman, P.L. Sabin, J.W. Ballard, and A.M. Chiodi (2001). Seaglider: A long-range autonomous underwater vehicle for oceanographic research, *IEEE J. Oceanic Eng.*, 26(4), 424-436.
- Glenn, S., C. Jones, M. Twardowski, L. Bowers, J. Kerfoot, J. Kohut, D. Webb, and O. Schofield (2008). Glider observations of sediment resuspension in a Middle Atlantic Bight fall transition storm. *Limnol. Oceanogr.*, 53(5, part 2), 2180-2196.
- Hines, S. (2005). Pairs of Seagliders set endurance records. *University of Washington - Office of News and Information*, [online] 5<sup>th</sup> April. Available at: <http://www.washington.edu/news/2005/04/05/pairs-of-seagliders-set-endurance-records/> [Accessed: 3rd July 2013].
- Johnson, K.S., W.M. Berelson, E.S. Boss, Z. Chase, H. Claustre, S.R. Emerson, N. Gruber, A. Körtzinger, M.J. Perry, and S.C. Riser (2009). Observing Biogeochemical Cycles at Global Scales with Profiling Floats and Gliders: Prospects for a Global Array. *Oceanography* 22(3) 216-225.
- Kongsberg, Inc. (2014). *Full scale production of KONGSBERG Seaglider begins.* <https://www.km.kongsberg.com/ks/web/nokbg0238.nsf/AllWeb/287D58C5F066C80AC1257C7C005B5BF4?OpenDocument>.
- Niewiadomska, K., H. Claustre, L. Prieur, and F. d’Ortenzio (2008). Submesoscale physical-biogeochemical coupling across the Ligurian Current (northwestern Mediterranean) using a bio-optical glider. *Limnol. Oceanogr.*, 53(5, part 2), 2210-2225.
- Osse, T. J. and C. C. Eriksen (2007). "The Deepglider: A Full Ocean Depth Glider for Oceanographic Research," *OCEANS 2007*, Vancouver, BC, Canada, 2007, pp. 1-12, doi: 10.1109/OCEANS.2007.4449125.
- Rudnick, D.L., R.E. Davis, C.C. Eriksen, D.M. Fratantoni, and M.J. Perry (2004). Underwater gliders for ocean research, *Mar. Tech. Soc. J.*, 38(2), 73-84.
- Sherman, J., R.E. Davis, W.B. Owens, and J. Valdes (2001). The autonomous underwater glider "Spray". *IEEE J Oceanic Eng.*, 26(4), 437-446.
- Simonetti, P.J., (1992). SLOCUM GLIDER, design and 1991 field trials, Webb Res. Corp., East Falmouth, MA, *Internal Rep.*, Sept. 1992.
- Steinberg, J., Ch. Eriksen, (2022). Eddy Vertical Structure and Variability: Deepglider Observations in the North Atlantic. *Journal of Physical Oceanography*. 52. DOI: 10.1175/JPO-D-21-0068.1
- Stommel, H. (1989). The Slocum Mission, *Oceanogr.*, 2(1), 22-25.
- Webb, D.C., and P.J. Simonetti (1997). A simplified approach to the prediction and optimization of performance of underwater gliders, In *Proc. 10<sup>th</sup> Int. Symp. on Unmanned Untethered Submersible Technology (USST)*, Durham, NH, Sept. 7-10, 1997, pp. 60-68.



Webb, D.C., P.J. Simonetti, and C.P. Jones (2001). SLOCUM: An underwater glider propelled by environmental energy, *IEEE J. Oceanic Eng.*, 26(4), 447-452

## CHAPTER 16: PLANKTON DATA

*Olga K. Baranova<sup>1</sup>, Todd D. O'Brien<sup>2</sup>, Tim P. Boyer<sup>1</sup>*

<sup>1</sup> *Ocean Climate Laboratory - National Centers for Environmental Information*

<sup>2</sup> *Office of Science and Technology - National Marine Fisheries Service*

*Silver Spring, Maryland, USA*

### 16.1. INTRODUCTION

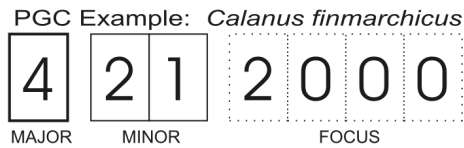
The term ‘plankton’ comes from the Greek ‘*planktos*’ (drifter). Plankton refers to floating or drifting organisms with limited powers of locomotion (Kennish, 1990). Planktonic organisms range in size from less than two microns to more than two centimeters (Levinton, 1995). The major plankton subdivisions include bacteria, phytoplankton, zooplankton, and temporary plankters which are planktonic only during some part of their life cycle, *e.g.*, eggs and larvae of fishes and other organisms (Kennish, 1990). Plankton participate across many levels of the pelagic ecosystem; from primary production and re-mineralization, to the transfer of materials and energies to higher trophic levels such as fish, birds, reptiles, and marine mammals (Harris *et al.* 2000). For these reasons it is important to have plankton observational data along with physical and chemical ocean profile data in the World Ocean Database. This opens up opportunities for finding interactions between plankton and other ocean variables (temperature, salinity, oxygen, nutrients, *etc.*) and for better understanding and preservation of pelagic ecosystems.

The plankton subset of the *World Ocean Database 2023* (WOD23) includes the content of the previously released *World Ocean Database 2018, 2013, 2009, 2005* (Baranova *et al.*, 2013, 2009, 2005), *World Ocean Database 2001* (O'Brien *et al.*, 2001), and *World Ocean Database 1998* (Conkright *et al.*, 1998). The WOD23 plankton data subset is a collection of measurements from serial bottle and plankton net-tow. The plankton measurements are represented in WOD23 as quantitative and qualitative abundance, and biomass data. The plankton measurements are stored in the OSD dataset (see Chapter 2) <sup>1</sup>.

Scientific taxonomic names in the WOD23 are stored using the corresponding ITIS (Integrated Taxonomic Information System, <https://www.itis.gov>) Taxonomic Serial Number (TSN). ITIS TSN's are not available for all plankton descriptions and biomass. WOD23 negative taxonomic codes (sequentially assigned numbers) were developed to preserve the original descriptions. In addition to ITIS or negative taxonomic codes, each plankton description has a *Plankton Grouping Code* (PGC) developed by O'Brien (2007). The PGC code follows the taxonomic hierarchy presented in *The Five Kingdoms* (Margulis & Schwartz 1998). The PGC is an ancillary code which places each taxon into broader groups (*e.g.*, *phytoplankton*, *diatoms*, *zooplankton*, *copepods*) and allows the WOD23 user access to hundreds of individual taxa by using a single PGC code. The PGC is 7-digit code divided into Major group (*e.g.* *Bacteria*,

---

<sup>1</sup> Disclaimer: The content of the WOD23 plankton subset has not been updated since the WOD18 release.



Phytoplankton, Zooplankton), Minor group (e.g., cyanobacteria, diatoms, crustaceans), and Focus group (e.g., copepods). For example, the copepod *Calanus finmarchicus* has a PGC code of “4212000”, specifying that it is in Major Group “4” (zooplankton), Minor Group

“21” (crustaceans), and Focus Group “2000” (copepods). Earlier versions of the *World Ocean Database* (2001, 2005) used a PGC precursor called the Biological Grouping Code, BGC (O’Brien *et al.* 2001). The PGC combines the BGC’s separate “protists” grouping with the “phytoplankton” group. From the WOD09 all BGC codes were replaced with their corresponding PGC codes.

The typical plankton cast, as represented in WOD23, stores taxon specific and/or biomass data in individual sets, called “Taxa-Record”. Figure 16.1 demonstrates an example of a plankton cast in WOD23.

Each “Taxa-Record” contains a taxonomic code (“Param\_number”), depth range (the upper and lower depth) of observation, the original measurements (e.g., abundance, biomass or volume), and all provided qualifiers (e.g., lifestage, sex, size, etc.) required to represent the plankton observation.

In addition to the observed data, a cast may include additional originator’s metadata information such as the “institution” which collected and identified the species of plankton, the “voucher institution” (institution which stores samples), sampling gear (e.g., Bongo Net, Continuous Plankton Recorder), net mesh size, sampling method (e.g., vertical, horizontal, or oblique haul), meteorology, and other general header information which are described in detail in [WOD23 Manual](#) (Garcia *et al.*, 2024).

The alternative way to receive plankton data is a “csv” (comma-separated value) output file (illustrated by Figure 16.2), which is available only through the [WODselect](#) – the online WOD23 database retrieval system.

**Longitude Latitude Year Month Day Time Cruise# CC Prof\_#**  
**-4.883 79.017 1991 6 9 ---- 10438 06 2087562**

Mesh\_size 200.000 Type\_tow 2.000 Lge\_removed 1.000  
 Gear\_code 118.000 net\_mouth\_area 0.300 Lge\_removed\_len 1.000  
 Tow\_speed\_avg 1.944

**Taxa-Record #1**

**Param\_number** 85263.000 upper\_depth 0 lower\_depth 100.000  
 Taxon\_lifestage 25.000 Taxon\_count 18.600 Taxon\_modifier 2.000  
 Units 70.000 CBV\_value 18.600 CBV\_calc\_meth 70.000  
 CBV\_flag 3.000 PGC\_group\_code 4282000.000

**Taxa-Record #2**

Param\_number -404.000000 upper\_depth 0 lower\_depth 100.000  
 int\_value 3100.000 Units 69.000 CBV\_value 31.000  
 CBV\_calc\_meth 69.100 CBV\_flag 3.000 PGC\_group\_code -404.000000

**Taxa-Record #3**

Param\_number 85263.000 upper\_depth 0 lower\_depth 100.000  
 Taxon\_lifestage 26.000 Taxon\_count 0.100 Taxon\_modifier 2.000  
 Units 70.000 CBV\_value 0.100 CBV\_calc\_meth 70.000  
 CBV\_flag 3.000 PGC\_group\_code 4282000.000 etc ....

**Access# 0000772**  
**Cast Number 9617720**  
**Orig\_Stat\_Num 7**  
**Bottom\_Depth 1413.000**  
**T-S\_Probe 7.000**  
**NODCorig 3.000**

**Figure 16.1. An example of a plankton cast in WOD23 (generated using *WODselect* data retrieval portal).**

```

CAST ,,9617720,WOD Unique Cast Number,WOD code,,,,,,,,,
NODC Cruise ID,,06-10438 ,,,,,,,,,,
Originators Station ID,,7,,,integer,,,,,,,,,
Originators Cruise ID,,,,,,,,,
Latitude,,79.0167,decimal degrees,,,,,,,,,
Longitude,,-4.8833,decimal degrees,,,,,,,,,
Year,,1991,,,,,,,,,
Month,,6,,,,,,,,,
Day,,9,,,,,,,,,
METADATA,,,,,,,,,
Country,,DE,NODC code,GERMANY, FEDERAL REPUBLIC OF,,,,,,,,,
Accession Number,,772,NODC code,,,,,,,,,
Project,,435,NODC code,IAPP (International Arctic Polynya Programme),,,,,,,,,,
Platform,,199,OCL code,POLARSTERN,,,,,,,,,
Institute,,892,NODC code,ALFRED-WEGENER-INSTITUTE (BREMERHAVEN),,,,,,,,,,
Bottom depth,,1413,meters,,,,,,,,,
Database origin,,3,WOD code,GODAR Project,,,,,,,,,
BIOLOGY METADATA,,,,,,,,,
Mesh size,,200,microns,,,,,,,,,
Type of tow,,2,WOD code,VERTICAL TOW,,,,,,,,,
Large plankters removed, ,1,WOD code,yes,,,,,,,,,
Gear,,118,WOD code,Bongo Net,,,,,,,,,
Net mouth area,,0.3,m2,,,,,,,,,
Min length removed,,1,cm,,,,,,,,,
Average tow speed,,2,knots,,,,,,,,,
BIOLOGY,Upper Z,Lower Z,Measuremnt Type,ORIGINAL VALUE ,F,Orig unit,WOD CBV
value ,F,_unit,_meth,WOD PGC,ITIS TSN,mod,lif,
1,0. meters,100. meters,Taxon_count,18.6,0,#/m3,18.6,3,
#/m3,70,4282010,CALANUS,MODIFIER=spp. (multiple species),LIFE STAGE=C1:
COPEPODITE I
2,0. meters,100. meters>Total Dry Mass,3100,0,mg/m2,31,3,mg/m3,69.1,-404,Zooplankton
Dry Mass (mg/unit),,,,,,,,,,
3,0. meters,100. meters,Taxon_count,0.1,0,#/m3,0.1,3,
#/m3,70,4282010,CALANUS,MODIFIER=spp. (multiple species),LIFE STAGE=C2:
COPEPODITE II
.....
END OF BIOLOGY SECTION

```

Figure 16.2. An example of a plankton cast in 'csv' output file available on-line through the *WODselect*.

## 16.2. BASIC QUALITY CONTROL

Plankton numerical abundance and total biomass measurements are stored with the data originator’s units in WOD23 (e.g., counts in units of “number per m<sup>3</sup>”, “wet mass per m<sup>2</sup>”, “displacement volume per haul”, “count per haul”, “count per ml”). To allow easier comparison of incoming measurements with different units, each numerical abundance or biomass measurement has been recalculated into a common unit named *Common Base-unit Value* (CBV). The CBV is calculated from the original value using sampling metadata (e.g., towing distance, water volume filtered) but does not account for differences in mesh size, gear efficiency, or sampling depth intervals. The calculation method used to create the CBV is stored in the *CBV calculation method* field and described in detail in [WOD23 Manual](#), Appendix 5.11, (Garcia *et al.*, 2024). Table 16.1 lists CBV units by data type.

**Table 16.1. Measurement Type and/or Groups and their corresponding CBV unit.**

Measurement Type or Group	CBV unit
Total Biomass (displacement volume, settled volume)	ml · m <sup>-3</sup>
Total Biomass (wet mass, dry mass, ash free dry mass)	mg · m <sup>-3</sup>
Zooplankton Abundance	# · m <sup>-3</sup>
Phytoplankton Abundance	# · ml
Bacterioplankton Abundance	# · µl
Ichthyoplankton Abundance	# · m <sup>-3</sup>

The addition of the PGC and CBV to each plankton measurement allows for individual value checks against broad, group-based ranges (O'Brien *et al.*, 2001). Grouped by major PGC groups (Table 16.2) and Total Biomass types (Table 16.3), these broad range checks are used to detect and flag extremely large or small values.

**Table 16.2. WOD23 broad group-based ranges for plankton abundance.**

Group	Min Value	Max Value	Units
Bacteria	0.001	5,000	# · µl
Phytoplankton	0.001	50,000	# · ml
Zooplankton	0.001	200,000	# · m <sup>-3</sup>
Ichthyoplankton	0.001	200,000	# · m <sup>-3</sup>

**Table 16.3. WOD23 broad group-based ranges for biomass.**

<b>Group</b>	<b>Min Value</b>	<b>Max Value</b>	<b>Units</b>
Total Displacement Volume	0.005	10	ml · m <sup>-3</sup>
Total Settled Volume	0.025	50	ml · m <sup>-3</sup>
Total Wet Mass	0.5	10,000	mg · m <sup>-3</sup>
Total Dry Mass	0.01	500	mg · m <sup>-3</sup>
Total Ash free Dry Mass	0.001	100	mg · m <sup>-3</sup>

WOD23 applied quality flags to Common Base-unit Values as follows:

- 0 - accepted value
- 1 - range outlier (outside of broad range check)
- 2 - questionable value\*

\* The contents from an entire net tow may be flagged as “questionable” in cases of gross gear failure (*e.g.*, a broken net or leaking bottle). Individual observations may also be flagged in cases of gear-incompatible capture (*e.g.*, phytoplankton cells snagged in a large mesh net, presence of a single copepod caught in a Nansen bottle).

### **16.3. DATA SOURCES**

The plankton data that comprise WOD23 have been contributed by 37 countries, 142 institutions and more than 50 projects. Significant amounts of data (104,740 casts) have no information about the project. Among them are data provided by the Instituto del Mar del Peru (IMARPE). This contribution (~23,000 casts) comes from a joint data rescue effort with the IMARPE and the Intergovernmental Oceanographic Commission’s Global Oceanographic Data Archaeology and Rescue project (GODAR), which digitized over forty-five years of IMARPE phytoplankton monitoring data. Substantial amounts of historical biomass and abundance data are from the archives of the National Centers for Environmental Information (NCEI), former National Oceanographic Data Center (NODC) and the World Data Center for oceanography, Silver Spring.

Table 16.4 summarizes data contributing countries. The top five contributors are the United States, Japan, Peru, Russia (Former Soviet Union), and the United Kingdom. Within the United States, the National Marine Fisheries Service (NMFS) has played a cooperative or leading role in major sampling and monitoring programs which were responsible for collecting ~70% of the US contribution, and 40% of the total global content. The NMFS-associated programs are indicated with asterisks in Table 16.5.

A considerable portion of biomass data (~69,500 casts) was received from Coastal and Oceanic Plankton Ecology Production and Observation Database (COPEPOD)<sup>2</sup> as a result of collaboration between NCEI and National Marine Fisheries Service (NMFS).

---

<sup>2</sup> Data acquired through the COPEPOD database were provided in COPEPOD’s format and mainly include data from CalCOFI, ECOMON, and SEAMAP projects.

**Table 16.4. National contributions of plankton casts in WOD23.**

<b>ISO Country Code</b>	<b>Country Name</b>	<b># Casts</b>	<b>% of Total</b>
US	United States	130,675	53.3
JP	Japan	41,372	16.9
PE	Peru	22,874	9.3
SU	Union of Soviet Socialist Republics	20,551	8.4
GB	Great Britain	16,253	6.6
ID	Indonesia	2,098	0.9
PT	Portugal	1,611	0.7
NO	Norway	1,422	0.6
FR	France	1,222	0.5
IN	India	970	0.4
DE	Germany	958	0.4
AU	Australia	763	0.3
CA	Canada	733	0.3
RU	Russian Federation	508	0.2
PL	Poland	405	0.2
ZA	South Africa	396	0.2
EC	Ecuador	352	0.2
MX	Mexico	293	0.1
BR	Brazil	216	0.1
KR	Korea Republic of	193	0.1
PH	Philippines	184	0.1
TW	Taiwan	141	0.1
NC	New Caledonia	136	0.1
DK	Denmark	133	0.1
IS	Iceland	133	0.1
CO	Colombia	97	> 0.1
ES	Spain	71	> 0.1
AR	Argentina	64	> 0.1
BE	Belgium	38	> 0.1
CI	Cote d'Ivoire	37	> 0.1
NL	Netherlands	36	> 0.1
SG	Singapore	35	> 0.1
CD	Congo, the Democratic Republic	29	> 0.1
PK	Pakistan	22	> 0.1
NG	Nigeria	12	> 0.1
SE	Sweden	11	> 0.1
TH	Thailand	10	> 0.1
<i>Total</i>		<i>245,059</i>	<i>100.0</i>

Another large portion (38,980 casts) of the zooplankton and biomass data was acquired through the California Cooperative Oceanic Fisheries Investigations (CalCOFI) project. The CalCOFI project was initiated in 1949 to study the collapse of the U.S. west coast sardine fishery.



Hydrographic casts have been occupied from 1950 to the present along cross-shelf transects. Additional information can be found on CalCOFI's Web Page, <http://www.calcofi.org> .

The Ecosystem Monitoring Program (EcoMon) and its predecessor Marine Resources Monitoring Assessment and Prediction (MARMAP) program is one of the important contributors of the plankton data (25,981 casts). The NMFS-wide EcoMon program maintains plankton sampling databases for the northeast U.S. continental shelf ecosystem. Sampling was conducted under a number of different programs with varying temporal and spatial scales. The largest and most comprehensive was the Marine Resources Monitoring Assessment and Prediction Program (MARMAP), which sampled the shelf from 1977 to 1987. Data collected over time includes biological surveys of fishes, fish eggs and larvae.

The Ecosystems and Fisheries-Oceanography Cooperative Investigations (EcoFOCI) program contributed a valuable amount of plankton biomass data (13,608 casts). The EcoFOCI is a joint research program between the Alaska Fisheries Science Center and the Pacific Marine Environmental Laboratory. Originally, the FOCI was established by NOAA in 1984 to study walleye pollock in the western Gulf of Alaska. Later on, the EcoFOCI has broadened its study to ecosystem research in the North Pacific and Alaskan waters <https://www.ecofoci.noaa.gov>.

A significant amount of data (11,620 casts) was received through the Southeast Area Monitoring and Assessment Program (SEAMAP). Since its beginning in 1981 SEAMAP monitoring of marine resources within the Gulf of Mexico, South Atlantic, and Caribbean regions <http://www.seamap.org/> .

The Outer Continental Shelf Environmental Assessment Program (OCSEAP) contributed another large portion of the plankton data (7,920 casts). The OCSEAP was established in 1984 by basic agreement between the U.S. Department of Commerce, National Oceanic and Atmospheric Administration (NOAA) and the U.S. Department of the Interior (USDOI), Minerals Management Service (MMS) for environmental studies of Alaskan Outer Continental Shelf waters considered for oil development (Truett, J.C., 1985).

Another source of data was the Eastern Tropical Pacific Ocean (EASTROPAC) program (5,544 casts). The first EASTROPAC survey (February 1967 through March 1968) was a cooperative effort towards the understanding of the oceanography of the eastern Tropical Pacific Ocean. Participating scientists were primarily from the NMFS, Scripps Institution of Oceanography, and the Inter-American Tropical Tuna Commission.

Kuroshio Exploitation and Utilization Research (KER) project provided 4,234 casts. KER was designed to study the subtropical circulation system, marine ecology, and fishery around Japan. The project was conducted in 1977 – 1995.

Table 16.5 gives project contributions of plankton casts sorted by percent contribution from each project.

**Table 16.5. Project contributions of plankton casts sorted by percent contribution from each project.**

<b>NCEI Project Code</b>	<b>Project Name</b>	<b># Casts</b>	<b>% of Total</b>
33	CalCOFI: California Cooperative Oceanic Fisheries Investigation	38,980	27.7
637	EcoMon: Ecosystem Monitoring Program/MARMAP: Marine Resources Monitoring Assessment and Prediction	25,981	18.5
174	EcoFOCI: Ecosystems and Fisheries-Oceanography Cooperative Investigations	13,608	9.7
121	SEAMAP: Southeast Area Monitoring and Assessment Program	11,620	8.3
81	OCSEAP: Outer continental shelf environmental assessment program	7,920	6.4
3	EASTROPAC (1967-1968)	5,544	4.5
526	General Fisheries Research (YugNIRO)	5,438	4.5
243	KER: Kuroshio exploitation and utilization research (1977 - 1995)	4,234	3.4
93	Brine Disposal	4,198	3.4
240	USAP or USARP: United States Antarctic Research Project	3,665	2.9
25	IIOE: International Indian Ocean Expedition	2,045	1.6
344	POFI: Pacific Oceanic Fisheries Investigations	1,310	1.1
372	OMEX: Ocean margin exchange project	1,234	1.0
367	GLOBEC: Georges Bank Program	951	0.8
361	JGOFS/AESOPS: US JGOFS Antarctic Environments Southern Ocean Process Study	943	0.8
30	ICNAF: International Commission for the Northwest Atlantic Fisheries	851	0.7
345	North Sea Project	827	0.7
241	BIOMASS: Biological Investigations of Marine Antarctic Systems and Stocks	712	0.6
322	SKIPJACK	684	0.6
365	JGOFS/ARABIAN: Arabian Sea Process Studies	657	0.5
31	CSK: Cooperative Study of the Kuroshio	599	0.5
83	OCS-SOUTH: Texas	533	0.4
275	JGOFS/BATS: Bermuda Atlantic Time Series	495	0.4
325	CINECA: Cooperative Investigations of Northern Part of Eastern Central Atlantic	400	0.3
82	PSERP: Mesa Puget Sound	396	0.3
645	Discovery Investigations	366	0.3
200	JGOFS: Joint Global Ocean Flux Study	363	0.3
273	EASTROPIC: Eastern Tropical Pacific 1955	323	0.3
410	TASC: Trans-Atlantic Study of <i>Calanus</i>	300	0.2
310	JGOFS/EQPAC: Equatorial Pacific basin study	279	0.2
450	SFRI Upwelling Cruise 1969	255	0.2
96	EPA: Buccaneer oil field	214	0.2

<b>NCEI Project Code</b>	<b>Project Name</b>	<b># Casts</b>	<b>% of Total</b>
321	BOFS: Biogeochemical Ocean Flux Study	180	0.2
422	ICITA - EQUALANT III	177	0.1
443	IMECOCAL: Investigaciones Mexicanas De La Corriente De California	174	0.1
421	ICITA - EQUALANT II	164	0.1
420	ICITA - EQUALANT I	163	0.1
34	MAZATLAN	119	0.1
255	CTZ: Coastal Transition Zone	100	> 0.1
246	BERPAC: Bering and Pacific Russian/US Cooperative Research Program	88	> 0.1
328	SIBEX: Second International Biomass Experiment - Fr	63	> 0.1
312	CEAREX: Coordinated Eastern Arctic Experiment	63	> 0.1
225	WOCE: World Ocean Circulation Experiment	41	> 0.1
435	IAPP: International Arctic Polynya Programme	41	> 0.1
90	ONR: Office of Naval Research	39	> 0.1
71	IDOE/CUEA	30	> 0.1
434	Arctic Ocean Section: Canada/U.S. joint expedition	18	> 0.1
77	SCOPE	11	> 0.1
447	Marine Food Chain Research Group	10	> 0.1
444	GSP: Greenland Sea Project	5	> 0.1
<i>Total</i>		245,059	100.00

#### **16.4. PLANKTON DATA DISTRIBUTIONS<sup>3</sup>**

The WOD23 plankton subset consists of 245,059 globally distributed casts (Figure 16.3). The temporal distribution of plankton sampling covers the period from 1900 to 2015 (Figure 16.4). Table 16.6 gives the yearly counts of plankton casts in the WOD23.

---

<sup>3</sup> Disclaimer: The content of the WOD23 plankton subset has not been updated since the WOD18 release.

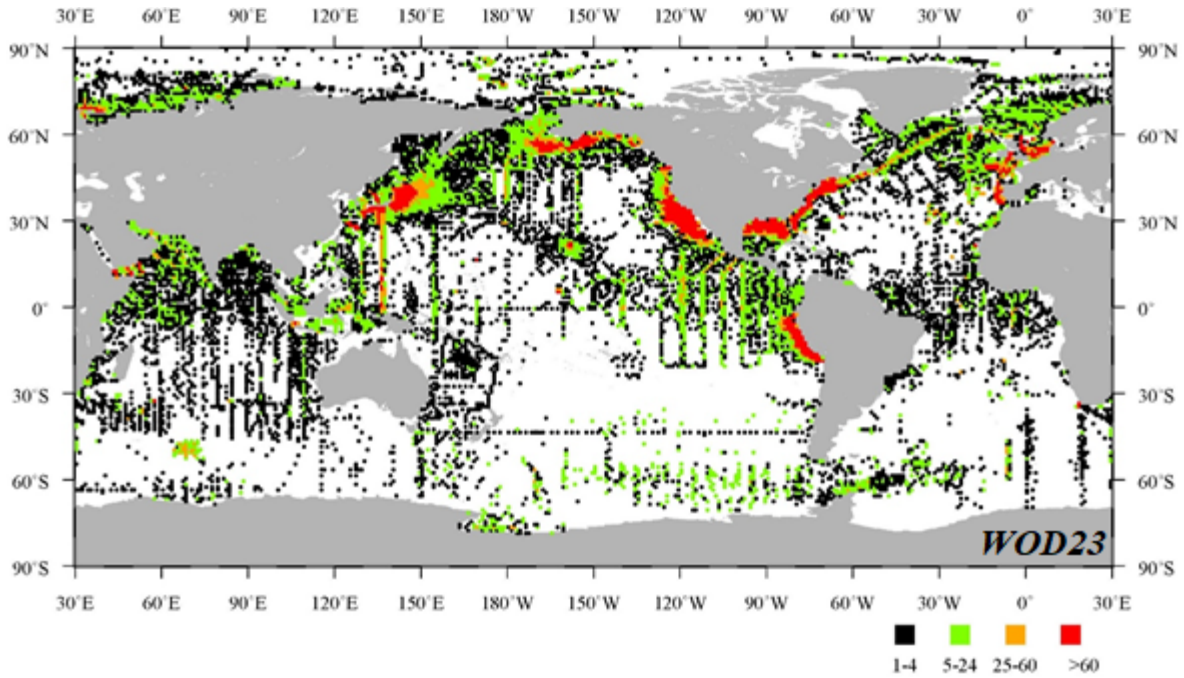


Figure 16.3. Geographic distribution of plankton (245,059 casts) in WOD23.

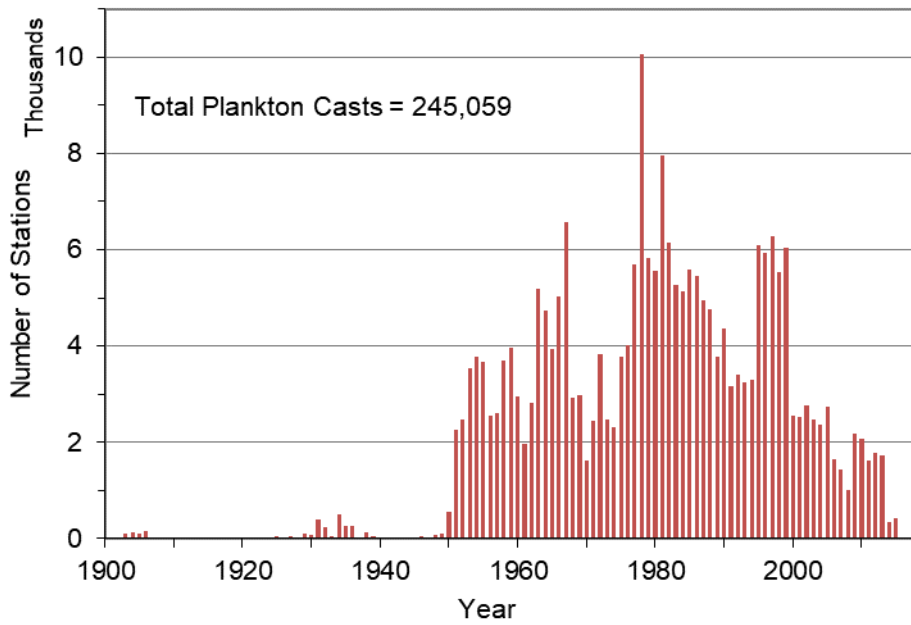


Figure 16.4. Temporal distributions of plankton casts in WOD23 as a function of year.

**Table 16.6. Number of plankton casts in WOD23 as a function of year**

Total Number of Casts = 245,059

<b>YEAR</b>	<b>CASTS</b>	<b>YEAR</b>	<b>CASTS</b>	<b>YEAR</b>	<b>CASTS</b>	<b>YEAR</b>	<b>CASTS</b>
1900	17	1929	93	1958	3,690	1987	4,949
1901	9	1930	88	1959	3,953	1988	4,759
1902	13	1931	409	1960	2,962	1989	3,767
1903	100	1932	247	1961	1,959	1990	4,370
1904	126	1933	62	1962	2,814	1991	3,162
1905	95	1934	490	1963	5,195	1992	3,399
1906	160	1935	273	1964	4,730	1993	3,254
1907	0	1936	262	1965	3,940	1994	3,308
1908	0	1937	7	1966	5,040	1995	6,082
1909	0	1938	134	1967	6,573	1996	5,937
1910	0	1939	51	1968	2,934	1997	6,275
1911	0	1940	2	1969	2,974	1998	5,527
1912	0	1941	0	1970	1,620	1999	6,045
1913	6	1942	2	1971	2,438	2000	2,562
1914	7	1943	0	1972	3,821	2001	2,531
1915	9	1944	0	1973	2,467	2002	2,768
1916	0	1945	0	1974	2,318	2003	2,483
1917	0	1946	54	1975	3,772	2004	2,367
1918	0	1947	36	1976	4,025	2005	2,745
1919	0	1948	67	1977	5,685	2006	1,644
1920	0	1949	98	1978	10,060	2007	1,444
1921	29	1950	558	1979	5,828	2008	1,002
1922	33	1951	2,266	1980	5,570	2009	2,176
1923	0	1952	2,468	1981	7,952	2010	2,067
1924	2	1953	3,531	1982	6,152	2011	1,627
1925	50	1954	3,783	1983	5,255	2012	1,793
1926	34	1955	3,676	1984	5,145	2013	1,729
1927	46	1956	2,566	1985	5,597	2014	334
1928	35	1957	2,602	1986	5,465	2015	423

## 16.5. PLANKTON CONTENT

The plankton measurements are represented in WOD23 as descriptive and numerical abundance, and biomass data. The majority (69%) of plankton measurements are total biomass. Contributions of plankton casts by measurement type are shown in Figure 16.5.

### 16.5.1. Abundance

The majority (83%) of plankton abundance measurements in WOD23 are numeric (*e.g.*, the number of individuals counted per sample or haul), while descriptive abundance measurements (*e.g.*, individual was "rare", "common", or "abundant" in sample or haul) are present in a smaller amount (17 %) of total abundance. The WOD23 plankton abundance content, listed by major plankton groups and sub-groups, is summarized in Table 16.7.

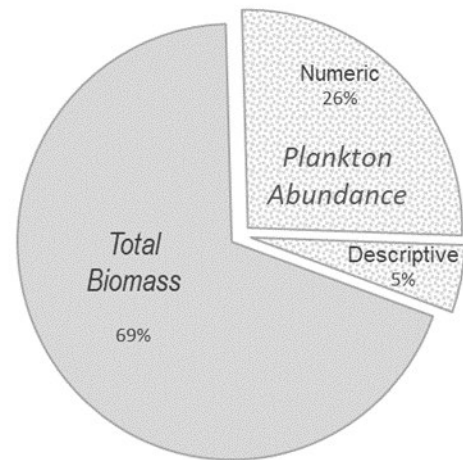


Figure 16.5 Contributions of Plankton casts by measurement type.

Table 16.7 WOD23 abundance measurements content.

PGC	Plankton Group	Numeric abundance (casts #)	Descriptive abundance (casts#)
<b>1000000</b>	<b>BACTERIA</b> ( <i>all sub-groups</i> )	1,986	28
1050000	Cyanobacteria	974	27
<b>2000000</b>	<b>PHYTOPLANKTON</b> ( <i>all sub-groups</i> )	37961	22,471
2030000	Amoebida	44	0
2040000	Granuloreticulosa (Foraminifera)	5,561	147
2070000	Dinomastigota (Dinoflagellata)	14,586	20,563
2080000	Ciliophora (ciliates)	4,893	7,769
2100000	Haptomonada (Coccolithophorids)	5,342	372
2110000	Cryptomonada (Chrytophyta)	1,910	8
2120000	Discomitochondria	1,333	242
2130000	Chrysomonada (Chrysophyta)	5,632	5,772
2160000	Diatoms (Bacillariophyta)	22,877	19,475
2270000	Actinopoda	3,817	436
2280000	Chlorophyta (green algae)	1,223	128
2300000	Ebriida	184	2

PGC	Plankton Group	Numeric abundance (casts #)	Descriptive abundance (casts#)
<b>4000000</b>	<b>ZOOPLANKTON</b> ( <i>all sub-groups</i> )	46,224	5,805
4020000	Porifera	1,941	3
4030000	Cnidaria (coelenterates)	16,103	2,676
4032000	Hydrozoa	13,843	665
4036000	Stauromedusae	2,381	28
4038000	Antipatharia	2,292	83
4040000	Ctenophora (comb jellies)	3,912	370
4050000	Platyhelminthes (flat worms)	2,042	0
4090000	Nemertina (ribbon worms)	2,352	44
4100000	Nematoda	2,053	7
4130000	Rotifera (rotifers)	2,182	623
4180000	Entoprocta	2,157	0
4190000	Arthropoda: Chelicerata	862	154
4200000	Arthropoda: Mandibulata ("insects")	4,881	14
4210000	Arthropoda: Crustacea ( <i>all sub-groups</i> )	49,737	6,553
4211000	<i>Crustacea</i> : Ostracoda	12,781	300
4212000	<i>Crustacea</i> : Copepoda	88,658	9,800
4213000	<i>Crustacea</i> : Cirripedia (barnacles)	7,091	783
4214000	<i>Crustacea</i> : Mysidacea	4,464	55
4216000	<i>Crustacea</i> : Isopoda	4,049	60
4217000	<i>Crustacea</i> : Amphipoda	22,309	1,718
4218000	<i>Crustacea</i> : Euphausiacea	17,227	1,728
4219000	<i>Crustacea</i> : Decapoda	15,008	1,159
4220000	Annelida (segmented worms)	26,547	4,716
4230000	Sipuncula	2,075	2
4260000	Mollusca ( <i>all sub-groups</i> )	19,337	1,896
4262500	<i>Mollusca</i> : Gastropoda (snails & slugs)	17,201	1,008
4265000	<i>Mollusca</i> : Bivalvia (bivalve molluscs)	3,627	593
4266000	<i>Mollusca</i> : Scaphopoda (tusk shell)	85	0
4267500	<i>Mollusca</i> : Cephalopoda	4,123	26
4290000	Bryozoa	3,203	137
4300000	Brachiopoda (lamp shells)	2,012	1
4310000	Phoronida	2,202	0
4320000	Chaetognatha (arrow worms)	26,878	3,400
4330000	Hemichordata	3,965	4
4340000	Echinodermata	6,614	1,040
4350000	Urochordata ( <i>all sub-groups</i> )	19,905	3,579
4352500	<i>Urochordata</i> : Ascidiacea (sea squirts)	870	0
4355000	<i>Urochordata</i> : Thaliacea (salps & doliolids)	11,385	101
4357500	<i>Urochordata</i> : Larvacea / Appendicularia	18,037	1,447
4360000	Cephalochordata / Leptocardia	2,458	17
<b>5000000</b>	<b>ICHTHYOPLANKTON</b>	54,286	217

The geographic distribution of numerical abundance casts of major plankton groups for WOD23 is shown in Figures 16.6 – 16.9.

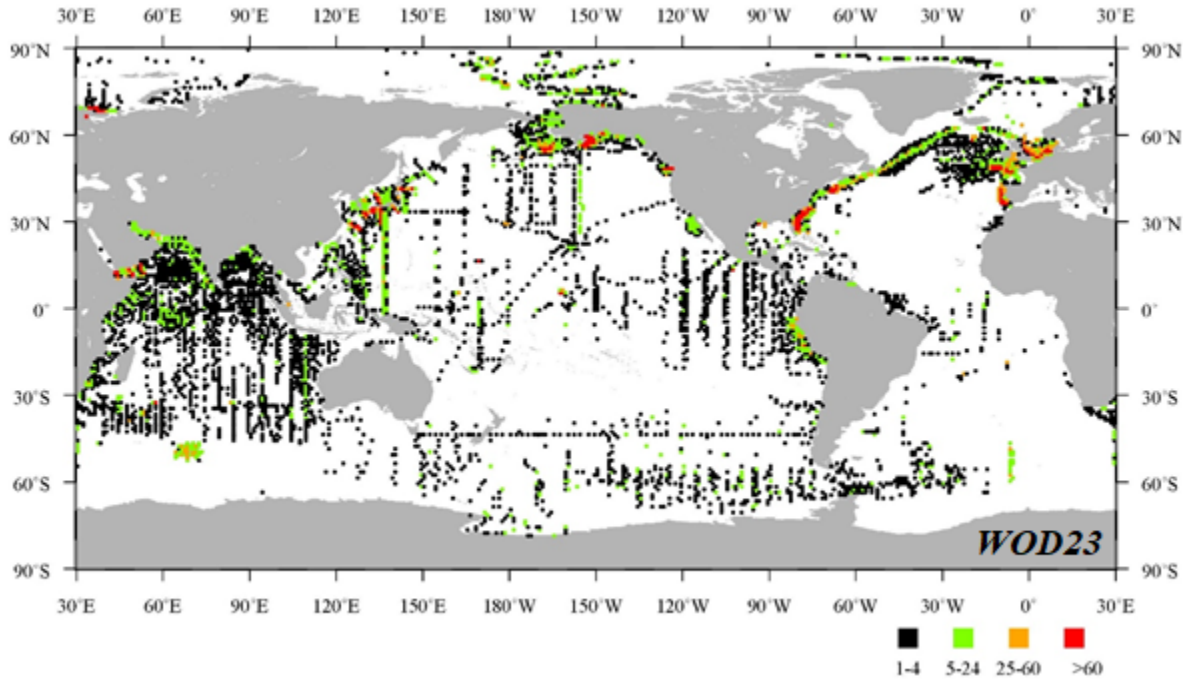


Figure 16.6. Geographic distribution of zooplankton numerical abundance (46,224 casts) in WOD23.

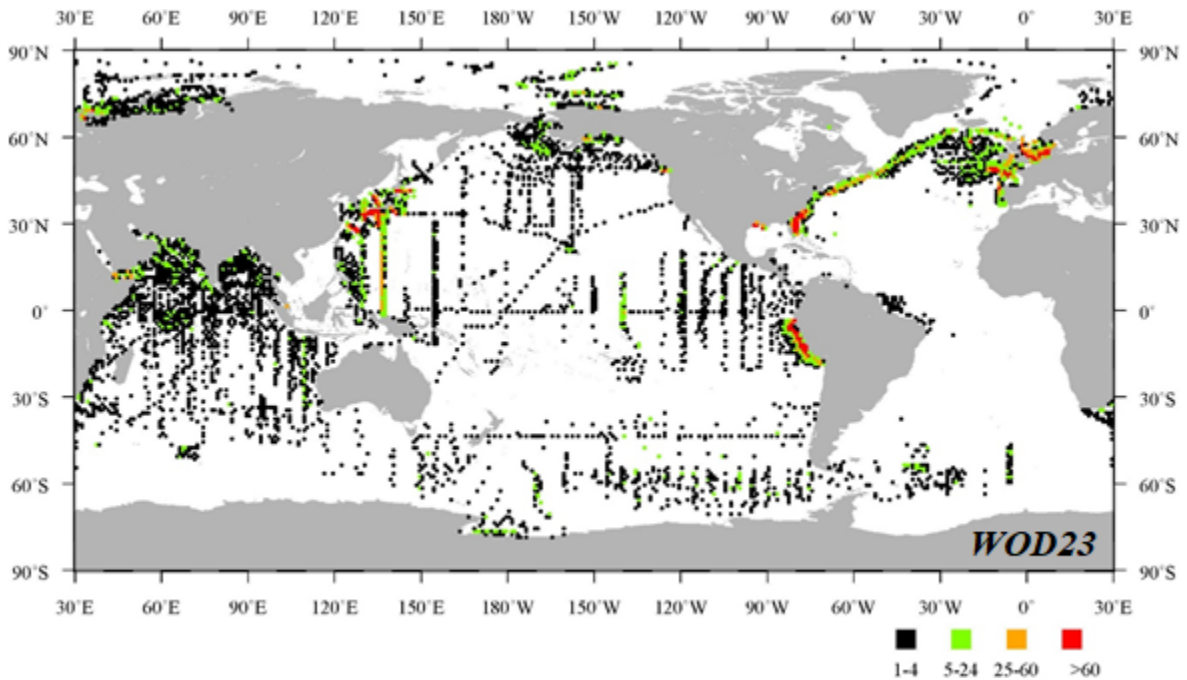


Figure 16.7. Geographic distribution of phytoplankton numerical abundance (37,961 casts) in WOD23.



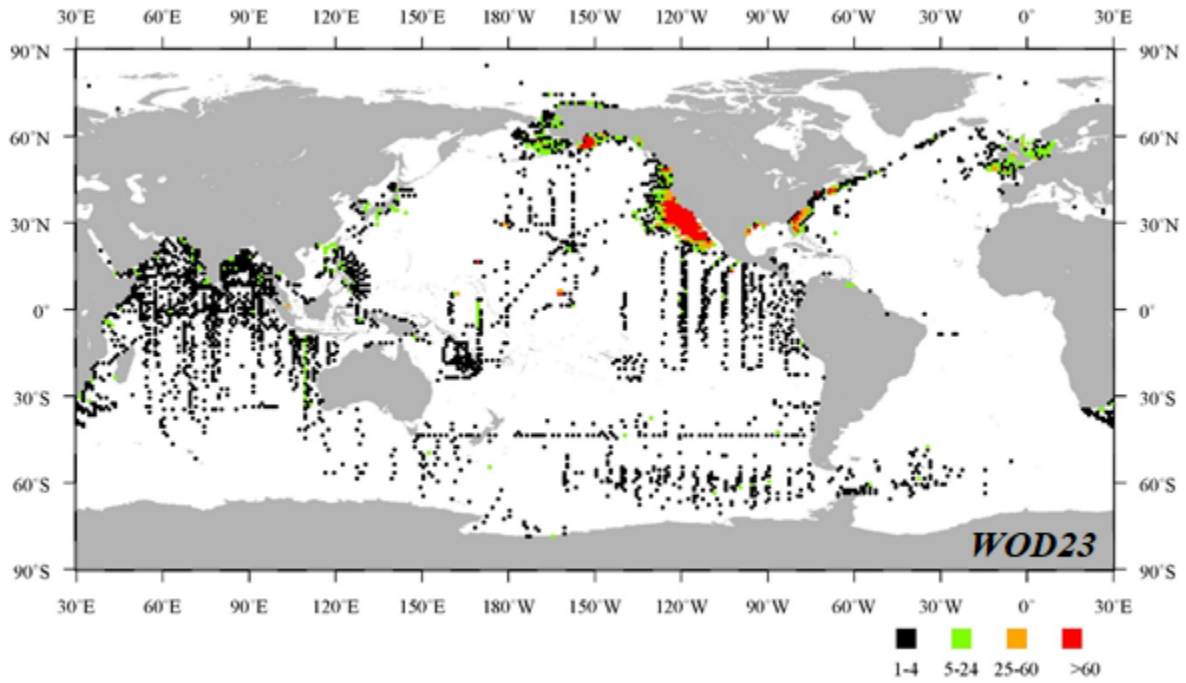


Figure 16.8. Geographic distribution of ichthyoplankton numerical abundance (54,286 casts) in WOD23.

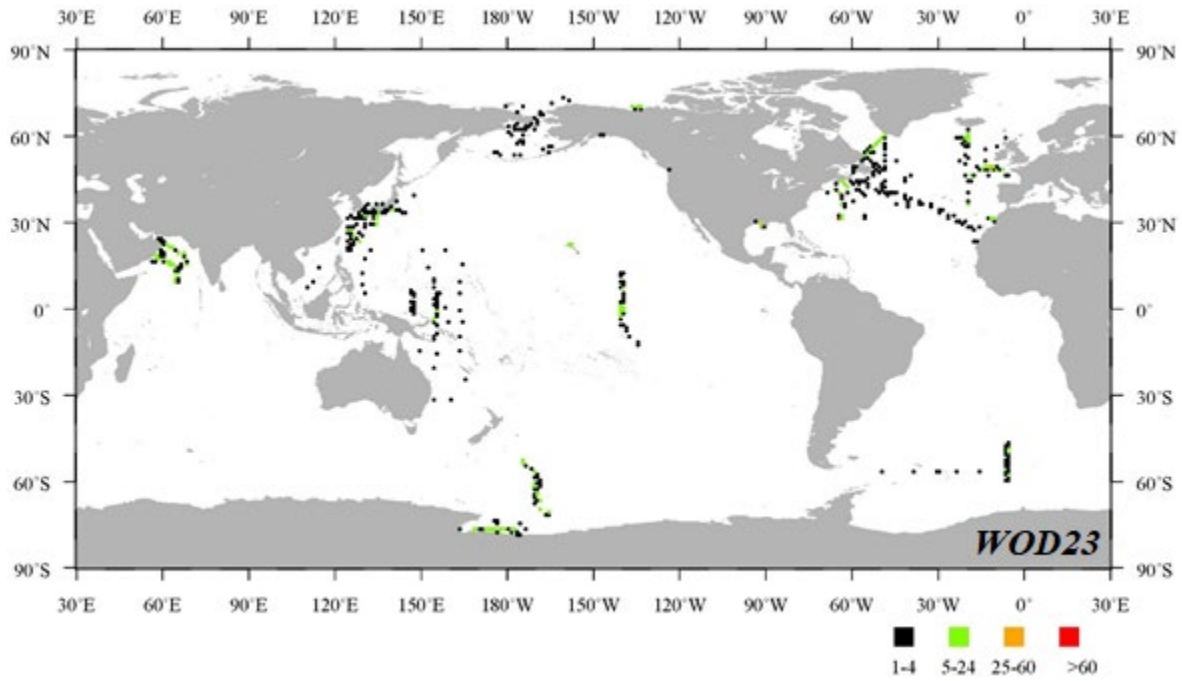


Figure 16.9. Geographic distribution of bacterioplankton numerical abundance (1,986 casts) in WOD23.

## 16.5.2. Total Biomass

The WOD23 total biomass data type represents measurement for which the entire contents of the plankton net are measured as a single, undifferentiated mass. This “mass” can be quantified by measuring the total settled volume, displacement volume, wet mass, dry mass, or ash-free dry mass of the entire sample. Although the sampling methods of total biomass data represented in the WOD23 may differ between projects and institutions, the general definitions and methods per Omori and Ikeda (1984) are:

*Total Settled volume:* the volume of a plankton sample poured into a graduated cylinder or sedimentation tube of 50-100 ml in volume and allowed to settle for 24 hours.

*Total Displacement volume:* the volume of plankton estimated by the volume of water displaced after adding the plankton sample into a graduated cylinder.

*Total Wet Mass:* the mass of plankton determined after eliminating as much surrounding water as possible.

*Total Dry Mass:* the mass of plankton determined after removal of all water and heat dried to a final mass at 60-70°C.

*Total Ash-free Dry Mass:* a known weight of the dry sample ashed to a final weight at 450-500°C.

**Table 16.8. WOD23 biomass measurements content.**

<b>PGC Code</b>	<b>Taxonomic Description</b>	<b># Casts</b>	<b>% of Total</b>
-401	Total Displacement Volume	125,022	72.68
-402	Total Settled Volume	9,926	5.77
-403	Total Wet Mass	34,075	19.80
-404	Total Dry Mass	2,554	1.49
-405	Total Ash-free Dry Mass	446	0.26

The majority of WOD23 plankton biomass measurements are total displacement volume and total wet mass (Table 16.8). Total biomass data were mostly sampled using nets ranging from 200 to 500 µm mesh size, predominantly with standard nets 333 µm mesh size. Samples within this mesh range might include fish eggs, larvae, and small amounts of large phytoplankton, such as diatoms.

Additional information about measurement methods, as well as the protocol followed for removing large organisms, is stored in the Biological Headers described in detail in [WOD23 Manual](#), Table. 6 (Garcia *et al.*, 2024).

The geographic distribution of biomass casts for WOD23 is shown in Figures 16.10. – 16.14.

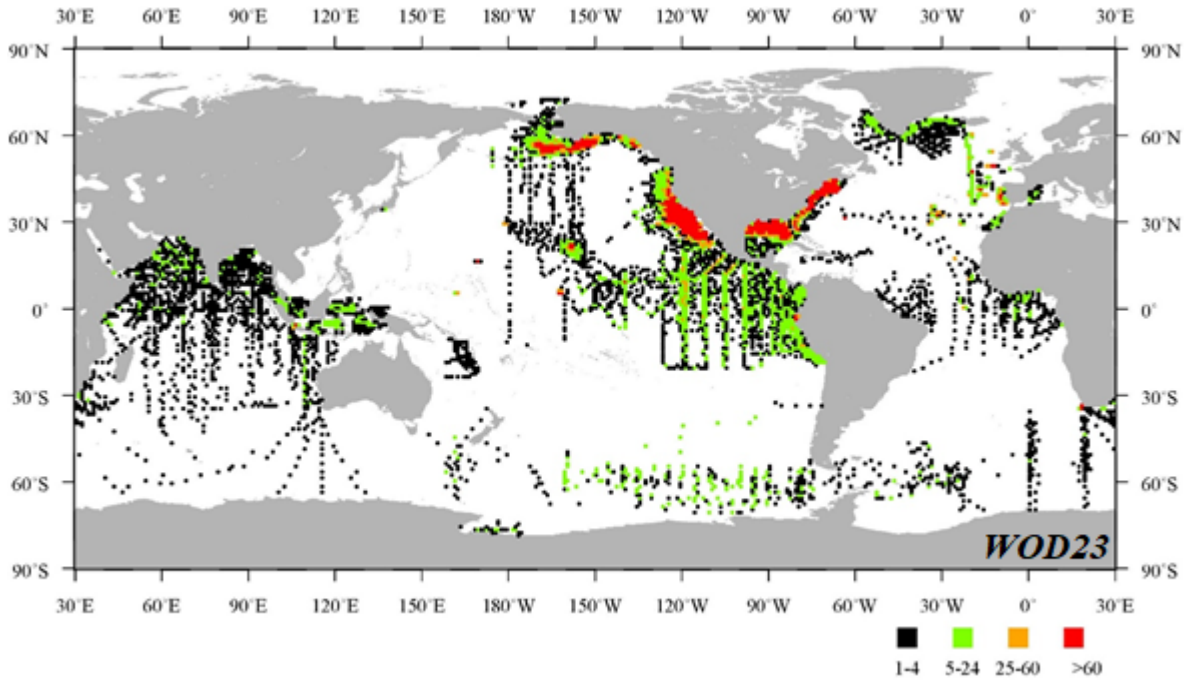


Figure 16.10. Geographic distribution of total displacement volume (125,022 casts) in WOD23.

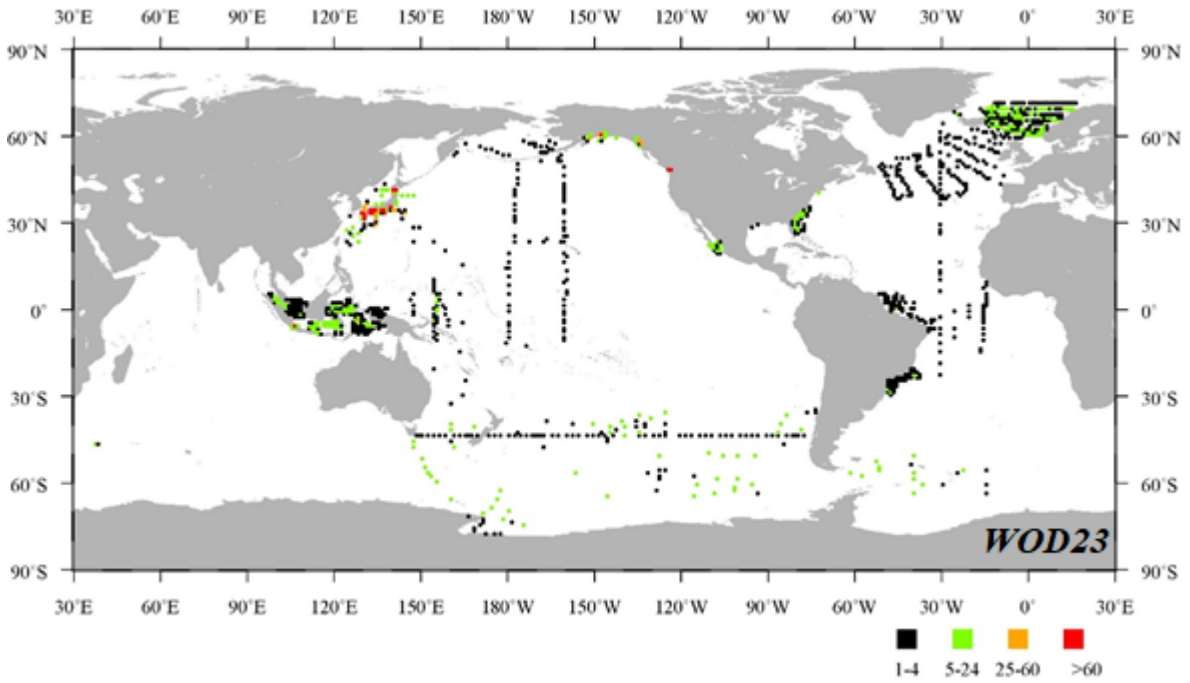


Figure 16.11. Geographic distribution of total settled volume (9,926 casts) in WOD23.

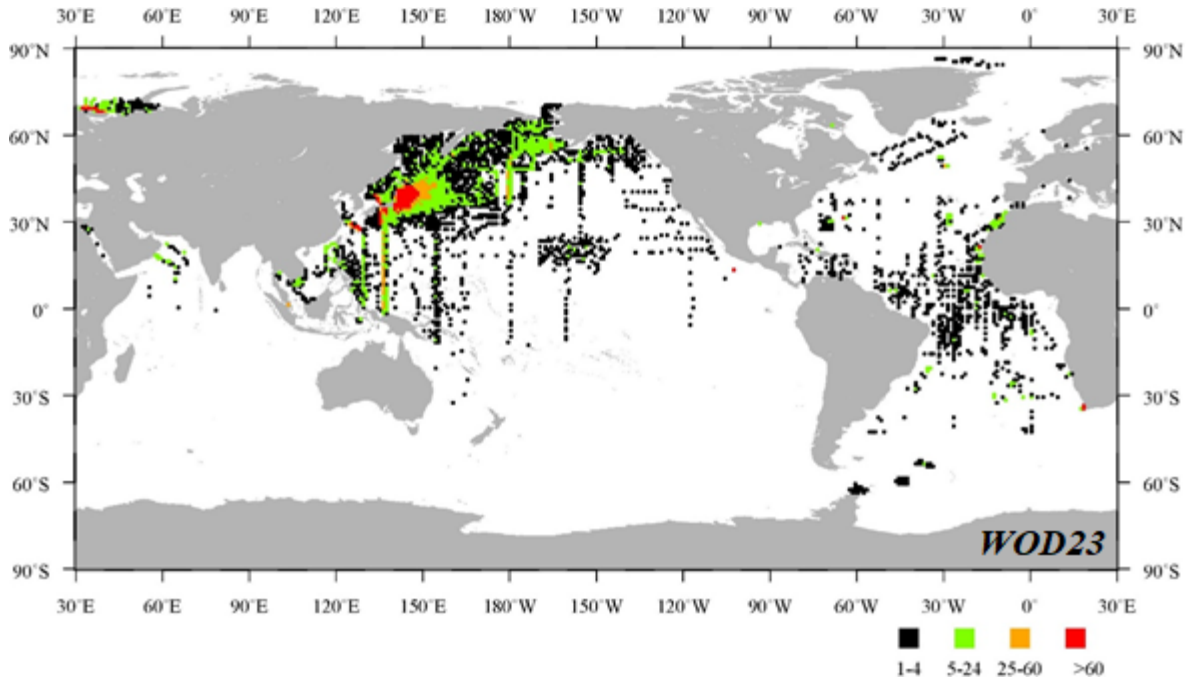


Figure 16.12. Geographic distribution of total wet mass (34,075 casts) in WOD23.

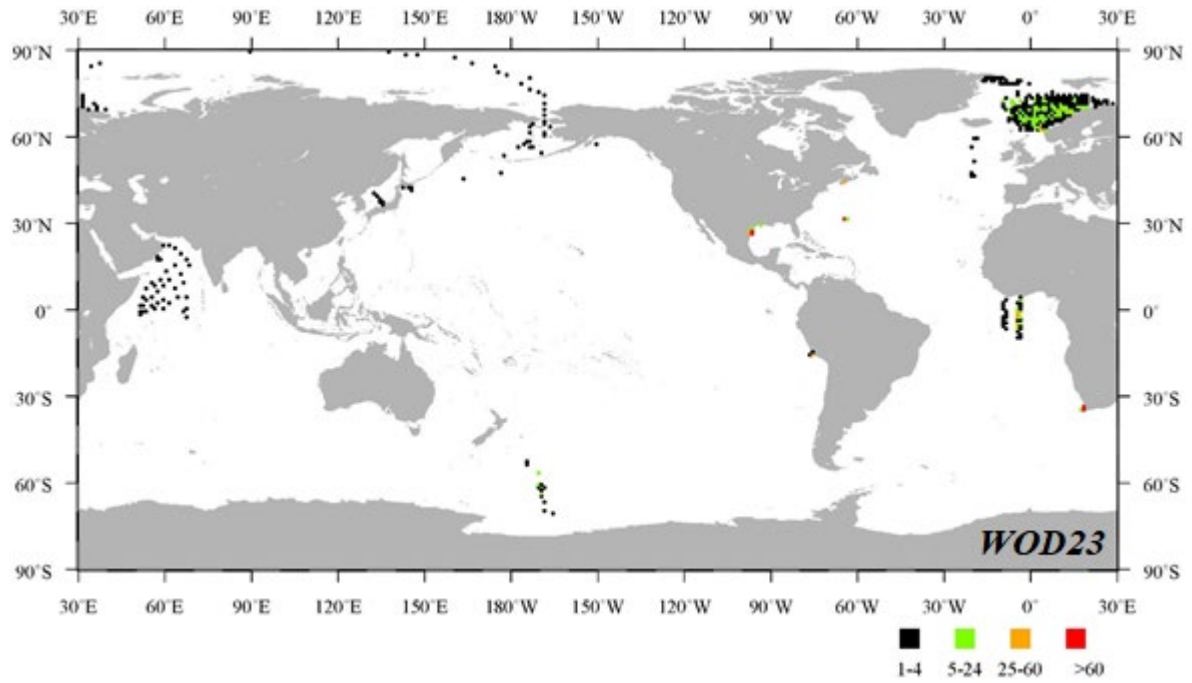


Figure 16.13. Geographic distribution of total dry mass (2,554 casts) in WOD23.

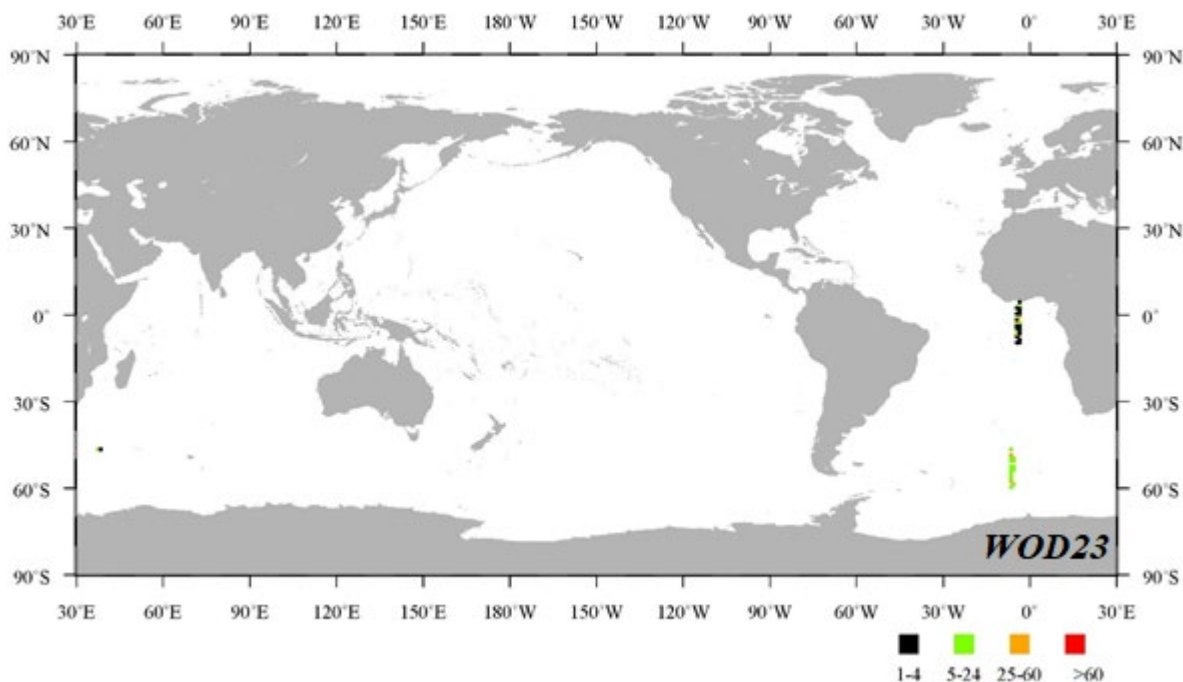


Figure 16.14. Geographic distribution of total ash-free dry mass (446 casts) in WOD23.

## 16.6. REFERENCES AND BIBLIOGRAPHY

- Baranova, O.K., T. O'Brien, T.P. Boyer (2018). Chapter 16. Plankton Data, In World Ocean Database 2018. A. Mishonov, Tech. Ed. *NOAA Atlas NESDIS 87*, pp. 187-206.
- Baranova, O.K., T. O'Brien, T.P. Boyer (2013). Chapter 16. Plankton Data, In World Ocean Database 2013, S. Levitus, Ed., A. Mishonov, Tech. Ed. *NOAA Atlas NESDIS 72*, pp. 189-208.
- Baranova, O.K., T. O'Brien, T.P. Boyer, I.V. Smolyar (2010). Chapter 16. Plankton Data, In World Ocean Database 2009, S. Levitus, Ed. *NOAA Atlas NESDIS 66*, U.S. Gov. Printing Office, Wash., D.C., pp. 192-210.
- Baranova, O.K., J.I. Antonov, T.P. Boyer, D.R. Johnson, H.E. García, R.A. Locarnini, A.V. Mishonov, M.T. Pitcher, I.V. Smolyar (2006). Chapter 14. Plankton Data, In World Ocean Database 2005, S. Levitus, Ed. *NOAA Atlas NESDIS 60*, U.S. Gov. Printing Office, Wash., D.C., pp. 150-169.
- Conkright, M.E., T. O'Brien, L. Stathoplos, C. Stephens, T.P. Boyer, D. Johnson, S. Levitus, R. Gelfeld (1998). World Ocean Database 1998, Volume 8: Temporal Distribution of Station Data Chlorophyll and Plankton Profiles, *NOAA Atlas NESDIS 25*, U.S. Gov. Printing Office, Wash., D.C., 129 pp.
- Harris, R.P., P.H. Wiebe, J. Lenz, H.R. Skjoldal, and M. Huntley (2000). ICES Zooplankton Methodology Manual, *Academic Press*, 684 pp.

- Garcia H.E., T.P. Boyer, R.A. Locarnini, J.R. Reagan, A.V. Mishonov, O.K. Baranova, C.R. Paver, Z. Wang, C.N. Bouchard, S.L. Cross, D. Seidov, and D. Dukhovskoy (2024). World Ocean Database 2023: User's Manual. A. Mishonov, Tech. Ed., *NOAA Atlas NESDIS 98*, pp 134. [https://doi.org/10.25923/j8gq\\_eee82](https://doi.org/10.25923/j8gq_eee82)
- Kennish, M.J. (Ed.) (1990). *Practical Handbook of Marine Science*, CRC Press, Boca Raton, Ann Arbor, Boston, 710 pp.
- Lalli, C.M. and T.R. Parsons (1997). *Biological Oceanography. Introduction*. University of British Columbia, Vancouver, Canada, 314 pp.
- Levington, J.S. (1995). *Marine Biology. Function, Biodiversity, Ecology*. Oxford University Press, New York, Oxford, 420 pp.
- Margulis, L. and K.V. Schwartz (1998). *Five Kingdoms: An Illustrated Guide to the Phyla of Life on Earth*. W.H. Freeman & Company (New York), 520 pp.
- O'Brien, T.D., M.E. Conkright, T.P. Boyer, C. Stephens, J.I. Antonov, R.A. Locarnini, H.E. Garcia (2002). *World Ocean Atlas 2001, Volume 5: Plankton*. S. Levitus, Ed. *NOAA Atlas NESDIS 53*, U.S. Gov. Printing Office, Wash., D.C., 89 pp., CD-ROMs.
- O'Brien, T.D. (2007). COPEPOD: The Global Plankton Database. A review of the 2007 database contents and new quality control methodology. U.S. Dep. Commerce, *NOAA Tech. Memo. NMFS-F/ST-34*, 28 p.
- Omori, M. and T. Ikeda (1984). *Methods in Marine Zooplankton Ecology*, Wiley & Sons, New York, 332 pp.
- Truett, J.C. (Ed.) (1985). *The Norton Basin Environment and Possible Consequences of Planned Offshore Oil and Gas Development*. A final report for the U.S. Department of the Interior, Minerals Management Service Alaska OCS Region, Anchorage, AK and the U.S. Department of Commerce, National Oceanic and Atmospheric Administration, OCS Environmental Assessment Program, Anchorage, AK. *NTIS No. PB86-200946/AS. MMS Report 85-0081*. 123 pp